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**The Evolution of Oceanic Gabbros:  
*In-situ* and Ancient Examples**

by

Astri Jæger Sweetman Kvassnes

June 2004

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Astri Jæger Sweetman Kvassnes

Massachusetts Institute of Technology  
Cambridge, Massachusetts 02139

and

Woods Hole Oceanographic Institution  
Woods Hole, Massachusetts 02543

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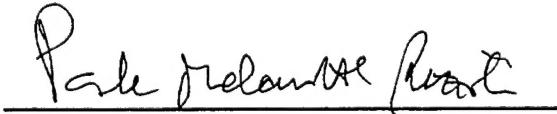
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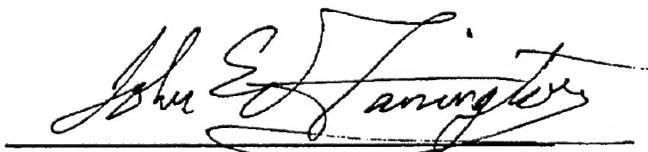
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\_\_\_\_\_  
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Department of Geology and Geophysics

  
\_\_\_\_\_  
Paola Malanotte-Rizzoli

MIT Director of Joint Program

  
\_\_\_\_\_  
John W. Farrington

WHOI Dean of Graduate Studies

**The Evolution of Oceanic Gabbros: *In-situ* and Ancient Examples**

By

Astri Jæger Sweetman Kvassnes

Candidatus Scientiarum, University of Bergen, 1997

Submitted in partial fulfillment of the requirements for the degree of  
Doctor of Philosophy

at the

**MASSACHUSETTS INSTITUTE OF TECHNOLOGY**

and the

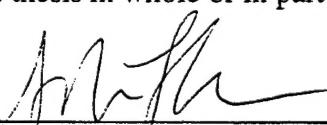
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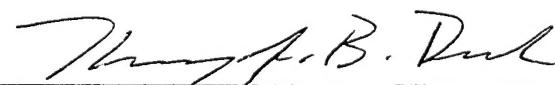
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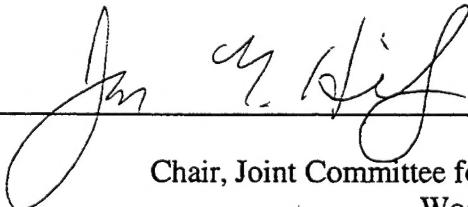
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and Woods Hole Oceanographic Institution  
April 2003

Certified by



Henry J. B. Dick  
Thesis Supervisor

Accepted by



J. Gregory Hirth  
Chair, Joint Committee for Marine Geology and Geophysics  
Woods Hole Oceanographic Institution

# The Evolution of Oceanic Gabbros: *In situ* and Ancient Examples

## The Evolution of Oceanic Gabbros: *In-situ* and Ancient Examples

by

Astri Jæger Sweetman Kvassnes

Submitted to the Department of Earth Atmosphere and Planetary Sciences  
on April 30<sup>th</sup>, 2004 in Partial Fulfillment of the  
Requirements for the degree of Doctor of Philosophy in  
Marine Geology

### Abstract

This study is a geochemical investigation into the accretion of lower oceanic crust and processes of shallow melt-rock reaction at mid-ocean ridges. Major-, trace-elements, and isotopes from whole-rocks and minerals from the Lyngen Gabbro, a 480-My old dismembered ophiolite from the Scandinavian Caledonides, indicate that this igneous complex was produced from hydrous supra-subduction zone magmas, a remnant of an incipient ocean-arc. Such ophiolites are better models for the structural evolution than the geochemical evolution of the lower oceanic crust at mid-ocean ridges.

Minerals in gabbros from Atlantis Bank, Southwest Indian Ridge, a modern, *in-situ* example of lower ocean-crust, were analyzed for major and trace-elements. The MELTS algorithm indicates that these gabbros formed by near-fractional crystallization at mid-crustal pressures. The gabbroic crust is more evolved than the lavas and represents melts fractionated 50-95% relative to a mantle-derived melt-composition, supported by trace-element models. This argues against the often-cited gabbro-glacier accretion model, where mantle-derived melts are transported to a shallow melt-lens and fractionates there before eruption. There remain >770-m of additional primitive cumulates below 1500-m deep Hole 735B or within the underlying mantle. Thus, the seismic Moho, beneath Hole 735B, could be the crust-mantle boundary, rather than an alteration front as suggested elsewhere.

The Atlantis Bank gabbros have augites that are more primitive than plagioclases and olivines with which they coexist. Melt-rock interaction, where ascending melts dissolve the pre-existing gabbroic rocks and create hybrid magma may have caused this. Dissolution-experiments for plagioclase-olivine and plagioclase-augite mineral pairs were performed at 1180°-1330°C and 20-min – 24hrs. Dissolution occurs rapidly and out of equilibrium, with the dissolution rates dependent on the  $\Delta T$  above the solidus. Rocks with small grain-boundary areas (coarse grained or nearly mono-mineralic) heat internally when enclosed in hot magma, causing xenoliths or wall-rock to melt and disaggregate. The dissolution-derived magma crystallizes minerals more refractory-looking than the melts that precipitated the original gabbroic rocks. Assimilation of gabbroic rocks increases the Na content and decreases the Fe content of the melt that digests it, thus basaltic glasses formed after this hybridization will falsely reflect a lower degree and pressure of mantle melting.

Thesis Supervisor: Henry JB Dick

Title: Senior Scientist at Woods Hole Oceanographic Institution

The Evolution of Oceanic Gabbros: *In situ* and Ancient Examples

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## The Evolution of Oceanic Gabbros: *In situ* and Ancient Examples

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## Dedication

I dedicate this thesis to the “Sweetnes”-Clan.

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## Chapter 1

### Introduction

The oceanic lower crust is dominated by olivine gabbro, generally composed of olivine, plagioclase and augite (Coleman (1977)). It is generally believed that this lower crust is formed by crystallization of basaltic magmas that ascend from the upper mantle. The lower crust is unfortunately difficult to study at fast-spreading ridges and ocean islands due to the 2+ km of dikes and lavas that cover it and the lack of tectonic windows. Therefore, many studies have used ophiolites as analogues for the origin and nature of the lower ocean crust. Many authors inferred that most ophiolites are the remnants of mid-oceanic crust (Nicolas, 1989, Elthon et al, 1992). However, it became apparent that the exact tectonomagmatic setting of any one ophiolite is uncertain, and it has been suggested that many ophiolites formed in *supra-subduction* settings (e.g. Miyashiro, 1975).

The extensive knowledge that rose from ophiolite studies, however, give us valuable structural, petrologic and geochemical concepts to understand and classify the origin of ophiolites and to understand crustal accretion at ocean ridges and how that might differ between tectonic settings such as mid-ocean ridges, ocean arcs and back-arc spreading centers. For example, it has been established that increasing activities of water in magma increase the An-content of plagioclase at a given Ca/Na ratio of the melt (Hough and Luhr, 1991, Arculus and Wills, 1990). In addition, the activity of augite increases with respect to olivine and plagioclase under wet conditions (Gaetani *et al.*, 1993), causing high-Ca pyroxene to crystallize before plagioclase. High-pressure (>5-kbar) have the opposite effect on the anorthite-content of plagioclase relative to melt composition (Panjasawatwong *et al.*, 1995), resulting in decreased An% content of plagioclase relative to lower pressure conditions. Variations in water content of magmas and pressure of crystallization are likely important across the accretionary spectrum from mature ocean ridges to ocean arc environments.

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There are two main lines of thought on the geodynamic method of accretion of the lower ocean crust (Figure 1-1 A and B). One is the so-called “gabbro glacier” model, the other is the “many sills” model. The former end-member is based on observations in ophiolites (Sleep, 1975; Dewey and Kidd, 1977 and Quick and Denlinger, 1993) and developed further to account for geophysical observations at the East Pacific Rise (EPR) and other magmatically robust spreading segments (Henstock et al., 1993; Phipps Morgan and Chen, 1993). The gabbro-glacier model assumes that melt arrives directly from the upper mantle to a shallow melt sill, where the magmas fractionate and the cumulates subsequently flows downwards in a near steady-state fashion. Thus, the lower crust at such a ridge should always be more primitive than the lavas from the upper crust. The many sills model (Nisbet and Fowler, 1978; Pedersen, 1986; Bédard et al, 1988, 1991, 1993; Boudier et al., 1996; Kelemen et al., 1997b; Korenaga and Kelemen, 1997) indicates that the lower crust is accreted by the injections by individual sills that fractionate *in-situ* and eject their melt to higher levels. Thus, the lowermost crust is largely accreted there, and the composition of the gabbros becomes gradually more evolved upwards.

The truth probably lies between the gabbro glacier and the many sills model (MacLennan et al., 2004), and many models have been proposed that combine the two. Nevertheless, the extent to which the lower ocean crust interacts with subsequent ascending magmas is largely unknown. The rate of assimilation is naturally dependent on the heat available to incorporate and melt pre-existing cumulates into magma, and the residence time of the magma in the lower crust and other variables including ambient lithospheric temperature. The effect of the composition and grain-size of the assimilant, thus the compositional effects on the magma and lower crust, has not yet been established for the oceans.

The potential temperature of the melts generated in the upper mantle at mid ocean ridges could range from 1180°-1510°C (Klein and Langmuir, 1987, Kinzler and Grove, 1992), although other authors would disagree, arguing that the temperature is constant around 1250°C (e.g. Green et al, 2000; Presnall, 2002). The upper end of the temperature

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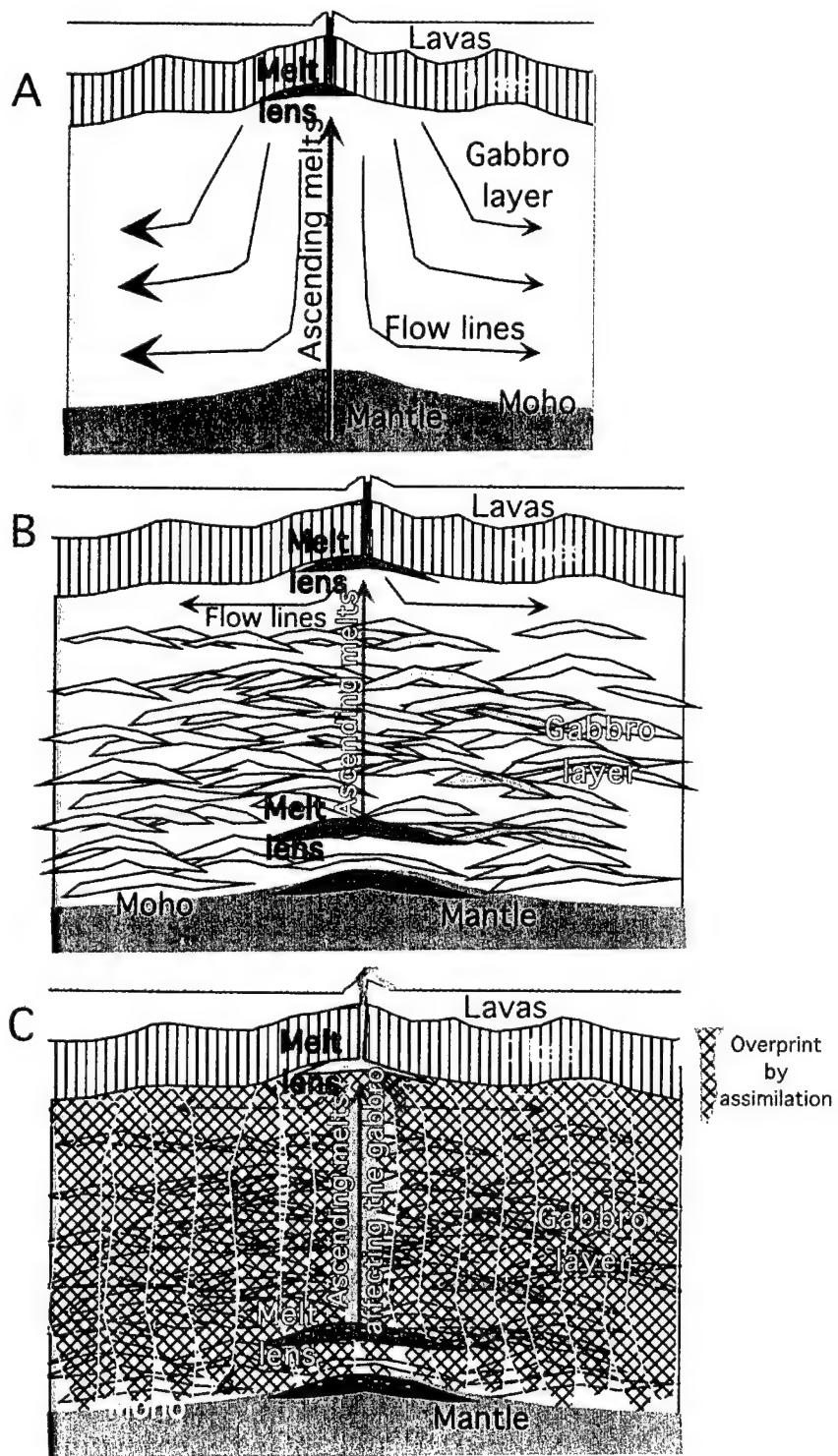


Figure 1-1: A: The Gabbro-glacier model. B: The many sills model. C: The model from this study, where the many-sills model is overprinted by melt-rock reaction, and xenoliths are being incorporated into the ascending magma.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

spectrum represents the fast spreading ridges where the EPR is the fastest and hottest. It has been documented that the EPR has a shallow axial magma chamber (AMC) (Detrick et al., 1987) that lies at low velocity zone (Harding et al., 1989; Vera et al., 1990). Dunn et al. (2000) showed the 5-8 km wide low velocity zone under the spreading center of the EPR has as much as 20 % melt fractions within it. A second melt lens has been documented near the seismic Moho transition (Crawford and Webb, 2002), and it has been argued that its stability is limited, as any crystallization in the chamber causes the surroundings to heat up and melt the surroundings in a manner that is not observed (Chen, 2001). The EPR therefore clearly has a heat-budget that is big enough to remelt the lower crustal material. At intermediate-spreading ridges, the AMC reflector is found at deeper levels in the crust (Phipps Morgan and Chen, 1993; Chen and Lin, in press). The existence of shallow steady-state magma-chambers at slower ridges is debated (Sinton and Detrick, 1992), and very deep magma chambers below the conductive boundary-layer in the mantle (Mével et al., 2002) would not be possible to detect with the seismic techniques available today (Detrick and Reeves-Sohn, pers comm. 2004). Erupted MORB show extensive differentiation and mixing that require that magma chambers exist at slow-spreading ridges, even if only ephemerally. The evidence of magma-chamber type processes happening before eruption at all ridges, suggest that there should be sites that are hot enough to dissolve preexisting crustal material even if melt bodies usually evade direct detection along the slow-spreading ocean ridges. Steady-state and ephemeral magma-chambers are likely to reside in the crust long enough for considerable interaction to occur between melts, cumulates and wallrocks.

Ocean islands are also sites of elevated magma-temperatures (Watson and McKenzie, 1991; Clague et al., 1991). In fact, there are Hawaiian picrites that record crystallization temperatures in excess of 1400°C (Natland, pers comm. 2004). Ocean - islands develop post-shield volcanism of very different compositions from the earlier shield-stage lavas. Their isotopes are more like those of mid-ocean crust, and they are frequently very alkaline in composition (Yang et al., 2003). The post-shield lavas often contain xenoliths made out of the pre-existing, tholeiitic, lower-crust that the hot-spot

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

basalts had to penetrate to get to the surface (Fodor and Galar, 1997; Neumann et al., 2000). This suggests that the post-shield stage magmas, which spend more time in the lower crust than the shield-stage magmas, actually may have digested some lower crustal material on ascent.

The purpose of this study is to constrain the processes by which the lower crust at ocean ridges accretes, and to identify how the melts and cumulates interact there. The six following chapters of this thesis progressively approach these subjects.

In Chapter 2, we used geochemical techniques to investigate the tectonomagmatic origin of the Ordovician Lyngen Magmatic Complex, Northern Norway. We present 419 major and trace element analyses together with selected major- and trace element mineral analyses and  $^{143}\text{Nd}/^{144}\text{Nd}$ -isotope whole-rock analyses of gabbroic to tonalitic plutonic rocks from the Ordovician Lyngen Magmatic Complex in Troms, Northern Norway. This dismembered ophiolite represents concurrent boninitic and tholeiitic magmatism in a fore-arc setting, possibly in an outer arc high. The gabbro-compositions are very different from those in silicic continental arcs. The cumulative rocks have extremely low levels of incompatible trace elements and high levels of compatible elements, reflecting melts derived from high degrees of melting in the mantle. The mineral compositions of the gabbros indicate hydrous conditions during fractionation and the transition between the boninitic and the tholeiitic cumulates is diffuse without tectonic contacts, suggesting that magmatism happened concurrently.

In Chapter 3, we utilize the large database of major-element mineral-compositions compiled for Hole 735B, at Atlantis Bank, Southwest Indian Ridge (Dick et al., 2002). We develop an extensive, quantitative and qualitative fractionation-model for the crustal section. We use a special version of the Melts program (Melts5) (Ghiorso and Sack, 1995) and in addition BASALT (1atm) (Weaver and Langmuir, 1990) and Yang et al. (1996) to evaluate the effect of melt-composition, oxygen fugacity, water, pressure, and crystallization processes that produced the lower crustal rocks at Atlantis Bank. This allows the extent of disequilibrium in the lower crustal material at Atlantis Bank to be

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

quantified. We find that the best fractionation-model is fractional crystallization of a near dry melt with  $\text{Na}_8 \sim 3$  and  $\text{Fe}_8 \sim 9$  at crustal pressure on the quartz-fayalite-magnetite oxygen fugacity buffer. The gabbros sampled at Atlantis Bank generally represent melts that have crystallized 50-95% relative to a primary, mantle-derived magma. The basaltic glasses sampled on and around Atlantis Bank represent a melt that crystallized 30-50%. Therefore, the gabbro-glacier model is not valid for this slow-spreading ridge. Our model predicts that approximately 770 vertical meters of the lower crustal section has not been sampled and that the total crustal thickness at Atlantis Bank was more than 4.4-km. Thus, the observed seismic Moho there may indeed be the crust-mantle boundary, and not an alteration front as previously suggested by Muller et al. (1997). However, the majority of the gabbros in the lower 1000 meters of the hole have plagioclase coexisting with augite that have Mg#'s up to 10% higher than required by fractional crystallization of likely parental magmas in the area. Thus, we establish that interstitial melts in a mush-zone have dissolved and reprecipitated pre-existing crystals, causing the melt to become more Mg and Ca rich as plagioclase and augite dissolve rather than re-equilibrate with the melt. In the many-sills model, the gabbros become progressively more evolved upwards while accreting by fractional crystallization. Since we see that the solid-lines of descent of the gabbros indicate assimilation-processes, the many-sills model is only valid if pervasive overprints by ascending melts are incorporated (Figure 1-1C).

In Chapter 4, we demonstrate the kinetics of melting of natural mineral-pairs representative of the minerals found in lower ocean-crustal rocks, and find rapid melting rates and unusual melt-compositions. Therefore, the melts produced in the experiments allow us to assess the chemical changes that incorporating lower crustal material may cause in the ascending magmas. Our experiments were carried out using olivine - plagioclase and clinopyroxene - plagioclase mineral pairs over the temperature range of 1330 - 1220 °C at the quartz-fayalite-magnetite buffer for durations of 0.25 to 24 hours. Isothermal experiments performed over the temperature range of 1230 – 1105 °C defined the equilibrium phase relations in the system. The mechanisms of dissolution are similar

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to those found in other experimental studies that utilized end-member mineral compositions. The experimental results were used to determine to what extent cumulate igneous rocks from the oceanic lower crust would produce disequilibrium melt-compositions. We find that the efficiency of the reaction is dependent on grain-size and mineral-distribution of a rock that is reheated. Coarse-grained rocks (>3mm) will disintegrate internally upon reheating, whereas finer grained rocks will melt from the outside in. Mixing between the dissolution-melts and the surrounding magma is aided by the disaggregation of the xenocrysts of different densities, and the viscosities of the interstitial melts are similar to the ascending magma. The new, hybrid magma will have higher Na<sub>2</sub>O and lower Fe<sub>2</sub>O than the original. In addition, lower crustal assimilation of gabbro will result in the destruction of troctolites in the lower crust, and will lead to the early onset of high-Mg# olivine-gabbros.

In Chapter 5, we have analyzed the major element compositions of the magmatic minerals of 99 gabbros from on and around Atlantis Bank, covering an area of almost 660-km<sup>2</sup>. Tectonic windows provide the necessary exposures of the lower crust at this slow-spreading ridge (Tucholke and Lin, 1998). Samples from the northern area have the most primitive magmatic minerals, extending to ferro-gabbros with evolved mineral-compositions. Gabbros from the Western Wall and the top of the bank have intermediate compositions, whereas the southern rocks are all evolved, including layered olivine-gabbros sampled at gabbro-peridotite contacts. Despite the large variations in mineral chemistry in our samples, the primitive and evolved ends of the geochemical spectrum of ODP Hole 735B have not been found in our data. Although we have sampled the area extensively, we find that, like the gabbros from Hole 735B, the surface samples represent melts that have crystallized 50-90% relative to mantle-derived compositions. Therefore, if the lower crust was accreted by fractional crystallization, large amounts of troctolite are expected to exist somewhere in the lowermost crust or intruded into the residual mantle at depths below what is exposed on the sea-floor. However, the co-existing mineral-compositions vary in such a way as to suggest extensive reaction, dissolution and reprecipitation within the cumulate pile. A simple fractional-crystallization model, like

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that above, does not fully describe the evolution of Atlantis Bank gabbros. We propose an alternative model where primitive cumulates at higher levels may have been re-homogenized and transformed by later, interstitial melts.

In Chapter 6, we use the trace-element compositions of clinopyroxene from gabbros on and around Atlantis Bank to re-evaluate the geochemical evolution model of the gabbros and incorporated assimilation-fractionation-crystallization models (AFC) into our crustal thickness estimates. The gabbros cannot have formed by fractional crystallization alone, as the LREE/HREE ratios of the augites in the gabbros increase significantly with fractionation. We find that even though the parental melts experienced assimilation and/or *in-situ* fractionation (DePaolo, 1981; Langmuir, 1989), overall the gabbro-suite reflects melts that crystallized >50-95% relative to a parental magma in equilibrium with the mantle. This agrees well with our previous results using major-elements in Chapter 3. We therefore conclude that the estimate for the crustal thickness at Atlantis Bank of at least 4.4-km is valid, and that the melt that formed the lower ocean crust at Atlantis Bank experienced extensive melt-rock reaction during ascent. In addition, we found that the augites from a single diabase dikes indicate depleted melt compositions different from the compositions in the augites. The dike may be the result of late stage melting of the already depleted mantle.

In Chapter 7, we developed a method for calculation of trapped melt-fractions in gabbros from the Southwest Indian Ridge from whole-rock analyses. The model is calibrated with the use of *in-situ* mineral analyses mass-balanced with whole-rock compositions. We find that the trapped melt-fraction of ODP Hole 735B is ~15% on average. Eu-anomalies are inversely related to the trapped melt-fraction. This suggests that the whole-rock positive Eu-anomalies in Hole 735B, the signature of cumulate plagioclase, are masked by other REE's, as seen in diabases.

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## Chapter 2

# The Lyngen Gabbro: The lower crust of an Ordovician Fore-Arc

Astri J. S. Kvassnes<sup>1</sup>, Anita Hetland Strand<sup>2</sup>, Heidi Moen-Eikeland<sup>3</sup>, Rolf B Pedersen<sup>4</sup>.

1: Woods Hole Oceanographic Institution, MS#8, Woods Hole, 02543 MA, USA,

[astri@mit.edu](mailto:astri@mit.edu)

2: Statoil, 4035 Stavanger, Norway

3: Nipefjellet 83, 5300 Kleppestø, Norway

4: Department of Earth Science, University of Bergen, Allegt 41, 5007 Bergen, Norway

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### ***Abstract:***

We present 419 major and trace element analyses together with selected major- and trace element mineral analyses and  $^{143}\text{Nd}/^{144}\text{Nd}$ -isotope whole-rock analyses of gabbroic to tonalitic plutonic rocks from the Ordovician Lyngen Magmatic Complex in Troms, Northern Norway. We find that this dismembered ophiolite represents concurrent boninitic and tholeiitic magmatism in a fore-arc setting, possibly in an outer arc high. Our geochemical data demonstrates the distribution of lower-crustal material present in such settings. The compositions are very different from silicic continental arcs. The cumulative rocks have extremely low levels of incompatible trace elements and high levels of compatible elements. Therefore, the magmas were likely the result of high degrees of melting in the mantle. The mineral compositions of the gabbros indicate hydrous conditions during fractionation. In at least one location the transition between the boninitic and the tholeiitic cumulates is diffuse without tectonic contacts, suggesting that magmatism happened concurrently.

### ***Introduction***

The lower ocean crust is generally inaccessible, and therefore its origin and nature have often been studied in ophiolites. The exact tectonomagmatic setting of any one ophiolite is, however, uncertain. Many authors have inferred that most ophiolites are the remnants of mid-oceanic crust (Nicolas, 1989, Elthon et al, 1994). Other authors have suggested that many ophiolites formed in *supra-subduction* settings (e.g. Miyashiro, 1975). The extensive knowledge that rose from this controversy gives us valuable petrologic and geochemical tools to understand and classify the origin of ophiolites. For example, it has been established that increasing activities of water in magma increase the An-content of plagioclase at a given Ca/Na ratio of the melt (Hough and Luhr, 1991, Arculus and Wills, 1990). In addition, the activity of augite increases with respect to olivine and plagioclase (Gaetani et al., 1993), causing high-Ca pyroxene to crystallize before plagioclase. High-pressure (>5-kbar) crystallization also results in this reaction series, but pressure has the opposite effect on the anorthite-content of plagioclase relative to the melt composition

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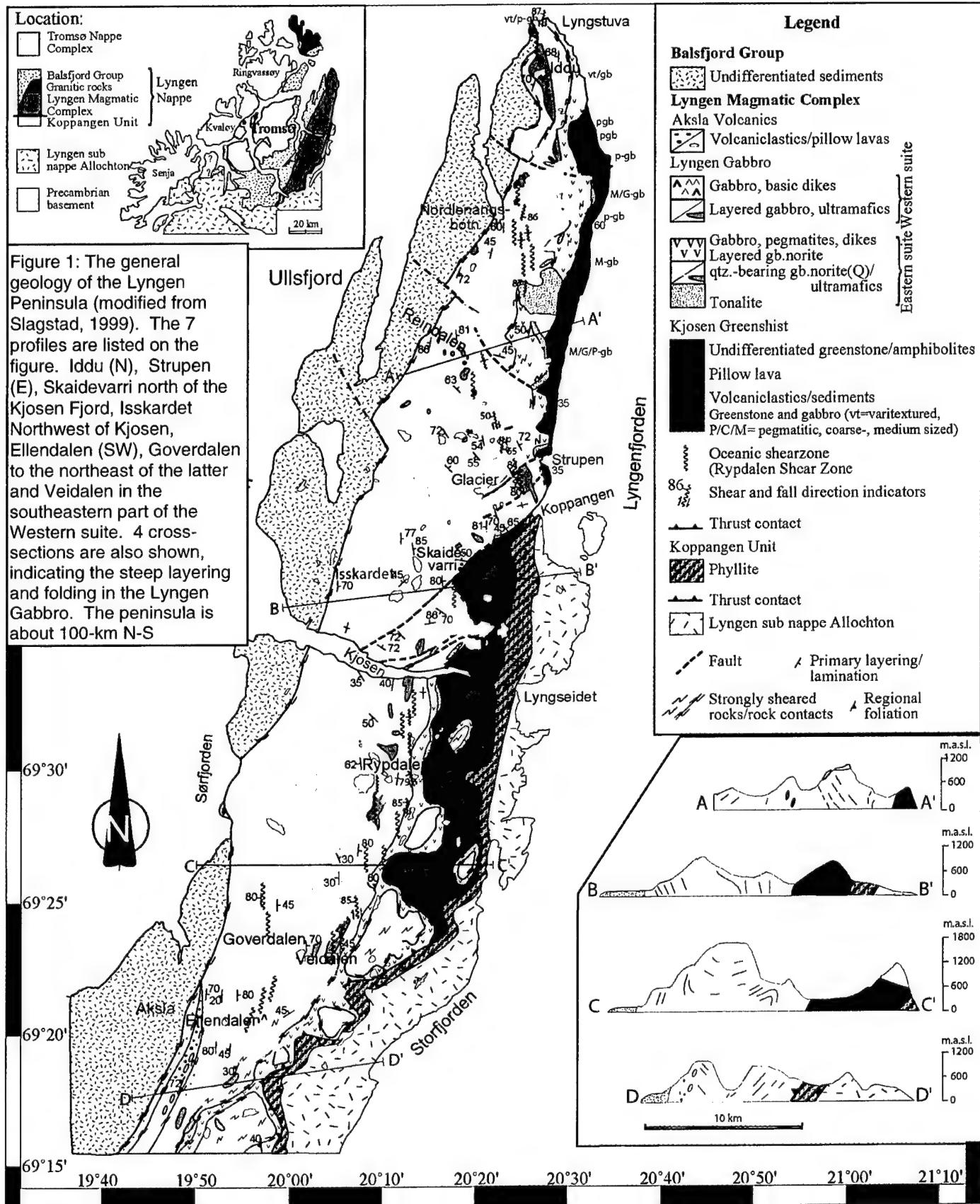
(Panjasawatwong *et al.*, 1995), resulting in decreased An% content of plagioclase relative to lower pressure conditions.

Hydrous magmatic conditions are common in arcs. Such a *supra-subduction* setting implies a tectonomagmatic environment ranging from back-arc basins, with magma sources similar to those from mid-ocean ridges re-melted with the aid of plate-derived water, through arcs and into the extremely depleted fore-arcs, often comprising boninitic magmatism. In addition, the outer-arc high represents the incipient arc made up of interspersed boninitic and tholeiitic magmatism. For instance, the Lau-Basin in the Southeast Pacific comprises a boninitic outer arc high and fore-arc, through a small arc and an actively spreading back-arc basin (Crawford *et al.*, 1981). Isotope-geochemistry and trace element geochemistry is used to support the theory that material derived from the slab is introduced into the subduction-zone mantle wedge that subsequently melt, and in addition, assimilation of existing mantle-rocks and crustal material may be assimilated into the ascending magma (e.g. AFC, DePaolo, 1981.) Consequently, dismembered magmatic complexes of oceanic origin can be put into a tectonic context using geochemistry. In this study, we used geochemical techniques to investigate the tectonomagmatic origin of the Ordovician Lyngen Magmatic Complex. We show that the Lyngen Gabbro represents the transition from MAR-like hydrous magmatism to a hydrous magma of a much more depleted source, like those in incipient arcs.

### The Lyngen Magmatic Complex

The Lyngen Magmatic Complex dominates the Lyngen Peninsula and contains the largest massif of gabbroic rocks known in the Scandinavian Caledonides (Fig 2-1). It is postulated to be a dismembered ophiolite complex (Minsaas, 1981), as no mantle section or sheeted-dike complex has been identified. The Lyngen Magmatic Complex comprises the Lyngen Gabbro, the Aksla Volcanics, and the Kjosen Greenschist (Furnes and Pedersen, 1995). Its minimum age is Llanvirnian-Arenigian (469 +/- 5Ma) (Oliver and Krogh, 1995), as determined from a tonalite intrusion in the southern part of the Kjosen Greenschist. The complex has undergone Caledonian deformation, and large low-angle thrust faults dipping to the west intersect the complex. Hence, there may be some repetition of units due to tectonic events.

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The Lyngen Gabbro is, by far, the dominant component of the Lyngen Magmatic Complex. A gravity profile has shown that the Lyngen Gabbro is wedge-shaped with the maximum thickness in the west (Chrostom, 1972). The Lyngen Magmatic Complex has been subdivided into a Western and Eastern suite, both trending NNE-SSW, based on different petrological and geochemical characteristics (Furnes and Pedersen, 1995). The Lyngen Gabbro contains large shear zones originally named the Rypdal Shear Zone. It separates the Western and Eastern suites in places and is accompanied by numerus variably sized dunite and wehrlite bodies. Later dikes with a wide range of compositions intersect the Lyngen Gabbro and Rypdal Shear Zone, leading Slagstad (1995) to interpret the shear zone to be of oceanic origin. The ultramafic rocks and the crosscutting dikes are not the subject of this geochemical study, and will only be described as a part of the general geology. Most of the rocks of the Lyngen Gabbro have been extensively uralitized and saussuritized (Randall, 1959).

The Aksla Volcanics trend NNE-SSW along a narrow zone ~15 km long and 400 meters wide, in the Southwestern part of the Lyngen Peninsula, near Lakselvbukta (Fig. 1). The rocks are mainly deformed, greenschist-facies pillow-lavas, hyaloclastite breccias and dikes (Furnes and Pedersen, 1995) of mid-ocean ridge (MORB) to Island Arc Tholeiite (IAT) affinities (Furnes and Pedersen, 1995). An imbricated slab of gabbro, up to 600 meters wide, has a sheared, subvertical contact to the overlying sedimentary rocks in the Balsfjord Group (Minsaas and Sturt, 1985, Kvassnes, 1997). The gabbro is varitextured and cut by numerous basic dikes and these dikes are henceforth assumed to be associated with the Aksla Volcanics. The Eastern, sub-vertical contact between the Aksla Volcanics and the layered and high-level gabbros of the Lyngen Gabbro trend N-S, and is strongly sheared.

The Kjosen Greenschist outcrops from Strupen southward, defining the easternmost part of the Lyngen Magmatic Complex. It comprises pillow lavas, volcaniclastics and undifferentiated greenschists of MORB to IAT affinities (Furnes and Pedersen, 1995). The rocks of the Kjosen Greenschist are strongly sheared and deformed.

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### Geological setting

The Norwegian Caledonides represent a stack of thin nappes thrust eastwards onto the Baltoscandian platform in late Silurian/early Devonian time (Dallmeyer and Andresen, 1992). The four major units are the Lower-, Middle-, Upper- and Uppermost Allochthon (Roberts and Gee, 1985). The Lower and Middle Allochthons and the lower part of the Upper Allochthon (Seve nappes) have Baltoscandian affinities. The remainder of the Upper- and especially the Uppermost Allochthon are exotic terrains thought to have been derived from spreading-ridges, rifted island arcs and marginal basins lying within the now consumed Iapetus ocean (Andresen and Stelthenpohl, 1994). These exotic magmatic complexes can be divided into two groups by U/Pb dating. Group 1 represents ophiolite complexes of Lower Ordovician (Tremadocian-Arenigian) age and comprises basaltic rocks with MOR-, IAT-, boninitic-, calc-alkaline- to alkaline affinities. Group 2 has an Upper-Ordovician (Ashgillian) age, defined by the Solund/Stavfjord Ophiolite Complex, and is predominantly N-MORB, but also contains local occurrences of basalts with E-MORB affinity (Dunning and Pedersen, 1988). The geochemical evolution of the ophiolites from the Norwegian Caledonides indicates their association with the development of island arcs (Pedersen and Furnes, 1991).

The central Troms area, northern Norway, has seven nappes. The lowermost three are the Divald Group (autochthon) overlain by the Målselv/ Kalak Nappe Complex, all interpreted as the Baltoscandian Margin. Above these nappes are the Vaddas Nappe, the Kåfjord Nappe, the Nordmannvik Nappe, the Lyngen Nappe Complex and the Tromsø Nappe Complex, are all considered to be exotic terrains relative to Baltica (Andresen et al., 1985).

The Lyngen Nappe, a part of the Upper Allochthon has been correlated with Gratangseidet mafic complex in the Ofoten area (S-W) and the Ullsfjord nappe complex (N-E) (Andresen and Bergh, 1985). Three major lithological units can be distinguished; from the base, upwards: The Koppangen Formation, the Lyngen Magmatic Complex and the Balsfjord Group (Fig 1, insert) (Randall, 1959 and 1971(a,b), Munday, 1970 and 1974 and Binns, 1978). The Koppangen Formation is dominated by phyllites, although a strongly foliated amphibolite zone marks its contact to the Lyngen Magmatic Complex. The upper Ordovician to Silurian Balsfjord Group is interpreted to rest with a

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stratigraphic unconformity upon the very irregular paleo-topography of the Lyngen Magmatic Complex (Minsaas, 1981). The Svensby formation, a part of the Balsfjord Group, has calc-alkaline pillow-basalts (Kvassnes, 1997). In the southwest, the contacts to the Lyngen Magmatic Complex are all strongly sheared. The group is locally inverted, but is considered a coherent unit that is getting younger westwards (Andresen and Bergh, 1985).

### *Sample localities*

Seven profiles throughout the Lyngen Gabbro have been sampled in detail. The areas north of the Kjosen Fjord are the Iddu area in the far north (81 samples, including a very detailed profile of 12 samples), Strupen in the east (128 samples), and Skaidevarri in the central part (91 samples). Isskardet is in the west, south of the Kjosen Fjord (24 samples), Ellendalen east of the contact to the volcanics (18 samples), Goverdalens is the valley to the northeast of Ellendalen (37 samples), and Veidalen is the valley in the central south (40 samples). The Norwegian Chart Service maps 1634IV, 1634V, 1633IV and 1533I in the series M711 cover the area.

#### **Iddu**

Iddu, on Lyngstuva, is the northernmost mountain of the Lyngen peninsula. The sample area is less than 1-km<sup>2</sup> and comprises gabbroic rocks and tonalites cut by basaltic dikes. Oceanic, amphibolite-grade, shear-zones and Caledonian, greenschist-grade, shear-zones crosscut the area. In addition, later, possibly Caledonian, high-angle normal faults crosscut the section.

The tonalites are classified as shear zone-related- or layered tonalites (Moen-Eikeland, 1999). The layered tonalites, interpreted to be the felsic fractionates of the original magma, are characterized by large gray-black, lens-shaped, quartz-grain-aggregates in a groundmass of quartz, plagioclase, and pyroxene. The quartz-aggregates can be up to 1.5-cm in size and their orientation follows the layering. The shear zone-related tonalites, interpreted by Slagstad (1995) to be the result of anatetic melting of amphibolized gabbro, are fine-grained and consists of plagioclase, quartz, pyroxene and amphibole.

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The gabbroic rocks are subdivided into massive and laminated gabbro, varitextured leucogabbro and gabbroic pegmatites (Moen-Eikeland, 1999). Massive and pegmatitic gabbros occur together. The massive gabbro is fine-grained with plagioclase, two pyroxenes, amphibole, and magnetite. Coarser patches often show larger grains of magnetite. The gabbroic pegmatites have plagioclase, pyroxenes, amphibole, magnetite, and quartz and the grain-sizes vary from 1- to 15-cm. The laminated gabbros are composed of plagioclase, pyroxenes, amphibole, and some quartz.

### **Strupen**

The area consists mainly of layered, gabbroic rocks that frequently contain quartz as a cumulate mineral. Five major gabbroic rock types have been described (Hetland, 1996). Stratified layers (Type 1) grade from leuco- to melagabbro, exhibiting grain-size lamination, non-oriented cumulate clinopyroxene, and 5-10-cm long clinopyroxene-lenses oriented parallel to the layering. The oxide-rich gabbroic layers (Type 2) with plagioclase-rich horizons have unusually sharp contacts with the adjacent layers. Stratified, oxide-rich gabbroic layers (Type 3) have increasing amounts of oxide upwards. Oxide and quartz-bearing gabbroic layers (Type 4) comprise distinctive lamination defined by quartz grains oriented parallel to the layering. Tonalite (Type 5) are found as to five to six-meter thick layers. The latter rock type often occurs above the oxide-rich layers, and tends to inter-finger with them. The quartz bearing gabbroic rocks are unique and characteristic features of the Eastern suite, although the volume of these rocks does not exceed 5% in the section. The profile sampled is 95-m long, and has been sampled perpendicular to the layering.

### **Skaidevarri**

This 160-m transect is an east-west profile along the south side of Skaidevarri. The rocks range from troctolitic gabbros through olivine-gabbros, gabbros, olivine-gabbronorites, olivine-oxide-gabbros and oxide-gabbros. The rocks are mostly uniform with some fine, near-vertical layering in places. The sampling was done perpendicular to the layering. Some of the gabbros have undergone greenschist-facies metamorphism, with saussuritized plagioclase, iddingsitized olivine, and clinopyroxene partly altered to

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

green amphiboles with only small amounts of relict clinopyroxene. The alteration grade ranges from 10-100%. The fresh rocks show poikilitic pink clinopyroxene, brown hornblende, and sometimes plagioclase occurs recrystallized or as inclusions in clinopyroxene. In fact, the magmatic minerals all have inclusions of the others minerals in them, suggesting a complex magmatic history.

### **Isskardet**

The area consists of gabbro showing pronounced compositional layering that is steeply dipping and strikes N-S and perpendicular to the valley. The layering is typically of 5-10 meter scale with smaller scale (10-50 cm) layering locally superimposed. The larger scale layers can be traced laterally for 1-2 km across the valley. The rocks are generally medium to fine-grained with granulitic textures, with appearances similar to those of Skaidevarri. They range from olivine-gabbros to oxide-gabbronorites, and no tonalites have been sampled.

### **Ellendalen**

This valley is described in the westernmost profile in figure 2-2. An imbricated gabbro-slab west of the Aksla Volcanics is composed of high-level varitextured and massive gabbros cut by mafic dikes. The dikes trend N-S, with a 60° and 80° dip to the west. At Lakselvdalstindane, the northern wall of Ellendalen is dominated by layered gabbros folded in a large drag-synform with an N-S trending fold-axis dipping 40° to the north (Randall, 1959). The layering can be followed for at least 2200 meters along the mountain scarp and can be observed from 300 to 1400 meters above sea level. The sequence in this mountain may therefore represent 850 meters of layered gabbro. At the innermost part of Ellendalen, an up to 1-km wide, sub-vertical, N-S trending anastomosing shear zone caused a large antiformal drag-fold to form in the layered gabbros. Isotropic and varitextured gabbros have been found in the lower part of these mountain-scars, where sampling ended. The gabbroic rocks have, in general, undergone greenschist alteration, and magmatic minerals have not been recovered except for rare cores of clinopyroxene.

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### Goverdalen

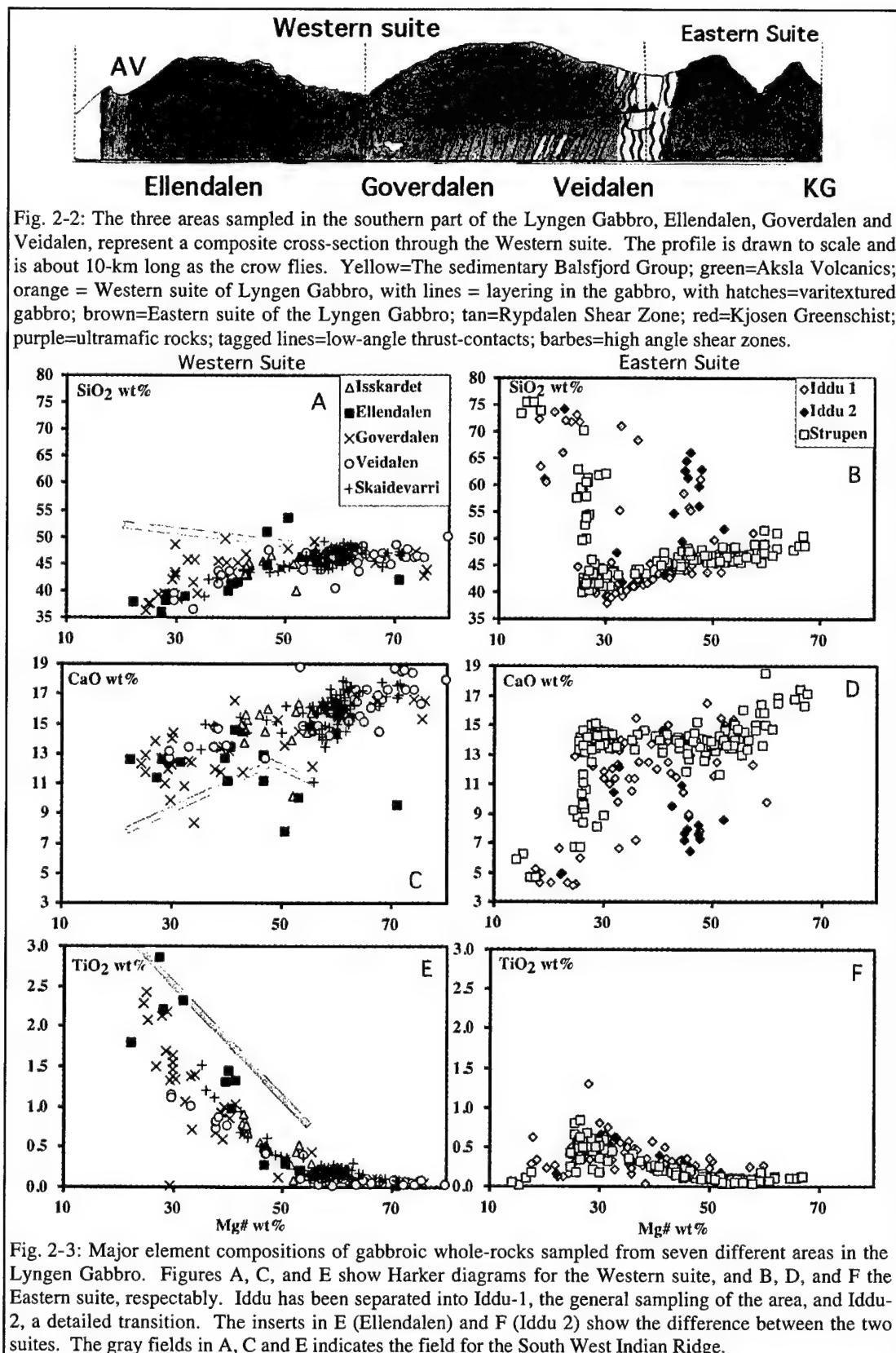
Following the shearzone northwards to Goverdalen, high-level gabbros continue into layered gabbroic rocks that can be followed eastwards through the valley and towards the Goverdalen Lake (Fig. 2-1 and 2-2). The gabbroic rocks of the Southwestern part of Mt. Nállangaisi host plagioclase-free wehrlite bodies close to the shear-zone. The steep walls of the valley display layered gabbros, and the shear-zone outcrops on the valley floor, displaying anatetic tonalites crosscut by inferred oceanic dikes (Selbekk, 1995, 1998). East of the shear-zone, layered and laminated gabbros were sampled towards the Goverdalen Lake. The rocks are generally two-pyroxene gabbros, and olivine is mostly absent. A very small outcrop of wehrlite was found. The most evolved gabbros are rich in magnetite and ilmenite. The rocks have frequently undergone greenschist alteration, and epidosites have been found in the west of this valley.

### Veidalen

East of the Goverdalen-lake, a very complicated layered sequence of wehrlites and gabbros outcrop. The layering appears to be deformed at a late magmatic stage, with multiple ductile folds, dunite pods and flame-structures. The gabbros are very fine grained, and plagioclases so dark that the rocks appear to be pyroxenites in outcrop. Some of the gabbros have high-Ca garnets and blue amphibole. The sequence looks similar to the Middle Series of the Rum layered intrusion on the Western coast of Scotland where it has been interpreted to represent the feeding zone for the intrusion (Emeleus *et al.*, 1996).

Mt. Balgesvarri, to the north of this valley, is dominated by layered gabbros, and the anastomosing shear-zone causes the layering to be folded into a synformal drag-fold. At the foot of the mountain, a large low-angle, westward dipping thrust-fault intersects the gabbro and an ultramafic body outcrops at the Eastern top. Entering Veidalen, an approximately two kilometer thick sequence of laminated meta-gabbro cut by basic and intermediate dikes can be followed towards the Rypdal shear-zone (Selbekk, 1998). Within the shear-zone a large ultramafic body of wehrlite and dunite (2 km long and 0.5 km wide) outcrops from Sydbreen in the north to Veidalsvatnet in the south. A smaller

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body of dunite is also seen below the small glacier in Nállancohka south of Veidalsvatnet.

From the middle of Veidalsvatnet and eastwards into Veidalen and Gaskacóhka quartz-bearing gabbros and quartz-rich tonalites typical of the Eastern suite of the Lyngen Gabbro appear. There are no geochemical data from this part of the section. The Kjosen greenschist outcrops at the Eastern side of Mt. Njállavarre.

### ***Methods***

#### **Whole rock analyses**

Major and trace element analyses were performed at the University of Bergen. Fist size samples were crushed in a jaw crusher, and 80-100 cm<sup>3</sup> were crushed to powder in an electrical agate mortar. The glass bead technique of Padfield and Gray (1971) was used for the major elements analyses and pressed-powder pellets for the trace elements analyses, using international basalt standards with recommended or certified values from Govindaraju (1994) for calibration. The analyses were carried out on a Phillips PW 1440 X-Ray fluorescence spectrometer. Instrumental precision for the major and trace elements for the glass-bead and pressed powder pellets have been documented by repeated analyses of representative samples. The relative standard deviations are close to 100% for low concentrations of the elements Zr, Y, Nb and P<sub>2</sub>O<sub>5</sub>. The whole-rock analyses are presented in Table 1.

Rare earth elements (REE), thorium, and tantalum have been analyzed at the Department of Earth Sciences, Memorial University of Newfoundland, Canada by inductively coupled plasma mass-spectrometry (ICP-MS). Analyses of international standards show that the relative uncertainties are better than 3% for Sr, Ba, Zr and the LREE, between 3 and 7% for Rb, Nb, Sc and HREE, and 10% for Ta and Th (Jenner et al., 1990). The data are presented in Table 1B.

#### **Mass spectrometry analyses**

73 samples have been measured for Sm and Nd-isotopes on a Finnegan MAT 262, 9-collector, fully automated mass spectrometer (MS) at the University of Bergen. All

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chemical processing were carried out in a clean-room environment with HEPA filtered air supply and positive pressure. The reagents were either purified in two-bottle Teflon stills or passed through ion-exchange columns. Samples were dissolved in a mixture of HF and HNO<sub>3</sub>. REE were separated by specific extraction chromatography using the method described by Pin *et al.* (1994). Sm and Nd were subsequently separated using a low-pressure ion-exchange chromatographic set-up with HDEHP coated Teflon powder (Richard *et al.* 1976). Sm and Nd were loaded on a double filament and analyzed in multidynamic mode. Nd isotopic ratios were corrected for mass fractionation using a <sup>146</sup>Nd/<sup>144</sup>Nd ratio of 0.7219. Sm and Nd concentrations were determined using a mixed <sup>150</sup>Nd/<sup>149</sup>Sm spike. Repeated measurements of the JM Nd-standard yielded an average <sup>143</sup>Nd/<sup>144</sup>Nd ratio of 0.511113 ± 15 (2σ) (n = 62). The typical Nd blank level in the laboratory is 5pg. The results are listed in Table 2.

### Mineral analyses

Major element analyses of clinopyroxene, orthopyroxene, plagioclase, and olivine have been performed using an ARL-SEMQ electron microprobe at the Nordic Volcanological Institute, Iceland. The microprobe analyses reported here are representative single-point analyses. All analyses were done with focused beam (2 μm in diameter). The analyses were performed with a beam potential of 15 kV, sample current of ~15 nA, 4 seconds counting times for peak, and MAN (mean atomic number) corrections for background. Standards used were basaltic glasses. Chemical analyses of clinopyroxene, orthopyroxene, plagioclase, olivine and ore-minerals of 79 additional samples were carried out at the University of Bergen. These were done as standardless EDS-analyses on a scanning electron microscope (JEOL scanning microscope, JSM-6400), using the Tracor Northern energy dispersive analysing system (TN 5600), with a 20 kV beam. A final set of mineral analyses was made on an ARL-SEMQ electron microprobe at the University of Bergen. All analyses were done with a defocused beam (10 μm in diameter). The analyses were performed with a beam potential of 15 kV, sample current of 10 nA, 40 seconds counting times for peak and 10 seconds for background (MAN (mean atomic number) background correction for Si, Ca, Al and Fe).

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Standards used were basaltic glasses, minerals, and metals. The results are listed in Table 3.

### **Ion-probe analyses**

Three samples were selected from Goverdalen and Veidalen, representing the geochemical spectrum of plagioclase-bearing rocks in the areas. Four samples were selected from Strupen, representing four of the rock types spread out in the lower half of the profile. Unaltered cores and rims of clinopyroxene from these rocks were analyzed on the CAMECA IMS 3f ion probe at Woods Hole Oceanographic Institution using the methods of Shimizu and Hart (1982). A primary beam of O<sup>-</sup> ions was focused to ~20 $\mu$ m for REE (La, Ce, Nd, Sm, Eu, Dy, Er, Yb) and ~10 $\mu$ m for other trace elements (Ti, V, Cr, Sr, Y, Zr). Molecular interferences were eliminated by energy filtering and a secondary voltage offset of -30 to -60V for the REE and -90 for the other trace elements. Uncertainties based on counting statistics were 5-10% (1 $\sigma$ ) for REE and 1-5% (1 $\sigma$ ) for the other trace elements. The data is presented in Table 4.

### **Mobility of elements**

The rocks of the Lyngen Gabbro have been exposed to several metamorphic events, ranging from lower greenschist-facies to upper amphibolite- and possibly granulite-facies metamorphism. Several studies have investigated how elements are leached or enriched due to metamorphism of magmatic rocks. Cann (1977) and Coish (1977) investigated the mobility of different elements during sea-floor metamorphism, concluding that Ti, P, Y, Zr, Nb, Cr, and Ni are stable during greenschist-facies metamorphism. Shervais (1982) showed Ti and V to be stable under a wide range of metamorphic conditions, ranging from sea-floor metamorphism to granulite facies metamorphism. Elliot (1972) investigated gabbros that occur as isolated masses within amphibolites, and concluded that TiO<sub>2</sub>, FeO<sup>l</sup>, MgO, MnO, and Na<sub>2</sub>O and possibly Al<sub>2</sub>O<sub>3</sub> were stable during amphibolite facies metamorphism. Weaver and Tarney (1981) analyzed mafic and ultramafic dikes cutting retrograded dikes, comparing dikes exposed to varying degree of metamorphism to fresh dikes. They concluded that Nd, P, Hf, Zr, Ti and the middle to heavy REE's are immobile even during strong metamorphism with

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

access to fluids. An investigation of gneisses from a mylonitic shear zone shows that  $\text{SiO}_2$ ,  $\text{K}_2\text{O}$ ,  $\text{Na}_2\text{O}$ ,  $\text{FeO}^t$  and  $\text{CaO}$  are mobile when fluid phases are present (Sinha *et al.*, 1986).

In general, it is difficult to determine whether an element is mobile or not, and whether it will be enriched or depleted during alteration and metamorphism. However, the above studies have shown that Zr, Y, P, Nb, Cr, Al, Ti, V and the REE will be relatively stable during the metamorphic conditions shown to be present along the composite cross-section. The ratios between these elements will hence be an accurate way to compare the rocks.  $\text{MgO}$  and  $\text{FeO}^t$  concentrations are also stable in the gabbros at higher metamorphic grades. As Sr is unstable during alteration, no Sr isotopic analyses have been attempted.

## Results

### Major elements

The gabbroic rocks from the Western suite of the Lyngen Gabbro are, in general, very low in silica (Fig 3) and large sections can be defined as ultrabasic (<45%  $\text{SiO}_2$ ). Very few samples from profiles that were analyzed from the Western suite have higher than 50-wt%  $\text{SiO}_2$ , and the average is 45.5 wt%. The Eastern suite follows a  $\text{SiO}_2$  enrichment trend resulting in tonalites as the most evolved rocks. However, less than 5% of the rocks of the Eastern suite are high-silica differentiates. The  $\text{Al}_2\text{O}_3$  contents in the Eastern suite are highest for the rocks with the lowest  $\text{SiO}_2$  contents, reflecting the importance of calcic plagioclase as a cumulate mineral. The highest  $\text{Al}_2\text{O}_3$  contents are found in the rocks with the highest  $\text{SiO}_2$  in the Western suite.  $\text{Na}_2\text{O}$  varies from traces to 4 wt%, but there is no clear correlation between  $\text{SiO}_2$  and  $\text{Na}_2\text{O}$  in the Western suite, whereas the sodium-content increases with  $\text{SiO}_2$  in the Eastern suite.  $\text{K}_2\text{O}$  and  $\text{P}_2\text{O}_5$  are extremely low, and rarely exceed 0.1 wt%.

As the  $\text{SiO}_2$  contents behave so differently between the suites, the degree of differentiation is explored using Mg#’s ( $100 * \text{MgO}/(\text{MgO}+\text{FeO})$ ) in wt-%.) The trend of decreasing  $\text{SiO}_2$  with increased differentiation is evident in the Western suite (Fig 2-4 A and B). The rocks of the Eastern suite partly indicate this trend, even if the most evolved

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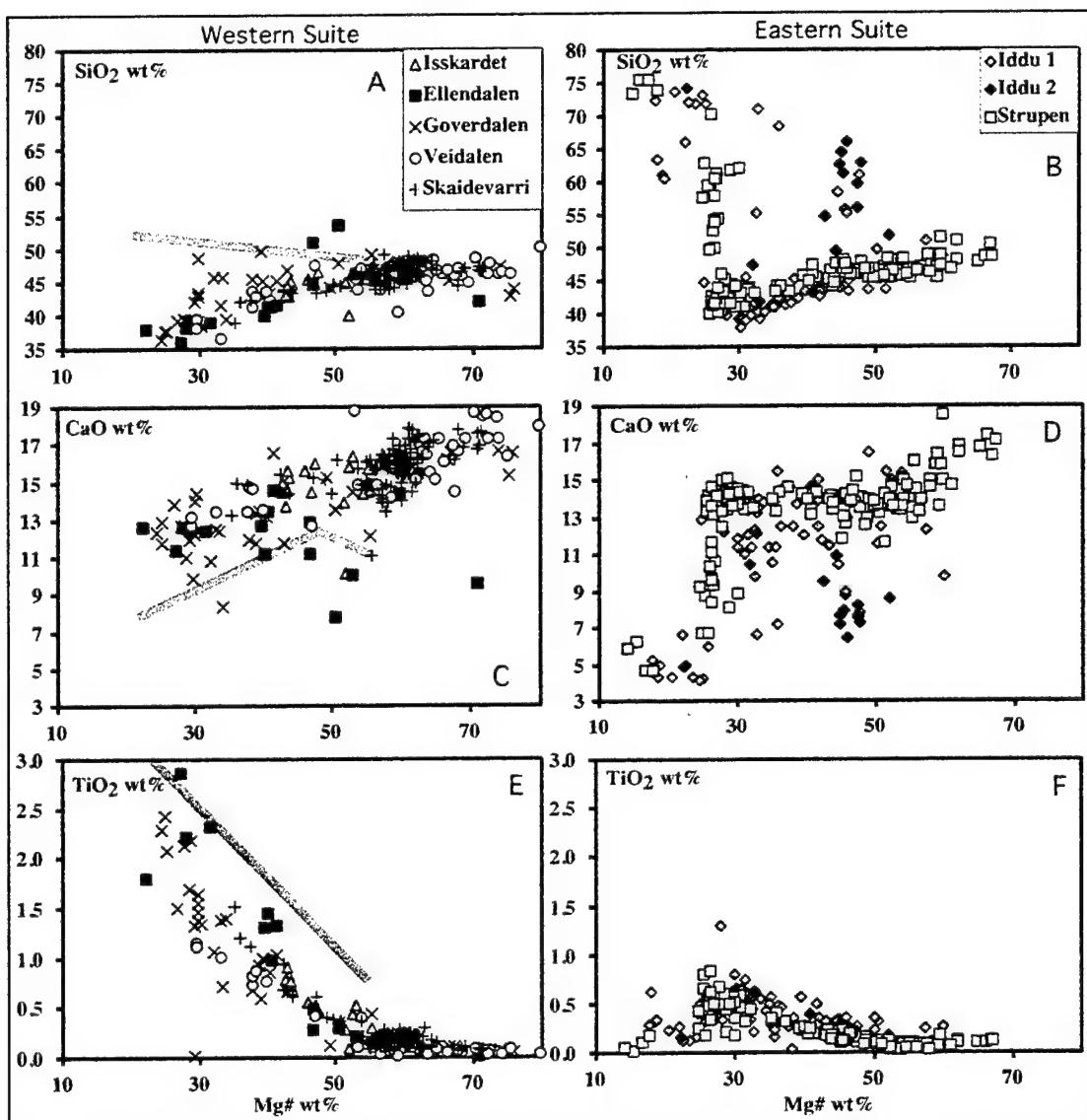


Fig. 2-4: Whole-rock Mg# ( $100 \cdot \text{MgO}/(\text{MgO}+\text{FeO}^*)$  ( $\text{FeO}^* = \text{all Fe as FeO}$ )) is plotted vs. SiO<sub>2</sub>, CaO and TiO<sub>2</sub> for the Western (left) and Eastern (right) suite. Note how the SiO<sub>2</sub> content in the gabbros correlates with Mg# which is the opposite of that of basalts (gray line). In the Eastern suite, the fractionation of magnetite causes the SiO<sub>2</sub> to increase.

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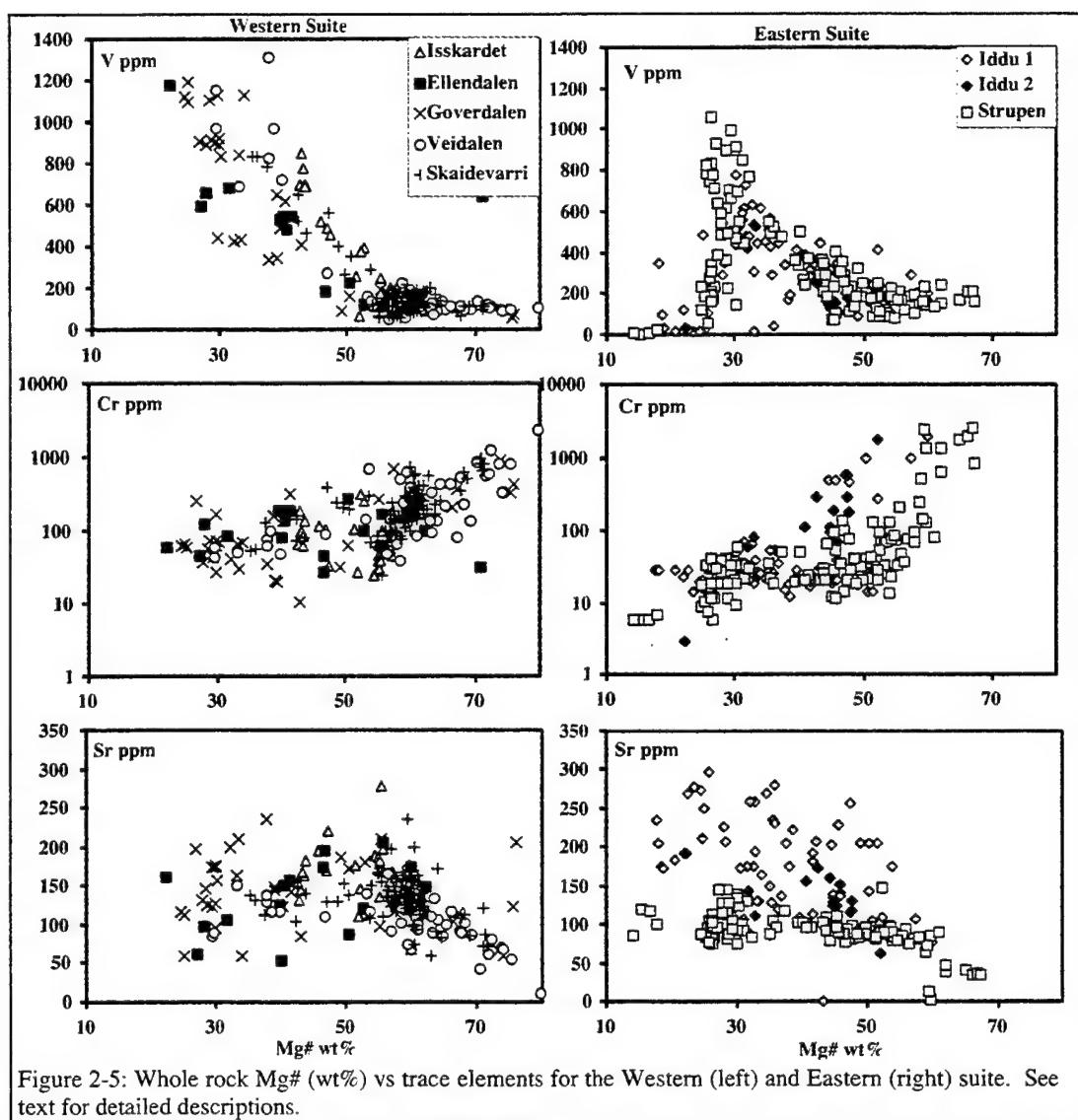


Figure 2-5: Whole rock Mg# (wt%) vs trace elements for the Western (left) and Eastern (right) suite. See text for detailed descriptions.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

rocks show increased SiO<sub>2</sub> levels and rocks from Iddu show early SiO<sub>2</sub> enrichment (at Mg# 45).

The CaO (Fig 2-4C and D) contents decrease gradually during fractionation in the Western suite. However, the Eastern suite, and in particular the Strupen profile, shows near constant CaO-levels until the onset of SiO<sub>2</sub> enrichment.

The TiO<sub>2</sub> contents are dramatically different between the two suites (Fig 2-4 E and F). The Eastern suite shows few rocks with TiO<sub>2</sub> higher than 1 wt%, whereas the Western suite has up to 3 wt% in the most SiO<sub>2</sub>-poor rocks. The TiO<sub>2</sub> trends indicate enrichment with decreasing Mg#'s for both suites. At around Mg# 30, the TiO<sub>2</sub> levels drop off for the Eastern suite, whereas they increase steeply for the Western suite, indicating that the evolved, iron-rich, melts precipitated iron-titanium oxides within the cumulates.

### Trace elements

The trace elements partly reflect what is observed in the major elements. It should be noted that many of the incompatible elements like Zr and Y are very close to the detection level of the XRF-method, especially in the Eastern suite.

Figure 5 demonstrates vanadium levels with similar patterns as the TiO<sub>2</sub>, although V indicates the presence of oxide-minerals in the Eastern suite as well as the Western suite. The V contents do, however, decrease suddenly at lower Mg#'s for the Eastern suite, possibly indicating melt-mineral segregation leaving oxide-minerals in the cumulate. In addition, the Cr concentrations are similar between the two suites but become much lower with fractionation in the Eastern suite than in the Western suite. The Sr contents show little in terms of a differentiation trend for either suite, but the concentrations are highest for the profile at Iddu.

Sm and Nd contents were measured by isotope dilution (Figure 2-6 and 2-7A). The trace elements are highest in the Western suite, where they show some correlation with Cr or Mg#. The Eastern suite has concentrations below 1ppm, and shows little correlation with either Cr or Mg#.

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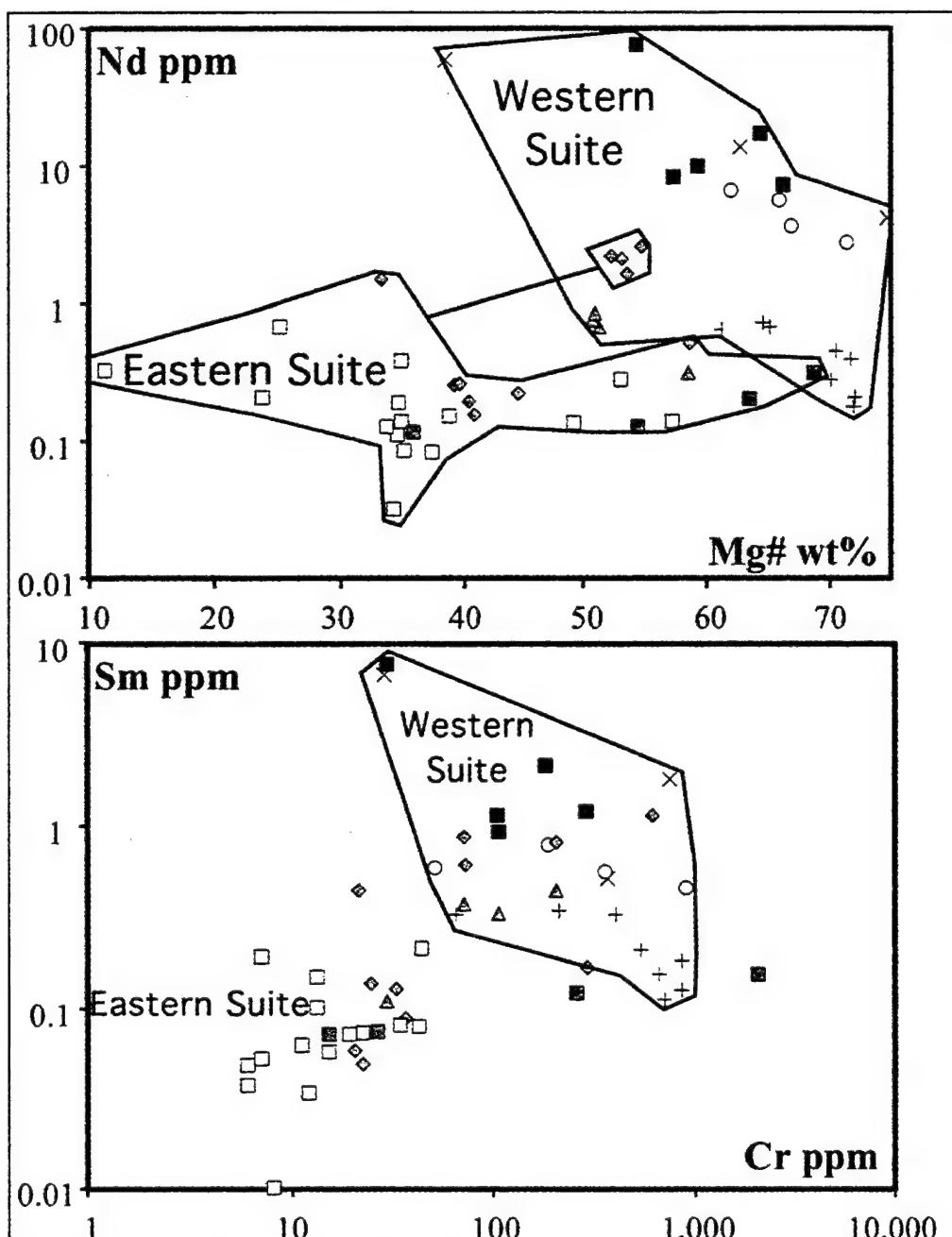


Fig. 2-6: Whole rock analyses of the Lyngen Gabbro. Symbols are the same as for Fig. 2-5, and in addition, the grey squares which represent the lower 50 meters from Strupen. Note the generally low contents of the incompatible trace elements in for the Eastern suite, whereas the Western suite has higher trace element concentrations that increase with differentiation.

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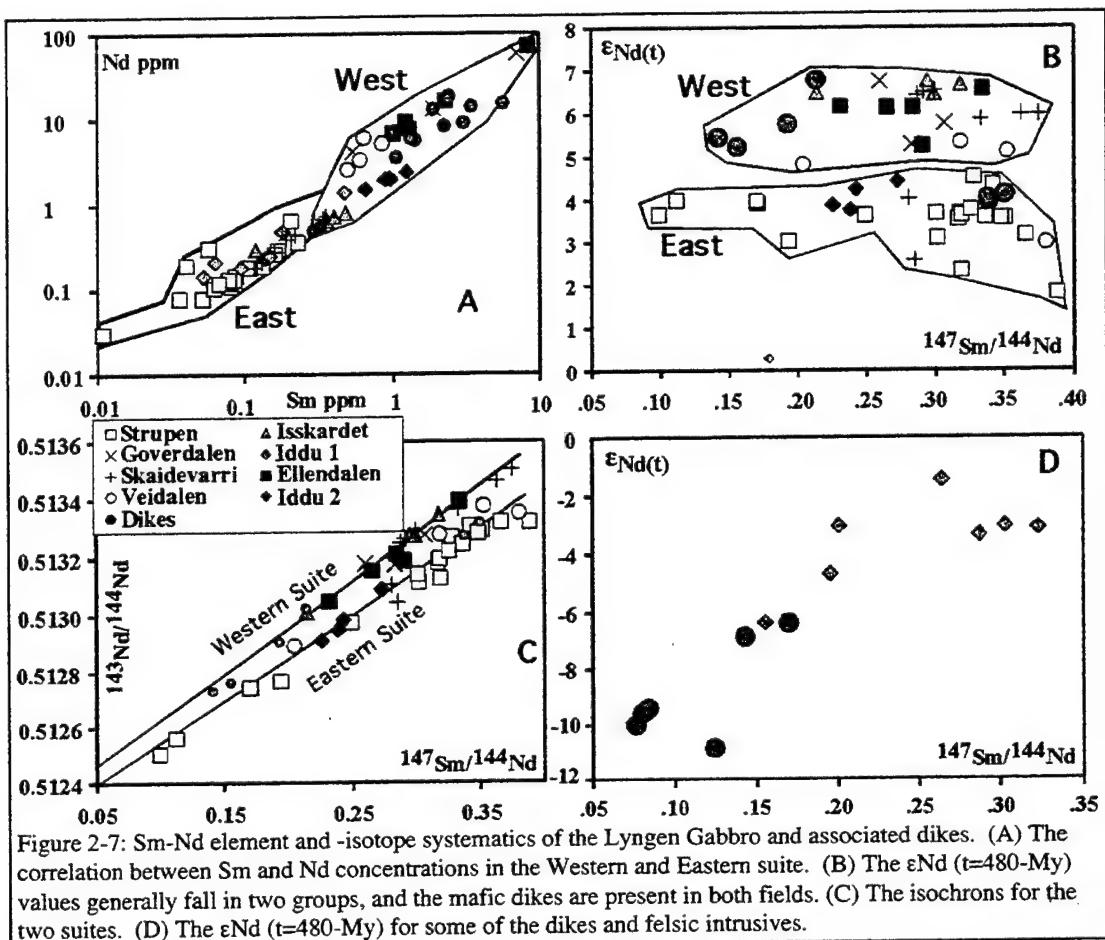


Figure 2-7: Sm-Nd element and -isotope systematics of the Lyngen Gabbro and associated dikes. (A) The correlation between Sm and Nd concentrations in the Western and Eastern suite. (B) The  $\epsilon_{\text{Nd}}$  ( $t=480\text{-My}$ ) values generally fall in two groups, and the mafic dikes are present in both fields. (C) The isochrons for the two suites. (D) The  $\epsilon_{\text{Nd}}$  ( $t=480\text{-My}$ ) for some of the dikes and felsic intrusives.

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### Nd –isotopes

The  $\epsilon_{\text{Nd}(t=480\text{My})}$  values for the profiles are shown in Figure 2-7. The profiles from Ellendalen, Isskardet, and Goverdalene have the highest values, around +6 (Fig. 2-7B). The Strupen samples and samples from the detailed profile at Iddu fall around +4. Veidalen samples fall in-between, whereas the Skaidevarri-profile shows values from both groups. The data plotted on an isochron-diagram indicate different initial values ( $\epsilon_{\text{Nd}(t)}$  is +6.39 and +3.95 respectively) (Fig. 2-7C). The dikes that cross both suites and Rypdal Shear Zone have  $\epsilon_{\text{Nd}}$  values similar to the Western and Eastern suite. Some of the dikes and shear-zone related tonalites from Iddu have negative  $\epsilon_{\text{Nd}(t)}$  values that become extremely low with differentiation (Fig. 2-7D) indicating AFC-processes (DePaolo, 1981).

### Mineral analyses: major elements

Figure 8b shows a plot of Mg# of clinopyroxene vs. An% of plagioclase in all the areas of this study (Mg#=100\*molar Mg/(Mg+Fe)). The Easternmost profile, from Strupen, has two groups, above and below 50 meters. Plagioclase An<sub>90.98</sub> coexists with Mg# 82-69 augite and Mg# 56-73 orthopyroxene. The upper half of the profile indicate the highest An contents of plagioclase found in the entire Lyngen Gabbro, and the lowest Mg#'s of clinopyroxene from the site. The lower half of the profile has clinopyroxene with Mg#'s as high as 82, but these also have An% above 90. The more An-rich trend in the upper half of the profile correlates with tonalites appearing in the layered gabbro sequence.

The next profile towards the west is Veidalen, with primitive rocks and An%'s in the higher eighties and into the 90's. The profiles in Veidalen show similar An contents to Strupen coexisting with higher Mg# augite (88 to 72). This profile is also associated with the aforementioned layered peridotites, which have augite Mg#'s up to 92 and Fo<sub>88</sub> olivine. West of this geochemical transition-zone, Goverdalene gabbros have An<sub>87-48</sub> plagioclase, while their augites change from Mg# 84 to 67. The olivines and orthopyroxenes follow similar trends (Fig. 2-8B and C).

The rocks of the Skaidevarre area have some plagioclase as high as An<sub>96</sub>, although the majority of the plagioclases are approximately An<sub>85</sub>, coexisting with Mg# 86-75

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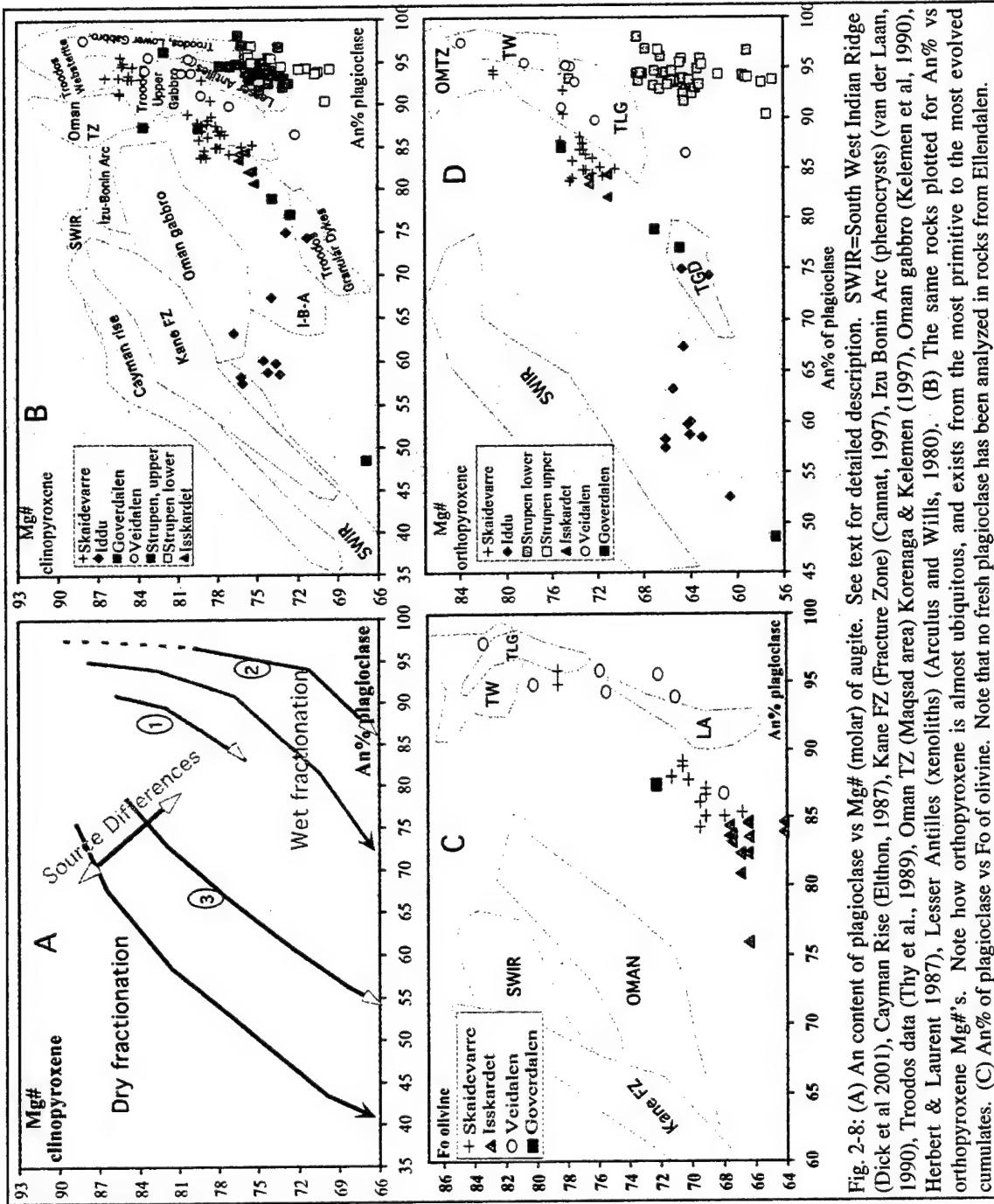
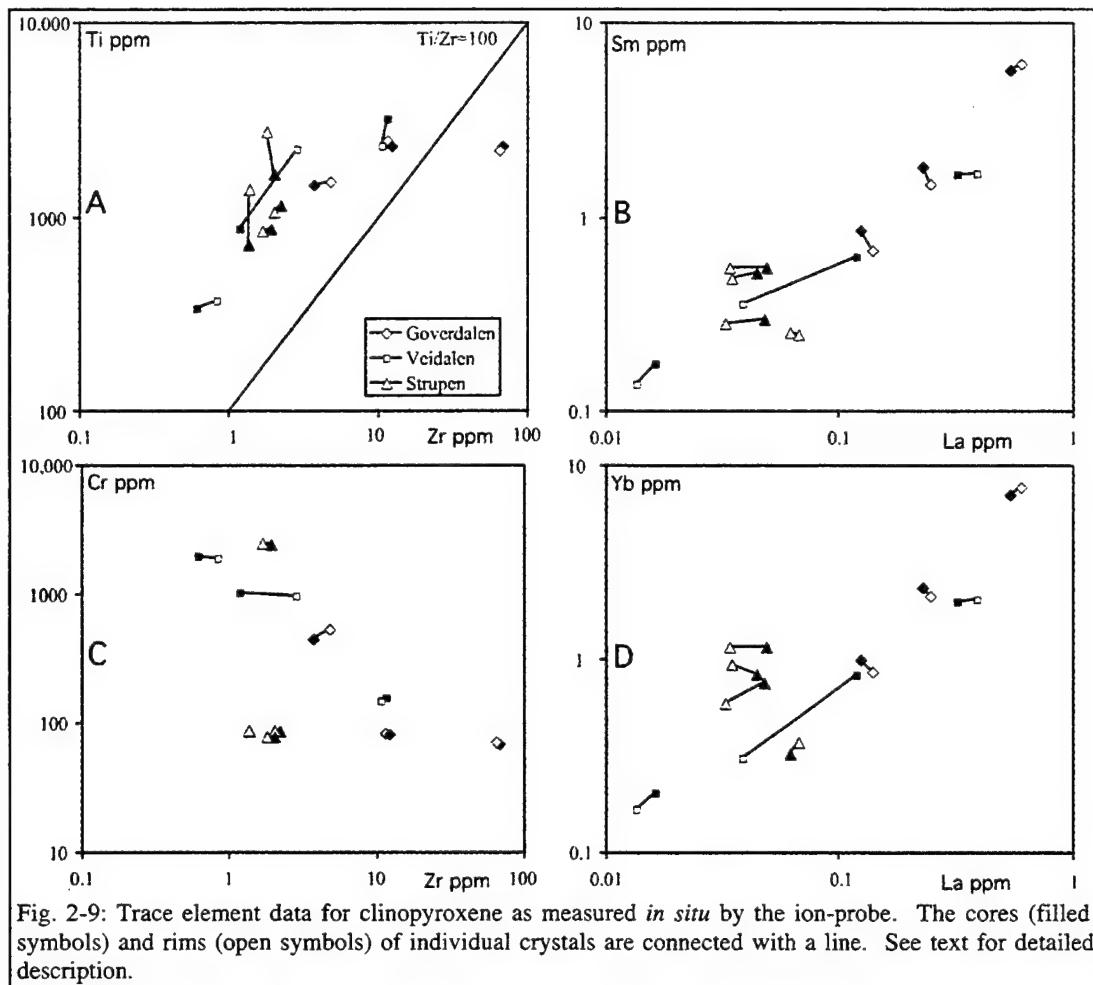


Fig. 2-8: (A) An content of plagioclase vs Mg# (molar) of augite. See text for detailed description. SWIR=South West Indian Ridge (Dick et al 2001), Cayman Rise (Elthon, 1987), Kane FZ (Fracture Zone) (Cannat, 1997), Izu Bonin Arc (phenocrysts) (van der Laan, 1990), Troodos data (Thy et al., 1989), Oman TZ (Maqṣad area) Korenaga & Kelemen (1997), Oman gabbro (Kelemen et al., 1990), Herbert & Laurent 1987), Lesser Antilles (xenoliths) (Arculus and Wills, 1980). (B) The same rocks plotted for An% vs orthopyroxene Mg#’s. Note how orthopyroxene is almost ubiquitous, and exists from the most primitive to the most evolved cumulates. (C) An% of plagioclase vs Fo of olivine. Note that no fresh plagioclase has been analyzed in rocks from Ellendalen.

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augite, Mg# 68-79 orthopyroxene and Fo<sub>66-78</sub> olivine. The Easternmost rocks are the most primitive cumulates, and have plagioclase An<sub>56-90</sub>. The rocks westwards of these are more evolved cumulates, although they do not fall below Mg# 70 together with plagioclase of An<sub>85</sub>. Isdalen rocks have cumulates that are similar to Goverdalén and the evolved Skaidevarre.

The rocks from Iddu vary greatly with regards to the plagioclase compositions (An<sub>56-75</sub>) with a near constant composition of augite and orthopyroxene, and have no olivine together with plagioclase.

### Mineral analyses: trace elements

Data from three areas, Goverdalén, Veidalen and Strupen, are shown in figures 2-9 and 2-10. Fig. 2-9 indicates that the fractionation-trend for each area is different. Rocks analyzed from Goverdalén and Veidalen have a general positive correlation for the incompatible elements. Strupen appears to have a very different distribution, however. Figure 2-10, indicate that the REE become more LREE/HREE depleted as the minerals become more evolved. There are no systematics regarding core-rim analyses, in fact, reverse zoning is common. The augites from Goverdalén and Veidalen develop negative Eu anomalies with differentiation, those from Strupen does not. The anomalies are absent in the most primitive rocks but are well developed in the evolved rocks, indicating closed system fractionation of plagioclase.

## *Discussion*

### Significance of the contrasting differentiation paths for the Eastern and Western suites

It is possible that the magmas that produced the Western suite and Eastern suite of the Lyngen Gabbro had different water-contents and/or had different mantle-sources. We have explored activity of water and source-affinity using the An%(pl)-Mg#(cpx) plot which distinguish the gabbroic rocks precipitated from a hydrous magma from a dryer one by the arc of the trend (Fig. 2-8A). Dry fractionation of basaltic melt produces a convex trend whereas those for crystallization of wet melts are concave.

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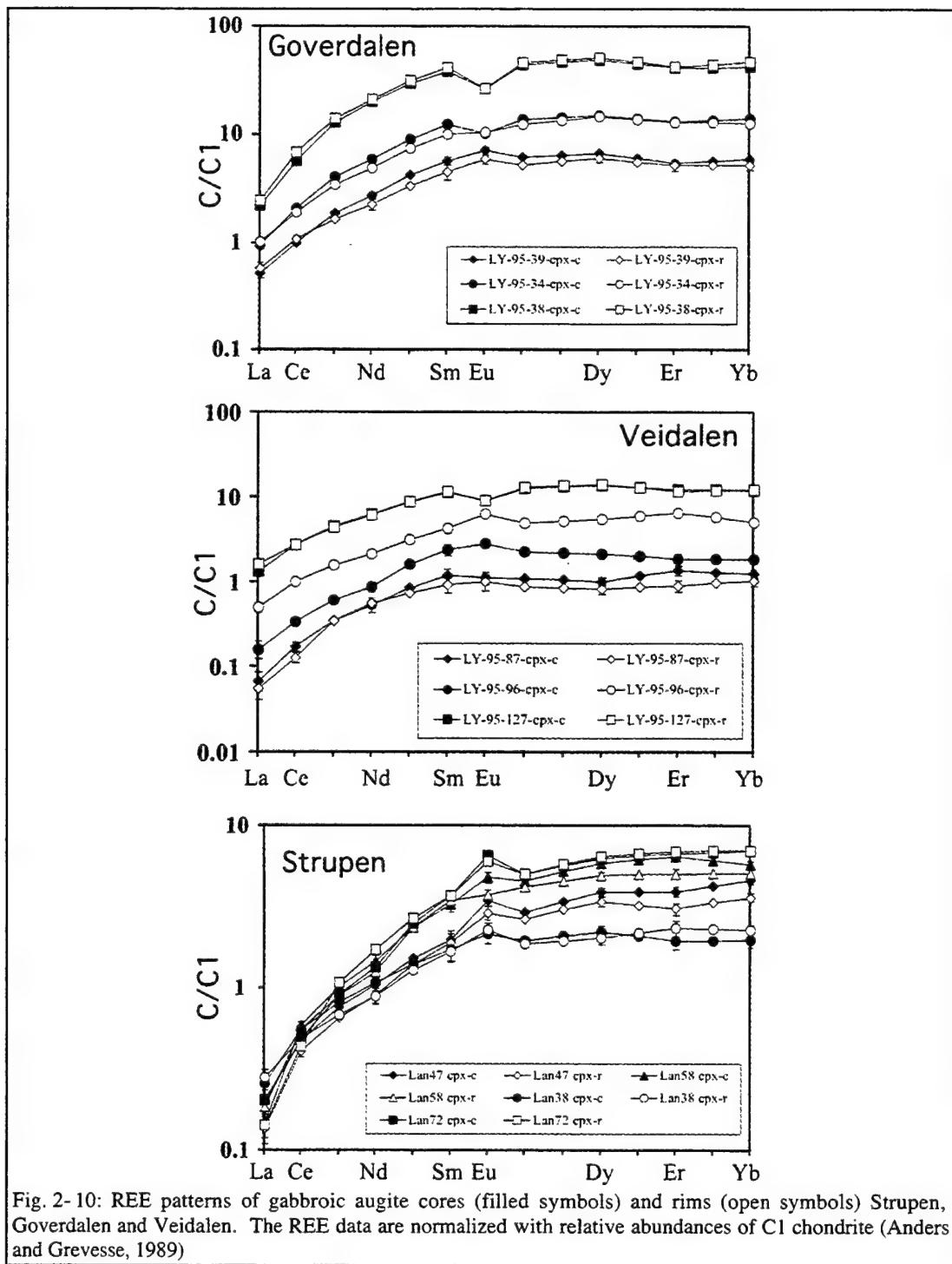


Fig. 2-10: REE patterns of gabbroic augite cores (filled symbols) and rims (open symbols) Strupen, Goverdalen and Veidalen. The REE data are normalized with relative abundances of C1 chondrite (Anders and Grevesse, 1989)

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The presence of water suppresses the plagioclase-olivine cotectic relative to the plagioclase-augite cotectic (Gaetani *et al.*, 1993) and changes the reaction series to augite and even orthopyroxene crystallization before plagioclase. In addition, a melt may gain or lose water so that  $K_d^{Ca/Na}_{Pl/L}$  changes, resulting in higher or lower An% plagioclase, respectively (Hough and Luhr, 1991). The effect is even more pronounced in silicic magmas (Arculus and Wills, 1990) and may cause An<sub>>95</sub> for more than 50% of the fractionating gabbros preceded by true ultramafic cumulates. Early fractionation of clinopyroxene before plagioclase results in a significant drop in Mg-content in the magma at a more rapid rate than if plagioclase was crystallized in concert with olivine only. Moreover, prevalent water variably depresses the liquidus and solidus of the silicic minerals, whereas there is little effect on the oxide-minerals in magmas (Gaetani *et al.*, 1993). The bend of the gabbro-trend produced by wet magma fractionation is caused by the onset of oxide fractionation. The rocks from the Western suite of the Lyngen Gabbro have this convex trend, similar to the Upper Gabbro from Troodos (Fig. 2-8B). Oxide minerals are common in primitive gabbros of both suites. The early oxide crystallization and fractionation is particularly pronounced at Iddu (Moen-Eikeland, 1999). The Iddu gabbro-compositions, however, are less systematic and range from the Izu-Bonin Arc to the field of Mid-Atlantic Ridge (MAR). The Eastern suite of the Lyngen Gabbro has rocks that have abundant quartz, coexisting with extremely An-rich plagioclase, two pyroxenes and magnetite, and the mineral-compositions are similar to the Lower Gabbros from Troodos (Fig. 2-8B). Overlapping with this trend are the Lesser Antilles lower crustal xenoliths, which Arculus and Wills (1990) demonstrated crystallized under hydrous conditions in a relatively silicic magma.

Dryer oceanic magmas, where plagioclase reaches the liquidus before clinopyroxene, experience a slower fractionation of Mg/Fe from the magma relative to Ca/Na in part because Ca is not depleted by augite-fractionation at an early stage. In addition,  $K_d^{Ca/Na}_{Pl/L} \approx 1$  (Grove *et al.*, 1992), and therefore reflect the Ca/Na ratio of the melt more directly. The vertical displacement of the curve varies with the Ca/Na and Mg/Fe of the parental melt (Fig. 2-8A). For instance, Mid-Cayman Rise gabbro compositions are similar to those from the South West Indian Ridge whereas gabbros from the Kane

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Fracture Zone region of the MAR where the Ca/Na is lower and Mg/Fe is higher, have lower An% plagioclase at a given Mg# of augite (Fig. 8B).

Increasing pressure lowers  $Kd_{PL}^{Ca/Na}$  and leads to the formation of a lower An% plagioclase (Panjasawatwong, *et al.*, 1995). In addition, the pressure depresses the plagioclase-olivine cotectic relative to the olivine-clinopyroxene cotectic, and the Ca-Tschermak-component of the clinopyroxene increases (Gaetani *et al.*, 1993). The result of fractionation at high pressures lead to ultramafic rocks with high-Al augite, followed by gabbroic cumulate with a convex trend, plagioclase starting at relatively low An-contents. Plagioclase-free wehrlites are common in the Western suite (Kvassnes, 1997), and cumulative, plagioclase-free harzburgites and lherzolites are found in the Eastern suite (Hetland, 1996). However, the high An% plagioclase and the concave fractionation-trends indicate crystallization from wet melts at relatively low pressures.

We used Melts 5 (Ghiorso and Sack, 1995) to estimate fractionation-trends for the rocks from this study (arrows 1-3 in figure 2-8A). A primitive MAR melt extracted from PetDB was used as an initial composition and partly reproduced fractionation-trends for Goverdalen and Skaidevarri by adding water to the melt at low crustal pressures (1-kbar) (arrow 1). The dry fractionation may therefore replicate some of the gabbros at Iddu. It has been argued that these calculations may be questionable for the Melts program (Yang *et al.*, 1996). On inspection, the high  $Kd_{PL}^{Ca/Na}$  values are comparable to those of hydrous melts (Housh and Luhr, 1991), and the partition-coefficients for the mafic minerals are reasonable. However, the addition of water to MAR-type melts or crystallization at higher pressures cannot reproduce the Strupen trend, as the Ca/Na of the melt is too low. The fractionation trend of the Stupen gabbros can, however, be reproduced by water-saturated crystallization of a dike ( $\epsilon_{Nd(t=480My)}$  of 4.1) that cross-cuts the Western suite in Goverdalen (sample 95LY71 as shown by arrow 2 in figure 2-8A). The difference between the upper and lower half of the Strupen profile can be explained by the addition of just a little more water to the magma that produced the upper half.

### Parental magma compositions

We have adapted Shervais' (1982)  $TiO_2$ -V tectonomagmatic discrimination-diagram for basalts to accommodate our gabbroic rocks and have compared our whole-

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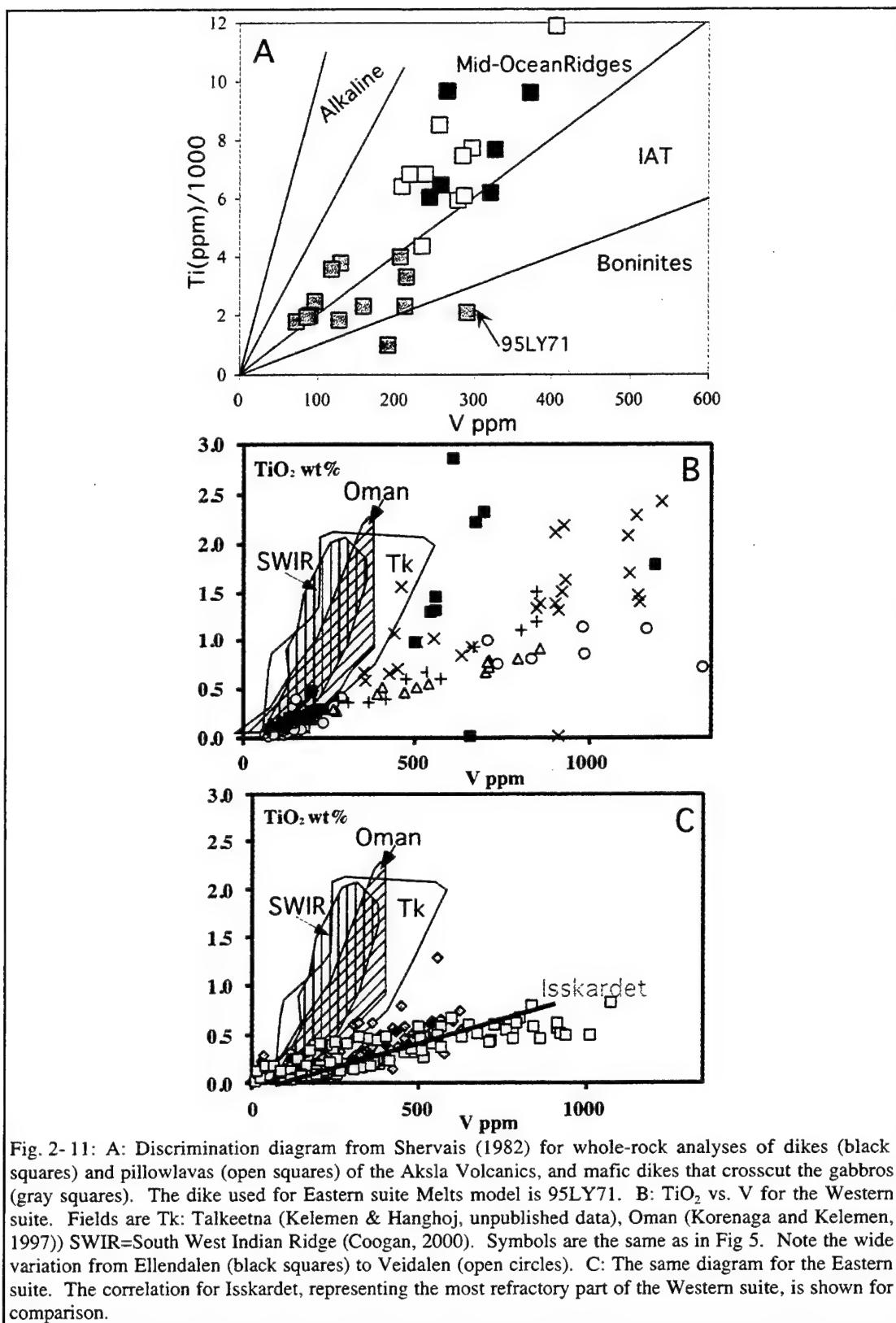


Fig. 2-11: A: Discrimination diagram from Shervais (1982) for whole-rock analyses of dikes (black squares) and pillowlavas (open squares) of the Aksla Volcanics, and mafic dikes that crosscut the gabbros (gray squares). The dike used for Eastern suite Melts model is 95LY71. B:  $\text{TiO}_2$  vs.  $V$  for the Western suite. Fields are Tk: Talkeetna (Kelemen & Hangoj, unpublished data), Oman (Korenaga and Kelemen, 1997)) SWIR=South West Indian Ridge (Coogan, 2000). Symbols are the same as in Fig 5. Note the wide variation from Ellendalen (black squares) to Veidalen (open circles). C: The same diagram for the Eastern suite. The correlation for Isskardet, representing the most refractory part of the Western suite, is shown for comparison.

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rock analyses to known sites. The Ti/V-ratio provides constraints on the tectonomagmatic conditions for the primary magmas, as both  $\text{TiO}_2$  and V follow the same magmatic processes. The source of the Eastern suite must have been derived from a more depleted mantle source than the Western suite, but the transition between the suites is gradational in the south (Fig. 2-11B and C). Most of the Aksla Volcanics dikes and pillow-lavas, however, are similar to those in oceanic ridge settings, whereas the dikes that crosscut the gabbros are more varied (Fig. 2-11A).

The REE-melt composition has been calculated using partition coefficients for the selected augite from Goverdalén, Veidalen and Strupen (Figure 2-12). The result is compared to N-MORB, IAT (Sun, 1980), the Aksla Volcanics (Furnes and Pedersen, 1995), boninites (Hickey and Frey, 1982) and Troodos lavas (Taylor and Nesbitt, 1988). The result supports that Goverdalén gabbros and Veidalen gabbros may have crystallized from magma similar to the Aksla Volcanics or an N-MORB seeing that their REE patterns are parallel. The Strupen augites, however, are very low in LREE/HREE and did not crystallize from “normal” boninitic magmas, but from ultra-depleted magmas not enriched in LREE, similar to boninitic lavas from Troodos. However, most boninitic lavas from Troodos have LREE enrichment and a U-shaped pattern (Taylor and Nesbitt, 1988). Like the Troodos magmas, the augites from Strupen indicate trends that are not parallel, suggesting source-evolution or crustal contamination (DePaolo, 1981). The patterns cross as the rocks become more evolved, suggesting that the chemical signature of the profile changes at about 50 meters above the base.

### Temporal and spatial relations between the Western and Eastern suite

The gabbros in the Eastern suite uniformly have lower  $\epsilon_{\text{Nd}}$  values than the Western suite and each suite follows separate, but parallel isochrons (Fig. 2-7), suggesting similar ages of the suites. Therefore, the parental magmas must have had different and relatively homogeneous compositions before emplacement. Slagstad (1995) and Selbekk (1998, 2002) suggested that the oceanic Rypdalén Shear Zone generally separates the two suites and there are several sites in the northern half of the Lyngen Peninsula that clearly supports this. However, in Skaidevarri the profile starts with low  $\epsilon_{\text{Nd}}$  values and An-rich

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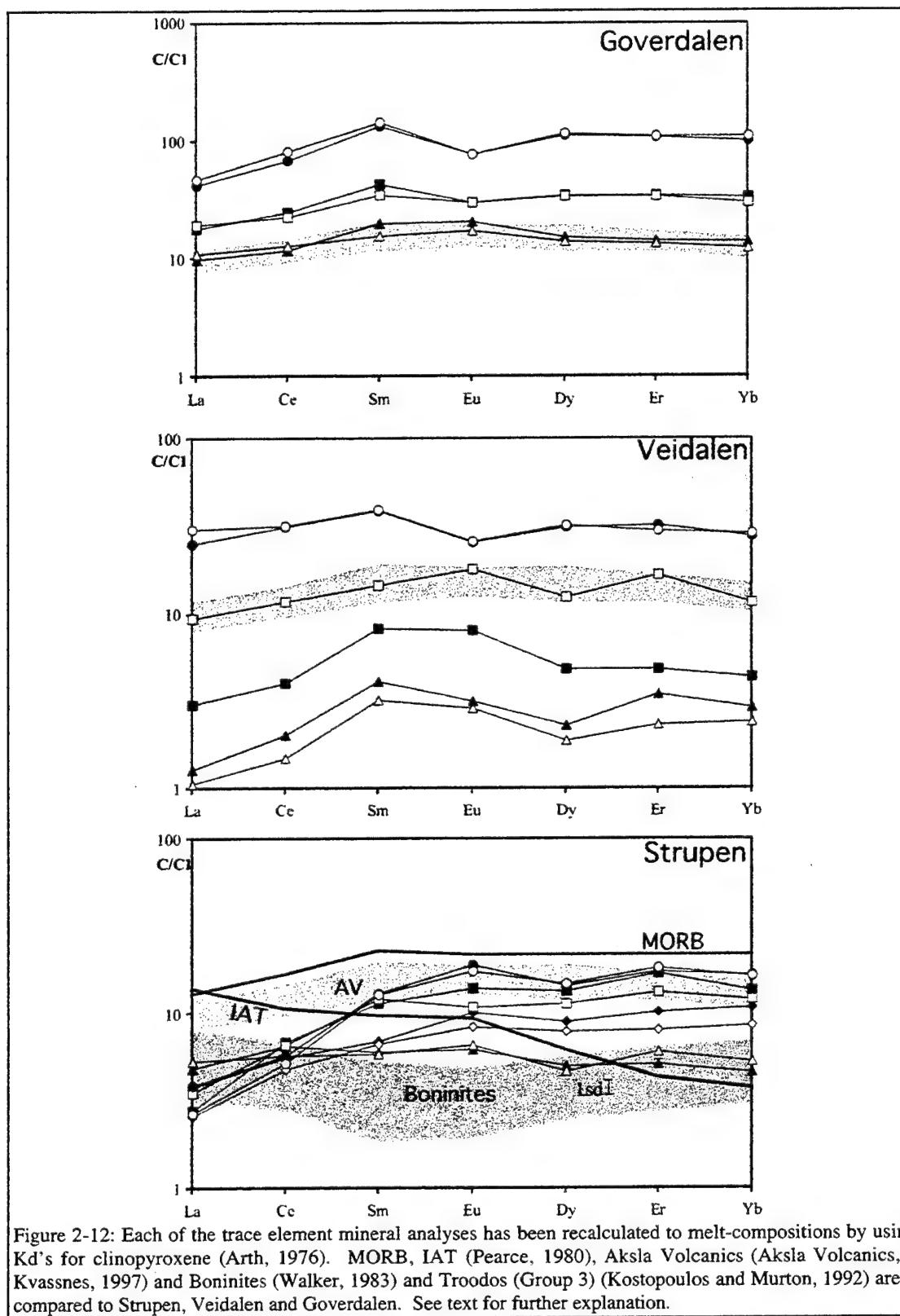


Figure 2-12: Each of the trace element mineral analyses has been recalculated to melt-compositions by using Kd's for clinopyroxene (Arth, 1976). MORB, IAT (Pearce, 1980), Aksla Volcanics (Aksla Volcanics, Kvassnes, 1997) and Boninites (Walker, 1983) and Troodos (Group 3) (Kostopoulos and Murton, 1992) are compared to Strupen, Veidalén and Goverdalén. See text for further explanation.

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plagioclase in the east. The trend changes over less than 10-m into higher  $\epsilon_{\text{Nd}}$  values and reduced plagioclase An% contents, whereas the Mg#'s of mafic minerals remain mostly unchanged. Since the most primitive rocks are those with low  $\epsilon_{\text{Nd}}$  it is not likely that the magma was contaminated by AFC-processes, and we suggest that the Skaidivarre-profile was sampled across a magmatic transition between the two suites. The layered series has no intrusive or tectonic contacts, indicating that the two magmatic systems coexisted. In Skaidivarre, the gabbros are cut by the Rypdalen Shear-Zone further to the west than the contact between the suites. It is therefore unlikely that the Rypdalen Shear-Zone caused the boninites to be emplaced together with the tholeiites, but rather that for some parts of the complex magmatism was concurrent and that the rest of the complex was produced with close proximity between the suites.

Several dikes crosscut the gabbros. On the basis of  $\epsilon_{\text{Nd}}$ , the mafic dikes represent magma from both suites. However, some of the felsic dikes that crosscut the gabbros have negative values, indicating enriched material either through assimilation or by source enrichment (Fig. 2-7D). These dikes have not been dated, and it uncertain if they were intruded at a later stage.

### Plate Tectonic Affinity and comparison to other sites.

Boninites are generally believed to be derived from a peridotitic source after MORB extraction (Crawford *et al.*, 1989; Kelemen *et al.*, 1990) and frequently show a LREE enrichment causing a U-shaped pattern. However, there are several categories of boninites, Low-Ca (3 types) and High-Ca boninites and the Eastern suite represents high-Ca type boninite as seen in the extremely An-rich plagioclase. Although the augites from the Eastern-suite gabbro do not reflect the U-shaped REE-pattern often seen in boninites, the cumulates are clearly crystallized from a hydrous, ultra-depleted magma while having lower  $\epsilon_{\text{Nd}}$  values than the Western suite. The Western suite gabbros were also crystallized from a hydrous magma, although the corresponding upper crustal material reflects tholeiites like those from Mid-Ocean Ridge type magmatism, making us conclude these are island arc tholeiites. In addition, the two suites were emplaced concurrently with a gradational transition between them in the south.

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It has been known for some time that a back-arc type, and arc and fore-arc (*supra-subduction zone*) type magmatic rocks occur in ophiolites. For example, Schouten and Kelemen, (2002) demonstrated that the lower (tholeiitic) and upper (boninitic) lavas in the Troodos Ophiolite are related to the Upper and Lower Gabbros, respectively, and they argue that the systems were contemporaneous. Taylor and Nesbitt (1988) showed REE-patterns that cross for the Troodos Upper Lavas, and Cameron *et al.* (1983) points out that the trace-element and isotope affinities between the Upper and Lower Lavas are so different that they must have had different magma sources.

Present-day systems, like the Izu-Bonin arc in the Southeast Pacific, comprise an outer-arc high, a boninitic fore-arc, an arc section and an actively spreading back-arc basin (Crawford *et al.*, 1981; Taylor *et al.*, 1992). It is less common to find coexisting back arc, arc, and fore-arc basin magmatism obducted in the same ophiolite (Taylor *et al.* 1992). In fact, there is no typical silicic or calc-alkaline arc-magmatism in the Lyngen Magmatic Complex, though an arc-section is present in the similar-age ophiolites in southern part of Norway (e. g. the Karmøy Ophiolite of Pedersen and Hertogen, 1990). Taylor *et al.* (1992) suggested that the tectonomagmatic conditions that produce boninites and tholeiites together is currently not active in the oceans today. It is often suggested that coexisting tholeiitic and boninitic magmatism is the result of incipient arcs, and part of the early arc of the Izu-Bonin arc is currently the outer-arc high. ODP-site 458, situated on the outer-arc high between the Mariana trench and –ridge (Hickey-Vargas, 1989) did find contemporaneously emplaced back-arc basin tholeiites and high-Ca boninite lavas. They suggest that the boninite parent-magmas were generated by high degrees of melting of a moderately depleted lithosphere residual from generating Philippine-sea MORB, and that hydrous melting at greater depth, possibly within the asthenosphere, generated the tholeiites. The close proximity to the trench make the oceanic crust more likely to be emplaced into the deeper levels on an orogenic belt, preserving them during erosion and orogenic collapse. The Svensby Formation, a part of the Balsfjord Group overlying the Lyngen Magmatic Complex, has calc-alkaline pillow-basalts dismembered from other arc rocks. It is possible that the closure of the large Iapetus Ocean may have caused several slivers of the supra-subduction zone to be

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obducted onto the continental margin, preserving only the Lyngen Magmatic Complex in this part of the Scandinavian Caledonides.

### ***Conclusion***

The Lyngen Gabbro represents the lower crustal section of an incipient arc or outer-arc high of an Ordovician oceanic *supra-subduction* zone. The Western suite was precipitated from magma that could have been derived from the same system as the associated Aksla Volcanics and Kjosen Greenschist. This primary magma was derived from a source similar to that of mid-ocean ridges, although hydrous crystal-fractionation suggests that the tectonomagmatic setting was an island-arc tholeiite produced at spreading-ridge with fluids derived from the subducted slab. The Eastern suite comprise cumulates that were crystallized from magmas similar to those of the ultra-depleted magmas of fore-arcs. Although the volcanic counterpart of the ultra-depleted Eastern suite has not been found, dikes crosscutting the gabbros may represent the magma that formed them. The oceanic Rypdal Shear Zone generally separates the two suites, but non-tectonic transitions from boninitic to island-arc tholeiitic affinities suggest the magmatism happened concurrently.

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Table 1: A: Whole-rock analyses of the Western suite of the Lyngen Gabbro. Fe<sub>2</sub>O<sub>3</sub>\* is all Fe measured as Fe<sub>2</sub>O<sub>3</sub>. Mg# is 100\*MgO/(MgO+0.899\*Fe<sub>2</sub>O<sub>3</sub>\*). FeO and Fe<sub>2</sub>O<sub>3</sub> was determined for some samples by wet chemistry. Abbreviations: gb=gabbro, tex=textured, pegm=pegmatite, ol=olivine, ox=magnetite and/or ilmenite, diss=dissolved, troct=troctolite, hbl=hornblende, qz=quartz, pyx=pyroxenite. “Place in profile” indicates the relative distance from east for each of the profiles. Strupen and Skaidevarri are individual profiles, Ellendalen, Goverdalen and Veidalen has a common profile.

Location	Isskardet											
Sample#	LY-IS-1	LY-IS-2	LY-IS-3	LY-IS-4	LY-IS-6	LY-IS-7	LY-IS-9	LY-IS-10	LY-IS-11	LY-IS-12	LY-IS-13	LY-IS-16
<b>Description</b>												
<b>Place in profile</b>												
<b>W&lt;=E</b>												
<b>Major elements (XRF)</b>												
SiO <sub>2</sub>	47.2	47.5	47.4	46.1	45.7	46.8	40.2	46.7	45.0	45.3	45.0	43.2
TiO <sub>2</sub>	.18	.23	.24	.20	.15	.54	.10	.21	.70	.29	.54	.83
Al <sub>2</sub> O <sub>3</sub>	21.4	17.7	17.9	19.5	21.3	12.3	16.8	21.5	19.7	21.0	20.7	18.1
Fe <sub>2</sub> O <sub>3</sub> *	7.9	7.7	7.7	8.7	8.3	11.3	14.3	7.0	10.8	9.6	10.4	13.7
MnO	0.14	0.16	0.15	0.15	0.14	0.14	0.19	0.19	0.13	0.14	0.14	0.16
MgO	8.6	9.3	9.1	9.6	8.9	11.3	13.8	7.8	7.5	9.1	8.2	9.2
CaO	14.6	16.1	16.2	14.7	14.5	16.4	10.3	15.8	15.3	14.0	14.6	13.8
Na <sub>2</sub> O	.93	.79	.84	.86	.91	.53	.56	.97	.91	.86	.87	.70
K <sub>2</sub> O	0	0	0	0	0	0	0	0	0	0	0	0
P <sub>2</sub> O <sub>5</sub>	.03	.01	.02	.02	.01	.02	.00	.01	.01	.00	.02	.00
LOI	0.8	1.4	0.6	0.9	0.4	0.8	1.3	0.9	1.2	1.0	0.4	0.8
Sum	101.8	100.8	100.1	100.6	100.2	100.2	97.6	101.0	101.2	101.2	100.8	100.5
<b>Trace elements (XRF)</b>												
V	136	194	209	151	122	401	72	190	697	265	498	787
Cr	33	47	58	32	26	266	29	108	144	111	113	105
Co	37	43	37	45	40	57	82	35	55	49	42	61
Ni	36	35	35	48	54	62	111	35	41	63	34	48
Cu	13	16	28	22	17	70	39	24	131	28	8	13
Zn	37	36	35	39	38	47	63	32	33	37	38	41
Rb	0	1	0	0	0	1	1	0	1	0	0	2
Sr	181	170	170	172	190	115	113	199	185	179	171	146
Y	5	7	6	6	4	10	4	6	6	6	4	6
Zr	15	13	15	14	12	18	13	13	14	12	14	13
Nb	3	2	3	2	2	2	2	2	2	3	2	2
Ba	0	0	0	1	1	1	0	3	6	0	4	0
La	10	14	18	8	8	20	3	11	15	12	9	12
Ce	9	7	7	6	1	3	3	6	10	6	5	4
Nd	68.7	70.5	70.2	68.8	68.1	66.6	65.8	69.1	57.9	65.4	61.1	57.4
Mg#	5.6	5.3	5.2	6.1	6.0	7.2	10.0	4.9	2.9	8.1	6.5	8.1
FeO	1.8	2.0	2.0	1.9	1.7	3.3	3.3	1.6	7.6	0.0	3.3	4.8
Fe <sub>2</sub> O <sub>3</sub>												

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Location	Isskardet													
Sample#	LY-IS-17	LY-IS-19	LY-IS-20	LY-IS-23	LY-IS-25	LY-IS-26	LY-IS-28	LY-IS-29	LY-IS-29a	LY-IS-30	LY-IS-31	LY-IS-33		
Description														
Place in profile W<=E														
<b>Major elements</b>														
SiO <sub>2</sub>	46.6	47.1	46.6	45.6	43.5	44.3	46.8	46.8	46.8	45.4	44.7	47.5		
TiO <sub>2</sub>	.32	.24	.46	.58	.81	.75	.19	.48	.20	.79	.93	.13		
Al <sub>2</sub> O <sub>3</sub>	16.3	17.1	16.2	21.2	18.4	19.2	20.3	22.8	19.6	18.2	15.3	31.6		
Fe <sub>2</sub> O <sub>3</sub> *	9.1	8.4	10.0	9.1	12.3	12.0	8.1	7.6	8.3	11.1	14.3	3.4		
MnO	0.16	0.16	0.17	0.13	0.15	0.14	0.12	0.15	0.14	0.17	0.17	0.06		
MgO	10.0	9.8	9.8	6.8	8.1	8.2	8.9	6.1	9.2	7.6	9.5	3.8		
CaO	15.8	16.1	15.9	15.7	14.5	14.7	15.0	16.1	15.0	15.7	15.1	15.0		
Na <sub>2</sub> O	.73	.74	.73	1.10	.79	.84	.99	1.05	.92	.83	.66	2.24		
K <sub>2</sub> O	0	0	0	0	0	0	0	0	0	0	0	0		
P <sub>2</sub> O <sub>5</sub>	.00	.02	.02	.03	.02	.02	.03	.02	.01	.02	.00	.01		
LOI	1.2	0.5	0.1	0.2	0.8	0.1	0.4	0.3	0.5	0.2	0.0	2.8		
<b>Sum</b>	100.2	100.0	99.9	100.5	99.3	100.2	100.8	101.3	100.6	99.9	100.5	106.5		
<b>Trace elements †</b>														
V	254	202	386	531	704	705	163	462	181	700	853	131		
Cr	70	76	329	122	70	67	47	36	41	92	201	41		
Co	45	44	46	37	53	49	43	31	44	46	64	23		
Ni	45	59	34	45	17	17	31	12	33	98	58	12		
Cu	20	204	22	15	13	14	13	14	17	449	221	11		
Zn	38	34	37	30	33	38	36	26	35	36	47	20		
Rb	0	0	1	0	0	0	0	0	0	0	0	0		
Sr	138	157	148	198	157	166	189	222	184	169	133	281		
Y	7	7	7	6	5	5	6	5	7	8	4			
Zr	15	15	16	12	15	13	12	11	12	14	14	7		
Nb	3	2	1	2	3	2	2	2	2	3	2	2		
Ba	5	0	6	1	1	0	0	0	11	0	0	1		
La														
S														
Ce	14	13	17	15	9	15	13	15	13	13	20	7		
Nd	6	4	3	3	7	7	12	7	10	9	5	4		
Mg#	68.8	70.1	66.2	60.0	56.7	57.7	68.6	61.5	68.8	57.6	57.0	69.0		
FeO	5.9	11.8	6.4	6.0	6.8	7.0	5.7	4.6	6.1	6.3	8.5	2.1		
Fe <sub>2</sub> O <sub>3</sub>	2.5	4.6	3.0	2.5	4.8	4.2	1.8	2.6	1.6	4.2	4.9	0.9		

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Location	Ellendalen										
Sample#	92LY93	92LY119	93LY20	93LY19	94LY20	94LY21	94LY23	94LY25	94LY26	94LY24	94LY39
Description	metagb	metagb	metagb	massiv gb	gabbro	gabbro	gabbro	varitex gb	varitex gb	gabbro	gabbro
Place in profile	W<=E	9350	9350	8600	8500	7450	7400	7150	7000	6950	6800
Major elements											
SiO <sub>2</sub>	54.0	51.3	48.4	47.9	47.7	45.5	42.4	36.2	38.4	47.1	46.6
TiO <sub>2</sub>	.31	.50	.09	.13	.25	.17	.03	2.88	2.24	.24	.22
Al <sub>2</sub> O <sub>3</sub>	12.8	15.1	20.2	20.3	15.3	20.3	3.8	8.2	10.6	14.8	19.6
Fe <sub>2</sub> O <sub>3</sub> *	10.4	10.1	5.2	5.4	7.5	7.0	13.0	27.7	24.1	8.0	10.3
MnO	.17	.19	.10	.10	.16	.14	.19	.24	.20	.16	.15
MgO	9.4	7.9	7.7	7.3	9.9	7.9	28.2	9.2	8.4	10.5	10.3
CaO	7.9	11.2	15.5	16.4	14.5	15.0	9.7	11.4	12.7	16.0	10.1
Na <sub>2</sub> O	3.95	2.78	1.11	1.20	1.00	1.17	0.00	1.44	1.13	0.39	2.33
K <sub>2</sub> O	.00	.16	.58	.00	.00	.00	.00	.03	.04	.00	.02
P <sub>2</sub> O <sub>5</sub>	.01	.08	.00	.01	.00	.02	.01	.02	.00	.01	.01
LOI	2.2	2.1	2.3	2.5	2.5	2.9	1.5	1.8	2.5	1.5	1.1
Sum	101.0	101.3	101.2	101.1	98.8	99.9	98.6	98.7	100.3	98.6	100.6
Trace elements †											
V	229	191	96	110	144	124	648	600	667	165	128
Cr	281	29	104	179	180	181	33	49	130	193	103
Co	32	43	34	37	48	47	47	68	78	57	45
Ni	128	37	44	47	49	51	11	25	64	71	65
Cu	5	41	29	17	9	7	25	120	164	55	77
Zn	75	58	37	32	79	48	83	118	98	72	65
Rb	0	1	18	1	0	0	0	0	2	0	0
Sr	89	198	150	177	137	207	435	64	100	120	122
Y	9	25	5	6	7	6	6	17	11	7	2
Zr	8	39	10	11	11	15	18	13	9	7	9
Nb	3	5	1	2	0	2	3	3	2	2	4
Ba	63	86	188	10	6	6	0	0	32	0	34
La	2	5					0	0	8	9	13
S							0.02	0.07		0.04	0.03
Ce	13	19	17	21	8	18	12	14	18	11	11
Nd	0	4	11	14	7	10	8	10	10	7	8
Mg#	64.4	60.9	74.7	72.8	72.6	69.2	81.2	39.8	40.9	72.4	66.6
FeO											33.7
Fe <sub>2</sub> O <sub>3</sub>											0.26

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

**Table 1A: Continued.**

Location	Ellendalen	Ellendalen	Ellendalen	Ellendalen	Ellendalen	Goverdalen									
Sample#	94LY38	94LY37	94LY36	94LY35	94LY34	94LY33	95LY1	95LY2	95LY3	95LY5	95LY24	95LY23			
Description	melagb	melagb	gabbro	gabbro	gabbro	leucogb	massiv gb	massiv gb	gabbro	microgb	melagb	leucogb			
Place in profile	W<=E	5250	5220	5190	5160	5130	5000	4550	4550	4550	3849	3839			
<b>Major elements</b>															
SiO <sub>2</sub>	41.7	39.2	40.2	41.9	41.7	44.9	44.1	43.0	47.7	39.6	36.7	48.0			
TiO <sub>2</sub>	1.47	2.34	1.32	1.34	1.00	.29	.07	.06	.08	2.14	2.31	.36			
Al <sub>2</sub> O <sub>3</sub>	9.0	9.8	13.3	16.6	15.9	23.3	16.1	15.1	15.6	13.6	13.3	19.3			
Fe <sub>2</sub> O <sub>3</sub> *	22.9	22.9	17.7	14.4	14.0	6.1	5.0	5.9	4.1	21.2	23.5	8.3			
MnO	.23	.23	.18	.18	.10	.08	.11	.08	.21	.18	.21	.16			
MgO	13.6	9.4	10.3	9.0	8.5	4.7	13.9	16.0	10.5	7.3	6.8	7.5			
CaO	11.3	12.5	12.8	14.6	13.5	13.0	16.7	15.5	16.7	12.8	12.4	13.7			
Na <sub>2</sub> O	0.96	0.97	1.06	1.12	1.66	2.94	0.39	0.28	1.19	0.43	0.64	1.80			
K <sub>2</sub> O	.03	.05	.02	.00	.00	.01	.00	.00	.01	.00	.01	.00			
P <sub>2</sub> O <sub>5</sub>	.02	.03	.00	.02	.01	.03	.01	.00	.01	.01	.02	.01			
LOI	3.5	2.6	3.2	2.7	2.9	2.9	3.6	3.6	2.6	3.1	2.8				
<b>Sum</b>	101.9	99.9	100.2	102.0	99.3	98.2	99.2	99.4	100.6	99.9	98.9	98.3			
<b>Trace elements †</b>															
V	550	691	536	553	492	194	80	63	101	895	1126	172			
Cr	84	92	194	200	145	48	450	359	971	40	65	65			
Co	70	73	67	69	63	33	44	46	32	73	62	44			
Ni	47	88	23	22	33	18	207	258	81	32	45	25			
Cu	6	306	14	55	11	15	13	5	8	147	174	14			
Zn	101	89	77	63	70	50	24	33	22	99	106	47			
Rb	0	2	0	0	0	2	0	0	0	0	0	0			
Sr	56	108	127	160	153	177	208	126	61	133	119	173			
Y	9	10	7	7	4	0	2	4	9	8	6				
Zr	12	14	13	14	12	12	10	9	7	11	12	11			
Nb	2	4	0	2	3	0	6	4	5	6	5	5			
Ba	0	71	5	85	0	69	57	0	55	42	66				
La	5	8	13	0	5	0	0	0	0	0	0	0			
S															
Ce	12	14	8	19	18	26	18	10	18	16	7	3			
Nd	8	6	0	8	11	8	12	5	4	6	6	0			
Mg#	54.2	45.0	53.8	55.5	54.7	60.8	84.8	84.5	83.5	40.7	36.5	64.5			
FeO															
Fe <sub>2</sub> O <sub>3</sub>															

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Location	Goverdalen			Goverdalen														
Sample#	95LY8	95LY6	94LY67	94LY68	95LY43	95LY42	95LY41	95LY29	95LY30	95LY31	95LY32	95LY33						
Description	leucogb	melagb	gabbro	gabbro	gabbro	melagb	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	leucogb	melagb	gabbro	pegmatite		
Place in profile	W<=E	3705	3701	3500	3500	3100	3100	3000	3000	2880	2740	2560	2550	2550	2547			
Major elements																		
SiO <sub>2</sub>	44.4	38.7	47.1	46.0	45.7	39.8	47.0	39.6	39.8	43.3	43.7	47.1						
TiO <sub>2</sub>	.14	1.36	.25	.13	.69	2.20	.17	.04	1.42	.95	.87	.68						
Al <sub>2</sub> O <sub>3</sub>	26.9	15.7	16.2	14.6	23.9	13.0	17.2	12.6	14.0	16.5	13.0	10.1						
Fe <sub>2</sub> O <sub>3</sub> *	5.8	18.1	7.5	9.0	9.3	22.2	9.0	20.9	22.1	14.2	15.7	16.8						
MnO	.09	.16	.15	.15	.13	.23	.16	.22	.20	.19	.20	.23						
MgO	5.0	7.0	9.1	12.0	5.0	8.0	10.0	7.7	10.1	8.0	9.4	11.3						
CaO	15.4	14.5	16.3	15.7	12.1	11.1	12.2	12.5	8.4	13.5	13.2	11.8						
Na <sub>2</sub> O	1.43	0.55	0.68	0.22	2.02	0.94	1.02	0.68	0.81	0.91	0.51	1.30						
K <sub>2</sub> O	.00	.00	.00	.00	.00	.00	.00	.00	.06	.04	.03	.02						
P <sub>2</sub> O <sub>5</sub>	.01	.01	.01	.01	.03	.01	.02	.00	.00	.01	.01	.01						
LOI	3.1	3.2	1.8	2.5	2.8	1.7	3.1	2.2	3.3	2.7	2.6	1.0						
Sum	102.2	99.2	99.0	100.2	101.5	99.2	100.0	98.6	100.2	100.3	99.2	100.3						
Trace elements †																		
V	93	844	148	172	343	923	130	905	1137	657	626	417						
Cr	33	74	741	402	37	76	67	46	73	170	141	11						
Co	41	68	62	61	43	55	45	67	41	62	67	64						
Ni	30	38	70	96	26	49	40	34	76	88	79	51						
Cu	9	56	62	184	56	159	29	139	86	123	170	13						
Zn	37	79	85	43	62	118	67	108	98	78	79	70						
Rb	0		0	0	0	2	0	0	0	0	0	0						
Sr	189	160	119	70	238	128	213	125	61	150	126	88						
Y	3	5	4	0	7	10	6	9	14	5	9	43						
Zr	11	10	8	5	14	14	13	11	14	9	10	47						
Nb	3	6	4	5	5	4	5	4	5	6	5	7						
Ba	31	54	28	61	81	54	58	93	30	36	36	53						
La	0	5	4	0	0	9	0	0	0	0	0	0						
S	0.02	0.11	0.02	0.03	0.14	0.37	0.02	0.23	0.09	0.16	0.36	0.02						
Ce	8	16	19	11	13	6	12	7	15	12	14	9						
Nd	8		7	6	8	5	7	4	5	3	0	6						
Mg#	63.3	43.5	70.7	72.7	51.9	41.7	68.9	42.4	47.7	53.0	54.5	57.2						
FeO																		
Fe <sub>2</sub> O <sub>3</sub>																		

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Location	Goverdalen						Goverdalen						Goverdalen						Veidalen						Veidalen					
Sample#	95LY34	95LY35	95LY36	95LY37	95LY38	95LY39	95LY40	95LY127	95LY126	95LY125	95LY124	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122	95LY122
Description	metagb	metagb	gabbro	gb pegm	gabbro	metagb	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	metagb	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	gabbro	
Place in profile	W<=E	253.5	2420	2290	2290	2280	2215	2190	2080	2050	2030	2000	1900	1900	1900	1900	1900	1900	1900	1900	1900	1900	1900	1900	1900	1900	1900	1900	1900	
Major elements																														
SiO <sub>2</sub>	43.2	43.5	46.1	46.1	48.8	45.6	47.1	46.40	46.37	47.75	43.82	47.38																		
TiO <sub>2</sub>	1.49	1.65	1.09	.73	1.57	.13	0.1	0.17	0.44	0.78	0.15																			
Al <sub>2</sub> O <sub>3</sub>	11.4	13.7	17.3	19.7	15.0	16.2	17.5	16.13	17.69	14.78	15.65	15.47																		
Fe <sub>2</sub> O <sub>3</sub> *	22.9	20.9	14.7	13.0	16.0	6.9	5.5	7.75	8.26	13.49	14.42	7.35																		
MnO	.22	.22	.22	.18	.22	.12	.11	0.13	0.15	0.24	0.19	0.15																		
MgO	8.6	7.9	6.2	5.8	6.1	12.2	10.1	11.26	10.19	10.58	8.45	10.28																		
CaO	12.4	12.4	11.0	12.5	10.0	16.3	16.6	15.24	15.97	12.75	13.59	15.73																		
Na <sub>2</sub> O	0.76	0.86	1.82	1.79	2.87	0.69	0.70	0.56	0.48	0.52	0.8	0.52																		
K <sub>2</sub> O	.00	.00	.00	.02	.01	.00	.02	.002	0	0	0.08	0																		
P <sub>2</sub> O <sub>5</sub>	.00	.02	.02	.00	.10	.01	.01	.01	.01	.01	0	0																		
LOI	0.9	0.9	2.1	2.0	0.4	1.4	2.2	2.34	1.22	0.27	1.95	2.25																		
Sum	101.8	102.0	100.5	101.6	100.9	99.5	100.0	99.95	100.51	100.82	99.73	99.75																		
Trace elements †																														
V	113.4	92.7	43.3	44.2	45.2	106	114	101	140	280	731	154																		
Cr	77	65	44	32	28	216	395	185	71	94	50	88																		
Co	71	65	51	48	47	49	43	52	55	56	57	48																		
Ni	77	26	27	22	15	101	126	73	38	38	24	44																		
Cu	242	150	104	132	64	178	192	135	78	107	148	33																		
Zn	60	77	83	79	81	33	29	33	46	80	70	49																		
Rb	0	0	0	0	0	0	0	2	0	0	0	0																		
Sr	97	130	202	212	178	107	116	119	125	112	118	101																		
Y	11	13	12	4	23	6	3	3	6	8	8	6																		
Zr	11	11	14	10	22	7	8	8	9	8	9	6																		
Nb	7	6	4	6	4	5	4	3	3	6	5	3																		
Ba	47	31	24	92	109	36	0	53	83	65	82	46																		
La	0	4	5	5	0	0	0	0	7	6	0	10																		
S	0.79	0.46	0.39	0.51	0.20	0.24	0.10	0.19	0.16	0.22	0.33	0.03																		
Ce	14	18	8	6	8	15	8	12	13	9	13	15																		
Nd	5	6	10	5	6	10	5	6	10	11	3	0																		
Mg#	42.9	43.0	45.8	47.1	43.0	77.9	78.4	74.3	71.1	61.0	53.9	73.6																		
FeO																														

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Location	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen
Sample#	95LY123	95LY110	95LY109	95LY108	95LY107	95LY106	95LY105	95LY104	95LY102	95LY99	95LY97	95LY96
Description	metagabbro	metagb	metagb	metagb	gabbro	gabbro	gabbro	gabbro	metagb	metagb	metagb	metagb
Place in profile	W<=E	1900	1598	1588	1576	1574	1560	1554	1552	1542	1497	1493
Major elements												
SiO <sub>2</sub>	39.81	46.40	48.14	46.85	45.32	48.79	46.90	46.62	46.60	46.16	46.19	46.52
TiO <sub>2</sub>	1.16	0.07	0.09	0.09	0.07	0.1	0.06	0.1	0.06	0.07	0.08	0.07
Al <sub>2</sub> O <sub>3</sub>	11.9	15.39	14.12	14.42	17.72	10.74	15.42	12.49	15.7	16.88	15.95	17.38
Fe <sub>2</sub> O <sub>3</sub> *	22.7	4.81	5.25	5.3	5.8	6.79	5.16	5.62	4.85	5.89	6.3	6.66
MnO	0.17	0.1	0.09	0.11	0.11	0.14	0.1	0.1	0.09	0.11	0.13	0.12
MgO	8.44	12	12.17	11.94	11.61	14.32	13.06	15.21	11.17	11.25	11.57	11.2
CaO	12.79	18.49	17.41	18.62	17.38	18.79	17.41	16.47	18.74	16.71	17.05	17.4
Na <sub>2</sub> O	0.57	0.23	0.3	0.26	0.38	0.15	0.3	0.14	0.23	0.24	0.2	0.36
K <sub>2</sub> O	0.11	0	0	0	0	0	0	0	0	0	0	0
P <sub>2</sub> O <sub>5</sub>	0.02	0.01	0.01	0.01	0.01	0.02	0	0	0.01	0.02	0.01	0.01
LOI	2.25	2.22	1.78	1.95	1.432	0.42	1.51	2.91	2	2.71	2.483	1.501
Sum	99.9	99.7	99.35	99.55	99.83	100.24	99.93	99.65	99.46	100.03	99.94	101.22
Trace elements (												)
V	976	105	124	129	108	146	100	104	112	107	114	105
Cr	46	861	1317	597	143	890	345	839	616	248	84	226
Co	78	41	42	44	43	51	42	44	41	45	48	49
Ni	12	103	87	72	54	57	56	204	121	86	41	113
Cu	66	77	73	98	159	49	80	211	137	52	135	173
Zn	78	21	14	21	16	18	25	17	15	29	25	21
Rb	0	0	0	0	0	0	0	0	0	0	0	0
Sr	87	72	68	63	89	45	69	57	83	104	94	101
Y	11	0	2	3	0	0	0	2	0	0	3	0
Zr	9	5	6	6	6	5	6	6	6	8	8	7
Nb	4	4	5	3	5	4	3	3	4	5	4	5
Ba	72	45	35	27	45	35	51	67	48	31	53	44
La	0	8	0	15	0	0	7	0	9	0	8	0
S	0.03	0.03	0.02	0.03	0.04	0.03	0.03	0.08	0.03	0.03	0.04	0.07
Ce	7	13	17	19	13	13	21	20	18	14	17	13
Nd	6	6	3	5	0	6	7	8	0	0	5	8
Mg#	42.6	83.3	82.2	81.8	80.0	80.8	83.5	84.4	82.1	79.2	78.6	77.0
FeO												
Fe <sub>2</sub> O <sub>3</sub>												

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Location	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen	Veidalen
Sample#	95LY95	95LY94	95LY93	95LY91	95LY90	95LY88	95LY87	95LY81	95LY79	95LY76	95LY75	95LY73	
Description	gabbro	gabbro	gabbro	metagabbro	gabbro	metagabbro	gabbro	gabbro	ferrogabbro	gabbro	ferrogabbro	leucogabbro	metagabbro
Place in profile	W<=E	1480	1468	1460	1452	1448	1432	1429	1370	1370	1350	1350	1340
Major elements													
SiO <sub>2</sub>	48.50	46.15	45.39	43.84	45.90	44.19	50.43	38.50	44.00	41.67	44.05	47.41	
TiO <sub>2</sub>	0.07	0.09	0.07	0.75	0.12	0.05	1.14	0.06	0.83	0.06	0.09		
Al <sub>2</sub> O <sub>3</sub>	19.19	14.65	18.9	15.46	14.83	21.26	4.81	14	21.77	15.72	25.08	17.22	
Fe <sub>2</sub> O <sub>3</sub> *	6.38	7.96	6.78	16.07	8.59	6.88	5.92	21.38	6.03	15.42	5.73	6.11	
MnO	0.13	0.16	0.13	0.2	0.15	0.11	0.11	0.16	0.1	0.16	0.1	0.12	
MgO	10.15	12.24	10.32	8.63	12.03	6.97	20.73	7.96	9.35	8.32	6.64	11.41	
CaO	15.58	17.38	16.77	13.54	16.12	18.87	18.04	13.26	16.51	14.83	14.96	14.59	
Na <sub>2</sub> O	0.34	0.39	0.26	1	0.25	0.28	0.12	0.6	0.68	0.39	1.01	0.54	
K <sub>2</sub> O	0	0.01	0	0.12	0	0.02	0.01	0.15	0.04	0.05	0	0.59	
P <sub>2</sub> O <sub>5</sub>	0.01	0.02	0.02	0	0	0.01	0	0.01	0.02	0	0.01	0.01	
LOI	0.345	1.72	1.16	1.296	2.178	2.14	0.6227	1.73	2.997	2.631	3.311	2.55	
Sum	100.68	100.77	99.78	100.9	100.16	100.86	100.85	98.88	101.56	100	100.094	100.65	
Trace elements (ppm)													
V	98	140	111	1321	136	169	111	1159	82	829	57	121	
Cr	143	350	142	81	259	156	2496	61	100	66	50	565	
Co	41	57	50	72	59	49	52	72	45	69	40	41	
Ni	57	80	77	38	103	40	260	39	63	44	28	102	
Cu	155	229	134	692	111	84	12	590	6	246	12	35	
Zn	26	41	35	59	38	18	20	43	30	45	26	33	
Rb	0	3	0	7	5	0	0	0	0	0	0	45	
Sr	108	91	112	132	95	143	13	92	136	140	168	91	
Y	2	3	3	3	3	3	0	0	0	4	2	4	
Zr	7	6	6	8	7	8	5	7	8	9	10	7	
Nb	5	5	4	4	4	4	4	6	3	7	4	5	
Ba	59	40	53	120	51	37	27	75	59	48	32	169	
La	0	5	4	0	5	0	6	0	5	9	5	0	
S	0.07	0.08	0.12	0.25	0.09	0.02	0.03	0.56	0.02	0.41	0.02	0.03	
Ce	21	16	21	14	15	17	18	10	21	7	12	10	
Nd	9	6	11	5	4	7	5	4	4	6	4	0	
Mg#	76.0	75.4	75.2	51.7	73.6	66.9	87.5	42.6	75.6	51.8	69.8	78.8	
FeO													

Table 1A: Continued.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Location	Veidalen	Skaidevarri										
Sample#	95LY70	95LY68	95LY67	95LY65	95LY64	95LY63	95LY61	95LY60	95LY59	95LY58	95LY57	90LY141
Description	metagb	Oligbnorite										
Place in profile	W<=E	990	960	910	860	810	770	720	620	570	550	10.5
<b>Major elements</b>												
SiO <sub>2</sub>	45.75	45.98	47.33	47.91	46.20	43.03	48.01	47.16	46.92	40.66	36.90	46.72
TiO <sub>2</sub>	0.12	0.13	0.41	0.11	0.08	0.89	0.17	0.1	0.07	0.04	1.03	0.17
Al <sub>2</sub> O <sub>3</sub>	15.38	14.56	11.83	12.28	22.31	16.3	13.83	15.51	16.03	22.79	19.51	18.74
Fe <sub>2</sub> O <sub>3</sub> *	9.11	8.98	11.16	9.11	6.35	15.25	8.74	7	6.82	7.33	16.7	7.19
MnO	0.16	0.15	0.27	0.19	0.13	0.15	0.18	0.14	0.14	0.11	0.19	0.12
MgO	10.8	12.01	11.54	11.93	7.97	8.45	11.07	11.47	11.86	9.45	7.33	9.65
CaO	16.07	15.68	14.92	15.98	14.28	14.73	15.76	15.25	16.17	15.91	13.55	15.4
Na <sub>2</sub> O	0.21	0.22	0.31	0.35	0.5	0.22	0.42	0.46	0.29	0.57	0.49	0.82
K <sub>2</sub> O	0.02	0	0	0.03	0.34	0.03	0.04	0.17	0.11	0.09	0.09	0
P <sub>2</sub> O <sub>5</sub>	0.01	0.01	0.01	0	0	0.01	0	0.01	0.01	0	0.01	0
LOI	2.3	2.49	2.2	1.46	3.31	2.57	1.621	2.595	1.98	3.05	3.592	1.78
<b>Sum</b>	99.93	100.2	99.98	99.34	101.55	101.63	99.84	99.88	100.39	100	99.39	100.6
<b>Trace elements (ppm)</b>												
V	174	172	148	163	111	980	229	141	123	68	698	105
Cr	151	402	741	672	42	108	533	449	449	260	53	207
Co	60	61	62	57	39	69	56	45	46	47	62	45
Ni	58	96	70	69	33	34	82	76	108	57	21	84
Cu	44	184	62	20	37	168	24	94	48	9	103	118
Zn	45	43	85	76	41	47	64	36	28	36	74	22
Rb	0	0	0	0	30	0	0	13	7	6	11	2
Sr	93	70	119	77	123	118	105	86	119	133	153	148
Y	0	0	4	5	0	0	4	2	2	0	0	6
Zr	8	5	8	6	8	7	7	7	7	9	9	11
Nb	4	5	4	4	4	4	5	4	4	5	5	0
Ba	0	61	28	37	138	29	52	39	40	46	74	32
La	0	0	4	0	0	7	0	0	5	0	0	0
S	0.03	0.03	0.02	0.02	0.03	0.33	0.02	0.03	0.03	0.02	0.07	0.07
Ce	9	11	19	14	6	10	17	16	13	10	12	18
Nd	0	6	7	10	0	3	7	7	6	5	3	
Mg#	70.3	72.7	67.3	72.3	71.5	52.5	71.6	76.6	77.6	72.0	46.7	72.81
FeO												3.57
Fe <sub>2</sub> O <sub>3</sub>												3.26

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Location	Skaidevarri										
Sample#	90LY142	90LY143	90LY144	90LY145	90LY146	90LY147	90LY148	90LY149	90LY150	90LY151	90LY152
Description	Oligonorite	Oligonorite	Oligb	gnorite	Altered	Gabbro	Gabbro	Gabbro	Gabbro	Gabbro	Gabbro
Place in profile	W<E	12	13.5	15	16.5	18	19.5	22	23.5	25	28
Major elements											
SiO <sub>2</sub>	47.26	47.16	47.67	45.57	48.69	46.29	47.14	46.51	48.33	46.97	47.26
TiO <sub>2</sub>	0.18	0.18	0.19	0.14	0.24	0.15	0.17	0.31	0.17	0.17	0.17
Al <sub>2</sub> O <sub>3</sub>	19.41	16.53	17.15	18.15	14.16	18.49	21.19	18.59	8.48	18.98	19.41
Fe <sub>2</sub> O <sub>3</sub> *	7.21	8.1	7.81	8.81	7.08	6.91	5.05	6.87	9	6.08	5.1
MnO	0.13	0.13	0.13	0.14	0.15	0.12	0.11	0.13	0.17	0.12	0.11
MgO	9.31	10.85	10.33	10.51	10.04	9.64	7.15	9.95	13.65	8.33	7.12
CaO	15.62	15.73	15.79	14.24	17.86	15.51	16.82	15.04	16.34	16.61	17.95
Na <sub>2</sub> O	0.83	0.5	0.56	0.62	0.78	0.66	0.83	0.85	0.33	0.97	0.91
K <sub>2</sub> O	0	0	0	0.01	0	0.01	0.01	0.04	0	0.02	0
P <sub>2</sub> O <sub>5</sub>	0.01	0.02	0.02	0.02	0	0.01	0	0.01	0	0.01	0
LOI	1.25	1.77	1.58	2.8	1.84	2.46	2.55	2.94	4.64	2.35	2.33
Sum	101.2	100.97	101.25	101.01	100.85	100.26	100.99	101.1	101.24	100.62	100.36
Trace elements †											
V	104	132	134	86	176	100	131	131	207	129	158
Cr	13.8	302	209	101	205	163	231	235	580	250	245
Co	46	52	51	52	46	43	34	47	58	42	36
Ni	81	89	106	48	102	78	56	92	104	81	52
Cu	11.5	197	140	44	195	47	173	114	201	156	155
Zn	26	22	27	28	26	20	27	33	31	17	12
Rb	0	0	0	0	0	2	0	3	2	0	0
Sr	147	115	118	127	115	132	157	143	61	143	166
Y	7	7	8	7	9	7	6	8	11	8	8
Zr	12	9	11	10	12	11	12	12	11	11	11
Nb	2	3	0	0	0	0	0	0	0	0	2
Ba	34	0	28	54	29	20	56	36	77	34	0
La											
S											
Ce	12	18	14	24	26	18	24	27	22	14	16
Nd											
Mg#	72.03	72.77	72.52	70.41	73.88	73.56	73.85	74.29	75.16	73.21	73.58
FeO	4.15	4.76	4.28	5.92	4.32	4.56	3.20	4.45	5.81	3.81	3.16
Fe <sub>2</sub> O <sub>3</sub>	2.65	2.86	3.10	2.30	2.33	1.89	1.53	1.98	2.61	1.89	1.62

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Location	Skaidevari											
Sample#	90LY153	90LY154	90LY155	90LY156	90LY157	90LY158	90LY159	90LY160	90LY161	90LY162	90LY163	
Description	Gabbro	Gabbro	Gabbro	Gabbro	Gabbro	Olgbnorite	gbnorite	gbnorite	Olgbnorite	Olgbnorite	Olgbnorite	Diss oxol gb
Place in profile	W<=E	29.5	31	32	33.5	35	36.5	38	39.5	41	42.5	44
<b>Major elements</b>												
SiO <sub>2</sub>	48.15	46.78	48.88	46.52	48.32	49.18	47.39	46.07	48.17	47.15	48.09	
TiO <sub>2</sub>	0.22	0.17	0.24	0.15	0.22	0.26	0.22	0.13	0.26	0.17	0.21	
Al <sub>2</sub> O <sub>3</sub>	14.79	17.18	13.84	16.93	17.23	11.16	14.91	16.53	12.79	18.36	14.77	
Fe <sub>i</sub> O <sub>3</sub> *	7.19	6.99	7.16	7.21	5.81	9.58	7.06	7.37	7.9	7.2	7.71	
MnO	0.14	0.14	0.17	0.14	0.13	0.19	0.14	0.13	0.15	0.12	0.15	
MgO	10.37	9.84	10.38	10.56	8.52	13.09	9.99	10.79	10.99	9.47	10.9	
CaO	17.27	16.38	17.18	16.08	17.47	15.52	16.66	16.31	17.42	15.6	15.97	
Na <sub>2</sub> O	0.45	0.67	0.98	1.04	1.22	0.42	0.77	0.48	0.51	0.61	0.55	
K <sub>2</sub> O	0	0.02	0.02	0.05	0.02	0	0.01	0	0	0	0	
P <sub>2</sub> O <sub>5</sub>	0	0	0	0	0	0	0	0	0	0	0	
LOI	1.94	2.35	1.8	2.74	1.92	1.11	1.4	2.68	1.01	1.81	1.72	
<b>Sum</b>	100.52	100.52	100.65	101.44	100.86	100.5	98.54	100.51	99.21	100.49	100.08	
<b>Trace elements</b>												
V	173	136	193	123	183	193	173	112	199	119	147	
Cr	275	265	334	266	439	607	340	175	617	293	428	
Co	50	47	46	50	40	55	48	51	51	49	50	
Ni	72	79	46	75	63	83	55	77	55	82	92	
Cu	107	44	13	120	66	103	80	52	23	130	101	
Zn	28	25	27	24	13	35	21	26	21	24	30	
Rb	0	0	0	3	0	2	0	0	0	0	0	
Sr	111	140	121	126	131	76	121	120	96	128	129	
Y	10	7	10	7	11	9	8	7	10	7	9	
Zr	10	10	11	10	12	9	10	10	9	9	11	
Nb	3	2	2	2	3	4	3	0	3	2	0	
Ba	0	62	41	77	30	93	0	58	30	30	52	
La												
Ce	24	15	23	21	28	19	21	24	16	16	22	
Nd												
Mg <sup>#</sup>	74.21	73.74	74.31	74.50	74.52	73.16	73.84	74.49	73.51	72.40	73.82	
FeO	4.76	4.52	4.62	4.62	3.67	5.71	4.62	4.93	4.76	4.39	4.69	
Fe <sub>2</sub> O <sub>3</sub>	1.95	2.02	2.08	2.13	1.77	3.30	1.98	1.95	2.66	2.37	2.55	

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Location Sample#	Skaidevarri 90LY164	Skaidevarri 90LY165	Skaidevarri 90LY166	Skaidevarri 90LY167	Skaidevarri 90LY168	Skaidevarri 90LY169	Skaidevarri 90LY170	Skaidevarri 90LY171	Skaidevarri 90LY172	Skaidevarri 90LY173	Skaidevarri 90LY174
Description	Gabbro	Oligonorite	Oligbonorite	Altered	Olivinegb	Oligbonrite	Altered	Altered	gnnorite	Oligbonrite	Altered
Place in profile W<=E	45.5	47	48.5	50	51.5	53	54.5	56	57.5	59.5	61
<b>Major elements</b>											
SiO <sub>2</sub>	47.13	47.56	47.09	47.78	45.49	47.4	47.98	47.17	47.33	44.58	47.04
TiO <sub>2</sub>	0.18	0.24	0.24	0.21	0.16	0.2	0.22	0.13	0.16	0.13	0.14
Al <sub>2</sub> O <sub>3</sub>	14.38	13.19	15.27	19.39	16.45	17.82	18.33	19.52	20.63	19.05	21.21
Fe <sub>2</sub> O <sub>3</sub> *	7.14	8.27	8.27	5.95	9.74	7.04	6.23	5.35	5.89	8.26	5.92
MnO	0.14	0.16	0.16	0.12	0.15	0.13	0.13	0.1	0.1	0.13	0.11
MgO	11.64	11.31	9.88	7.55	12.04	9.19	8.25	8.47	7.69	9.83	7.92
CaO	17.3	16.68	16.53	16.85	14.04	15.99	17.06	17.21	16.08	14.03	14.14
Na <sub>2</sub> O	0.4	0.46	0.69	1.15	0.49	0.72	0.8	0.86	0.88	0.59	0.85
K <sub>2</sub> O	0	0	0.01	0.03	0	0.02	0.03	0.03	0.02	0.02	0.02
P <sub>2</sub> O <sub>5</sub>	0.03	0.02	0	0	0.01	0	0	0.01	0.02	0.01	0
LOI	2.56	1.88	1.38	1.92	2.14	2.03	1.93	2.52	1.84	2.93	3.14
<b>Sum</b>	100.9	99.77	99.52	100.95	100.7	100.55	100.96	101.38	100.64	98.56	100.48
<b>Trace elements</b>											
V	147	169	160	145	94	132	171	128	112	68	96
Cr	271	421	256	83	98	153	158	174	149	137	108
Co	51	53	52	39	58	46	43	37	39	49	34
Ni	84	83	90	47	85	71	40	44	41	80	25
Cu	118	116	96	109	112	95	31	25	17	85	9
Zn	21	27	24	18	32	22	15	13	15	26	22
Rb	2	0	3	3	0	3	3	2	0	2	0
Sr	85	102	130	160	131	142	179	174	150	132	161
Y	7	8	9	8	6	7	9	5	5	6	6
Zr	10	10	11	12	9	12	13	12	11	12	11
Nb	2	2	0	3	0	0	2	0	0	0	0
Ba	43	34	0	21	73	0	0	79	0	66	38
La											
S											
Ce	21	13	23	9	8	32	25	22	14	18	17
Nd											
Mg#	76.48	73.18	70.44	71.68	71.15	72.25	72.54	75.95	72.26	70.36	72.74
FeO	4.42	5.17	4.76	3.81	6.02	4.28	4.42	3.50	3.88	5.20	3.88
Fe <sub>2</sub> O <sub>3</sub>	2.28	2.58	3.03	1.76	3.12	2.33	1.37	1.50	1.62	2.54	1.65

Table 1A: Continued.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Location	Skaidevarri											
Sample#	90LY175	90LY176	90LY177	90LY178	90LY179	90LY180	90LY181	90LY182	90LY183	90LY184	90LY185	
Description	Altered											
Place in profile	W<=E	62.5	65	67	69	71	73	75	77	79	81	83
<b>Major elements</b>												
SiO <sub>2</sub>	44.52	45.98	48.34	46.13	48.33	44.5	48.09	47.29	45.63	47.53	45.78	
TiO <sub>2</sub>	0.11	0.14	0.2	0.18	0.26	0.18	0.24	0.24	0.15	0.24	0.17	
Al <sub>2</sub> O <sub>3</sub>	21.59	20.49	13.03	17.41	16.38	15.62	15.93	14.87	20.62	17.8	18.47	
Fe <sub>2</sub> O <sub>3</sub> *	6.68	6.42	7.52	7.37	7.09	10.33	7.35	8.76	6	6.4	7.3	
MnO	0.11	0.12	0.15	0.13	0.15	0.16	0.14	0.14	0.11	0.12	0.12	
MgO	8.41	8.66	11.87	10.32	8.83	12.59	9.47	10.57	7.75	8.78	9.74	
CaO	14.83	16.16	16.93	15.9	16.59	13.53	16.48	16.2	16.58	17.01	15.07	
Na <sub>2</sub> O	0.69	0.84	0.55	0.65	0.77	0.35	0.78	0.62	1.07	0.55	0.79	
K <sub>2</sub> O	0.03	0	0.01	0.02	0.01	0	0	0.02	0	0	0	
P <sub>2</sub> O <sub>5</sub>	0.02	0	0.02	0.01	0	0	0.01	0.02	0	0	0.01	
LOI	3.53	2.74	2.18	2.54	1.92	3.6	2.06	1.74	2.42	2.15	2.99	
<b>Sum</b>	100.51	101.54	100.8	100.66	100.32	100.86	100.57	100.47	100.33	100.58	100.45	
<b>Trace elements †</b>												
V	67	104	175	125	195	93	181	168	108	166	133	
Cr	89	313	247	154	165	101	162	197	152	183	194	
Co	44	44	55	50	48	59	50	56	42	43	49	
Ni	61	67	75	81	54	82	66	58	67	42	58	
Cu	26	38	143	46	99	83	117	92	48	61	18	
Zn	27	32	25	25	22	41	20	18	21	19	26	
Rb	0	0	0	0	0	0	0	0	0	0	0	
Sr	150	156	94	152	135	118	128	144	166	141	131	
Y	5	6	8	7	9	7	9	10	7	9	7	
Zr	12	12	10	11	11	7	12	10	11	12	10	
Nb	0	2	2	0	0	0	0	0	0	0	1	
Ba	74	43	32	0	35	0	31	29	41	0	13	
La												
Ce	24	19	14	21	24	20	17	19	16	23	13	
Nd												
Mg#	71.52	72.91	75.90	73.64	71.30	70.86	71.99	70.65	72.04	73.24	72.69	
FeO	4.18	3.91	4.73	4.66	4.45	6.56	4.56	5.61	3.74	3.98	4.45	
Fe <sub>2</sub> O <sub>3</sub>	2.08	2.12	2.32	2.24	2.20	3.11	2.33	2.59	1.89	2.02	2.41	

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Location	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri
Sample#	90LY186	90LY187	90LY188	90LY189	90LY190	90LY191	90LY192	90LY193	90LY194	90LY195	90LY196
Description	Altered	Altered	Altered	Altered	Gabbro	gabbronorite	Oxgbnorite	Altered	Altered	Altered	Altered
Place in profile W<=E	85	87	89	91	93	94.5	96	98	100	102	104
<b>Major elements</b>											
SiO <sub>2</sub>	48.39	44.27	44.05	43.89	46.9	49.28	44.1	44.2	46.51	46.94	45.86
TiO <sub>2</sub>	0.22	0.09	0.16	0.41	0.38	0.25	0.96	0.13	0.22	0.19	0.1
Al <sub>2</sub> O <sub>3</sub>	17.43	23.21	15.79	17.85	13.79	16.45	14.84	19.59	17.74	18.01	24.23
Fe <sub>2</sub> O <sub>3</sub> *	5.8	5.36	11.34	10.53	9.95	7.68	14.72	8.39	7.82	7.17	4.64
MnO	0.12	0.09	0.17	0.13	0.15	0.14	0.16	0.14	0.14	0.13	0.09
MgO	8.65	6.57	12.71	8.94	10.29	9.14	9.59	9.77	8.98	8.76	5.06
CaO	17.62	15.77	11.22	15.24	15.84	16.3	14.56	14.36	15.75	16.01	16.18
Na <sub>2</sub> O	0.49	1.21	0.53	0.69	0.63	0.78	0.67	0.72	0.95	1.07	1.66
K <sub>2</sub> O	0.01	0	0	0.02	0.02	0.01	0.02	0.03	0.02	0.02	0
P <sub>2</sub> O <sub>5</sub>	0.02	0	0	0.01	0.01	0.01	0	0	0.01	0	0.01
LOI	1.85	3.22	4.1	2.85	1.9	0.7	1.15	3.19	2.34	2.07	2.59
<b>Sum</b>	100.6	99.78	100.07	100.57	99.87	100.74	100.79	100.53	100.48	100.37	100.42
<b>Trace elements (ppm)</b>											
V	191	73	112	407	300	180	659	74	152	139	66
Cr	210	105	26	255	311	140	150	74	106	135	91
Co	41	39	53	59	62	48	65	51	51	48	28
Ni	31	41	71	86	67	87	67	69	65	81	42
Cu	25	9	119	180	205	215	187	17	113	59	38
Zn	11	19	47	40	32	19	29	40	19	16	7
Rb	1	1	0	1	4	1	3	3	2	1	1
Sr	119	162	112	132	111	125	107	146	138	156	167
Y	9	4	4	6	11	8	8	7	8	8	6
Zr	11	10	8	10	10	11	9	12	12	12	11
Nb	1	2	2	0	1	3	1	1	3	2	1
Ba	37	18	53	52	109	23	125	36	37	13	66
La											
S	19	18	16	16	20	17	24	20	25	22	12
Ce											
Nd											
Mg#	74.84	70.97	69.10	62.87	67.35	70.36	56.51	69.91	70.91	68.51	
FeO	3.47	3.23	7.68	5.95	5.37	4.73	7.38	5.17	4.90	4.69	2.58
Fe <sub>2</sub> O <sub>3</sub>	1.98	1.81	2.89	3.99	4.04	2.48	6.60	2.70	2.43	2.01	1.80

Table 1A: Continued.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Location	Skaidevarri										
Sample#	90LY197	90LY198	90LY199	90LY200	90LY201	90LY202	90LY203	90LY204	90LY205	90LY206	90LY207
Description	Altered	Olivinegb									
Place in profile	W<=E	106	109	111	113	115	116.5	117.5	119.5	121.5	125
<b>Major elements</b>											
SiO <sub>2</sub>	48.18	44.97	46.02	48.45	45.21	47.62	44.59	44.84	47.05	46.18	43.73
TiO <sub>2</sub>	0.24	0.14	0.18	0.3	0.16	0.17	0.38	0.14	0.26	0.23	0.62
Al <sub>2</sub> O <sub>3</sub>	16.42	18.87	18.94	14.9	19.49	21.83	16.81	20.11	18.38	17.14	15.31
Fe <sub>2</sub> O <sub>3</sub> *	6.94	7.19	6.95	7.67	7.4	4.68	10.28	7.43	6.69	8.24	12.24
MnO	0.15	0.12	0.11	0.14	0.13	0.1	0.15	0.12	0.13	0.14	0.15
MgO	9.19	9.8	9.3	9.98	9.67	5.93	9.4	8.83	8.1	9.54	9.72
CaO	15.82	16.04	15.57	16.45	15.55	17.42	16.29	14.8	16.5	15.82	15.37
Na <sub>2</sub> O	0.85	0.69	0.83	0.6	0.68	1.34	0.32	1.15	0.94	0.54	0.75
K <sub>2</sub> O	0.01	0.01	0.01	0.01	0	0	0.01	0	0.01	0	0.01
P <sub>2</sub> O <sub>5</sub>	0.01	0	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.02
LOI	2.05	2.72	2.71	1.75	2.93	1.83	2.52	2.87	2.15	2.53	2.53
<b>Sum</b>	99.84	100.55	100.64	100.24	101.22	100.57	100.76	100.3	100.21	100.36	100.44
<b>Trace elements (ppm)</b>											
V	177	92	116	223	100	146	359	97	200	155	569
Cr	142	143	183	318	120	82	212	96	151	77	407
Co	44	50	46	49	49	32	61	45	45	52	67
Ni	53	79	60	46	53	28	86	56	45	62	104
Cu	130	36	21	65	35	21	175	22	82	120	140
Zn	22	26	18	21	24	9	42	26	22	30	35
Rb	1	0	1	0	0	1	0	0	0	0	1
Sr	122	159	129	124	137	169	140	199	144	138	132
Y	8	5	7	8	7	6	7	6	9	9	7
Zr	10	11	10	11	11	10	8	11	12	10	11
Nb	1	3	2	2	1	1	1	1	2	1	0
Ba	60	42	65	15	50	51	59	20	23	30	60
La											
S											
Ce	18	23	15	14	17	23	17	10	21	16	24
Nd											
Mg#	72.54	73.11	72.75	72.19	72.27	71.65	64.59	70.33	70.72	69.78	61.30
FeO	4.52	4.52	4.35	4.90	4.62	3.03	5.92	4.62	4.15	4.96	0.15
Fe <sub>2</sub> O <sub>3</sub>	1.97	2.22	2.17	2.28	2.32	1.35	3.77	2.35	2.13	2.78	4.61

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1A: Continued.

Location	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri	Skaidevarri
Sample#	90LY208	90LY209	90LY210	90LY211	90LY212	90LY213	90LY214	90LY215	90LY216	90LY217	90LY218
Description	Oligboniorite	Gabbro	Oligbonorite	Altered	Altered	Gabbro	Gabbro	Gabbro	Gabbro	Oloox	Oloox
Place in profile	W<=E	126.5	128	130	135	137	139	141	143	145.5	147
Major elements											149
SiO <sub>2</sub>	45.89	47.36	46.92	46.02	46.24	45.62	46.99	47.48	42.36	44.96	46.49
TiO <sub>2</sub>	0.21	0.25	0.29	0.25	0.22	0.26	0.25	0.27	1.22	0.4	0.15
Al <sub>2</sub> O <sub>3</sub>	17.99	18.88	15.61	14.16	16.37	13.39	17.72	17.09	17.62	19.24	21.12
Fe <sub>2</sub> O <sub>3</sub> *	8.47	6.16	8.64	9.04	7.86	9.72	6.66	7.05	14.98	10.37	7.96
MnO	0.14	0.12	0.15	0.16	0.13	0.15	0.12	0.14	0.14	0.15	0.14
MgO	10.36	8.08	10.24	12.91	10.92	13.12	9.14	9.34	7.49	9.17	8.68
CaO	15	15.49	16.21	14.61	15.97	15.39	15.75	16.71	15.04	14.45	15.12
Na <sub>2</sub> O	0.69	1.26	0.72	0.74	0.58	0.5	1.08	0.72	0.71	0.74	0.86
K <sub>2</sub> O	0.01	0.06	0	0.01	0.01	0.03	0.09	0.01	0	0	0
P <sub>2</sub> O <sub>5</sub>	0.01	0.01	0	0	0.01	0	0.01	0.01	0.01	0.01	0.02
LOI	1.72	2.65	0.97	2.47	2.51	2.82	2.54	1.87	1.07	0.94	0.46
Sum	100.47	100.31	99.75	100.37	100.82	101.02	100.36	100.71	100.64	100.42	101
Trace elements †											
V	117	185	186	157	144	169	184	160	841	273	71
Cr	145	170	180	289	396	778	349	195	58	218	63
Co	52	41	55	56	50	63	44	43	66	54	47
Ni	77	45	83	85	75	146	65	51	45	53	36
Cu	49	31	109	36	67	118	92	35	140	46	14
Zn	27	17	27	38	22	23	37	19	27	30	26
Rb	0	5	0	2	1	3	8	2	1	0	0
Sr	141	237	131	133	125	100	202	164	134	155	169
Y	7	8	9	8	7	9	7	8	7	6	7
Zr	11	12	11	10	10	9	12	13	11	13	11
Nb	3	1	2	2	1	3	1	0	2	1	0
Ba	30	105	14	38	41	49	80	14	39	57	17
La											
S	16	16	26	16	17	22	23	19	21	20	15
Ce											
Nd											
Mg#	70.93	72.35	70.27	74.02	73.48	72.92	73.24	72.55	49.93	63.82	68.51
FeO	5.10	4.15	5.47	6.09	4.79	5.41	4.25	4.49	6.43	5.00	4.22
Fe <sub>2</sub> O <sub>3</sub>	2.86	1.60	2.62	2.34	2.59	3.77	1.99	2.11	7.91	4.87	3.32

Table 1A: Continued.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Location	Skaidevari	Skaidevari	Skaidevari	Skaidevari	Skaidevari	Skaidevari
Sample#	90LY219	90LY220	90LY221	90LY222	90LY223	90LY224
Description	Oloxbg	Oxidegb	Oxide-	Oxidegb	Oloxbg	Oxgb
Place in profile						
W<=E	151	153	155	157	159	160.5
<b>Major elements</b>						
SiO <sub>2</sub>	47.8	39.25	43.68	42.59	45.9	43.87
TiO <sub>2</sub>	0.21	1.53	0.63	1.13	0.22	0.69
Al <sub>2</sub> O <sub>3</sub>	20.46	13.71	19.25	15.81	19.68	18.69
Fe <sub>2</sub> O <sub>3</sub> *	7.23	19.84	11.88	16.16	8.61	11.17
MnO	0.14	0.19	0.14	0.16	0.15	0.14
MgO	8.04	9.59	8.21	8.63	9.05	7.31
CaO	16.31	13.34	14.48	14.95	15	15.49
Na <sub>2</sub> O	0.84	0.35	0.7	0.6	0.62	1.07
K <sub>2</sub> O	0	0.01	0.01	0	0.01	0.03
P <sub>2</sub> O <sub>5</sub>	0.01	0	0	0.01	0.02	0.01
LOI	0.003	3.09	1.82	0.64	1.78	2.13
<b>Sum</b>	101.05	100.9	100.81	100.66	101.03	100.59
<b>Trace elements (</b>						
V	130	842	470	796	127	529
Cr	64	55	84	138	72	76
Co	45	74	60	68	51	56
Ni	28	46	45	65	50	39
Cu	72	233	88	200	27	152
Zn	17	52	26	26	25	29
Rb	1	2	1	2	1	3
Sr	156	141	142	114	152	151
Y	6	6	6	9	7	8
Zr	11	9	11	10	10	9
Nb	2	2	1	3	3	2
Ba	4	60	70	54	6	50
La						
Ce	16	11	25	25	20	7
Nd						
Mg#	68.93	49.09	57.96	51.58	67.71	56.62
FeO	4.28	10.81	5.78	6.87	4.52	6.19
Fe <sub>2</sub> O <sub>3</sub>	2.52	7.95	5.52	8.60	3.64	4.36

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1B: The Eastern suite of the Lyngen Gabbro. The data are similar to those in Table 1A.

Location Sample#	Skaidevarri 90-LY-134	Skaidevarri 90-LY-135	Skaidevarri 90-LY-136	Skaidevarri 90-LY-137	Skaidevarri 90-LY-138	Skaidevarri 90-LY-139	Skaidevarri 90-LY-140	Skaidevarri HLY-93	Iddu HLY-93	Iddu HLY-12-93	Iddu HLY-13-93	Iddu HLY-14-93	Iddu HLY-16-93	
Description	Gbnorite	Gbnorite	Oxgabbro	Gabbro		Altered	Olgabnorite	Olgabborio	qz-gabbro	qz-granit.	gb pegm.	pyxite	gb pegm. boundary	
Place in profile W<=E (m)	0	1.5	3	4.5	6	7.5	9						center	
<b>Major elements (XRF)</b>														
SiO <sub>2</sub>	47.3	47.0	47.4	46.7	44.7	46.9	47.3	45.8	61.3	45.6	51.3	49.0	43.0	
TiO <sub>2</sub>	.13	.11	.11	.08	.08	.10	.07	.67	.12	.34	.26	.29	.50	
Al <sub>2</sub> O <sub>3</sub>	15.4	14.7	15.4	17.8	19.7	16.6	14.8	14.0	11.5	19.0	8.5	9.0	16.9	
Fe <sub>2</sub> O <sub>3</sub> *	6.2	5.9	5.6	5.3	5.9	5.5	5.7	18.3	9.9	13.3	11.7	12.6	16.9	
MnO	.11	.11	.10	.10	.09	.10	.10	.34	.19	.18	.19	.20	.19	
MgO	12.3	12.8	12.1	10.1	11.0	12.0	12.5	7.4	8.1	9.0	14.0	16.8	8.6	
CaO	17.0	16.8	17.7	17.9	16.2	17.1	17.8	11.1	7.9	11.6	12.4	9.9	12.6	
Na <sub>2</sub> O	.35	.28	.41	.49	.27	.26	.19	1.33	1.50	1.12	1.00	1.08	1.01	
K <sub>2</sub> O	.03	.00	.05	.02	.00	.01	.00	.14	.00	.25	.32	.22	.05	
P <sub>2</sub> O <sub>5</sub>	.00	.00	.00	.00	.00	.02	.00	.01	.02	.01	.03	.06	.01	
LOI	2.5	2.9	2.6	2.6	2.2	2.9	2.5	1.4	0.6	1.4	1.3	1.5	2.0	
Sum	101.4	100.7	101.5	101.0	100.1	101.5	101.0	100.5	101.2	101.0	100.6	100.6	101.7	
<b>Trace elements (XRF)</b>														
V	11.8	11.8	11.8	12.0	71	101	123	563	180	453	299	208	528	
Cr	534	853	991	660	372	703	849	31	499	32	1028	1994	57	
Co	46	43	39	45	42	42	58	33	56	58	55	70		
Ni	143	121	138	82	78	139	141	10	92	8	243	429	41	
Cu	217	69	195	66	49	179	170	75	13	63	162	39	134	
Zn	17	26	15	11	14	23	8	141	90	71	76	75	71	
Rb	3	2	4	0	0	0	0	6	2	182	12	5	4	
Sr	87	74	122	115	111	89	74	109	119	3	108	78	122	
Y	6	5	6	3	4	4	15	12	13	8	10	2		
Zr	9	9	10	10	8	9	15	13	1	25	36	10		
Nb	0	0	0	2	3	0	2	1	71	1	2	0		
Ba	33	0	47	76	24	0	58	145	0	13	49	52	0	
La								10	4	0.155	8	13	13	
S								0.31	0.03	0.04	0.05	0.03	0.28	
Ce	21	12	21	21	18	16	19	32	0	2	40	5	4	
Nd								8	0	3	4	0	0	
Mg#	66.3	68.4	68.4	65.6	65.1	68.6	68.9	28.6	44.7	40.5	54.5	57.1	33.7	
FeO	3.6	3.2	3.1	3.1	3.3	3.0	3.0							
Fe <sub>2</sub> O <sub>3</sub>	2.3	2.5	2.1	1.8	2.3	2.2	2.3							

Table 1B: Continued

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Location	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu
Sample#	HLY-18-93 HLY-19-93 HLY-20-93 HLY-22b-93 HLY-23-93 HLY-24-93 HLY-25-93 HLY-26-93 HLY-27-93 HLY-31-93 HLY-35-93 HLY-38-93 HLY-39-93 HLY-40-93 HLY-41-93													
Description	tonalitic- gb pegm	tonalitic- qz-granit.	tonalitic- gb pegm	tonalitic- melagb	tonalitic- gb pegm	tonalitic- qz-granit.								
Place in profile														
W<=E (m)														
<b>Major elements</b>														
SiO <sub>2</sub>	42.2	74.1	38.3	41.3	47.0	71.3	68.8	61.4	41.5	40.3	41.0	66.3	72.7	41.3
TiO <sub>2</sub>	.31	.24	.67	.63	.59	.16	.23	.17	.34	.46	.60	.76	.29	.30
Al <sub>2</sub> O <sub>3</sub>	16.9	14.6	17.6	14.7	14.7	18.3	13.7	14.8	10.9	15.2	17.3	16.2	14.8	.61
Fe <sub>2</sub> O <sub>3</sub> *	16.7	4.6	18.1	20.2	20.9	9.8	5.6	5.7	4.4	21.0	18.4	2.4	6.7	13.2
MnO	.13	.04	.14	.17	.20	.13	.10	.11	.18	.16	.15	.17	.13	.10
MgO	8.2	1.0	7.0	8.4	10.1	9.5	2.4	2.8	0.9	9.0	8.2	8.4	1.7	.24
CaO	13.9	4.4	11.9	13.4	10.6	15.1	6.7	7.3	4.4	12.4	13.7	12.2	6.7	5.5
Na <sub>2</sub> O	1.13	2.40	.99	.80	.64	1.19	1.90	2.42	3.15	1.19	.70	1.11	2.03	14.0
K <sub>2</sub> O	.02	.00	.18	.00	.00	.07	.11	.13	.00	.02	.00	.18	.03	.05
P <sub>2</sub> O <sub>5</sub>	.01	.10	.01	.02	.01	.02	.12	.10	.06	.00	.01	.02	.05	.04
LOI	2.7	1.5	2.4	2.0	3.1	0.9	0.7	0.7	1.0	1.4	2.3	1.5	.08	.17
<b>Sum</b>	102.2	103.0	97.2	101.7	102.2	102.1	103.0	102.9	86.7	102.3	101.7	101.8	101.8	101.5
<b>Trace elements (ppm)</b>														
V	574	26	789	734	515	420	27	51	109	456	621	128	33	302
Cr	55	30	36	36	28	286	22	26	16	28	72	24	30	29
Co	65	11	66	75	53	59	20	22	12	69	70	21	13	57
Ni	21	7	27	10	9	40	7	8	6	10	27	5	7	8
Cu	86	12	105	165	73	139	14	23	8	91	163	21	12	34
Zn	55	24	83	64	104	30	32	38	70	77	72	85	56	101
Rb	3	2	13	3	3	2	2	4	2	2	9	4	4	1
Sr	130	186	136	94	88	112	261	282	177	133	139	194	237	229
Y	4	10	2	2	3	3	36	29	8	3	4	8	4	10
Zr	11	96	11	10	9	11	94	82	11	11	12	17	12	14
Nb	0	2	2	0	2	1	5	6	0	1	2	1	0	3
Ba	5	66	68	47	35	0	55	87	44	80	62	13	60	1
La	5	13	12	10	7	10	6	11	7	10	8	8	0	10
S	0.22	0.03	0.19	0.38	0.20	0.13	0.03	0.03	0.03	0.50	0.46	0.03	0.03	0.08
Ce	8	24	0	0	3	22	24	0	19	14	17	12	15	
Nd	6	9	2	4	5	0	15	17	2	7	9	8	8	12
Mg#	32.8	18.6	27.8	29.3	32.7	49.3	30.3	33.3	16.7	30.1	30.9	77.7	19.9	25.8
FeO														
Fe <sub>2</sub> O <sub>3</sub>														

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1B: Continued

Location	Iddu HLY-42-93	Iddu HLY-43-93	Iddu HLY-45-93	Iddu HLY-46-93	Iddu HLY-48-93	Iddu HLY1.1	Iddu HLY2.1	Iddu HLY3.1	Iddu HLY4.1	Iddu HLY1.2	Iddu HLY2.2	Iddu HLY3.2	Iddu HLY4.2	Iddu HLY5.2	Iddu HLY6.2																		
Description	gb pegm	gb pegm	gb pegm	gb pegm	qz-granit.	gb pegm	qz-gabbro	qz-gabbro	melagb	qz-gabbro	gb pegm	melagb	melagb	melagb	melagb	melagb	melagb	melagb	melagb	melagb	melagb	melagb	melagb	melagb	melagb	melagb	melagb						
Place in profile W<=E (m)	Major elements	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Fe <sub>2</sub> O <sub>3</sub> *	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	P <sub>2</sub> O <sub>5</sub>	LOI	Sum	Trace elements	V	Cr	Co	Ni	Cu	Zn	Rb	Sr	Y	Zr	Nb	Ba	La	S	Ce	Nd	Mg# <sup>#</sup>	FeO	Fe <sub>2</sub> O <sub>3</sub>
	39.5	58.6	45.6	60.8	55.6	43.7	45.5	45.5	46.1	40.0	44.3	43.8	45.6	44.2	44.0																		
	.82	.33	.05	.36	.59	.30	.24	.35	.13	.31	.52	.58	.33	.33	.14																		
	14.7	13.8	10.1	12.4	18.0	20.1	19.5	18.3	25.8	15.6	16.0	15.6	15.3	18.4	25.7																		
	22.8	8.4	14.1	4.8	10.7	11.7	10.8	12.8	6.0	22.3	15.3	16.8	13.3	11.8	5.6																		
	.19	.16	.27	.09	.11	.15	.17	.19	.11	.18	.21	.22	.20	.16	.09																		
	8.7	6.0	7.8	1.0	4.6	8.9	9.7	9.3	6.3	7.8	9.9	9.8	12.2	8.5	5.3																		
	11.5	10.5	12.6	5.1	9.9	13.9	14.0	13.6	15.5	12.3	12.6	12.1	12.6	14.3	15.6																		
	1.10	1.19	.64	2.24	1.43	.83	.85	.00	.02	.24	.43	.76	.09	.32	1.16																		
	.03	.08	.00	.36	.00	.05	.06	.15	.13	.33	.39	.39	.35	.30																			
	.02	.06	.03	.10	.14	.00	.01	.01	.00	.02	.00	.01	.02	.01																			
	2.5	2.6	0.1	1.8	0.0	0.8	1.0	0.0	0.4	1.0	0.8	0.8	1.0	1.3	1.1																		
	101.8	101.8	101.7	89.0	101.0	100.5	101.8	100.5	102.0	100.9	100.5	101.0	102.0	102.0	99.1																		
	444	156	178	44	638	309	249	327	123	553	384	419	262	325	156																		
	49	529	18	13	70	41	29	19	26	38	23	29	15	21	15																		
	61	36	50	16	68	60	61	66	39	75	68	70	67	64	40																		
	44	92	6	5	36	17	14	17	22	16	18	16	18	10	8																		
	89	25	9	10	134	179	198	150	133	143	105	110	146	111	133																		
	94	77	101	50	71	52	69	69	52	105	98	128	113	58	47																		
	2	3	1	13	8	4	7	7	9	9	12	8	10	14	16																		
	94	205	178	174	178	142	146	137	177	118	115	110	106	141	206																		
	15	4	11	4	2	3	2	2	5	2	2	0	3	4	0																		
	11	48	11	24	12	9	9	10	11	11	11	11	10	10	11																		
	1	4	1	4	1	0	0	0	0	3	0	0	0	0	0																		
	37	21	27	207	50	40	21	29	24	58	54	114	85	102	43																		
	17	9	12	7	7	0	7	5	5	7	6	6	7	0																			
	0.15	0.04	0.04	0.03	0.18																												
	16	18	6	14	22	13	20	19	13	21	16	18	33	22																			
	3	14	4	7	5	10	6	8	5	6	17	5	10	11	10																		
	27.7	41.7	35.6	17.0	30.1	43.1	47.3	42.1	51.0	25.8	39.1	36.9	47.8	42.0	48.5																		
	FeO																																
	Fe <sub>2</sub> O <sub>3</sub>																																

Table 1B: Continued

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Location	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu	Iddu
Sample#	HLY1.3	HLY2.3	HLY3.3	HLY1.4	HLY2.4	HLY3.4	HLY1.5	HLY3.5	HLY4.5	HLY5.5	HLY6.5	HLY1.6	HLY2.6
Description	gb pegm	qz-gabbro	qz-granit.	tonalite-									
Place in profile	W<-E (m)			tonalite-									
Major elements													
SiO <sub>2</sub>	60.1	56.0	55.5	40.1	41.2	44.0	50.0	43.4	41.3	41.6	43.6	59.5	72.1
TiO <sub>2</sub>	.29	.33	.35	.53	.35	.11	.37	.45	.38	.30	.34	.21	.64
Al <sub>2</sub> O <sub>3</sub>	13.5	14.4	14.3	20.7	26.0	24.5	14.2	21.2	22.7	20.5	18.6	14.2	14.8
Fe <sub>2</sub> O <sub>3</sub> *	8.4	9.3	9.5	15.3	12.0	5.3	12.2	13.0	13.1	12.0	10.9	5.8	3.2
MnO	.15	.18	.18	.13	.09	.09	.19	.14	.12	.12	.13	.05	.03
MgO	6.7	6.9	7.1	6.3	4.2	4.5	11.0	8.5	6.2	5.9	7.0	1.8	1.0
CaO	7.7	8.8	9.0	14.2	14.5	16.6	11.6	11.9	11.4	11.5	13.9	6.1	4.9
Na <sub>2</sub> O	1.61	1.60	1.27	.02	.15	.20	.59	.16	.03	.04	.27	2.45	2.30
K <sub>2</sub> O	.74	.75	.84	.23	.32	.40	.35	.43	.206	.1.75	.66	1.39	.64
P <sub>2</sub> O <sub>5</sub>	.07	.07	.08	.00	.02	.01	.06	.00	.01	.02	.01	.05	.04
LOI	1.0	1.2	1.4	1.9	1.4	2.1	1.1	2.0	2.3	2.4	1.9	1.4	0.8
Sum	100.2	99.6	99.5	100.3	97.8	101.6	99.5	98.2	98.4	99.2	97.4	98.7	98.9
Trace elements													
V	121	164	161	483	353	95	215	348	458	434	321	114	35
Cr	614	518	525	26	19	18	1035	24	23	20	17	18	16
Co	32	39	37	63	57	34	52	55	50	50	56	19	9
Ni	119	99	100	7	6	6	148	6	7	5	6	10	12
Cu	30	32	33	69	112	99	24	111	62	53	111	28	13
Zn	74	76	74	66	46	237	83	57	56	57	50	26	50
Rb	24	21	23	13	17	23	9	65	92	81	30	60	28
Sr	258	230	230	177	210	208	207	210	272	236	184	298	252
Y	19	16	16	3	2	2	10	2	1	2	2	6	7
Zr	61	35	40	10	12	11	39	12	12	13	11	62	99
Nb	6	5	5	2	1	1	3	1	1	2	1	11	5
Ba	226	156	284	41	23	142	66	411	1073	742	231	804	361
La	10	14	16	13	6	6	14	6	10	3	6	18	21
S	35	28	35	9	18	12	33	17	14	1	18	38	46
Ce	18	14	18	10	8	10	17	9	10	7	6	14	19
Nd	44.5	42.7	42.8	29.2	26.0	46.0	47.3	39.4	32.1	33.0	39.1	23.6	22.9
FeO													
Fe <sub>2</sub> O <sub>3</sub>													

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1B: Continued

Location Sample#	HLY5.6	Iddu HLY6.6	Iddu HLY1.7	Iddu HLY2.7	Iddu HLY3.7	Iddu HLY4.7	Iddu HLY1.8	Iddu HLY2.8	Iddu HLY1.9	Iddu HLY2.9	Iddu HLY20-94	Iddu HLY21-94	Iddu HLY22-94	Iddu HLY24-94	Iddu HLY25-94
Description	tonalite- tonalite- qz-granit.	melagb	melagb	melagb	melagb	melagb									
Place in profile W<=E (m)															
<b>Major elements</b>															
SiO <sub>2</sub>	73.5	72.3	39.4	41.7	41.8	42.9	39.6	42.6	40.7	42.7	40.5	41.2	45.1	41.8	39.2
TiO <sub>2</sub>	.17	.14	.64	.48	.28	.22	.66	.25	.64	.38	.57	.52	.38	.36	.66
Al <sub>2</sub> O <sub>3</sub>	14.3	15.1	17.0	17.8	24.4	24.7	18.5	26.1	20.6	23.7	17.8	18.7	15.2	24.6	17.6
Fe <sub>2</sub> O <sub>3</sub> *	2.6	2.4	19.0	15.2	9.5	8.0	18.4	7.8	15.2	10.3	17.5	16.0	13.2	10.9	19.0
MnO	.02	.02	.16	.15	.10	.09	.14	.10	.16	.14	.16	.14	.19	.11	.16
MgO	0.7	0.6	8.4	7.9	5.1	7.2	3.9	6.6	5.8	8.0	7.7	10.0	4.6	7.6	7.6
CaO	4.2	5.1	14.1	14.7	14.8	15.1	14.2	15.5	13.3	13.8	13.9	13.9	13.1	11.5	14.0
Na <sub>2</sub> O	2.69	2.27	.00	.00	.11	.13	.00	.19	.36	.15	.00	.00	.23	.24	.00
K <sub>2</sub> O	.62	.47	.13	.14	.24	.22	.07	.26	.43	.48	.23	.25	.38	.1.91	.10
P <sub>2</sub> O <sub>5</sub>	.07	.07	.02	.01	.01	.01	.02	.03	.03	.16	.01	.01	.04	.01	.01
LOI	0.4	0.9	1.5	1.8	1.6	1.7	1.9	1.2	1.4	1.1	2.2	2.0	1.9	2.5	2.4
Sum	99.3	99.4	100.2	100.0	97.9	98.0	100.7	98.0	99.4	98.7	100.8	100.1	99.7	98.5	100.6
<b>Trace elements</b>															
V	22	32	538	456	352	291	536	301	316	202	622	559	360	486	596
Cr	20	30	40	37	16	20	32	20	19	13	26	21	21	22	24
Co	7	9	76	70	51	46	73	43	63	49	72	68	61	43	72
Ni	6	5	9	8	6	5	8	7	6	4	9	8	7	7	6
Cu	12	17	75	100	33	41	91	31	64	38	96	99	87	17	103
Zn	21	24	67	55	51	52	46	69	64	65	69	70	48	72	72
Rb	26	18	7	9	14	13	4	14	21	22	11	14	16	97	6
Sr	276	272	132	138	207	194	174	232	196	224	167	151	138	261	139
Y	6	6	2	2	2	1	1	1	2	2	1	0	3	1	1
Zr	96	80	9	11	10	10	11	14	14.5	20	18	22	21	19	
Nb	5	5	2	1	1	1	2	1	2	1	4	4	4	4	3
Ba	368	256	48	7	101	25	84	105	124	66	90	93	102	884	118
La	15	19	11	10	1	1	10	8	14	10	0	0	0	0	0
S															
Ce	45	50	11	10	16	14	15	14	15	13	8	15	14	3	9
Nd	19	17	15	2	9	2	7	6	11	9	0	0	3	4	3
Mg#	22.4	20.5	30.6	34.3	35.1	39.0	28.1	33.3	30.4	36.0	31.3	32.6	43.0	29.7	28.6
FeO															
Fe <sub>2</sub> O <sub>3</sub>															

Table 1B: Continued

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Location Sample#	HLY26.94	Iddu										
Description	tonalite- qz-granit.											
Place in profile W<=E (m)												
<b>Major elements</b>												
SiO <sub>2</sub>	45.0	66.4	62.9	56.3	52.2	61.6	64.8	49.6	47.6	74.4	43.4	42.0
TiO <sub>2</sub>	.51	.21	.15	.20	.19	.15	.14	.19	.54	.18	.41	.64
Al <sub>2</sub> O <sub>3</sub>	22.1	12.9	13.1	10.8	8.5	13.4	12.7	17.0	16.7	13.8	16.9	18.2
Fe <sub>2</sub> O <sub>3</sub> *	11.5	7.5	8.7	12.6	15.0	8.9	7.8	11.8	15.3	3.6	13.1	16.4
MnO	.10	.14	.16	.23	.33	.15	.14	.21	.25	.06	.24	.23
MgO	3.4	5.6	6.3	10.1	14.5	6.6	5.7	8.4	6.4	.9	8.1	7.2
CaO	13.0	6.5	7.8	8.3	8.7	8.0	7.3	11.0	10.5	4.9	14.7	12.2
Na <sub>2</sub> O	.97	1.53	1.53	1.13	.80	1.43	1.41	1.19	1.17	2.54	.34	.37
K <sub>2</sub> O	.68	.00	.00	.00	.00	.00	.00	.01	.06	.05	.02	.04
P <sub>2</sub> O <sub>5</sub>	.26	.04	.03	.01	.01	.02	.02	.00	.00	.01	.01	.01
LiO	1.2	0.4	0.3	0.7	0.4	0.4	0.7	0.8	1.3	0.7	1.5	2.0
Sum	98.7	101.2	101.1	100.3	100.7	100.6	100.6	100.2	99.9	101.2	98.8	99.4
<b>Trace elements (ppm)</b>												
V	495	111	134	209	172	168	152	173	429	42	399	539
Cr	21	71	118	604	1858	201	99	70	62	3	116	84
Co	53	23	33	41	42	34	31	50	47	10	59	48
Ni	19	44	47	80	194	47	43	55	15	7	29	14
Cu	157	23	15	19	12	21	135	90	18	160	113	82
Zn	56	84	74	105	147	75	71	97	116	34	76	112
Rb	33	3	4	1	1	1	3	2	3	3	2	8
Sr	214	153	125	95	63	126	131	162	145	194	158	114
Y	4	13	15	19	24	13	14	16	21	8	7	13
Zr	23	17	20	19	20	18	19	20	26	19	14	14
Nb	4	2	3	4	4	2	1	3	4	2	0	2
Ba	163	8	8	6	47	96	29	97	52	0	81	134
La	0	7	6	9	10	8	9	8	12	8	7	5
S	0.03	0.03	0.03	0.02	0.03	0.02	0.03	0.08	0.08	0.03	0.12	0.12
Ce	9	5	9	6	8	0	0	5	20	0	4	0
Nd	8	0	5	0	12	4	0	0	3	0	0	0
Mg <sup>#</sup>	22.7	43.0	42.0	44.4	49.1	42.5	42.1	41.4	29.5	20.1	38.1	30.4
FeO												
Fe <sub>2</sub> O <sub>3</sub>												

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1B: Continued

Location Sample#	Strupen Lan-1	Strupen Lan-2	Strupen Lan-3	Strupen Lan-4	Strupen Lan-5	Strupen Lan-6	Strupen Lan-7	Strupen Lan-8	Strupen Lan-9	Strupen Lan-10	Strupen Lan-11	Strupen Lan-12	Strupen Lan-13	Strupen Lan-14	Strupen Lan-15	Strupen Lan-16	Strupen Lan-17
<b>Description</b>																	
Place in profile W<=E (m)	1.25	2.5	3.25	3.6	4.25	4.9	5.5	6.25	6.6	7.25	8.25	9.1	9.75	10.25	11	11.25	11.5
<b>Major elements</b>																	
SiO <sub>2</sub>	46.6	47.3	46.8	41.8	41.4	42.9	47.3	47.3	46.8	48.1	48.2	46.4	46.6	45.8	42.9	44.5	
TiO <sub>2</sub>	.09	.13	.10	.46	.60	.49	.10	.08	.12	.07	.15	.20	.13	.22	.56	.17	
Al <sub>2</sub> O <sub>3</sub>	16.5	13.5	17.5	18.9	17.2	19.1	17.7	16.9	17.6	18.4	17.6	16.5	18.7	18.1	18.2	19.4	
Fe <sub>2</sub> O <sub>3</sub> *	7.6	8.9	9.8	15.0	16.9	16.3	9.4	9.8	10.1	10.0	9.1	10.9	11.6	10.8	12.7	18.8	
MnO	.16	.17	.18	.17	.21	.20	.19	.20	.17	.18	.20	.19	.19	.17	.19	.19	
MgO	9.6	11.3	9.9	5.5	6.7	5.8	9.3	10.1	10.2	9.5	9.3	9.0	9.2	8.7	8.0	5.8	
CaO	16.0	16.5	13.7	14.8	14.5	13.6	14.9	14.5	13.8	13.8	13.9	14.2	15.3	14.2	14.4	14.0	
Na <sub>2</sub> O	.16	.22	.23	.24	.26	.41	.33	.27	.26	.25	.09	.14	.25	.28	.31	.27	
K <sub>2</sub> O	.01	.01	.00	.02	.01	.01	.00	.00	.00	.00	.00	.00	.00	.00	.00	.01	
P <sub>2</sub> O <sub>5</sub>	.01	.00	.02	.01	.01	.02	.01	.01	.00	.01	.01	.01	.01	.01	.00	.01	
LOI	2.3	1.4	0.1	1.8	1.8	2.2	0.2	0.2	0.0	0.3	0.3	0.0	0.3	0.4	0.2	2.4	
Sum	98.9	99.3	98.3	98.7	99.5	101.0	99.4	99.4	99.6	99.4	100.4	99.8	99.9	99.9	101.4	98.4	
<b>Trace elements †</b>																	
V	153	207	187	709	562	624	162	133	217	241	121	242	303	227	380	776	
Cr	254	552	74	41	36	27	76	71	72	33	27	29	105	22	25	31	
Co	46	53	51	65	69	66	57	51	58	58	52	60	65	56	57	56	
Ni	71	79	57	20	13	9	50	51	54	49	36	39	34	33	24	10	
Cu	114	74	59	203	123	163	98	107	101	196	135	150	120	110	176	219	
Zn	48	50	57	58	85	90	62	57	62	58	62	61	61	57	53	50	
Rb	2	0	4	2	0	0	0	0	0	0	2	2	1	0	0	2	
Sr	83	67	95	111	102	117	93	88	89	94	101	97	87	101	98	98	
Y	2	3	2	2	1	2	3	1	1	2	2	2	1	1	2	1	
Zr	17	19	17	19	19	18	18	18	16	17	22	19	16	17	16	15	
Nb	4	2	5	3	3	4	4	4	5	4	2	2	4	4	4	2	
Ba	6	3	10	2	2	5	2	4	8	2	3	5	4	6	3	7	
La	20	10	17	6	5	3	22	10	8	13	13	14	8	15	14	13	
S	0.05	0.03	0.04	0.34	0.42	0.38	0.13	0.06	0.05	0.05	0.05	0.09	0.09	0.09	0.23	0.35	
Ce	53	48	31	2	47	64	63	42	49	23	11	37	52	66	55	59	
Nd	6	2	2	2	2	7	7	8	5	7	2	2	2	2	10	2	
Mg#	55.8	56.1	50.3	26.7	28.3	26.3	49.7	50.6	50.2	48.6	50.5	45.1	44.2	44.7	38.7	23.7	
FeO	5.03	5.81	6.02	8.06	9.04	8.23	6.87	5.54	4.96	6.94	7.48	7.89	7.38	6.29	5.81	6.77	
Fe <sub>2</sub> O <sub>3</sub>	2.05	2.48	3.19	6.15	6.91	7.23	1.85	3.74	4.65	2.40	0.89	2.22	3.47	6.32	11.32	2.63	

Table 1B: Continued

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Location Sample#	Strupen Lan-18	Strupen Lan-19	Strupen Lan-21	Strupen Lan-22	Strupen Lan-23	Strupen Lan-24	Strupen Lan-25	Strupen Lan-26	Strupen Lan-27	Strupen Lan-28	Strupen Lan-29	Strupen Lan-30	Strupen Lan-31	Strupen Lan-32	Strupen Lan-33	Strupen Lan-34	Strupen Lan-35
Description																	
Place in profile																	
W<=E (m)	11.9	12.1	13	13.4	14	14.1	14.5	14.75	15	15.3	16.25	17	17.5	17.9	18.1	18.25	18.4
Major elements																	
SiO <sub>2</sub>	42.2	44.1	47.3	47.1	40.8	47.4	48.3	41.2	46.6	47.2	47.4	49.1	51.3	47.7	43.7	45.8	49.1
TiO <sub>2</sub>	.56	.34	.08	.11	.62	.10	.07	.53	.25	.09	.10	.15	.08	.33	.05	.17	
Al <sub>2</sub> O <sub>3</sub>	18.3	22.6	16.8	18.2	17.6	14.8	16.7	19.8	16.8	17.3	16.5	16.3	9.9	16.6	19.3	27.1	2.8
Fe <sub>2</sub> O <sub>3</sub> *	17.5	12.7	10.4	10.6	17.4	11.4	9.9	15.4	12.1	9.6	8.3	9.2	8.4	13.1	5.6	11.9	
MnO	.17	.16	.20	.19	.18	.27	.21	.14	.19	.18	.16	.18	.16	.18	.11	.26	
MgO	5.4	5.2	10.4	9.6	5.8	10.7	10.1	5.7	9.0	10.4	11.1	10.1	13.3	11.0	7.0	5.5	15.6
CaO	14.3	14.5	13.5	13.4	14.4	14.1	13.5	14.4	14.3	14.1	14.1	14.1	15.1	17.0	15.1	14.7	18.6
Na <sub>2</sub> O	.41	.37	.26	.28	.35	.28	.11	.21	.13	.24	.23	.07	.03	.24	.29	.59	.06
K <sub>2</sub> O	.01	.00	.00	.00	.02	.00	.00	.00	.00	.00	.00	.00	.00	.00	.01	.00	
P <sub>2</sub> O <sub>5</sub>	.00	.02	.01	.00	.01	.00	.01	.00	.02	.01	.01	.01	.00	.01	.01	.01	
LOI	2.3	1.0	0.1	0.0	2.1	0.5	0.4	0.7	0.2	0.0	0.1	0.3	0.6	0.0	0.2	1.7	0.0
Sum	101.2	101.0	99.1	99.4	99.2	99.5	99.3	98.2	99.7	99.1	99.3	99.6	101.6	99.3	100.8	101.2	98.5
Trace elements †																	
V	755	457	130	198	646	97	113	669	409	196	195	203	250	158	487	96	155
Cr	34	23	34	36	26	134	74	28	56	75	81	101	670	140	53	98	1443
Co	64	54	58	52	78	59	60	63	69	56	56	58	60	54	62	36	72
Ni	7	12	53	39	7	37	51	12	51	54	61	75	88	68	25	19	109
Cu	230	104	22	74	293	52	73	259	200	46	153	98	54	23	242	11	24
Zn	68	50	71	57	80	69	47	67	57	55	51	47	51	62	48	67	
Rb	0	0	1	0	3	0	2	3	0	0	2	2	2	0	2	0	0
Sr	107	132	93	98	105	84	98	107	91	89	83	83	41	87	119	149	5
Y	2	1	4	1	2	2	2	2	3	1	2	3	2	3	2	0	9
Zr	18	18	18	17	18	17	17	18	18	18	17	18	17	18	18	19	17
Nb	6	4	5	4	5	5	2	2	2	4	4	4	2	4	5	4	4
Ba	5	10	8	6	2	9	8	4	3	4	7	5	6	5	2	4	7
La	17	12	8	16	8	13	17	11	10	21	15	24	22	19	10	13	23
S	0.49	0.16	0.03	0.07	0.49	0.09	0.08	0.27	0.19	0.03	0.04	0.03	0.03	0.03	0.26	0.03	0.03
Ce	58	69	43	9	23	59	4	2	45	16	58	2	9	39	80	38	26
Nd	2	5	2	8	6	2	2	2	2	2	2	2	2	2	5	2	8
Mg#	23.7	29.1	50.1	47.5	25.0	48.4	50.5	27.1	42.7	52.0	53.8	54.9	59.1	56.8	34.7	49.4	56.7
FeO	8.43	3.81	7.14	4.32	9.18	7.96	7.68	8.26	8.50	3.64	1.87	6.15	6.86	5.54	9.69	4.52	8.33
Fe <sub>2</sub> O <sub>3</sub>	8.21	8.48	2.52	5.85	7.31	2.65	1.41	6.34	2.77	5.56	7.49	1.56	2.26	2.41	0.64	2.75	

### Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1B: Continued

Location Sample#	Strupen Lan-36	Strupen Lan-37	Strupen Lan-38	Strupen Lan-39	Strupen Lan-40	Strupen Lan-41	Strupen Lan-42	Strupen Lan-43	Strupen Lan-44	Strupen Lan-45	Strupen Lan-46	Strupen Lan-47	Strupen Lan-48	Strupen Lan-49	Strupen Lan-50	Strupen Lan-51	Strupen Lan-52		
Description	Place in profile	W<=E (m)	18.5	18.9	19.75	20.25	20.75	21.4	22	22.75	23.5	24.25	25.4	26.1	27	27.5	28.5	29	29.05
<b>Major elements</b>																			
SiO <sub>2</sub>	48.3	48.1	49.1	51.8	50.8	48.8	42.8	42.8	46.8	46.7	46.7	46.6	40.2	40.5	41.8	44.2	46.9	46.9	
TiO <sub>2</sub>	.12	.13	.13	.19	.14	.14	.44	.50	.13	.11	.24	.17	.68	.62	.60	.46	.17		
Al <sub>2</sub> O <sub>3</sub>	9.8	8.9	8.4	5.2	8.4	8.1	19.6	21.1	17.4	17.6	18.3	18.0	16.8	18.8	21.8	23.8	13.8		
Fe <sub>2</sub> O <sub>3</sub> *	9.6	8.8	8.1	12.1	8.1	8.4	15.4	14.0	10.3	10.5	11.9	11.8	19.7	17.9	15.3	12.9	13.6		
MnO	.20	.18	.17	.25	.16	.17	.16	.20	.19	.20	.20	.19	.19	.19	.19	.17	.35		
MgO	13.9	14.4	14.1	15.8	14.6	15.2	5.9	5.3	9.6	10.1	8.2	9.3	6.0	5.8	5.2	4.6	11.0		
CaO	16.5	16.9	17.5	15.9	16.4	17.2	14.0	14.6	14.3	13.7	14.3	13.9	14.0	14.6	15.1	15.2	13.6		
Na <sub>2</sub> O	.27	.23	.05	.00	.06	.17	.53	.33	.37	.32	.09	.26	.14	.16	.21	.36	.25		
K <sub>2</sub> O	.00	.00	.01	.01	.02	.00	.02	.04	.00	.00	.00	.00	.00	.00	.00	.00	.00		
P <sub>2</sub> O <sub>5</sub>	.02	.01	.01	.00	.01	.01	.01	.02	.02	.01	.01	.01	.00	.00	.02	.00	.02		
LOI	0.2	1.0	1.0	0.2	1.6	0.0	2.3	2.9	0.7	0.2	0.2	0.0	0.4	0.5	0.6	0.3	0.0		
<b>Sum</b>	99.0	98.6	98.5	101.5	100.2	98.2	101.3	101.8	99.8	99.5	100.0	100.0	98.1	99.0	100.8	102.0	99.7		
<b>Trace elements<sup>t</sup></b>																			
V	159	181	215	246	219	169	708	512	182	195	373	271	796	720	497	375	120		
Cr	1388	1851	2035	2592	2721	865	63	22	23	30	21	15	34	32	21	19	82		
Co	57	58	55	73	58	53	62	61	56	58	67	61	71	74	60	48	61		
Ni	103	128	129	145	166	141	25	9	39	50	34	33	8	7	6	12	33		
Cu	33	44	28	74	39	12	120	229	118	149	216	126	254	142	75	57	33		
Zn	59	54	41	67	46	42	82	52	58	63	64	49	62	53	49	95			
Rb	0	0	2	2	0	0	0	1	0	0	2	0	2	2	2	0			
Sr	49	43	36	15	39	36	126	140	94	94	103	97	98	105	131	147	83		
Y	6	4	2	2	4	2	3	1	3	2	2	2	2	2	2	2	5		
Zr	17	17	19	18	18	17	19	19	16	18	18	17	17	19	19	17	18		
Nb	4	4	2	4	3	4	5	4	4	4	3	5	5	3	2	3	4		
Ba	3	6	5	5	6	3	8	7	2	2	11	2	7	2	8	6			
La	23	16	18	10	13	14	10	11	16	13	8	20	4	16	8	15	4		
S	0.04	0.04	0.03	0.06	0.03	0.03	0.26	0.32	0.12	0.09	0.10	0.08	0.65	0.49	0.30	0.11	0.08		
Ce	69	29	2	2	10	60	106	74	74	59	2	57	86	16	2	2	69		
Nd	5	2	2	2	3	6	2	2	4	2	3	3	2	8	2	2	2		
<b>Mg#</b>	59.1	62.1	63.4	56.6	64.2	64.5	27.9	27.6	48.2	49.1	40.7	44.0	23.3	24.5	25.4	26.2	44.8		
FeO	6.94	6.73	5.91	8.23	5.89	5.78	8.26	8.26	6.32	3.67	7.75	2.38	10.00	9.08	7.82	6.26	5.71		
Fe <sub>2</sub> O <sub>3</sub>	2.00	1.37	1.62	3.09	1.63	2.00	6.29	4.88	3.35	6.44	3.35	9.14	8.73	7.91	6.68	6.05	7.31		

Table 1B: Continued

Location Sample#	Strupen Lan-53	Strupen Lan-54	Strupen Lan-55	Strupen Lan-56	Strupen Lan-57	Strupen Lan-58	Strupen Lan-59	Strupen Lan-60	Strupen Lan-61	Strupen Lan-62	Strupen Lan-63	Strupen Lan-64	Strupen Lan-65	Strupen Lan-66	Strupen Lan-67	Strupen Lan-68	Strupen Lan-69	
Description	Place in profile W<=E (m)	29.5	31	31.75	32.5	33.5	34.75	35	35.4	35.6	36.25	36.6	37.75	38.5	39.5	39.75	39.85	42
<b>Major elements</b>																		
SiO <sub>2</sub>	48.8	43.8	41.7	46.1	46.0	46.2	41.7	44.4	47.9	46.7	48.4	48.7	47.1	46.6	45.2	46.2	46.2	
TiO <sub>2</sub>	.09	.58	.63	.24	.27	.22	.60	.37	.19	.12	.12	.10	.16	.28	.23	.21	.21	
Al <sub>2</sub> O <sub>3</sub>	16.6	16.0	15.6	18.7	18.3	19.1	18.4	23.7	14.7	18.0	17.0	17.7	19.9	20.4	15.0	17.8	17.8	
Fe <sub>2</sub> O <sub>3</sub> *	10.3	19.2	13.0	13.2	12.9	17.5	12.6	13.4	8.4	9.1	8.8	8.5	10.5	12.7	13.2	12.2	12.2	
MnO	.22	.21	.21	.21	.21	.22	.17	.15	.28	.16	.17	.16	.18	.18	.22	.20	.20	
MgO	9.9	7.3	6.8	7.7	7.7	7.9	5.5	4.8	9.5	9.4	9.5	9.2	7.7	7.6	10.3	9.4	9.4	
CaO	13.8	14.0	14.1	14.2	14.4	13.6	14.7	14.5	13.6	16.1	15.2	15.1	14.6	13.6	14.1	14.5	14.0	
Na <sub>2</sub> O	.09	.31	.11	.35	.37	.39	.17	.45	.14	.22	.05	.08	.20	.06	.22	.21	.25	
K <sub>2</sub> O	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	
P <sub>2</sub> O <sub>5</sub>	.01	.01	.02	.01	.01	.00	.02	.01	.02	.01	.00	.00	.00	.02	.01	.00	.00	
LOI	0.1	0.2	0.0	0.0	0.0	0.0	0.2	0.1	0.6	<0.01	1.1	0.1	0.1	0.1	0.0	0.0	0.0	
<b>Sum</b>	99.9	101.5	98.5	100.5	100.4	100.6	98.8	101.4	99.7	100.1	99.6	99.9	100.4	99.2	100.6	99.8	100.3	
<b>Trace elements †</b>																		
V	138	910	908	349	371	273	839	476	198	215	197	231	178	241	513	365	317	
Cr	45	39	43	22	20	22	31	19	68	222	140	80	51	30	52	143	37	
Co	58	78	71	62	65	57	77	60	65	53	51	58	52	56	61	67	61	
Ni	52	27	22	19	21	11	8	15	54	44	55	49	22	31	50	41	41	
Cu	60	220	193	126	136	109	196	131	78	38	34	137	169	125	230	185	75	
Zn	71	69	55	64	67	70	57	56	78	49	44	54	51	57	58	66	69	
Rb	2	0	2	0	0	0	2	1	2	0	2	2	0	2	0	1	1	
Sr	94	87	83	104	105	103	103	136	81	86	82	88	97	114	104	78	90	
Y	3	2	2	2	2	2	2	1	2	1	2	2	3	2	2	4	2	
Zr	19	17	18	18	17	16	18	17	19	16	19	17	17	18	16	17	16	
Nb	2	5	4	4	4	3	4	2	4	2	4	2	5	2	3	5	4	
Ba	3	5	2	6	7	3	5	4	5	5	5	5	9	4	4	6	7	
La	10	22	2	15	16	8	3	9	5	13	21	14	19	9	15	14	14	
S	0.07	0.40	0.32	0.17	0.19	0.18	0.36	0.21	0.12	0.04	0.03	0.04	0.04	0.09	0.12	0.09	0.05	
Ce	2	52	2	45	31	41	47	16	5	46	13	2	52	84	49	56	61	
Nd	2	6	2	2	3	7	2	7	6	6	2	2	3	2	5	10	4	
Mg#	48.9	27.5	26.2	37.2	36.8	38.0	23.9	27.6	41.5	52.7	51.0	51.2	52.9	42.5	37.5	43.8	43.5	
FeO	8.33	11.02	10.74	8.84	3.91	3.50	8.94	2.89	9.72	2.24	6.80	6.73	6.05	7.34	5.64	3.03	4.79	
Fe <sub>2</sub> O <sub>3</sub>	1.14	7.12	7.43	3.24	8.92	9.01	7.65	9.42	2.69	5.96	1.64	1.36	1.87	2.40	6.49	9.82	6.93	

Table 1B: Continued

Location Sample#	Strupen Lan-70	Strupen Lan-71	Strupen Lan-72	Strupen Lan-73	Strupen Lan-74	Strupen Lan-75	Strupen Lan-76	Strupen Lan-77	Strupen Lan-78	Strupen Lan-79	Strupen Lan-80	Strupen Lan-81	Strupen Lan-82	Strupen Lan-83	Strupen Lan-84	Strupen Lan-85	Strupen Lan-86
Description																	
Place in profile W<=E (m)	42.25	42.75	43.25	43.35	43.6	44.6	45.76	46.25	46.5	47.5	48	49.5	50	50.5	51.25	52	52.5
<b>Major elements</b>																	
SiO <sub>2</sub>	44.6	50.4	60.8	61.6	47.6	46.0	45.9	45.4	44.7	54.8	45.6	49.9	44.1	43.5	44.4	45.8	
TiO <sub>2</sub>	.36	.47	.40	.38	.27	.10	.11	.06	.20	.47	.57	.69	.81	.48	.53	.43	.18
Al <sub>2</sub> O <sub>3</sub>	18.7	18.3	14.9	15.1	17.1	17.6	18.2	17.9	19.2	22.2	17.5	15.9	15.3	16.1	17.8	17.3	
Fe <sub>2</sub> O <sub>3</sub> *	15.1	12.6	9.6	9.1	13.5	10.2	10.4	9.6	12.1	14.6	10.2	17.4	22.7	18.6	18.6	15.5	11.9
MnO	.23	.17	.13	.12	.27	.19	.18	.19	.20	.18	.12	.22	.23	.19	.18	.22	.19
MgO	7.5	4.0	3.1	2.9	8.3	10.7	10.1	10.0	8.2	5.3	3.3	6.0	6.9	7.5	7.1	7.8	10.1
CaO	13.4	11.8	9.4	9.4	13.2	14.0	13.9	13.7	14.1	14.3	10.8	12.6	13.5	13.5	13.8	14.5	14.1
Na <sub>2</sub> O	.11	.39	.21	.11	.14	.33	.36	.35	.43	.46	.39	.38	.46	.44	.37	.34	.35
K <sub>2</sub> O	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
P <sub>2</sub> O <sub>5</sub>	.01	.00	.01	.01	.01	.01	.01	.01	.01	.02	.01	.01	.01	.01	.01	.01	.01
LOI	<0.01	0.8	0.1	0.1	0.0	0.4	0.5	1.5	0.7	0.2	0.1	1.5	0.2	0.0	0.6	0.6	0.0
Sum	100.0	98.9	98.7	98.9	100.4	99.5	99.8	99.3	100.6	102.4	97.6	100.4	103.0	101.0	102.4	101.5	99.9
<b>Trace elements<sup>a</sup></b>																	
V	504	345	186	238	250	194	249	101	296	500	241	595	831	861	923	538	328
Cr	23	17	6	13	21	24	21	14	21	19	12	41	34	34	34	19	22
Co	69	48	32	33	66	57	59	55	58	64	34	54	73	69	77	68	64
Ni	18	7	6	6	33	46	44	43	25	9	7	14	7	24	38	19	38
Cu	165	90	55	70	98	239	180	116	146	96	50	95	163	173	205	136	92
Zn	71	82	51	45	92	57	65	67	63	63	40	88	79	57	61	64	62
Rb	2	0	0	2	2	0	0	0	0	0	0	0	1	0	0	0	0
Sr	107	116	100	94	99	86	91	92	105	130	106	102	89	93	94	98	91
Y	2	2	2	2	2	1	1	2	1	0	0	2	3	3	2	2	3
Zr	16	18	17	16	18	6	5	7	8	5	7	7	5	6	7	7	7
Nb	4	4	3	4	5	5	5	4	5	4	5	6	4	6	4	4	4
Ba	2	5	4	2	69	2	7	5	7	5	8	3	6	4	7	2	3
La	2	17	23	3	16	17	14	20	12	11	7	9	22	17	8	17	17
S	0.18	0.14	0.09	0.08	0.07	0.06	0.08	0.17	0.24	0.10	0.31	0.67	0.40	0.38	0.26	0.08	
Ce	2	85	31	75	2	41	39	31	44	44	38	48	65	57	57	61	55
Nd	4	4	2	2	2	0	4	5	0	4	0	0	4	6	5	2	
Mg#	33.0	24.0	24.3	24.2	38.1	51.2	49.3	51.0	40.5	26.4	24.6	23.4	28.7	27.7	33.5	46.0	
FeO	8.87	5.51	5.88	4.93	9.76	5.26	6.12	6.48	6.08	4.50	4.36	8.10	3.02	3.31	12.64	10.87	7.99
Fe <sub>2</sub> O <sub>3</sub>	5.38	6.54	3.17	3.68	2.80	4.36	3.69	2.49	5.42	9.65	5.44	8.53	19.33	15.00	4.69	3.55	3.06

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1B: Continued

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Location Sample#	Strupen Lan-87	Strupen Lan-88	Strupen Lan-89	Strupen Lan-90	Strupen Lan-91	Strupen Lan-92	Strupen Lan-93	Strupen Lan-94	Strupen Lan-95	Strupen Lan-96	Strupen Lan-97	Strupen Lan-98	Strupen Lan-99	Strupen Lan-100	Strupen Lan-101	Strupen Lan-102	Strupen Lan-103	
Description	Place in profile W<=E (m)	53	55.3	55.8	57	58.8	60.25	61	61.8	62	62.5	62.8	63.2	64	65	66	68	68.3
Major elements																		
SiO <sub>2</sub>	45.8	45.9	45.7	43.4	46.4	45.8	43.7	46.4	54.6	75.9	75.9	63.2	46.7	45.7	45.9	45.5	44.9	
TiO <sub>2</sub>	.14	.10	.15	.47	.07	.09	.38	.50	.41	.03	.13	.26	.07	.14	.13	.19	.16	
Al <sub>2</sub> O <sub>3</sub>	17.7	17.0	18.5	16.0	15.9	14.9	17.0	19.4	16.1	13.0	13.1	15.7	16.9	17.6	18.5	19.0	19.0	
Fe <sub>2</sub> O <sub>3</sub> *	11.2	10.6	11.4	18.7	10.6	10.1	16.6	13.6	10.4	0.9	1.9	7.3	10.7	11.4	11.5	12.4	11.4	
MnO	.19	.19	.19	.19	.24	.18	.19	.14	.02	.04	.12	.22	.21	.21	.19	.18	.18	
MgO	10.8	11.6	10.1	7.9	11.3	13.0	8.1	4.5	3.3	0.1	0.3	2.1	10.5	10.3	10.0	8.6	8.0	
CaO	14.1	13.0	13.7	13.6	14.8	13.7	14.1	13.5	9.8	6.4	4.8	6.8	13.7	13.8	14.0	13.5	13.9	
Na <sub>2</sub> O	.34	.35	.34	.35	.33	.34	.37	.78	.56	1.08	2.15	1.46	.55	.36	.40	.38	.36	
K <sub>2</sub> O	.00	.00	.00	.00	.00	.00	.00	.00	.00	.01	.00	.00	.00	.00	.00	.00	.00	
P <sub>2</sub> O <sub>5</sub>	.00	.01	.00	.00	.01	.02	.01	.01	.00	.02	.03	.02	.01	.03	.01	.02	.01	
LOI	0.0	0.7	0.1	0.1	0.0	0.8	1.9	2.0	2.5	1.1	0.8	0.4	0.0	0.7	0.4	0.5	2.3	
Sum	100.2	99.4	100.2	100.7	99.6	98.8	101.3	100.7	97.6	98.5	99.2	97.2	99.3	99.4	100.0	99.9	100.3	
Trace elements†																		
V	255	181	259	778	93	166	561	397	220	7	13	132	121	212	191	355	302	
Cr	22	34	22	32	92	152	37	19	13	6	6	9	55	42	33	27	21	
Co	59	54	59	70	55	56	63	50	30	0	4	17	56	57	56	60	60	
Ni	38	45	30	19	54	76	27	10	6	5	5	17	54	48	35	24	24	
Cu	144	248	174	232	26	38	217	62	42	8	8	59	70	139	138	136	133	
Zn	57	56	56	55	73	62	65	93	58	2	13	87	78	65	62	77	74	
Rb	0	0	0	0	0	0	0	0	0	0	0	72	0	0	0	0	0	
Sr	91	87	98	85	82	74	90	148	107	122	119	679	91	87	89	94	93	
Y	0	1	0	2	3	2	1	10	6	3	3	22	3	2	2	1	2	
Zr	5	6	6	7	7	6	9	8	14	14	14	243	7	6	6	6	7	
Nb	4	4	4	5	4	5	4	5	5	3	3	14	6	4	4	3	3	
Ba	7	4	8	7	7	11	6	10	0	1	1	30	5	3	1	8	6	
La	11	11	13	9	21	17	13	17	5	7	0	54	16	22	5	16	16	
S	0.07	0.07	0.13	0.34	0.04	0.03	0.29	0.15	0.09	0.02	0.03	0.04	0.06	0.07	0.08	0.16	0.14	
Ce	38	58	72	70	59	31	32	40	91	27	54	1169	67	52	85	17	29	
Nd	0	6	4	8	0	2	2	0	1	0	0	36	0	5	0	3	0	
Mg#	49.1	52.3	46.9	29.8	51.8	56.3	32.7	25.1	24.1	13.7	14.9	22.5	49.6	47.4	46.7	41.1	41.1	
FeO	6.77	7.49	8.71	10.51	5.98	6.19	7.78	6.88	5.26	0.54	1.01	3.06	6.62	3.60	5.26	6.48	8.50	
Fe <sub>2</sub> O <sub>3</sub>	3.76	2.32	1.82	7.11	3.98	3.26	8.06	6.01	4.56	0.29	0.83	3.88	3.40	7.42	5.67	5.22	2.08	

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1B: Continued

Location Sample#	Strupen Lan-104	Strupen Lan-105	Strupen Lan-106	Strupen Lan-107	Strupen Lan-108	Strupen Lan-109	Strupen Lan-110	Strupen Lan-111	Strupen Lan-112	Strupen Lan-113	Strupen Lan-114	Strupen Lan-115	Strupen Lan-116	Strupen Lan-117	Strupen Lan-118	Strupen Lan-119	Strupen Lan-120	Strupen Lan-121
Description																		
Place in profile W<=E (m)	69	70.5	71.6	72.5	74	74.2	74.8	75.2	77.8	77.8	75.7	79.5	80.3	81	81.9	83	83.8	
Major elements																		
SiO <sub>2</sub>	42.6	43.7	52.8	58.1	73.6	75.3	74.2	41.7	62.1	59.8	54.1	48.4	47.2	47.0	44.6	44.3	62.5	
TiO <sub>2</sub>	.64	.52	.43	.40	.07	.14	.20	.85	.23	.42	.49	.12	.14	.14	.52	.52	.19	
Al <sub>2</sub> O <sub>3</sub>	16.1	19.9	15.6	14.0	12.2	12.3	12.9	14.4	14.5	13.8	16.3	15.2	15.4	16.6	18.6	18.6	15.7	
Fe <sub>2</sub> O <sub>3</sub> *	20.9	16.0	11.5	10.1	1.7	1.9	3.3	23.4	8.4	9.8	12.8	13.7	12.8	12.7	18.5	18.4	7.4	
MnO	.21	.20	.14	.14	.03	.03	.05	.23	.10	.12	.14	.35	.28	.25	.17	.16	.08	
MgO	6.7	5.5	3.6	3.2	0.2	0.1	0.6	7.5	3.0	4.0	10.0	9.6	10.8	6.9	6.0	2.8		
CaO	13.2	13.4	10.5	9.7	6.0	4.5	4.8	13.7	8.3	8.8	11.2	11.9	12.8	13.1	13.7	14.2	8.9	
Na <sub>2</sub> O	.38	.64	.37	.47	1.46	2.89	2.30	.44	.54	.34	.50	.53	.42	.32	.42	.42	.25	
K <sub>2</sub> O	.00	.00	.00	.00	.02	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	
P <sub>2</sub> O <sub>5</sub>	.01	.01	.02	.01	.02	.02	.01	.01	.00	.01	.02	.02	.02	.01	.00	.03	.01	
LOI	1.6	2.1	2.2	1.9	1.0	1.0	0.9	1.6	0.9	0.9	0.0	0.0	0.0	0.0	0.0	1.1	0.3	
Sum	102.3	101.9	97.0	98.0	96.2	98.2	99.3	103.8	97.7	97.0	99.5	100.2	98.7	100.9	103.6	103.7	98.2	
Trace elements †																		
V	787	552	280	202	18	10	34	1070	232	197	319	77	114	191	1006	937	156	
Cr	33	28	15	19	6	7	7	43	12	11	13	12	12	19	34	31	10	
Co	68	58	37	26	3	3	7	69	25	25	37	47	52	56	70	73	24	
Ni	12	13	7	7	5	5	5	25	7	7	26	34	32	32	14	12	7	
Cu	135	75	60	44	9	7	9	222	37	37	68	34	83	75	75	347	220	
Zn	73	115	70	52	6	6	13	109	36	50	43	98	91	73	51	60	25	
Rb	0	0	0	0	0	0	0	0	0	1	0	0	0	0	0	0	0	
Sr	85	118	99	84	88	137	103	77	89	85	100	93	92	85	96	99	76	
Y	1	1	2	1	0	3	3	6	1	1	2	3	2	1	2	1	0	
Zr	6	8	7	7	7	15	10	9	9	7	6	6	6	6	6	6	5	
Nb	5	5	5	1	2	2	5	4	4	5	5	4	5	4	4	6	4	
Ba	7	3	6	4	3	2	1	4	1	3	4	6	4	8	4	5	0	
La	14	10	7	4	2	2	4	10	0	9	11	12	7	15	4	15	6	
S	0.45	0.29	0.17	0.10	0.02	0.03	0.60	0.06	0.09	0.24	0.09	0.40	0.10	0.40	0.39	0.07		
Ce	57	68	32	63	48	44	48	97	10	55	88	35	37	33	37	42	60	
Nd	0	0	0	0	0	0	0	4	3	0	3	0	7	9	7	1	0	
Mg#	24.2	25.4	23.9	24.1	12.7	6.0	16.0	24.1	26.4	23.1	24.0	42.3	42.8	45.9	27.0	24.7	27.7	
FeO	9.50	8.35	5.54	5.11	0.72	0.43	1.55	10.51	3.96	4.64	5.33	6.44	6.08	6.55	7.45	6.30	3.24	
Fe <sub>2</sub> O <sub>3</sub>	10.45	6.79	5.37	4.48	0.86	1.40	1.61	11.85	4.01	4.73	6.89	6.59	6.10	5.54	10.35	11.43	3.87	

Table 1B: Continued

Location Sample#		Strupen Lan-123	Strupen Lan-124	Strupen Lan-125	Strupen Lan-126	Strupen Lan-127	Strupen Lan-128	Strupen Lan-129	Strupen Lan-130	Strupen Lan-131
Description	Place in profile W<=E (m)	84.8	86	87.3	89	90.5	91	91.5	92.3	93
<b>Major elements</b>										
SiO <sub>2</sub>	47.1	46.9	46.6	47.1	46.0	57.8	60.9	70.5	48.0	
TiO <sub>2</sub>	.10	.07	.06	.08	.11	.45	.36	.19	.15	
Al <sub>2</sub> O <sub>3</sub>	17.5	16.0	16.5	19.2	16.0	14.0	14.3	13.1	15.2	
Fe <sub>2</sub> O <sub>3</sub> *	12.2	10.5	9.3	7.6	12.7	11.1	9.1	5.4	13.3	
MnO	.25	.20	.19	.14	.23	.13	.12	.08	.35	
MgO	10.3	12.0	11.4	10.6	11.9	3.2	2.9	1.7	9.9	
CaO	12.7	13.4	14.7	14.8	11.7	9.3	8.5	6.8	13.1	
Na <sub>2</sub> O	.31	.27	.31	.42	.29	.43	.43	.31	.13	
K <sub>2</sub> O	.00	.00	.00	.00	.00	.00	.00	.00	.00	
P <sub>2</sub> O <sub>5</sub>	.02	.01	.01	.02	.02	.00	.00	.01	.01	
LOI	0.4	1.1	1.6	1.1	1.7	1.6	1.4	0.5	0.0	
Sum	100.8	100.4	100.7	101.1	100.7	98.1	97.9	98.5	100.2	
<b>Trace elements †</b>										
V	145	129	115	142	189	244	168	63	84	
Cr	43	39	74	85	46	18	12	8	101	
Co	56	53	52	49	52	29	26	12	59	
Ni	47	55	59	82	53	7	6	5	25	
Cu	75	207	75	72	168	80	43	13	37	
Zn	80	59	69	52	81	58	62	27	112	
Rb	0	0	0	0	0	1	0	0	0	
Sr	90	76	86	92	85	90	105	79	99	
Y	1	1	2	1	1	2	1	1	4	
Zr	7	6	7	6	6	7	8	7	6	
Nb	5	5	5	5	4	4	6	3	4	
Ba	4	10	8	7	6	5	7	4	2	
La	17	18	14	19	16	9	8	6	12	
S	0.08	0.06	0.05	0.03	0.15	0.13	0.09	0.03	0.05	
Ce	46	66	39	53	59	68	54	44	30	
Nd	0	2	2	0	0	0	5	0	3	
Mg#	45.6	53.3	55.0	58.1	48.3	22.4	24.1	23.5	42.6	
FeO	6.88	5.36	6.01	4.93	6.59	5.51	4.72	2.81	7.52	
Fe <sub>2</sub> O <sub>3</sub>	4.67	4.60	2.69	2.21	5.48	5.03	3.87	2.27	5.06	

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 2: Isotope chemistry for the Lyngen Gabbro. In addition to data from each area, dikes crosscutting the gabbros in Goverdalens and at Iddu are included for comparison.

Sample	Sm (ppm)	Nd(ppm)	$^{147}\text{Sm} / ^{144}\text{Nd}$	$^{143}\text{Nd} / ^{144}\text{Nd}$	$^{143}\text{Nd} / ^{144}\text{Nd}$	Initial $^{143}\text{Nd}$	$\epsilon_{(^{143}\text{Nd})}$	$\epsilon_{(^{143}\text{Nd})}$
					CHUR	$/ ^{144}\text{Nd}$		
<b>Skaidevarri</b>								
90-LY-134	0.21	0.46	0.284	0.513055	0.51203	0.512162	7.962	2.60
90-LY-135	0.19	0.40	0.280	0.513117	0.51203	0.512236	9.162	4.05
90-LY-137	0.16	0.28	0.332	0.513374	0.51203	0.512330	14.178	5.89
90-LY-139	0.11	0.18	0.374	0.513511	0.51203	0.512335	16.862	5.99
90-LY-140	0.13	0.21	0.362	0.513472	0.51203	0.512335	16.096	5.99
90-LY-141	0.35	0.74	0.286	0.513259	0.51203	0.512359	11.941	6.44
90-LY-212	0.33	0.69	0.290	0.513276	0.51203	0.512362	12.261	6.52
90-LY-219	0.33	0.67	0.299	0.513303	0.51203	0.512364	12.792	6.55
<b>Strupen</b>								
Lan-1	0.12	0.21	0.363	0.513336	0.51203	0.512195	13.446	3.24
Lan-23	0.08	0.12	0.386	0.513337	0.51203	0.512125	13.467	1.88
Lan-38	0.16	0.32	0.299	0.513131	0.51203	0.512191	9.446	3.18
Lan-47	0.07	0.13	0.340	0.513323	0.51203	0.512255	13.181	4.42
Lan-58	0.07	0.14	0.326	0.513288	0.51203	0.512264	12.502	4.60
Lan-72	0.05	0.09	0.349	0.513311	0.51203	0.512215	12.946	3.64
Lan-84	0.08	0.16	0.316	0.513147	0.51203	0.512153	9.758	2.42
Lan-97	0.04	0.21	0.110	0.512582	0.51203	0.512237	-1.274	4.08
Lan-100	0.08	0.14	0.345	0.513302	0.51203	0.512216	12.768	3.65
Lan-106	0.059	0.113	0.315100	0.513211	0.51203	0.512221	11.007	3.75
Lan-107	0.073	0.141	0.3139	0.513201	0.51203	0.512214	10.806	3.62
Lan-109	0.054	0.336	0.0970	0.512525	0.51203	0.512220	-2.389	3.73
Lan-110	0.195	0.703	0.1680	0.512762	0.51203	0.512234	2.240	4.01
Lan-110	0.20	0.70	0.167	0.512762	0.51203	0.512235	2.241	4.04
Lan-111	0.217	0.394	0.3337	0.513265	0.51203	0.512216	12.063	3.66
Lan-112	0.035	0.085	0.2464	0.512993	0.51203	0.512219	6.758	3.72
Lan-113	0.064	0.129	0.2983	0.513159	0.51203	0.512221	9.987	3.76
Lan-114	0.103	0.197	0.3160	0.513213	0.51203	0.512219	11.032	3.72
Lan-116	0.152	0.284	0.3229	0.513240	0.51203	0.512225	11.570	3.84
Lan-130	0.01	0.03	0.191	0.512788	0.51203	0.512188	2.745	3.11
<b>Isskardet</b>								
LY-IS-9	0.11	0.32	0.211	0.513030	0.51203	0.512366	7.478	6.58
LY-IS-16	0.34	0.70	0.292	0.513296	0.51203	0.512379	12.653	6.84
LY-IS-25	0.38	0.78	0.297	0.513295	0.51203	0.512363	12.643	6.53
LY-IS-31	0.46	0.87	0.315	0.513365	0.51203	0.512374	14.007	6.75
<b>Ellendale</b>								
92LY119	7.92	78.15	0.2289	0.513068	0.51203	0.512348	8.212	6.24
92LY93	1.24	8.47	0.3314	0.513411	0.51203	0.512369	14.903	6.65
93LY20	0.95	7.42	0.2882	0.513208	0.51203	0.512302	10.943	5.34
94LY24	2.18	17.55	0.281	0.513231	0.51203	0.512347	11.392	6.23
94LY39	1.17	10.07	0.2625	0.513173	0.51203	0.512348	10.260	6.23
<b>Goverdalens</b>								
95LY2	0.53	4.24	0.2818	0.513187	0.51203	0.512301	10.534	5.32
94LY67	1.87	13.79	0.3057	0.513287	0.51203	0.512326	12.484	5.81
95LY38	6.92	60.35	0.2589	0.513190	0.51203	0.512376	10.592	6.79
<b>Veidalen</b>								
95LY127	0.80	5.72	0.3156	0.513298	0.51203	0.512306	12.699	5.41
95LY106	0.47	2.81	0.3775	0.513373	0.51203	0.512186	14.162	3.08
95LY94	0.57	3.70	0.3497	0.513394	0.51203	0.512294	14.571	5.19

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 2: Continued

Sample	Sm (ppm)	Nd(ppm)	$^{147}\text{Sm} / ^{144}\text{Nd}$	$^{143}\text{Nd} / ^{144}\text{Nd}$	Initial $^{143}\text{Nd}$		$\epsilon_{(^{143}\text{Nd})}$ (480My)	
					CHUR	/ $^{144}\text{Nd}$		
95LY75	0.60	6.73	0.2016	0.512914	0.51203	0.512280	5.208	4.91
<b>Iddu 1</b>								
Hly 2.8	0.06	0.23	0.154	0.512193	0.51203	0.511709	-8.856	-6.24
Hly 24-93	0.17	0.53	0.194	0.512404	0.51203	0.511794	-4.740	-4.58
Hly 24-94	0.05	0.16	0.199	0.512505	0.51203	0.511879	-2.770	-2.91
Hly 26-94	0.45	1.53	0.178	0.512604	0.51203	0.512044	-0.839	0.31
Hly 22b-93	0.09	0.2	0.285	0.512759	0.51203	0.511863	2.185	-3.23
Hly 11-94	0.92	2.13	0.262	0.512783	0.51203	0.511959	2.653	-1.35
Hly 1.8	0.13	0.26	0.301	0.512826	0.51203	0.511880	3.492	-2.91
Hly 25-94	0.14	0.27	0.321	0.512885	0.51203	0.511876	4.643	-2.98
<b>Iddu 2</b>								
Hly 5 PA	0.84	2.15	0.236	0.512967	0.51203	0.512225	6.242	3.84
Hly 1PA	0.62	1.67	0.223	0.512932	0.51203	0.512231	5.559	3.95
Hly 3PA	1.18	2.65	0.27	0.513108	0.51203	0.512259	8.993	4.50
Hly 7 PA	0.9	2.27	0.24	0.513005	0.51203	0.512250	6.983	4.33
<b>Dikes, Goverdalens</b>								
93LY30	2.976	9.469	0.1900	0.512925	0.51203	0.512328	5.423	5.84
95LY72	1.400	6.006	0.1409	0.512121	0.51203	0.511678	-10.254	-6.84
95LY72	1.396	5.995	0.1408	0.512125	0.51203	0.511683	-10.176	-6.75
95LY71	0.314	0.566	0.3357	0.513290	0.51203	0.512235	12.544	4.02
95LY71	0.284	0.494	0.3475	0.513336	0.51203	0.512244	13.441	4.20
95LY71	0.289	0.522	0.3347	0.513292	0.51203	0.512240	12.581	4.12
<b>Dikes, Iddu</b>								
HLY 2993	2.23	8.8	0.153	0.512781	0.51203	0.512300	2.614	5.30
Hly 33-93	3.38	14.73	0.139	0.512749	0.51203	0.512312	1.990	5.54
Hly 28-93	5.54	15.86	0.211	0.513043	0.51203	0.512380	7.725	6.86
Hly 4.6	2.4	19.49	0.074	0.511753	0.51203	0.511520	-17.439	-9.93
Hly 6.5	2.36	18.28	0.078	0.511789	0.51203	0.511544	-16.737	-9.47
Hly 27-93	1	7.53	0.081	0.511805	0.51203	0.511550	-16.425	-9.34
Hly 5.6	1.84	13.61	0.082	0.511810	0.51203	0.511552	-16.327	-9.30
Hly 36-93	1.29	6.39	0.122	0.511862	0.51203	0.511478	-15.313	-10.74
HLY 22a-93	1.04	3.76	0.168	0.512235	0.51203	0.511707	-8.037	-6.28

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 3: Mineral chemistry of the Lyngen Gabbro. All areas are represented, except for Ellendale. Abbreviations: pl = plagioclase, c = core, r=rim, cpx = augite, opx= orthopyroxene, Fo=forsterite% of olivine. The minerals are listed so cores are aligned with cores, rims with rims. There is therefore some repetition in the table so that the representative mineral-pairings are obvious. Orthopyroxene and olivine show internal variation, and the 1 sigma is therefore equal to the analytical error.

Sample#	An%	1σ	Sample#	Mg#		Mg#		Fo
				cpx	1σ	Sample#	opx	
<b>Skaidevarri</b>								
90LY134 plc (5)	93.0	0.71	90LY134 cpxc (5)	84.5	0.20			
90LY134 plrtocpx (5)	93.5	1.03	90LY134 cpxrtopl (5)	84.7	0.70			
90LY134 plc (5)	93.0	0.71	90LY134 cpx2- (6)	84.7	0.30			
90LY136-1-plrtocpx (4)	93.2	0.28	90LY136-1 cpxc (3)	85.4	0.24			
90LY136-1-plrtocpx (4)	93.2	0.28	90LY136-1 cpxrtopl (5)	86.4	1.11			
90LY136-2-plc (3)	91.3	0.28	90LY136-2 cpxc (4)	85.3	0.68			
90LY136-2-plrtocpx (5)	91.5	1.62	90LY136-2 cpxc (4)	85.3	0.68			
90LY138-1-pl (7)	94.7	0.85	90LY138-1 cpxc (5)	85.1	0.63	90LY138-1-ol (6)	78.7	
90LY138-1-plrtocpx (2)	95.7	3.82	90LY138-1 cpxr (5)	85.3	0.74	90LY138-1-ol (6)	78.7	
90LY138-2-plc (5)	94.6	0.67	90LY138-2 cpxc (5)	84.4	0.28	90LY138-2 opx (6)	81.1	
90LY138-2-plr (5)	95.0	1.19	90LY138-2 cpxrtopl (5)	85.0	0.64	90LY138-2 opx (6)	81.1	
90LY141-1-plr (5)	84.1	0.98	90LY141-1 cpxc (5)	79.0	0.42	90LY141-1 opx (6)	74.2	90LY141-1-ol (6) 69.6
90LY141-1-plrtocpx (5)	85.9	1.97	90LY141-1 cpxr (5)	79.4	1.23	90LY141-1 opx (6)	74.2	90LY141-1-ol (6) 69.6
90LY141-2-plc (5)	83.9	0.42	90LY141-2 cpxc (5)	78.8	0.33	90LY141-2 opx (6)	74.5	
90LY141-2-plr (5)	83.8	0.64	90LY141-2 cpxrtopl (6)	79.2	1.08	90LY141-2 opx (6)	74.5	
90LY142-2 plc (5)	85.0	0.67	90LY142-1cpx (5)	78.0	0.79	90LY142-1 opx (6)	73.0	90LY142-2ol (6) 69.2
90LY142-2 plrtocpx (5)	86.5	1.10	90LY142-1cpx (5)	77.7	0.63	90LY142-1 opx (6)	73.0	90LY142-2ol (6) 69.2
90LY142-2 plc (5)	85.0	0.67	90LY142-2cpxc (5)	77.7	0.93			
90LY142-2 plrtocpx (5)	86.5	1.10	90LY142-2cpxrtopl (5)	77.4	0.66			
90LY143-1-plc (5)	87.8	0.81	90LY143-1 cpxc (5)	78.7	0.73	90LY143-1 opx (6)	75.3	90LY143-1-ol (6) 71.3
90LY143-1-plrtocpx (5)	87.9	0.88	90LY143-1 cpxrtopl (5)	79.5	0.16	90LY143-1 opx (6)	75.3	90LY143-1-ol (6) 71.3
90LY143-2-plc (3)	88.6	1.38	90LY143-2 cpxc (5)	78.4	1.81			90LY143-2-ol (6) 70.6
90LY143-2-plrtocpx (4)	88.9	0.66	90LY143-2 cpxrtopl (5)	80.2	0.49			90LY143-2-ol (6) 70.6
90LY144-1-plc (5)	87.0	0.62	90LY144-1 cpxc (5)	79.2	0.63	90LY144-1 opx (5)	73.6	
90LY144-1-plrtocpx (5)	88.3	1.04	90LY144-1 cpxrtopl (5)	78.6	1.25	90LY144-1 opx (5)	73.6	
90LY144-2-plc (5)	90.6	1.45	90LY144-2 cpxc (5)	78.5	1.19	90LY144-2 opx (6)	75.0	
90LY144-2-plrtocpx (5)	92.9	1.41	90LY144-2 cpxrtopl (5)	79.1	0.92	90LY144-2 opx (6)	75.0	
90LY166 plag (7)	87.6	0.67	90LY166 cpxc (4)	79.1	1.64	90LY166 opx (10)	73.4	90LY166-ol (11) 70.3
90LY166 plag (7)	87.6	0.67	90LY166 cpxrtopl (4)	78.0	0.68	90LY166 opx (10)	73.4	90LY166-ol (11) 70.3
90LY168-2 plc (5)	85.0	0.28	90LY168-1 cpxc (4)	78.1	0.52			90LY168-1ol (6) 69.2
90LY168-2 plrtocpx (5)	87.0	0.89	90LY168-1 cpxrtopl (5)	77.9	0.76			90LY168-1ol (6) 69.2
90LY168-2 plc (5)	85.0	0.28	90LY168-2 cpxc (4)	78.0	0.84	90LY168-2 opx (4)	73.3	
90LY168-2 plrtocpx (5)	87.0	0.89	90LY168-2 cpxrtopl (4)	77.6	0.39	90LY168-2 opx (4)	73.3	
90LY169-1 plc (5)	84.6	0.46	90LY169-1 cpxc (5)	78.9	0.67	90LY169-1 opx (6)	72.5	
90LY169-1 plrtocpx (5)	86.2	0.68	90LY169-1 cpxrtopl (5)	78.7	0.89	90LY169-1 opx (6)	72.5	
90LY169-2 pl (6)	84.9	0.49						90LY169-2ol (6) 68.1
90ly 192 pl (6)	85.1	1.00	90ly 192 cpx rage(6)	76.1	1.93	90ly192 opx (6)	70.5	
90ly-216 plag (6)	84.3	0.95	90ly-216 cpx (6)	77.1	1.20	90ly-216 opx (6)	71.6	
90ly221 plag (5)	85.3	0.55	90ly221 cpx (5)	75.3	1.26	90ly221 opx (5)	71.8	90ly221-ol (6) 66.9
<b>Iddu</b>								
1pa-plag	52.6	2.03				1pa-opx	60.7	
2pa-plag	59.8	2.57	2pa-cpx	73.5	0.89	2pa-opx	64.3	
3pa-plag	60.2	1.70	3pa-cpx	74.5	0.84	3pa-opx	64.0	
4pa-plag	58.4	1.90	4pa-cpx	76.1	1.06	4pa-opx	66.1	
5pa-plag	58.9	0.59	5pa-cpx	74.2	1.09	5pa-opx	64.1	
6pa-plag	58.6	1.19	6pa-cpx	73.3	1.26	6pa-opx	63.0	
7pa-plag	74.4	2.28	7pa-cpx	71.2	1.89	7pa-opx	62.4	
11pb			11pb-cpx	72.6	1.69			
pb16-plag	75.0	1.16	16pb-cpx	72.8	1.26	16pb-opx	64.8	
pb18-plag	67.4	1.24	18pb-cpx	73.9	1.26	18pb-opx	64.6	
pb19-plag	57.6	1.19	19pb-cpx	76.0	1.15	19pb-opx	66.1	
pb21-plag	63.4	2.20	Pb21-cpx	76.7	0.89	Pb21-opx	65.4	

### Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 3: Continued

Sample#	An%	1σ	Sample#	Mg#		Mg#		Fo
				cpx	1σ	Sample#	opx	
<b>Strupen</b>								
Lan-1			Lan-1	76.9		Lan-1	67.8	
Lan-2	97.1		Lan-2	76.1		Lan-2	68.6	
Lan-3	94.6		Lan-3	75.6		Lan-3	66.4	
Lan-4			Lan-4	73.1		Lan-4	59.3	
Lan-5	96.9		Lan-5	73.4		Lan-5	63.6	
Lan-7	93.1		Lan-7	72.9		Lan-7	64.1	
Lan-8	93.7		Lan-8	75.2		Lan-8	64.9	
Lan-9	94.1		Lan-9	75.3		Lan-9	66.5	
Lan-10	94.6		Lan-10	75.7		Lan-10	67.1	
Lan-11	93.4		Lan-11	75.5		Lan-11	66.6	
Lan-12	93.0		Lan-12	75.5		Lan-12	65.0	
Lan-13	93.7		Lan-13	73.9		Lan-13	65.6	
Lan-14	93.7		Lan-14	74.7		Lan-14	65.2	
Lan-15	93.9		Lan-15	73.3		Lan-15	64.2	
Lan-16	94.0		Lan-16	74.3		Lan-16	64.8	
Lan-17			Lan-17			Lan-17	63.2	
Lan-18	95.5		Lan-18	74.1		Lan-18		
Lan-19	92.2		Lan-19	73.0		Lan-19	63.9	
Lan-20	92.6		Lan-20	72.6		Lan-20	63.7	
Lan-21	93.2		Lan-21	75.5		Lan-21	64.6	
Lan-22	92.6		Lan-22	74.0		Lan-22	65.7	
Lan-23	93.9		Lan-23	73.3		Lan-23	63.3	
Lan-24	93.5		Lan-24	73.3		Lan-24	64.6	
Lan-25	91.9		Lan-25	74.8		Lan-25	65.9	
Lan-26	93.5		Lan-26	75.3		Lan-26	67.0	
Lan-27	94.8		Lan-27	76.5		Lan-27	66.0	
Lan-28	94.7		Lan-28	77.8		Lan-28	68.4	
Lan-29	93.9		Lan-29	75.7		Lan-29	68.1	
Lan-30	94.6		Lan-30	77.0		Lan-30	68.6	
Lan-31	98.2		Lan-31	76.3		Lan-31	68.3	
Lan-32	94.7		Lan-32	77.0		Lan-32	68.7	
Lan-33			Lan-33	74.6		Lan-33	66.5	
Lan-38	96.3		Lan-38	82.0		Lan-38	74.5	
Lan-44	94.1		Lan-44	74.8		Lan-44	64.6	
Lan-45	94.2		Lan-45	73.9		Lan-45	65.5	
Lan-47	94.9		Lan-47	74.6		Lan-47	64.1	
Lan-48	93.9		Lan-48	74.5		Lan-48	63.4	
Lan-49	95.1		Lan-49			Lan-49		
Lan-50	94.3		Lan-50	74.3		Lan-50	62.8	
Lan-51	95.0		Lan-51			Lan-51		
Lan-52	94.5		Lan-52	73.1		Lan-52	61.7	
Lan-54	95.6		Lan-54			Lan-54		
Lan-56	95.6		Lan-56			Lan-56		
Lan-58	94.4		Lan-58	71.3		Lan-58	59.6	
Lan-60	95.6		Lan-60			Lan-60		
Lan-61	94.2		Lan-61	71.8		Lan-61	59.4	
Lan-62			Lan-62			Lan-62	65.9	
Lan-64	96.9		Lan-64	75.4		Lan-64	66.7	
Lan-65			Lan-65	75.0		Lan-65	67.4	
Lan-66			Lan-66	76.4		Lan-66	66.4	
Lan-67	96.0		Lan-67	75.9		Lan-67	63.2	
Lan-68	96.1		Lan-68	75.6		Lan-68	64.8	
Lan-69	95.7		Lan-69	75.6		Lan-69	64.9	
Lan-70	95.5		Lan-70	73.9		Lan-70	63.0	
Lan-71	94.2		Lan-71	69.6		Lan-71	59.1	
Lan-72	93.9		Lan-72	70.4		Lan-72	57.0	
Lan-73	90.5		Lan-73	69.9		Lan-73	57.5	
Lan-74	93.7		Lan-74	70.6		Lan-74	58.0	
Lan-76	95.0		Lan-76	75.3		Lan-76		

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 3: Continued

Sample#	An%	$1\sigma$	Sample#	Mg# cpx	$1\sigma$	Sample#	Mg# opx	Sample#	Fo
<b>Goverdalén</b>									
95ly34	77.2		95LY34	72.5		95LY34	64.9		
95ly35	79.0		95ly35	73.8		95ly35	67.0		
95ly38	48.6		95ly38	66.9		95ly38	56.8		
95ly39	87.2		95ly39	79.4		95ly39	75.2	95ly39	72.3
95ly40	87.6		95ly40	83.6					
<b>Veidalén</b>									
95ly127	86.6		95ly127	72.1		95ly127	64.4	95LY126	68.1
95LY126	89.9		95LY126	77.0		95LY126	72.3		
95ly107	94.2		95ly107	83.4			95ly107		75.5
95ly106	95.7		95ly106	83.2		95ly106	78.4	95ly106	75.9
95LY105	94.7						95LY105		80.3
95LY90	91.2		95LY90	79.2		95LY90	75.2		
95LY93	93.8		95LY93	80.7					
95LY94	95.3		95LY94	79.9		95LY94	74.3		
95LY95	93.8		95LY95	79.9		95LY95	74.0	95LY95	71.1
95LY96	95.4		95LY96	80.2		95LY96	74.8	95LY96	72.2
95LY87	97.7		95LY87	88.0		95LY87	83.8	95LY87	83.5
<b>Isskardet</b>									
IS-2 plagc (5)	83.8	0.69	IS-2 cpxc (5)	75.9	0.61		IS-2ol (6)		64.2
IS-2 plagr (5)	84.5	1.94	IS-2 cpxr (5)	75.8	0.70		IS-2ol (6)		64.2
IS-6 plagc (5)	75.7	7.29	IS-6 cpxc (5)	75.7	0.18		IS-6ol (6)		66.5
IS-6 plagr (5)	83.4	1.25	IS-6 cpxr (5)	76.4	0.25		IS-6ol (6)		66.5
IS-9 plagc (5)	83.0	0.92					IS-9ol (6)		67.6
IS-9 plagr (5)	83.6	0.63					IS-9ol (6)		67.6
IS-16 plagc (5)	84.2	0.96	IS-16 cpxc (5)	76.3	1.02	IS-16 opx (6)	72.7	IS-16ol (6)	67.8
IS-16 plagr (5)	83.5	1.45	IS-16 cpxr (5)	76.3	0.42	IS-16 opx (6)	72.7	IS-16ol (6)	67.8
IS-20 plag1-c (5)	82.2	0.59	IS-20 cpxc (5)	75.5	0.71	IS-20 opx (6)	71.2	IS-20ol (6)	66.6
IS-20 plag1-r (5)	84.5	1.12	IS-20 cpxr (5)	75.8	0.46	IS-20 opx (6)	71.2	IS-20ol (6)	66.6
IS-20 plag2- (4)	84.4	0.92	IS-20 cpxr (5)	75.8	0.46	IS-20 opx (6)	71.2	IS-20ol (6)	66.6
IS-26 plagc (5)	80.8	1.65	IS-26 cpxc (5)	75.2	0.43		IS-26ol (6)		67.1
IS-26 plagr (5)	82.2	2.55	IS-26 cpxr (5)	75.3	0.20		IS-26ol (6)		67.1

Table 4: Trace-element and REE-mineral chemistry for ten samples from the Lyngen Gabbro. All the analyses are listed as ppm. Cpx=augite, c=core, r=rim. The errors are less than 10% for the REE, and less than 5% for the trace elements.

Sample	La ppm	Ce ppm	Nd ppm	Sm ppm	Eu ppm	Dy ppm	Er ppm	Yb ppm	Tb ppm	V ppm	Cr ppm	Sr ppm	Y ppm	Zr ppm
<b>Goverdalen</b>														
LY-95-39-cpx-c	0.12	0.62	1.25	0.86	0.41	1.67	0.89	0.99	1486	260	449	8.7	8.7	3.7
LY-95-39-cpx-r	0.14	0.67	1.03	0.68	0.34	1.51	0.84	0.87	1537	262	541	7.8	8.9	4.7
LY-95-34-cpx-c	0.23	1.29	2.76	1.85	0.59	3.75	2.14	2.34	2508	437	85	10.0	22.6	11.3
LY-95-34-cpx-r	0.24	1.17	2.28	1.50	0.60	3.65	2.11	2.12	2338	413	83	9.3	23.5	12.0
LY-95-38-cpx-c	0.54	3.53	9.32	5.76	1.52	12.18	6.64	7.04	2352	265	70	11.6	67.5	67.3
LY-95-38-cpx-r	0.60	4.21	9.75	6.20	1.52	12.62	6.82	7.72	2221	258	72	11.2	74.2	64.2
<b>Verdalens</b>														
LY-95-87-cpx-c	0.02	0.11	0.24	0.18	0.06	0.25	0.22	0.21	339	236	1994	8.6	1.4	0.6
LY-95-87-cpx-r	0.01	0.08	0.26	0.14	0.06	0.20	0.14	0.17	371	255	1924	6.1	1.6	0.8
LY-95-96-cpx-c	0.04	0.21	0.40	0.35	0.16	0.52	0.30	0.31	872	299	1036	8.1	4.1	1.2
LY-95-96-cpx-r	0.12	0.62	0.96	0.63	0.36	1.36	1.05	0.83	2244	401	965	28.5	15.2	2.8
LY-95-127-cpx-c	0.32	1.65	2.86	1.69	0.51	3.41	1.99	1.98	3219	441	160	9.5	22.4	11.4
LY-95-127-cpx-r	0.39	1.66	2.81	1.71	0.51	3.50	1.85	2.04	2343	357	151	9.8	20.7	10.5
<b>Sirupen</b>														
Lan38 cpx-c	0.06	0.34	0.49	0.26	0.12	0.54	0.32	0.32	869	310	2437	6.5	3.5	1.9
Lan38 cpx-r	0.07	0.30	0.40	0.25	0.13	0.50	0.37	0.37	857	312	2486	6.7	3.5	1.7
Lan47 cpx-c	0.05	0.29	0.48	0.30	0.20	0.95	0.63	0.76	1411	380	88	6.5	6.4	1.4
Lan47 cpx-r	0.03	0.25	0.41	0.28	0.16	0.84	0.49	0.60	723	248	90	5.5	3.3	1.3
Lan58 cpx-c	0.03	0.35	0.67	0.48	0.27	1.44	1.04	0.95	1158	280	88	6.0	7.2	2.2
Lan58 cpx-r	0.04	0.34	0.57	0.51	0.21	1.21	0.81	0.84	1067	283	89	6.6	6.1	2.0
Lan72 cpx-c	0.05	0.29	0.61	0.55	0.37	1.55	1.07	1.16	1675	202	80	8.6	11.2	2.0
Lan72 cpx-r	0.03	0.27	0.78	0.55	0.34	1.57	1.11	1.15	2779	198	80	8.1	10.8	1.8

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

## Chapter 3

### Accretion of the lower crust at Atlantis Bank: The results of geochemical modeling

#### Abstract

We have developed a quantitative and qualitative model for the accretion of the lower crustal rocks at Atlantis Bank, Southwest Indian Ridge. We use the Melts-program to establish the mode of accretion and extent of reaction for gabbros sampled in Hole 735B, and find that fractional crystallization of a near dry melt with  $\text{Na}_8 \sim 3$  and  $\text{Fe}_8 \sim 9$  at crustal pressure on the quartz-fayalite-magnetite oxygen fugacity buffer fits the general evolution of the solid- and liquid-line of descent at Atlantis Bank. Our model predicts that approximately 770 vertical meters of the lower crustal section has not been sampled and that the total crustal thickness at Atlantis Bank was more than 4.4-km. Thus, the observed seismic Moho may indeed be the crust-mantle boundary, and not an alteration front as previously suggested.

Slow-spreading ocean ridges typically have gabbros with unusually high Mg-augites. We use our models and the degree of disequilibrium between the minerals to establish how far off the fractionation-trend the rocks plot. The majority of the gabbros in the lower 1000 meters of the hole have augite with elevated Mg#'s, up to 10 mol% higher than what is required by fractional crystallization. Thus, we propose that interstitial melts in a mush-zone have dissolved pre-existing minerals, causing the melt to become more Mg and Ca rich. Plagioclase and augite do not re-equilibrate with the melt, as solid-state diffusion is very slow. Thus, when the hybrid melt crystallizes, this will result in higher Mg# augites and higher anorthite plagioclase.

#### 1 *Introduction*

Coexisting phases in Atlantis Bank gabbros, and in other gabbros produced from elevated  $\text{Na}_8$  basalts, are often out of equilibrium with each other, i.e. they could not have been crystallized directly from mid-ocean ridge basalts at crustal pressures (Bloomer et

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al., 1989, 1991; Meyer et al., 1989; Elthon et al., 1987, 1992, Michael and Cornell, 1998, Dick et al., 2002). High Mg# augite is easily identifiable in the lower crust at these slow-spreading. Bloomer et al. (1989) showed that most gabbros from Indian Ocean fracture zones crystallized from more fractionated liquids than spatially associated erupted pillow basalts. It has been suggested that the primitive magmas are emplaced towards the center of each ridge-segment, which is why we do not see primitive gabbros in the fracture zones (Whitehead et al., 1984; Dick 1989; Magde, et al., 1997). Bloomer et al. (1989) also found that augite coexisting with plagioclase was more magnesian than expected for plagioclase anorthite content, and that augite starts crystallizing earlier relative to plagioclase than in 1-atm experiments. Bloomer et al. (*op. cit.*) suggested that they formed by in-situ crystallization (as parameterized by Langmuir, 1989) along the edges of small magma-bodies while higher-T magmas were repeatedly injected along them. These observations effectively require that there must be olivine-rich cumulates representing higher-T melts somewhere in the crust or shallow mantle. Elthon et al. (1987, 1992) observed the same early-onset crystallization of primitive-looking clinopyroxene along the Mid-Cayman rise, and suggested that they formed by high-pressure (3-10 kbar) crystallization, whereas Meyer et al. (1989) also found early onset of high-Mg pyroxene in dredged gabbros between the Shaka- and Islas Orcadas Fracture Zones, Southwest Indian Ridge. They argued that the high-Mg augites in the gabbros are due to small amounts of trapped liquid. Solid-state Fe/Mg diffusion is fast in olivine, therefore the interstitial melt gains Mg and loses Fe to the olivine as the melt evolves. As a result, the Mg# of the augite subsequently precipitated from this melt has constant Mg#'s, but very varied incompatible element levels. Bédard et al., (2000 and references therein) concurred, but argued that syntaxis (the processes, mechanical and chemical, by which magmas react with and assimilate their host rocks) causes the shift to high Mg augite relative to the anorthite content of the coexisting plagioclase. In this view, the chemical evolution of primary MORB are as dependent on assimilation and reaction with older cumulates as on fractional crystallization, as the crust acts as a reactive filter during melt transport.

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ODP Hole 735B was drilled at the center of a ridge-segment immediately to the east of Atlantis II Fracture Zone at 700-m water depth on top of Atlantis Bank, Southwest Indian Ridge (Shipboard Scientific Party, 1989; Dick et al., 1991a). The relatively shallow drill site is an uplifted oceanic core complex, and flexural uplift from the detachment and transform-parallel faults partly explains the shallow bathymetry (Baines et al., 2003). The seismic crustal thickness at the bank is  $5 \pm 1$  km thick (Muller et al., 1997). Two ODP Legs, 118 and 176, drilled a total of ~1500-m mostly gabbroic rocks with unusually high recovery (87%). In addition, another site, 158-m deep Hole 1105A was drilled 1-km to the north of Hole 735B (Shipboard Scientific Party, 1999) in very similar rocks.

There is a very large major-element mineral-, and major-element-, trace-element-, and isotope- whole-rock database for the entire core, as compiled by Dick et al. (2002) and Natland and Dick (2002). An extensive, quantitative and qualitative fractionation-model has not been demonstrated for the entire crustal section, although attempts have been made for Hole 1105A (Thy, 2003). In this chapter, we use a special version of the Melts program (Melts5) (Ghiorso and Sack, 1995) to evaluate the solid-lines of descent for the lower crustal rocks at Atlantis Bank to quantify the extent of disequilibrium. The effect of melt-composition, oxygen fugacity, water, pressure, and crystallization processes are evaluated. In addition, BASALT (1atm) (Weaver and Langmuir, 1990) and Yang et al. (1996) was used to assess the accuracy of the Melts algorithm in this system.

### 2 Methods

Melts5 (Ghiorso and Sack, 1995; Asimow and Ghiorso, 1998; Asimow pers comm. 2002) provides oxide-compositions, components, and thermodynamic properties of the silicate-, spinel (including ulvöspinel) and some accessory minerals, together with the properties of the evolving melt. This offers an opportunity to directly compare electron-microprobe mineral-analyses of the Atlantis Bank gabbros to model-minerals produced from several starting-compositions (Table 3-1) under different starting conditions. We varied the water-content from dry to 0.5 wt%, the oxygen fugacity from QFM+1 to QFM-1 and the pressure from 1kbar (0.1GPa) to 3kbar (0.3GPa). The original

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**Table 1: The starting compositions of the melts used for our models  
Cp10 had 6% olivine removed and 10% gabbro added (See chapter 3).**

	SWIR PMORB	Site 3 5/2 SWIR	Ridge glass RC27-9-44-14	Cp10, assimilated 10% gabbro	AII107-D20 MAR PMORB
<b>SiO<sub>2</sub></b>	50.3	50.3	49.9	51.1	49.6
<b>TiO<sub>2</sub></b>	0.91	1.94	1.67	0.96	0.63
<b>Al<sub>2</sub>O<sub>3</sub></b>	16.4	15.3	15.9	17.6	17.5
<b>Fe<sub>2</sub>O<sub>3</sub></b>		1.2	1.6	1.1	1.3
<b>FeO</b>		9.0	8.4	5.3	6.3
<b>FeO*</b>	7.45	10.0	9.9	6.3	7.5
<b>MnO</b>	N.d.	0.2	0.2	0.01	N.d.
<b>MgO</b>	10.9	7.6	8.2	8.7	9.9
<b>CaO</b>	11.4	10.6	10.9	12.6	13.0
<b>Na<sub>2</sub>O</b>	2.52	3.35	2.92	2.62	1.65
<b>K<sub>2</sub>O</b>	0.07	0.2	0.11	0.01	0.02
<b>P<sub>2</sub>O<sub>5</sub></b>	0.12	0.2	0.19	0.2	0.12
<b>H<sub>2</sub>O</b>	Variable	0.3	0.05	0.0	0.0
<b>Mg#</b>	72.3	57.4	59.8	71.2	70.4
<b>Ca#</b>	71.4	63.5	67.3	72.7	81.3
<b>Reference</b>	Dick et al, 2000	Robinson, 2000	Johnson and Dick, 1992	Modified from Kinzler and Grove, example 6 (1992)	PetDB, sample location is near Tristan da Cunha Fracture Zone

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compositions were evolved by fractionation of solids in 1°C intervals. Thus, the models represent an approximation to fractional crystallization, with large jumps produced where the model meets phase-boundary reaction-points. We performed batch equilibrium crystallization in only one model-run. We tested our results by using BASALT (Weaver and Langmuir, 1990, "W&L") and the algorithms from Yang et al., (1996), and compared the modeled mineral-compositions to the available data from Hole 735B. Each model run is evaluated using coexisting mafic minerals with plagioclase, together with the composition of the evolving melt.

### 3 Results

Dick et al. (2000) selected a primary melt composition ("SWIR PMORB", Table 3-1) from Kinzler and Grove (1992) as best fitting the bulk composition of Hole 735B. SWIR PMORB has  $\text{Na}_8 = 2.95$  and a  $\text{Fe}_8 = 9.2$  (using the algorithms from Klein and Langmuir, 1987), similar to published glass compositions from the Atlantis II Fracture Zone (Johnson and Dick, 1992). We test whether this is a valid starting composition by theoretically fractionating it at 1kbar, which represent mid- to lower crustal levels at Atlantis Bank. The liquid-line of descent and solid-lines of descent are shown in Figure 3-1. Our starting-composition has no  $\text{Cr}_2\text{O}_3$ , so chrome-spinel is not produced. The primary melt has olivine on its liquidus at 1253°C (1259°C, W&L) and starts precipitating plagioclase after only 6% fractionation at 1220°C (1224°C, 4% W&L). This troctolitic assemblage is precipitated until 26% crystallization (1207°C) (1186°C, 32% W&L), where augite falls on the liquidus. The olivine-plagioclase-augite assemblage is crystallized together to ~1183°C (49% fractionated) (1000°C, 90% W&L), where olivine is replaced by low-Ca pyroxene. BASALT was run at 1atm, and does not crystallize a spinel-phase ("spinel" is ulvöspinel), accounting for this calculated discrepancy at low temperatures. The magnetite/ilmenite boundary is found at 1108°C (78% crystallized), late-stage olivine appears at 1034°C (90% crystallized), and apatite starts precipitating late, at 950°C (97.5% crystallized). The algorithm developed by Yang et al. (1996), indicates temperatures ~50°C lower than Melts5 (Fig. 3-11B), and we will include this

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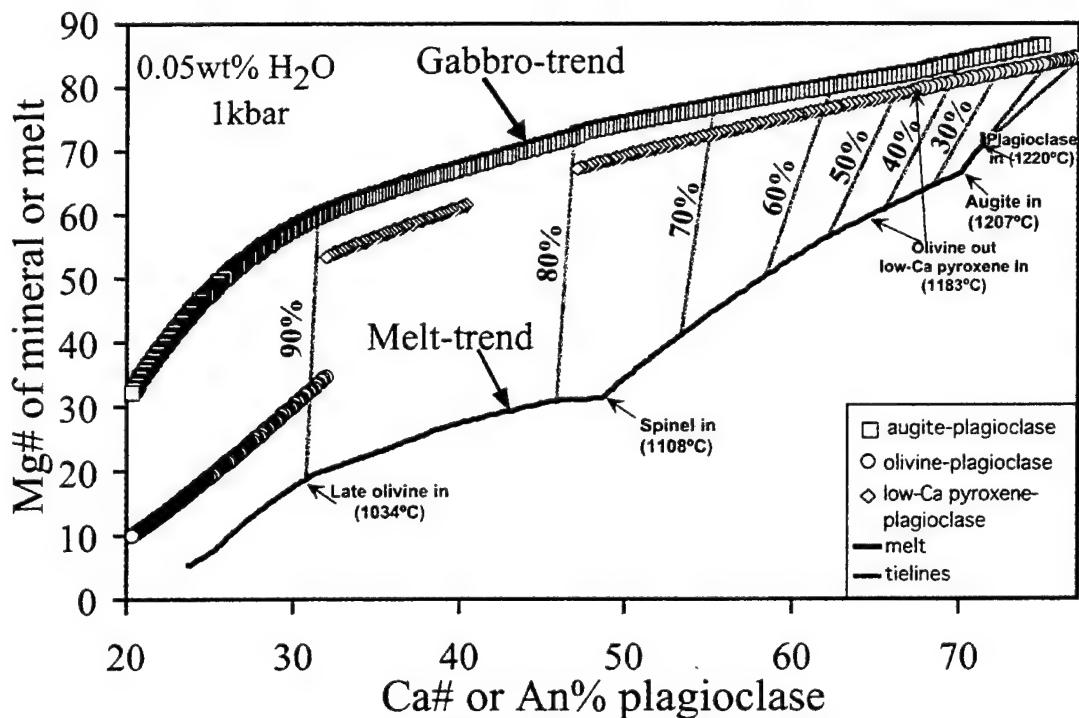


Figure 3-1: The complete model run for SWIR PMORB with 0.05wt% H<sub>2</sub>O at 1kbar and QFM. The symbols represent the An% (=100\*Ca/(Ca+Na), K is near absent from the system) of plagioclase co-precipitated with augite (squares), olivine (circles) and pigeonite (diamonds) as represented by their Mg#'s (100\*Mg/(Mg+Fe)). The black line represents the Mg# and Ca# of the evolving melt, and the gray lines connect the melt and mineral-pairs at 10% crystallization intervals relative to the original melt.

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discrepancy in the later evaluation, but will continue to use the 1kbar-model as a baseline-model.

Crystallization of the SWIR PMORB (damp and dry) at 3kbar and QFM (Fig. 3-2) has subtle differences between them. Three kbar at Atlantis Bank represents depths equal to crystallization within the mantle (e.g. 9-km). In the dry model (Fig. 3-2A), olivine is on the liquidus early ( $1270^{\circ}\text{C}$ ), but ceases to crystallize ( $1241^{\circ}\text{C}$ ) before plagioclase starts and is replaced by high and low-Ca pyroxene. Thus, the primitive, plagioclase-bearing cumulates are gabbronoritic in composition and olivine does not show in the plot. Olivine comes back on the solidus early ( $1176^{\circ}\text{C}$ , after  $\sim 70\%$  crystallization) and the ulvöspinel-phase starts crystallizing at  $1119^{\circ}\text{C}$  (85%). The damp melt (Fig. 3-2B) has two main differences from the dry melt: 1) Late-stage olivine comes back on the liquidus at  $40^{\circ}\text{C}$  below that of the dry melt. 2) The plagioclase has higher An% at any given mafic mineral Mg# and melt Ca#, as is clear from the right-slanted tilt in the gray lines that joins the crystallization steps of melt and the cumulates.

The basaltic carapace that once overlay Atlantis Bank is exposed on a lithospheric flow-line in the rift mountains  $\sim 55$ -km north of the ridge. Robinson et al. (2001) analyzed pillow-basalt glasses from this reciprocal basaltic carapace, and we have used one of these samples for a test of the models of the sensitivity to the oxygen fugacity (Table 3-1). It should be noted that since there are no Cr in these analyses, crome-spinel is not a part of these calculations. The glasses contained some water (0.27 wt%), which we use in the model shown in Figure 3-3. At QFM+1 (more oxidized) the model gives very high forsterite-contents for olivine relative to augite, and augite crystallizes early (Fig. 3-3A). This occurs as the olivine only takes up  $\text{Fe}^{2+}$  ions, making the effective Mg# higher than at lower oxygen fugacities. The high oxygen-fugacity model fails to lower the Mg# of the solids and liquids as the melt evolves, as a large fraction of the ulvöspinel-phase crystallizes (it uses 5 wt% of the original liquid to crystallize), keeping the Fe low in the melt. Thus, there is no late-stage olivine. The reduced case (QFM-1) sees the opposite effect (Fig. 3-3B). The spinel-phase is nearly suppressed, with lower melt and solid Mg#'s. relative to plagioclase. Augite crystallizes after 30% fractionation.

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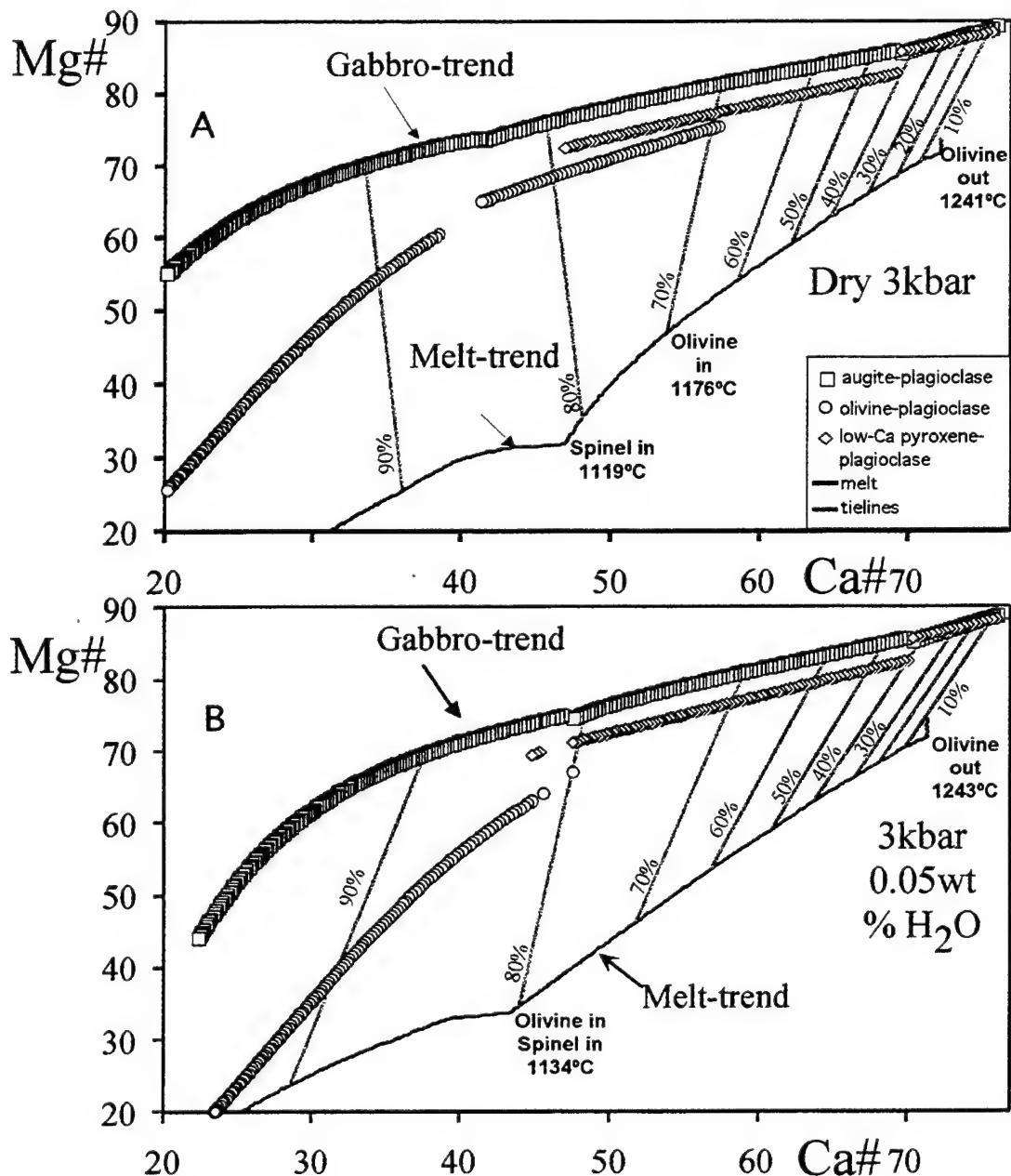


Figure 3-2: Higher pressure crystallization. The symbols are the same as in Figure 1. **A:** The crystallization sequence of SWIR PMORB at dry conditions and 3kbar pressure. **B:** The same melt crystallized with 0.05wt% H<sub>2</sub>O in the starting composition.

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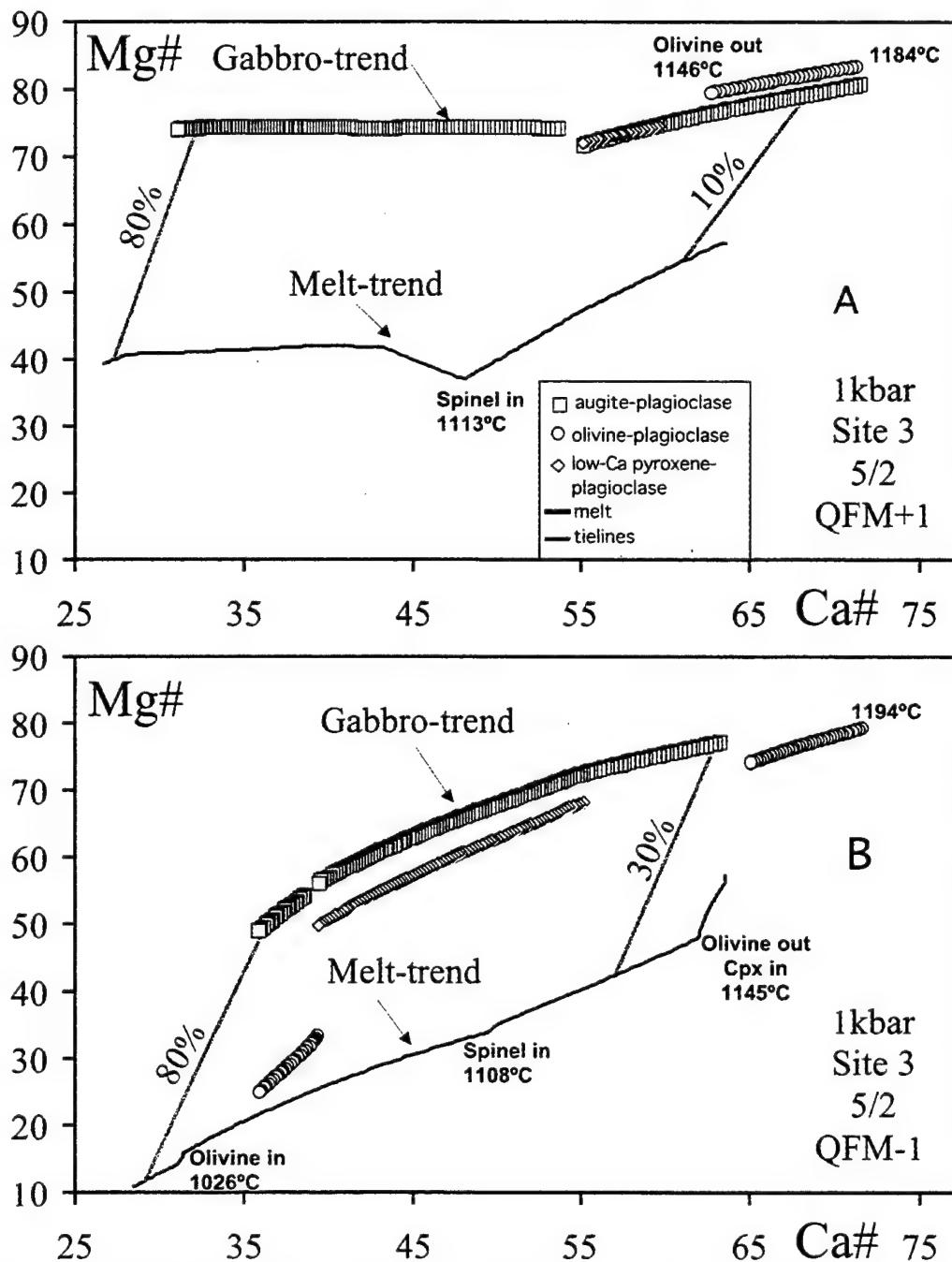


Figure 3-3: Crystallization of a glass from north of the ridge segment, at the upper crust produced reciprocally to Atlantis Bank, Sample Site3-5-2 (Robinson et al., 2001, Table 3-1). **A:** 1kbar, at QFM+1 log-unit. Only the tie-line at 10% crystallized and 80% crystallized are shown. At QFM+1 the melt is near saturated in plagioclase-augite-olivine. **B:** The same melt crystallized at QFM-1. Augite is suppressed until 34% crystallized.

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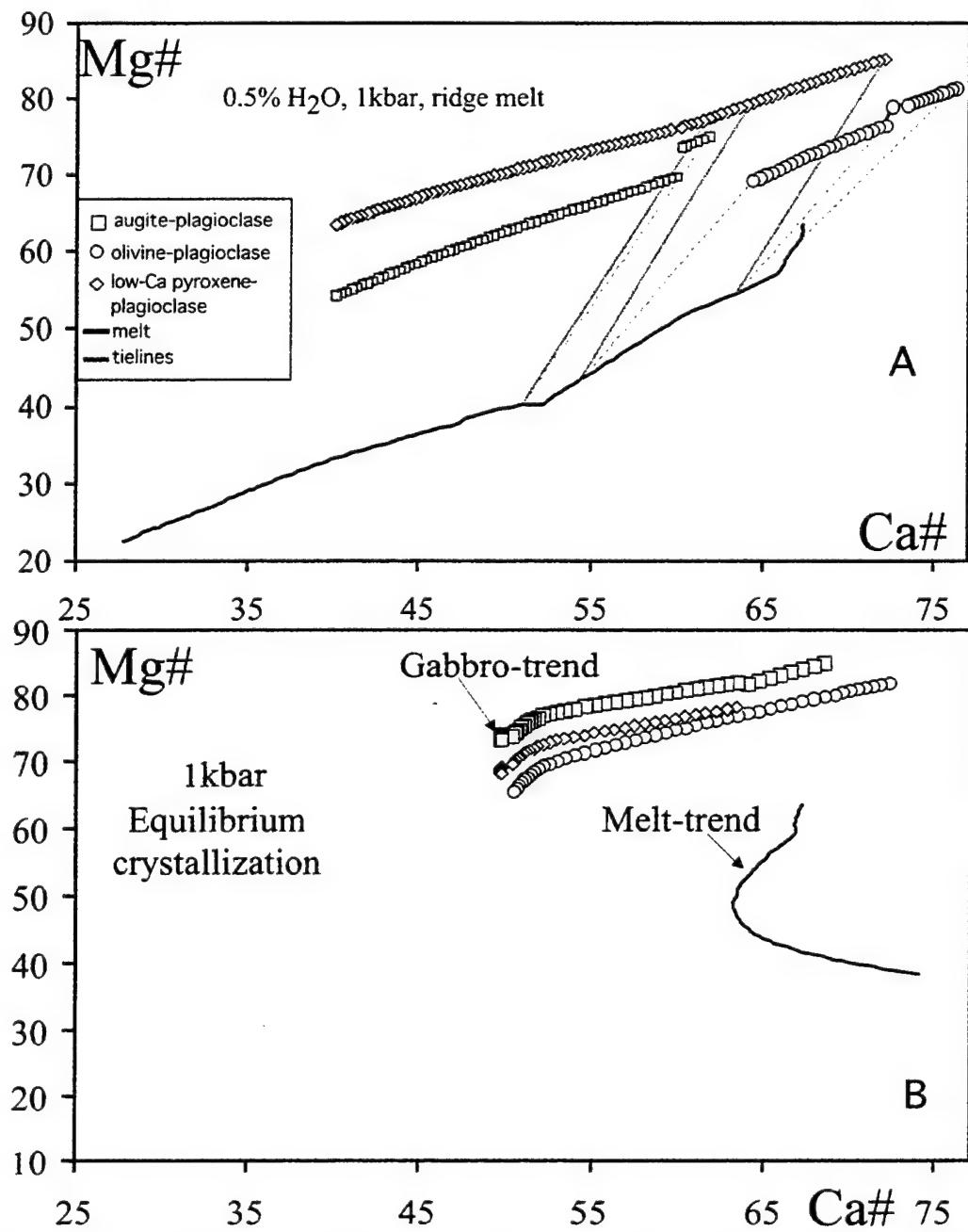


Figure 3-4: **A:** Crystallization of RC27-9-44-14 with 0.5 wt% H<sub>2</sub>O at 1kbar and QFM. **B:** Bulk crystallization of the same melt, where the solid is not separated from the liquid. Note how the melt compositions turns around and becomes higher in Ca# with crystallization.

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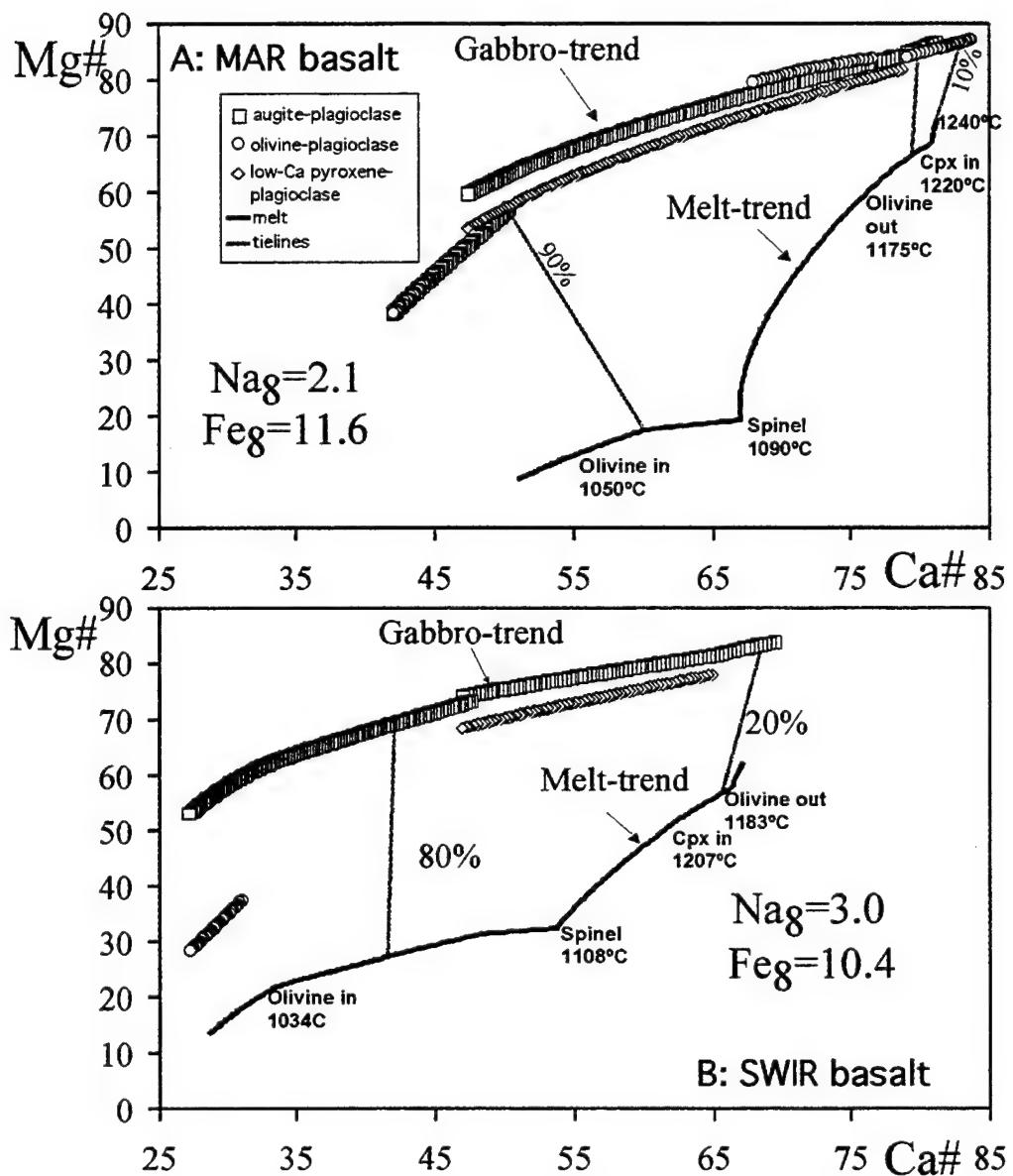


Figure 3-5: Note that the axes are slightly different from Figure 1-4. A: Crystallization of AII107-D20 at QFM and 1kbar. Tie-lines for 10 and 90% crystallized are shown, together with augite-in (23% crystallized). B: Crystallization of dry RC27-9-44-14 at 1kbar. Tie-lines for 20 and 80% crystallized (with respect to the starting-composition) are shown.

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Water has a strong effect on the solid-line of descent. Figure 3-4A shows the fractionation of glass RC27-9-44-14 collected from Atlantis II Fracture Zone at 1kbar and QFM with 0.5% of water added (Table 3-1). Compared to the previous models, it has elevated plagioclase anorthite and no late-stage olivine as Ca partitions more strongly into the plagioclase than the melt at such hydrous conditions (Hough and Luhr, 1991). Nonetheless, the Mg# of the augite is still high, while the Fo-content of olivine is low relative to drier model runs.

Batch equilibrium crystallization, where the solid and the liquid stay in equilibrium throughout solidification, is very different from the previous models (Fig. 3-4B). The model terminates for plagioclase around An<sub>50</sub> and no mafic mineral has Mg#'s below 65.

The composition of the original melt strongly influences how the composition of the solid will evolve. Figure 3-5A shows a dry fractionation of a southern Mid-Atlantic ridge basalt glass from near Tristan De Cunha (PetDB) compared to that for the Atlantis II Fracture Zone glass above (Fig. 3-5B). Note that the abscissa is wider in this figure than in the others. The Atlantic basalt is more depleted than the Southwest Indian Ridge basalt, with much lower Na<sub>8</sub> and higher Fe<sub>8</sub>. This produces a model result with composition of the solid having higher plagioclase anorthite contents at the same Mg# of the solids, and little plagioclase forming below An<sub>50</sub>. In the Southwest Indian Ridge-sample, however, the plagioclase has much lower An-contents, and 1/3 of the plagioclase crystallizes below An<sub>50</sub>.

## 4 Discussion

### 4.1 Mode of accretion of the gabbros in Hole 735B

We now compare our models to the existing data from Hole 735B (Dick et al., 2002) (Fig. 3-6) using co-existing minerals in the gabbros. Many authors have described the compositional variations in the minerals in Hole 735B (Bloomer et al., 1991; Dick et al., 1991a; Natland et al., 1991; Ozawa et al., 1991; Hébert et al., 1991; Dick et al., 1992; Dick et al., 2000; Natland and Dick, 2001; Thy and Dilek, 2000). Hole 735B has 953

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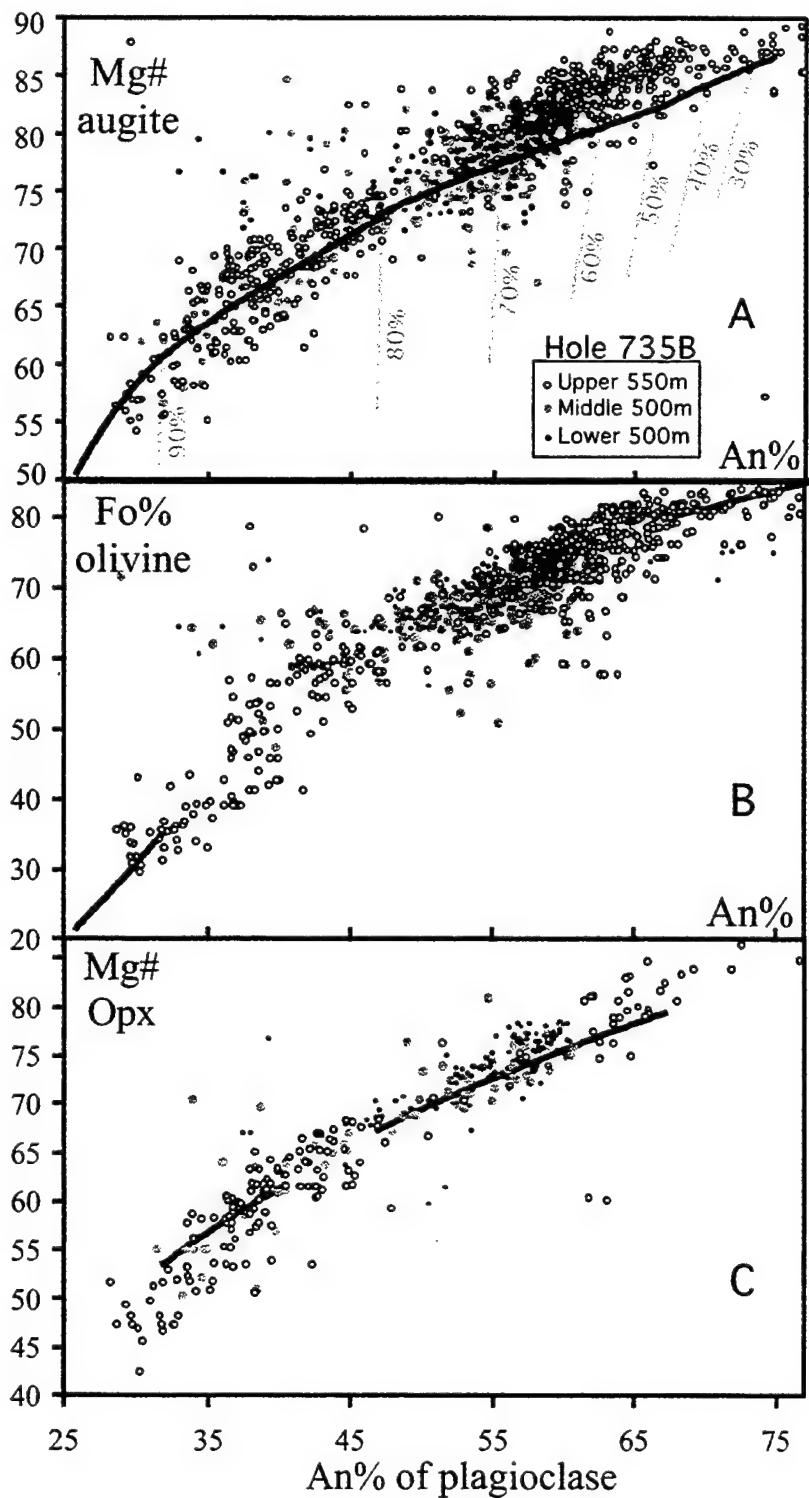


Figure 3-6: Average compositions of coexisting minerals from Hole 735B compared to the model (black model-line and gray tie-lines to the melts) in Figure 1. Open symbols are the top 550 meters, gray symbols are the middle 500 meters and the black symbols are the bottom 500 meters of Hole 735B. A: The Mg# of augite with the An% of coexisting plagioclase. If made possible by the dataset (Dick et al., 2000), cores are plotted with cores, and rims with rims. B: The forsterite content of olivine with the An% of the coexisting plagioclase. Note how the model calls for a break in olivine crystallization, whereas olivine is near ubiquitous in the core. C: The Mg# of low-Ca pyroxene in the core compared to the co-existing An% of plagioclase.

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discrete described igneous intervals, with additional subdivision possible on centimeter scales (Dick et al., 2000). Plagioclase and clinopyroxene are ubiquitous, and the gabbro-types are distinguished largely by variations in grain size and amount of olivine, oxides (magmatic ilmenite and titanomagnetite) and orthopyroxene. The average grain-size varies from 5-15 mm to 15-30 mm (the definition of very coarse grain-size is >30 mm) and crystals in excess of 10-cm have been found. Only a small fraction of the lithologic contacts are simple faults rather than involving a magmatic origin. The igneous stratigraphy is divided into 12 polygenetic units, but are separated into three major associations: 1) the dominant, oxide poor, olivine-bearing gabbros (and coarse-grained troctolites), 2) cross-cutting microgabbros and fine- to medium-grained troctolite, and 3) ferrogabbros often highly deformed relative to the olivine gabbro that encompass oxide-bearing gabbros and gabbronorites. The nature, origin and number of geochemical units are much debated. Of the studies that involve the entire core, Dick et al (2000) demonstrate five major olivine-gabbro units from 200 to 600m thick, crosscut by many ferrogabbro units that diminish in size and number down-hole. The ferrogabbros were interpreted to represent the intrusion of late interstitial melts that migrated from the cumulate into shear-zones during the detachment faulting that exposed the core-complex. The fraction of oxides down-hole is shown in Figure 3-7. Natland and Dick (2001) identified 97 separate olivine gabbro sequences, arguing that each represent an individual injection of magma forming two principal bodies of olivine gabbro. Thy and Dilek (2000) suggested the core represents four individual blocks of gabbro.

For the purpose of this study, the mineral-data is separated into the upper 550-m, the middle 450-m and the bottom 500-m, and is compared to the 1-kbar SWIR-PMORB model (Fig. 3-6). The upper 1000 meters of Hole 735B have a range of compositions best described by fractional crystallization, whereas the lower 500 meters terminate at the plagioclase anorthite contents similar to the batch-equilibrium crystallization-model (Fig. 3-4B) making it difficult to distinguish these processes there based on major elements alone. The more than 1500 plagioclase and clinopyroxene-averages have been meticulously sorted, so that cores are plotted with cores and rims with rims. This, and as

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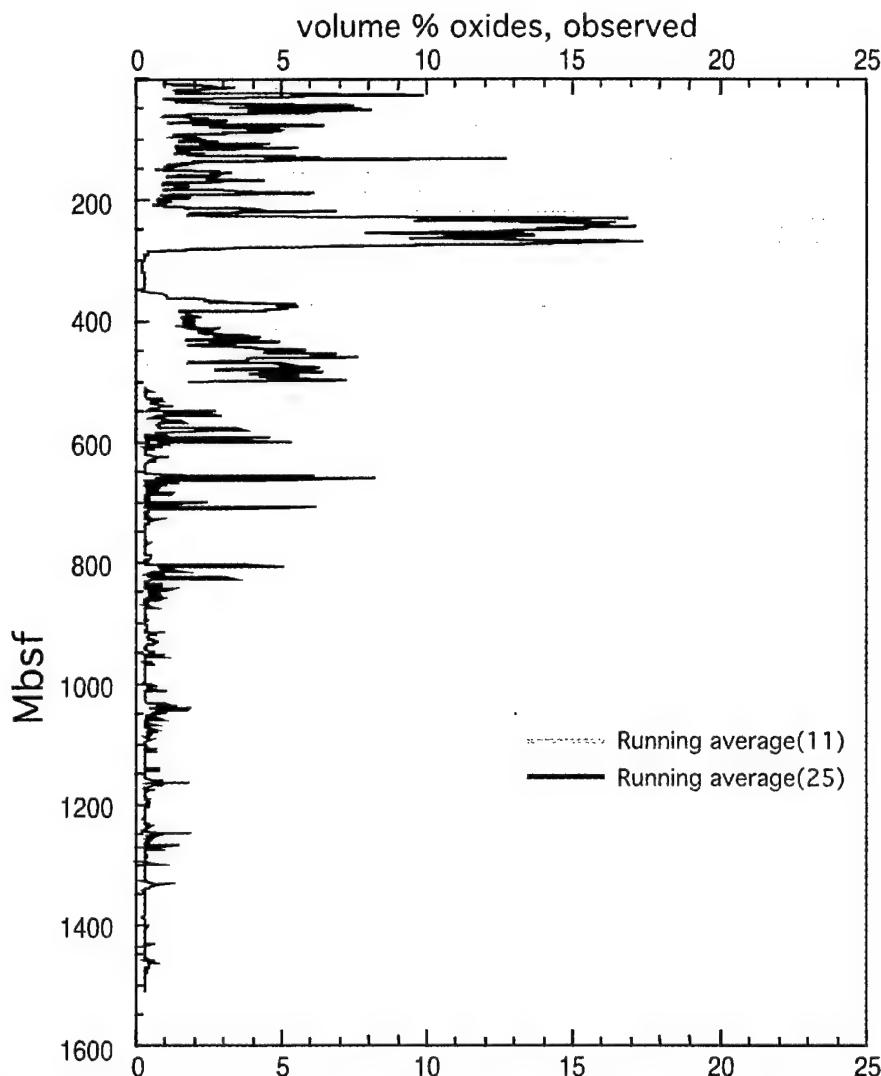


Figure 3-7: Observed volume% of oxides in Hole 735B. The lower 1000 meters are from Dick et al. (2000), whereas the upper 500 meters were logged by inspection with a hand-lens at TAMU. Variations in the volume of oxides were logged so that a total of ~1500 observations were made. The gray lines represent running averages of 11 observations, the black lines represent running averages of 25. As much as 50% oxides were observed in the core.

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errors in the analyses are small ( $\pm 0.5$  mol % An, Fo or Mg#), the scatter seen in the plot is not due to errors in the large data-set. A bimodal mineral-composition distribution is obvious in the upper third of the hole, as described by Bloomer et al (1991) and the many others mentioned above. The lower ~1000 meters do not have this bimodality, with plagioclase < An<sub>45</sub> rare, and the sections become progressively more homogeneous downwards. In the upper 550-m, the cross-cutting fine to medium-grained troctolites represent 25-45% crystallization, the olivine gabbros 45-70% crystallization, and the ferrogabbros represent 80-90% crystallization (by volume or weight) relative to the modeled primary melt composition (SWIR PMORB). The cross-cutting troctolites are the most primitive rocks in Hole 735B and differ significantly from the coarser, conformable troctolites deeper in the Hole. The olivine gabbros from the lower 1000-m were produced from melts fractionated > 60% elsewhere, and crystallized more than 80%. Thus, if mantle-derived melts do form these gabbros, and crystallization shielded from mantle olivine has not occurred in the mantle, there are at least 700 vertical meters of primitive crust somewhere else in the section not yet drilled (Table 3-2). Crystallization of melts in the presence of mantle olivine would, like Meyer et al. (1989) suggested, keep a constant Mg# of the magma while the Ca# of the melt evolves. The 3 % troctolites in the upper 550-m make up far less volume than the 20% expected from the model, suggesting ~70 meters of missing troctolites, making the total missing cumulates 770 vertical meters. If the gabbro complex becomes more differentiated towards the ends of the paleo-ridge segment, which is likely, then the undrilled cumulate-section would become even longer. In addition, escape of differentiated melts from the cumulate-pile, similar to erupted MORB's in the region would also increase the estimated thickness of the undrilled –section proportionally.

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**Table 3-2: Amount crystallized relative to a mantle derived melt for Hole 735B and the calculated undrilled section on the basis of the fractional crystallization model.**

Unit:	Troctolite	Olivine gabbro	Ferro gabbro	Undrilled section
Upper 550m	25-40%	40-80%	80-90% (+)	137m + 12% (troct) (=66m) ~200m
Lower 950m		60-80% (+)		570m
Total undrilled section				770m

As is described in Chapter 4, augite is texturally younger than the coexisting plagioclase in the rocks. Mineral-pairs that plot below the solid line of descent, towards more refractory plagioclase or less refractory augite, are consistent with this textural observation and the simple progressive crystallization model. Major parts of Hole 735B plot above the model solid-line of descent, however, and thus late crystallization of more evolved interstitial melts cannot explain them. The model also predicts a pause in olivine crystallization, but olivine is close to ubiquitous in the core, and there is no compositional discontinuity in forsterite content. It is important to note that the olivine – plagioclase pairs in the evolved oxide-gabbros generally do follow the same path as the model, indicating that the ferrogabbros have internal equilibrium (Fig. 3-6b). Low-Ca interstitial pyroxene (orthopyroxene) is also common throughout the core, at both the evolved and primitive end of the spectrum with a much wider range than the model predicts. It is therefore possible that there are reactions between pyroxene, olivine and silica over- and undersaturated interstitial melt, due to the peritectic reaction between olivine and melt that produce pyroxene and vice versa.

The primitive end of the model indicates that olivine is fairly close to equilibrium with plagioclase. However, if the system is in equilibrium, the coexisting mafic minerals also should be in equilibrium. Figure 3-8 compares Hole 735B data to the solid line of descent for SWIR PMORB. Again, if the system is well described by a fractional crystallization model, there should be no points plotting significantly above these lines whereas most of the refractory gabbros do. The extremely evolved oxide-olivine gabbros, however, follow the trend beautifully, and the iron-rich late-stage olivine defines the QFM-buffer with coprecipitating fayalite and magnetite. According to our model calculations, these late-stage rocks must be crystallized from melts that have fractionated

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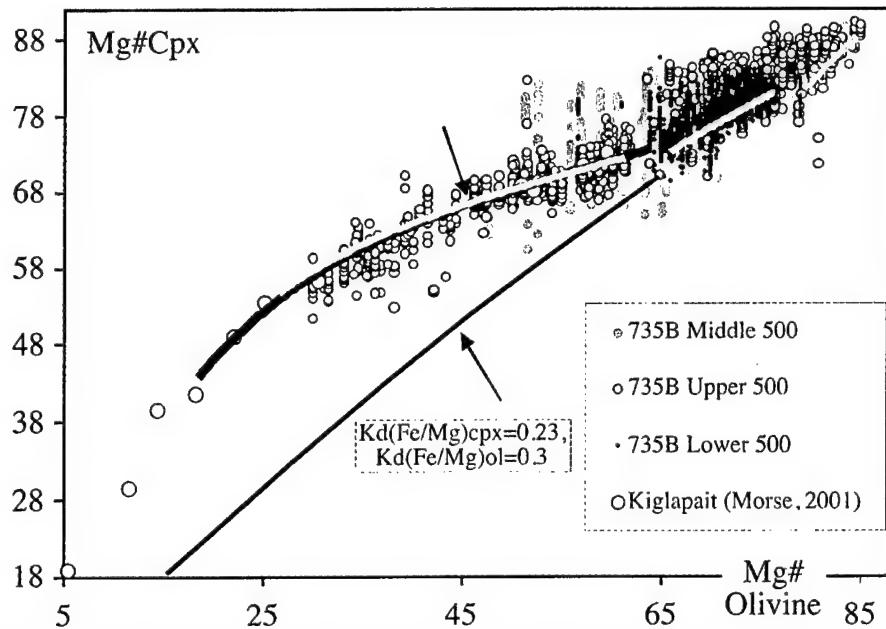


Figure 3-8: The augite coexisting with olivine of Hole 735B, as represented by their respective Mg#’s and Fo%. The augite-analyses represent all the analyses for each crystal, the olivines are averages. The thin black line represent a hypothetical olivine-augite pair in equilibrium with a melt as calculated by often-used partition-coefficients (as indicated). The thick gray line to the upper right represent the 1kbar model, the two to the left represent the 3kbar dry model, whereas the thick black curve represents the damp 3kbar model. Also included are data from the layered intrusion in Kiglapait (Morse, 2001), demonstrating how the ferrogabbros from Hole 735B are similar to the sandwich-horizons in layered intrusions.

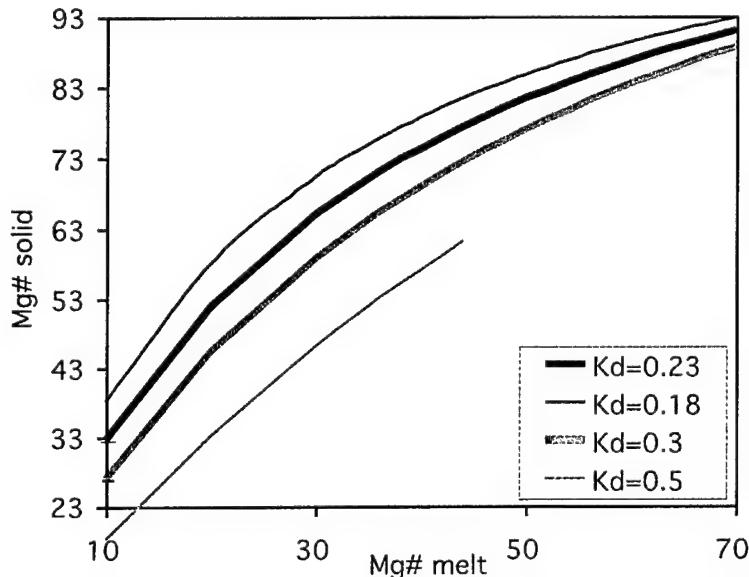


Figure 3-9: This figure demonstrates how the partition-coefficients for iron and magnesium between the crystals and melt control the variation of the Mg#’s of the crystals. Higher magma-temperatures makes the Mg#’s more similar (thick lines), whereas lower temperatures make them diverge (thin lines).

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more than 90% relative to SWIR PMORB. Thus the liquid-line of descent must have been controlled by near fractional crystallization at crustal pressures on the QFM-buffer.

### 4.2 The extent of disequilibrium in the refractory gabbros

The partition-coefficient dependency between the minerals and melts can be described by:

$$Kd_{Fe/Mg}^{xtl/melt} = \frac{X_{FeO}^{xtl}}{X_{FeO}^{liq}} * \frac{X_{MgO}^{liq}}{X_{MgO}^{xtl}}; \quad (1) \text{ so}$$

$$Kd_{Fe/Mg}^{xtl/melt} = \frac{Mg\#_{melt} - Mg\#_{melt} * X_{Mg}}{X_{Mg} - X_{Mg} * Mg\#_{melt}}, \quad (2a), \text{ therefore}$$

$$X_{Mg} = \frac{Mg\#_{melt}}{Kd_{Fe/Mg}^{xtl/melt} - Mg\#_{melt} * Kd_{Fe/Mg}^{xtl/melt} + Mg\#_{melt}}, \quad (2b)$$

$$Mg\#_{melt} = \frac{X_{Mg} * Kd_{Fe/Mg}^{xtl/melt}}{1 - X_{Mg} + X_{Mg} * Kd_{Fe/Mg}^{xtl/melt}}, \quad (2c), \text{ and}$$

$$Kd_{Fe/Mg}^{xtl1/xtl2} = \frac{X_{Mg}^{xtl1} - X_{Mg}^{xtl1} * X_{Mg}^{xtl2}}{X_{Mg}^{xtl2} - X_{Mg}^{xtl1} * X_{Mg}^{xtl2}}, \quad (2d).$$

(Roeder and Emslie, 1970; Wood and Blundy, 1997) where  $X_{Mg}^{phase}$  is the Mg# (molar Mg/(Mg+Fe) for these formulas only, 100\* Mg/(Mg+Fe) for the rest of this chapter) of a phase (usually the solid unless mentioned), and  $Kd_{solid/melt}^{Fe/Mg}$  are the partition coefficients between a solid and a liquid. Equation 2d describes the partitioning of Fe/Mg between two minerals. It is known that  $Kd_{augite/melt}^{Fe/Mg}$  evolves with fractionation from 0.23 towards 0.17 (all Fe as FeO) (Grove and Bence, 1977).  $Kd_{solid/melt}^{Fe/Mg}$  is generally  $0.3 \pm 0.03$  for melts that have Mg# > 45 (Roeder and Emslie, 1970). The primitive end of the fractionation-trend is not very sensitive to the partition-coefficients (Figure 3-9), in fact the entire spectrum of  $Kd_{augite/melt}^{Fe/Mg}$ , 0.23-0.17 (Grove and Bence, 1977), only increases the Mg# by ~2.5 mol%. Toplis and Carrol (1995) described  $Kd_{ol/melt}^{Fe/Mg}$  as high as 0.5 in ferro-basalts. Therefore, the evolved melts are far more sensitive to partition-changes, as  $Kd_{augite/melt}^{Fe/Mg}$  is much lower and  $Kd_{olivine/melt}^{Fe/Mg}$  much higher.

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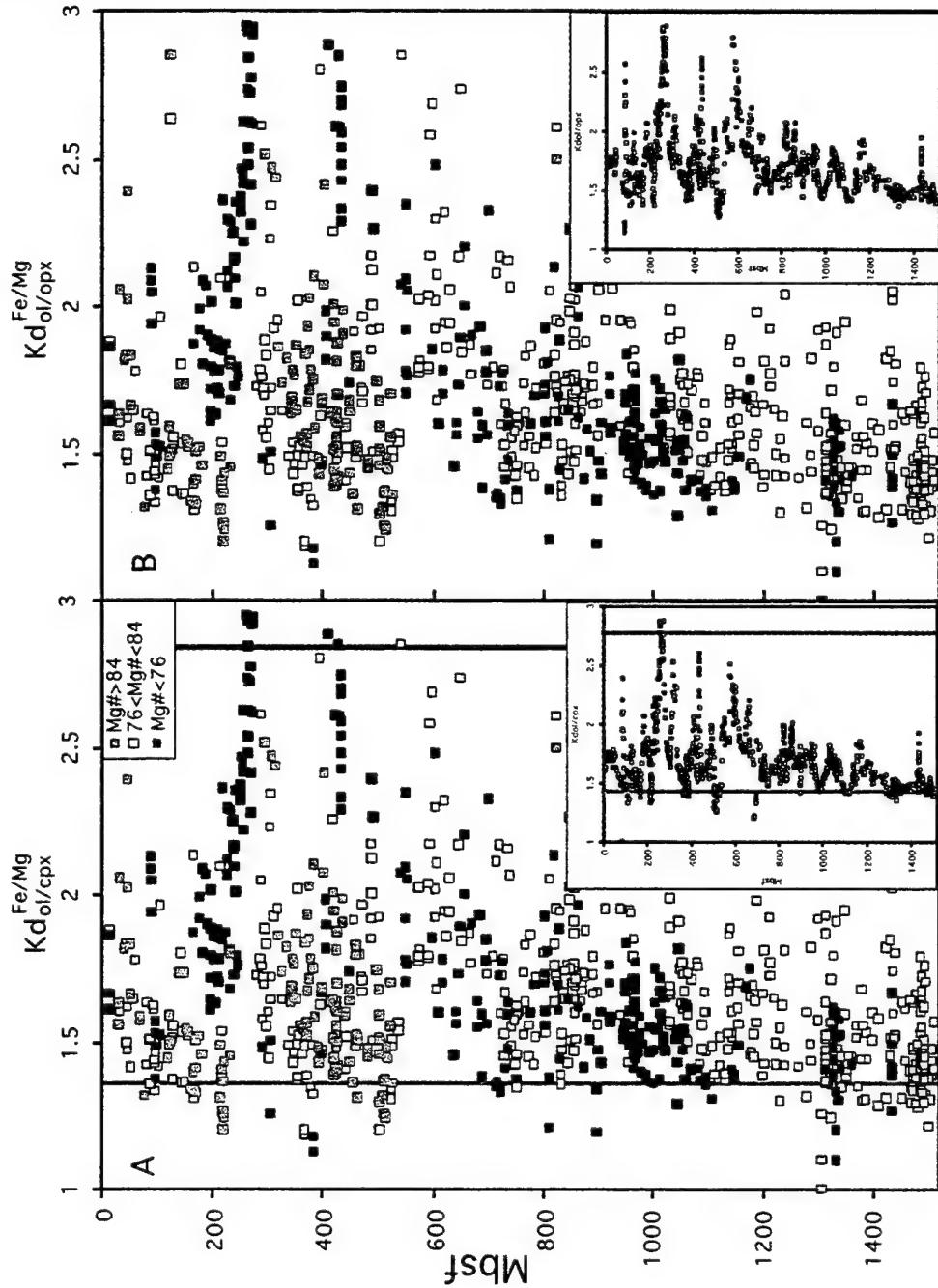
Both  $Kd_{\text{olivine/augite}}^{\text{Fe/Mg}}$  and  $Kd_{\text{olivine/oxp}}^{\text{Fe/Mg}}$  vary inversely with temperature (Loucks, 1996; Braun, pers. comm. 2004) and is used to subtract the effect of temperature of crystallization from the actual geochemical offset provided by the post-cumulus processes and quantify the disequilibrium with depth in Hole 735B relative to the Melts-model. Figure 3-10 shows the partitioning between coexisting olivine-augite and olivine-orthopyroxene, so they can be compared to the equilibrium-lines for the primitive and evolved gabbros. For  $Kd_{\text{olivine/augite}}^{\text{Fe/Mg}}$  (Fig. 3-10A) the gray line to the left indicates a  $Kd_{\text{olivine/augite}}^{\text{Fe/Mg}}$  of ~1.4 produced by  $Kd_{\text{augite/melt}}^{\text{Fe/Mg}} = 0.23$  and  $Kd_{\text{ol/melt}}^{\text{Fe/Mg}} = 0.33$ , the right line represents a  $Kd_{\text{olivine/augite}}^{\text{Fe/Mg}}$  of ~2.8 calculated from  $Kd_{\text{augite/melt}}^{\text{Fe/Mg}} = 0.18$  and  $Kd_{\text{ol/melt}}^{\text{Fe/Mg}} = 0.5$ , respectively (op. cit.). The geothermometer of Loucks (1996) provide  $Kd_{\text{olivine/augite}}^{\text{Fe/Mg}}$ 's that plot non-linearly with temperature.

The expected decrease in crystallization temperature for Mg# 89-84 augite increases in  $Kd_{\text{olivine/augite}}^{\text{Fe/Mg}}$  by 0.15. We used the baseline-model of PMORB SWIR as a reference for augite Mg#'s (Fig. 3-11) and recalculated the crystallization-temperatures of the melts using the algorithm from Yang et al. (1996) as Melts appears to give crystallization-temperatures ~50°C too high.

Similarly, over the compositional range of the intermediate gabbros (white squares in Fig. 3-10) (Mg#84-76) the  $Kd_{\text{olivine/augite}}^{\text{Fe/Mg}}$  can increase by 0.51. Examination of Figure 3-10 shows that many olivine-gabbros with augite > Mg# 84 plot far to the right of the 0.15 shift permitted by equilibrium, and many of the intermediate gabbros plot to the right of the maximum 0.51 shift. The oxide-gabbros (black squares in Fig. 3-10), however, should and do extend to the far right side of the diagram, representing crystallization of extremely evolved ferro-basalts. If we assume that the lowest value for  $Kd_{\text{olivine/augite}}^{\text{Fe/Mg}} = 1.35$  is correct for the most primitive gabbros and troctolites the temperature-dependence will only increase the Mg# of augite by 1.5 mol% for Fo<sub>80</sub> or 6 mol% for Fo<sub>70</sub>. Thus, the olivine-augite system is out of equilibrium in the gabbros: more than 5 mol% Mg# for the troctolites and 10 mol% Mg# for the olivine gabbros, leaving ~3.5 and ~4 mol% Mg# unaccounted for, respectively. It could be suggested that some

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Figure 3-10:  
 Partitioning of Fe and Mg between olivine and augite (A), and olivine and low-Ca pyroxene (B). The gray symbols indicate augites with  $Mg\# > 84$ , the black symbols have  $Mg\# < 76$  and the white squares are the intermediate compositions. The insert represents a running average of 9 analyses. The gray lines represent refractory (left) and evolved (right) partition-coefficients, as described in the text.



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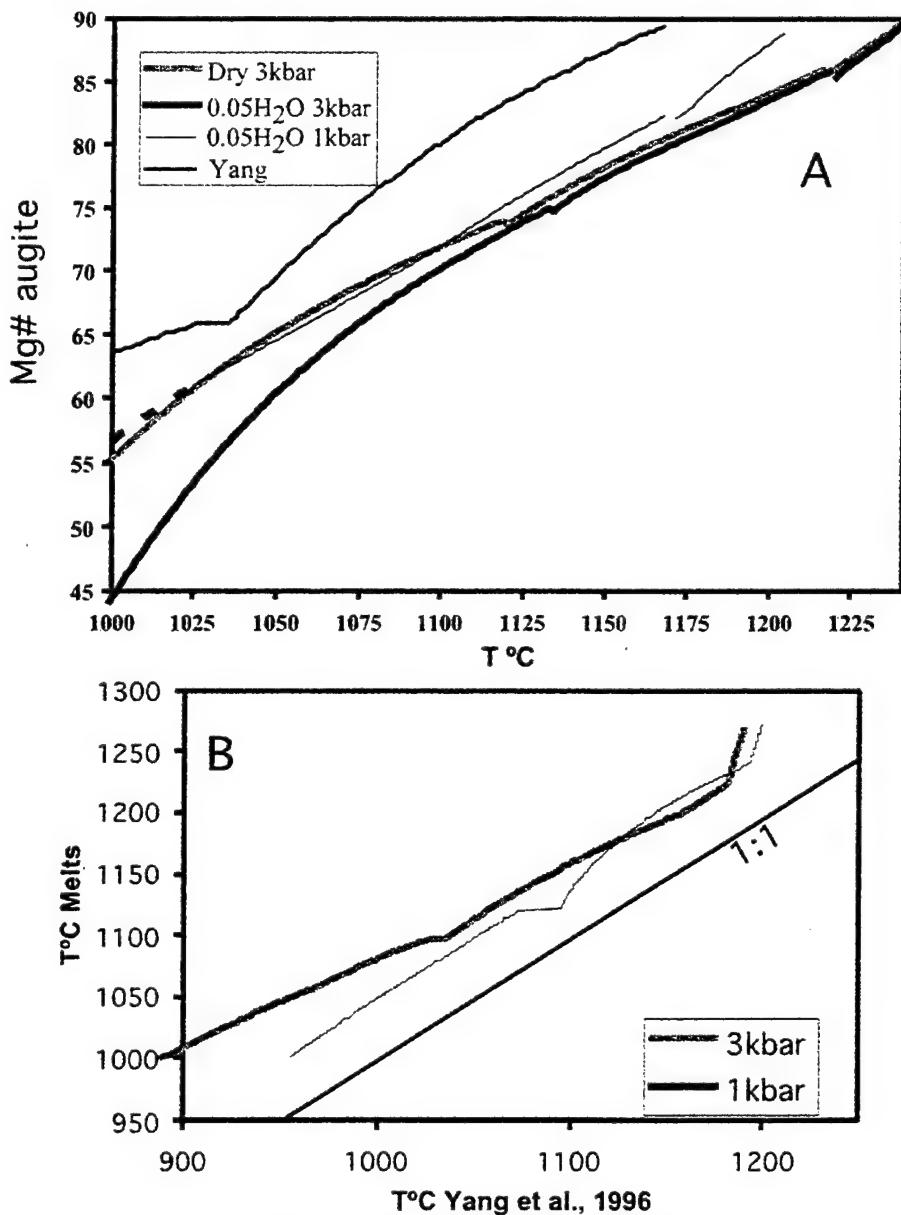


Figure 3-11: **A:** The variation of Mg# of augite in the models with respect to temperature. The line for Yang et al. (1996) is calculated on the 1kbar model melts using a Fe/Mg partition coefficient for augite/melt of 0.23. **B:** Comparison of the temperatures resolved using Melts5 and Yang et al (1996), indicating that there is a discrepancy of ~50°C between the models.

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of these effects were due to closure temperature effects, however the extent of disequilibrium of augite with respect to plagioclase in a rock is unaffected by the presence or absence of olivine. Likewise, the analyses of the augite have purposefully included exsolution-lamellae of orthopyroxene within the clinopyroxene in order to approximate the igneous composition.

Therefore, factors other than fractional crystallization at 1kbar must have affected the rocks. Note that some of the temperature-dependence discussed above is incorporated into the Melts algorithm, as will become apparent when the model is compared to those by other authors' models.

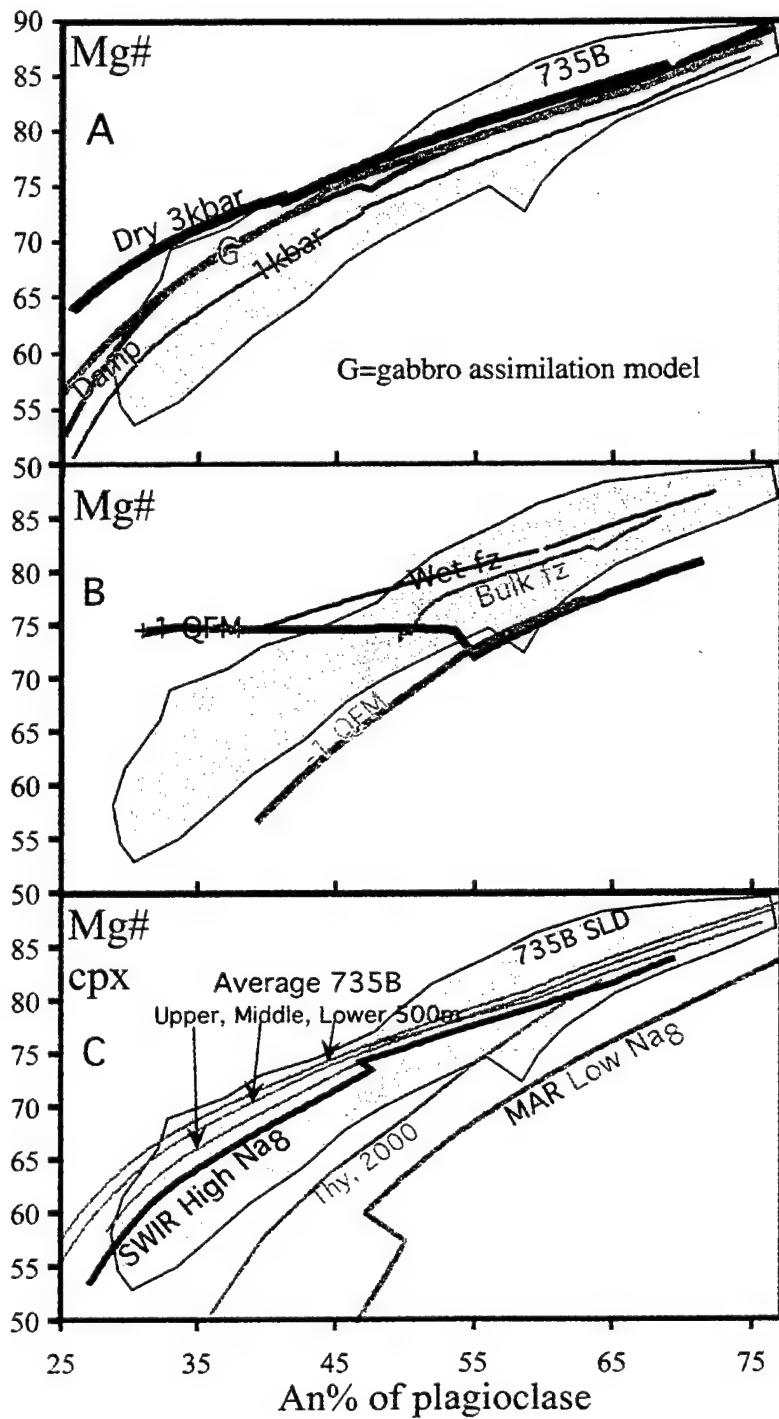
### 4.3 What caused the high-Mg augite?

Consistently, then, augites of the Hole 735B olivine-gabbros are more Mg-rich than predicted from the composition of the coexisting olivine and plagioclase. The formation of cumulus rocks, however, invokes two separate processes: accumulation of crystals; and postcumulus solidification and consolidation of the crystal pile (Wager et al., 1960; Hess, 1960). Although these processes may be related, they have different effects on the resulting rocks and should be treated separately, particularly for potentially open magmatic systems.

#### 4.3.1 Accumulation of crystals, the solids:

The conditions that potentially can affect the composition of the accumulating crystals during crystallization are shown in Figure 3-12. This figure compares solid lines of descent calculated from fractionation-models for different conditions and melt compositions with the coexisting plagioclase and augite of Hole 735B. We have also included a model derived from the work in Chapter 3, where 10 wt% of gabbro is assimilated into a primitive melt before emplacement. Near-fractional crystallization at 3kbar creates a higher Mg# of augite relative to the field for Hole 735B, while the 1kbar model for the SWIR PMORB composition best fits the observed gabbro-compositions overall. The damp melt, containing 0.05% H<sub>2</sub>O, clearly fits the evolved end of the spectrum better than the dry melt. However, if Southwest Indian Ridge-type magma

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assimilated only 10% gabbroic material on ascent and subsequently crystallized at 1kbar, the solid-line of descent is similar to the one produced by the uncontaminated 3-kbar model (Fig. 3-12a). Troctolite-assimilation follows the same trend, although augite is suppressed until 27% crystallization, compared to 18% for gabbro assimilation.

The effect of bulk-composition is illustrated in the Figure 3-12C, where the AIIFZ-glass (“High  $\text{Na}_8$ ”) plots within the field for Hole 735B, while the solid line of descent for a MAR-glass (Low  $\text{Na}_8$ ) plots below it. The model of Thy (2000), where the author used a glass from AIIFZ at constant  $K_d$ ’s and 1atm crystallization, fits the data poorly, as it does not account for temperature-dependence. Thy used another starting-composition with extremely high  $\text{Na}_8$  (~3.7), a basalt-glass from west of the Rodrigues Triple Junction (“Type 3 basalt”, Natland, 1991b). This model (not shown) fits the data much better, and plots higher than the AIIFZ glass, but the lava is very uncharacteristic of magmas found west of the Melville FZ (Mével et al., 2003) and is an unrealistic parental magma. We also modeled the solid-lines of descent for the bulk-composition of the lower, middle and upper 500-m of Hole 735B. The  $\text{Na}_8$  of the bulk becomes higher while  $\text{Fe}_8$  decreases deeper in the Hole (3.1, 3.3, 3.3 and 10.3, 9.3, 8.4 respectively), resulting in fractional crystallization models that have higher Mg# augites at plagioclase with the same anorthite content. However, the Mg#’s are not as high as those in Hole 735B. In fact, only the average upper 500-m of Hole 735B falls on the fractionation-trend of the basalts in the area.

The wet melt (0.5%  $\text{H}_2\text{O}$ ) model fits the primitive end of the gabbro-field, but overestimates the Mg#’s of augite in towards the evolved cumulates. Varying oxygen-fugacities using Site 3 lavas demonstrates that higher oxygen fugacities stabilize augite earlier on the liquidus relative to more reducing conditions (Fig. 3-12B), and when ulvöspinel becomes stable, the Mg#’s of augite become near constant with further crystallization. Thus, very locally, increasing oxygen fugacity due to crystallization could produce the Hole 735B gabbros plotting above the solid-line of descent at the evolved fractionation-trend.

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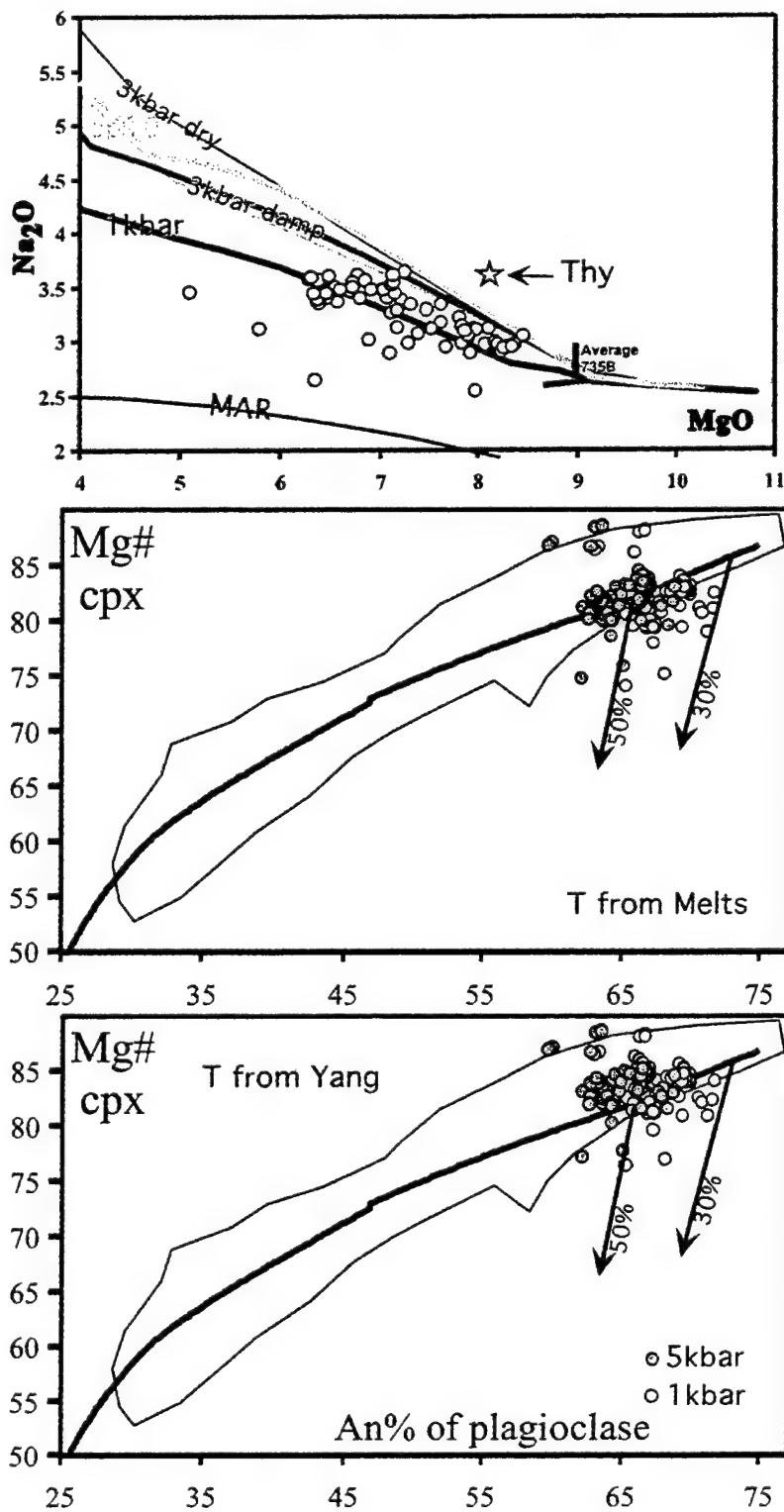


Figure 3-13: Liquid line of descent of the different Southwest Indian Ridge melts and one MAR-melt, and their relation to the solids. **A:** A suite of glasses from the ridge (open symbols, Johnson et al., 1990), compared to the starting composition from near the Rodrigues Triple Junction (star; Thy, 2003), and some of the models: “MAR”= AII107-D20 at QFM and 1kbar; “1kbar”=SWIR PMORB at 0.05wt%  $\text{H}_2\text{O}$ , QFM and 1kbar fractionation; “3kbar dry”=same melt crystallized at 3kbar, “3kbar damp” = SWIR PMORB with 0.05wt%  $\text{H}_2\text{O}$ ; “cp10” (thick gray line)=the melt that assimilated 10 wt% gabbro, thin gray line= Site 3 5/2 at QFM+1. **B:** Comparison of the field for Hole 735B gabbros and the 1kbar model to the calculated minerals that will be in equilibrium with the glasses at the ridge at 1kbar (white circles) or 5kbar (gray circles). The temperatures and melt-compositions from Melts5 were used to calculate the partition coefficients using Grove et al. (1992) for plagioclase and Putirka (1999) for augite. The tie-lines at 30 and 50% crystallized are indicated on the model. **C:** the same calculations using temperatures from Yang et al. (1996).

#### 4.3.2 Accumulation of crystals: the liquids:

Even if a crystallization model fits the solids, it may not fit the liquids erupted along the ridge as we suggested for the high Na<sub>8</sub> model of Thy (2003). Figure 3-13 shows some of the models compared to a glass-suite from AIIFZ. The most efficient liquid line of descent-modifier is clearly the composition of the original magma (Fig. 3-13A). The MAR-melt clearly underestimates the Na-content in the melts, whereas Thy (2003)'s starting composition (star) is too high in Na<sub>2</sub>O. The 1kbar model fits the data best, although the higher pressure-crystallization model, and the Melts-model with assimilated gabbro lie close to the distribution as well. We have calculated augite and plagioclase compositions in equilibrium with the ridge-glass compositions in order to directly compare them to those in the gabbroic rocks. The partition coefficients (Grove et al., 1992; Putirka, 1999) depend on melt composition and temperature, thus we have used both the temperatures of Yang et al. (1996) and those of Melts5 (Fig. 3-13 B,C). The ridge glasses recalculated at 1kbar represent melts crystallized ~30-50% respective to the SWIR PMORB melt, and most of the Hole 735B gabbros have more evolved compositions than these (as was pointed out by Bloomer et al. (1991) and Dick et al. (1991)). Based on fluid inclusions, Vanko and Stakes (1991) found that Unit V of Hole 735B formed 2 km below the seafloor, indicating that the upper crust was at least 1.65 km thick before the detachment exposed the lower crust. Since the upper crust in the area is shown to have crystallized 30% before emplacement, this adds another ~500 meters to the lower crustal budget. Increasing the pressure to 5-kbar elevates the augite Mg#'s as much as the Melts model does at 3-kbar (at 5-kbar plagioclase starts being suppressed with respect to augite in the Melts-model) and therefore, simply increasing the pressure more does not produce the high-Mg olivine-gabbros either.

The observed variation in the Hole 735B gabbros, then, may not be attributed to simple accumulation of crystals during fractional crystallization or batch equilibrium crystallization. The most evolved rocks are also difficult to model except by using a locally appropriate melt compositions. This also requires relatively low pressures, low water-contents at the fayalite-quartz-magnetite oxygen buffer to accrete these gabbros,

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whereas the elevated Mg#’s require post-cumulus processes as discussed in the next section. However, the damp SWIR-PMORB fractional crystallization model at 1-kbar explain the overall solid-line of descent of the data.

### 4.3.3 Post-cumulus processes

Meyer et al. (1989), Dick et al. (2000, 2002) and Coogan et al. (2001) cited processes within the crystal mush as the cause for the enormous variability of Southwest Indian Ridge augite compositions at constant forsterite or anorthite in primitive olivine-gabbros. Meyer et al. (op. cit.) described the unusually TiO<sub>2</sub>-enriched high-Mg# augite common in the primitive Southwest Indian Ridge gabbros as buffered by the olivine in mush while plagioclase and augite are slow to react with the melt, allowing the melts to evolve with respect to incompatible elements but not Mg, Fe and Ni.

Coogan et al. (2001) took a more radical approach, and suggested that the entire section of gabbros were made up from 50/50 melt and crystals. Moreover, as elements with different partition-coefficients indicate different amounts of trapped melt, they argue that the trapped melt moves within the mush and later returns to the eruptible magma reservoir, even *after* fractionation of a spinel-phase. We find that the melts evolved in our models have ~50 wt% SiO<sub>2</sub> until ~80% fractionation (when ulvöspinel is produced), as the silicates have SiO<sub>2</sub> contents similar to the melt. The melts become denser and retain a low viscosity (Fig. 3-14) due to increasing Na<sub>2</sub>O and FeO, and the viscosity increases abruptly after 80% crystallization. Plagioclase is also lighter than most of these melts. The evolved, oxide-rich, cumulates at Atlantis Bank are indeed abundant, and more than 200 felsic veins were described in the lower 1000-m of the section (Dick et al., 2000) suggesting that fractional crystallization went to completion. Although the evolved melts escaped into veins and shear-zones and crystallized there, we find it unlikely that the highly evolved, interstitial melts would return from the mush into the main magma-chamber, due to the abundance of oxide-minerals in Hole 735B suggesting that the last melt had high viscosities.

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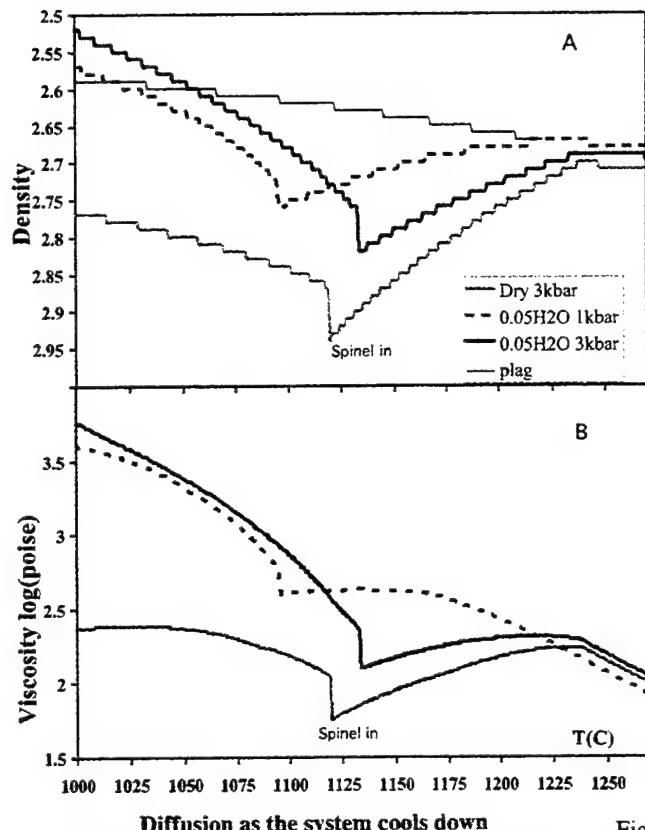


Figure 3-14: **A:** The density (g/cc) of the melts and plagioclase (plагe) as determined by Melts. The models are as indicated in previous figures. **B:** The viscosity of the melts with fractionation as determined by melts for the same models.

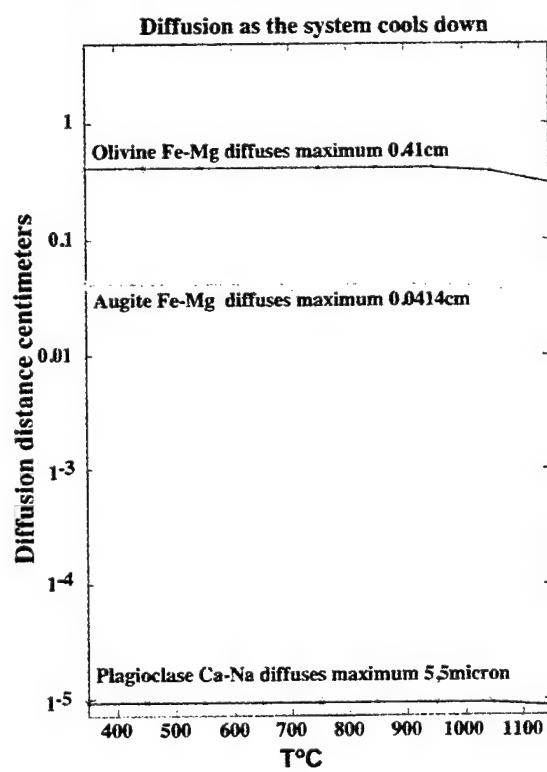


Figure 3-15: The longest diffusion distance for Ca-Na in plagioclase and Fe-Mg in olivine and augite that possible during cooling of Atlantis Bank, as determined by the cooling model of John et al (in press). Two scenarios, starting at 1350°C and 1150°C, respectively, demonstrate the temperature dependency on the diffusion. Thus, the lower temperatures ("postmagmatic") are the most applicable to this study. The rates are from Gaetani and Watson (2000) (Fe-Mg olivine), Gaetani (pers. comm. 2000; Fe-Mg augite) and Grove et al. (1984) (Ca-Na plagioclase).

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Dick et al. (2000) investigated the contacts between lithologic units in Hole 735B and found mineral-composition discontinuities that are inconsistent with fractional crystallization. They suggested that these were the result of more evolved late-magmatic liquids migrating upwards in the section across the contacts. Angeloni and Dick (1990) showed that cross-cutting fine-grained troctolites consistently have reversely-zoned plagioclase, indicating melt percolation and overgrowth on plagioclase by later, more *primitive* melts, but that the coexisting augites have higher levels of incompatible elements than would be expected at their Mg#’s. Bloomer et al. (1991) described Unit V of ODP Hole 735B to be typified by reversely zoned plagioclase chadacrysts surrounded by high-Mg augite oikocrysts, whereas the shallower gabbros have normal zoning. This observation is unfortunately not documented by published mineral analyses. Unit V (274-382mbsl), a massive, almost undeformed and oxide-free (e.g. Fig. 3-7) olivine-gabbro section, underlies the most massive oxide-gabbro sequence of the Hole. Natland and Dick (2002) describe the Unit IV-V contact as cross-cut by a hydrothermally altered dioritic intrusive breccia. Their description is similar to that of the migmatized aureole around gabbro-bodies that are found at Fuerteventura, Canary Islands (Hobson et al., 1998) suggesting that this contact is indeed intrusive. The intrusive boundary-theory is supported by John et al. (in press), who found evidence for re-heating events at 8.5-9-Ma related to off-axis magmatism.

In order to evaluate the many views of the post-cumulus evolution of the Hole 735B section, it is necessary to know the elemental diffusion-distances for each of the minerals in the gabbros (as shown in Figure 3-15), i.e. which minerals will be affected by major-element solid-state diffusion with melts percolating through a crystal mush. In addition, solid-state diffusion also shows the likelihood of disequilibrium augite crystallization. We use cooling rates determined for Atlantis Bank from the data from John et al. (in press) to estimate the potential diffusion-distances. The highest temperatures indicate the earliest magmatic stage for each mineral, with a diffusion-distance for the given time at this temperature. As the mineral cools, the accumulative diffusion-distance grows, and the resulting distance at the lowest temperature is the

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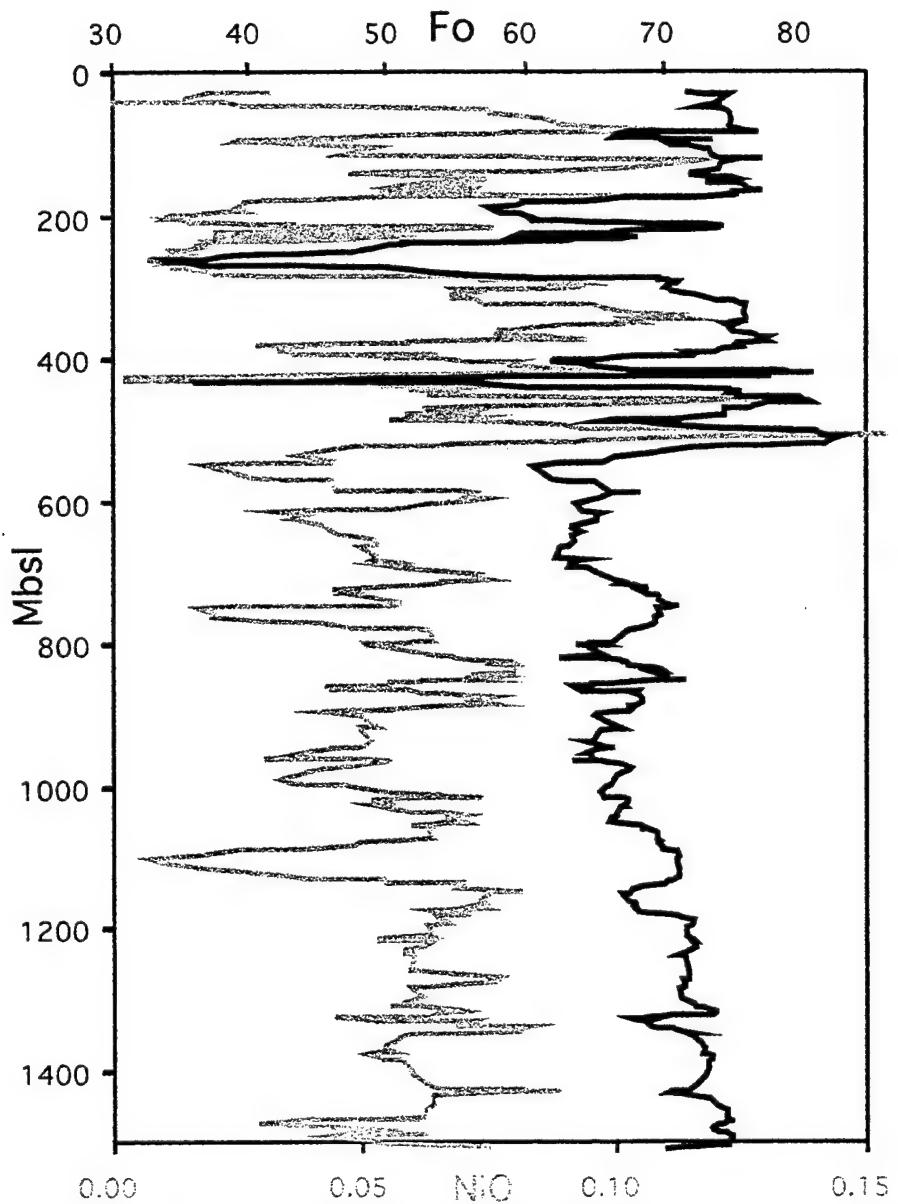


Figure 3-16: Running averages (of 9) of the NiO (wt%, gray) and Fo (black) content of olivine in Hole 735B.

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largest possible diffusion distance for a given mineral if diffusion started at the highest temperature. The Ca/Na of plagioclase, once crystallized, will not re-equilibrate with a melt in its solid state, the Fe/Mg of augite may equilibrate somewhat, whereas the Fe/Mg of olivine is easily re-equilibrated. Meyer et al. (1989), who studied very primitive cumulates from 7°E on the Southwest Indian Ridge, assumed that the olivines re-equilibrated with the interstitial melts to form the high-Mg# augites with elevated incompatible elements. If crustal accretion is assumed to occur by fractional crystallization, no olivines are formed at the intermediate temperatures in the model where the olivine-gabbros actually occur in Hole 735B, and there should be no olivine to react with. On the other hand, the magma could crystallize dunitic and troctolitic assemblages at the crust-mantle boundary and intrude as a mush (with the residues) into the crust. In such a mush, synneusis (the tendency of plagioclase crystals to cluster together due to surface-tension effects) will cause glomerophyric (or cumulophyric) clusters of plagioclase-phenocrysts, and these plagioclase-chains can create a semi-rigid network (Philpotts et al., 1998). If olivine settles out as the chains form, plagioclase will rise towards the ceiling of the magma-chamber. When there are 25% crystals present as chains, a semi-rigid framework may be completed, in which cumulates can mature. As major-element solid-state diffusion is slow in plagioclase and augite, they may only interact with the interstitial magma by dissolution, and once a plagioclase-plagioclase grain-boundary is formed, it is nearly impossible to disintegrate by dissolution (e.g. Chapter 4). Olivines, however, are free to exchange Fe and Mg with the melt. Refractory melts that intrude from below will be less dense than the melts in the network, and should therefore be able to infiltrate the mush.

Figure 3-16 shows running averages of NiO and forsterite content of olivines down-hole. Although the scatter is misleading, the lower 1000 meters of the hole has more NiO in olivine (average 0.057 wt% NiO with average Fo of 69.1) compared to the upper 550 meters (average 0.0515 wt% with average Fo of 68.7). If only the olivine-gabbros are taken into account, however, the numbers are 0.057 wt% NiO and average Fo<sub>69.6</sub>, and 0.074 wt% NiO and average Fo<sub>75.3</sub>. The Melts model indicates that 8%

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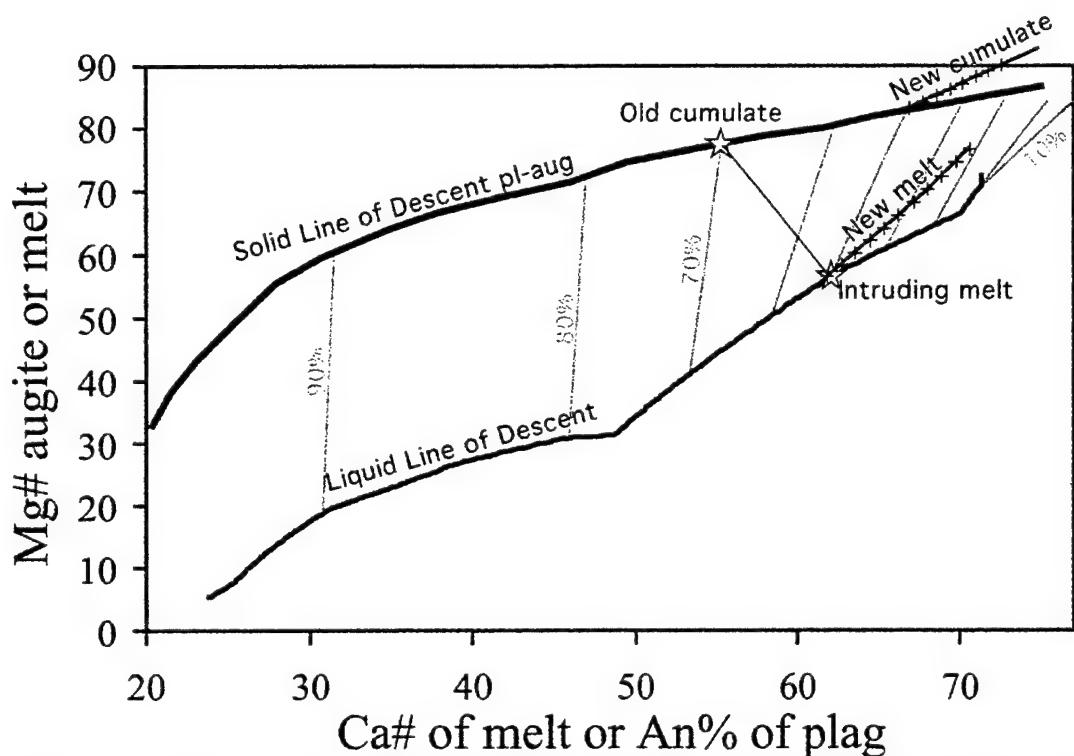


Figure 3-17: Compare to Figure 1: The upper black line represents coexisting augite and plagioclase, the lower black line is the Mg# and Ca# of the evolving melt, and the gray lines are the 10% incremental crystallization-tielines. The consequence of assimilating a cumulate representing 70% fractionated (upper star, 1140°C) into a 50% crystallized melt (lower star, 1180°C) is to increase the Mg# and Ca# of the melt (along the lower line where 10% increments are indicated with crosses), resulting in new cumulates following the upper respective line. Thus, the new cumulate will have higher Mg# augite and more anorthitic plagioclase than the original melt.

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fractionation (of plagioclase, clinopyroxene and orthopyroxene) is required to reduce the Mg# of the melt sufficiently from the average equilibrium liquids of the upper 550-m-s to those of the lower 1000-m. We can therefore contrast this amount of crystallization to what is required to reduce NiO from 0.074 to 0.057 wt%. The partitioning of Ni in olivine is dependent on the MgO of the melt (Hart and Davis, 1978), and we find  $D^{Ni} \sim 21$  (we used the MgO of the liquids in the 1kbar-model at the correct Mg#'s relative to the forsterite content of the average sections). Thus, a melt in equilibrium with the olivines in the upper 550 meters will have to fractionate 5.5 wt% olivine (plus other minerals) in order to reduce the NiO to that of the lower 1000 meters. If the parental melts of the upper and lower part of Hole 735B were similar, it is not fair to assume that the olivines we analyze today have the same composition as the olivines that initially crystallized and injected with the mush.

If we assume that early olivine are dispersed in a plagioclase mush-network, and that there is melt-flow through this mush, the olivines will re-equilibrate quickly with the interstitial melt. The interstitial olivines are sources of Mg and sinks for Fe for the evolving interstitial melt as Meyer et al (1989) suggested. Moreover, any melt that dissolves a pre-existing cumulate (i.e. the melt has a higher temperature than the solidus of what is melted) will produce a hybrid melt that has higher Mg# (and lower TiO<sub>2</sub>) (see Chapter 4). This is illustrated in Figure 3-17, where each tie-line connects a cumulate to its equilibrium melt. Therefore, a melt at slightly higher Mg# and Ca# will be able to dissolve rocks lower on the reaction-series (See Chapter 4 for details). The mixing-line between the cumulate-composition and the melt composition is indicated on the figure. Note that augite is Ca-rich (up to 22 wt% CaO), so the Ca# also increases. The hybrid melt initially crystallizes higher anorthite-content plagioclase and Mg# augite. This creates reverse zoning in plagioclase and augite, like Bloomer et al. (1991) observed in Unit V and the cross-cutting troctolites. It may also additionally create high-Mg augite with high levels of incompatible elements in them, as Meyer et al. (1989) found. As the hybrid melt crystallizes, it will produce gradually more evolved cumulates, just like the Melts-models outline.

## 5 Conclusion

The rocks sampled in Hole 735B represent 25-90% crystallization of a primary mantle-derived melt with  $\text{Na}_8 \sim 3$  and  $\text{Fe}_8 \sim 9$  at crustal pressure on the quartz-fayalite-magnetite oxygen fugacity buffer. The majority of the gabbros in the lower 1000 meters of the Hole represent 60-80% crystallization. The upper 550 meters have a very bimodal distribution with cross-cutting troctolites and massive olivine gabbros representing 25-45% and 45-70% crystallization respectively, whereas the ferrogabbros represent 80-90% crystallization. The glasses sampled along the ridge, on the other hand, represent 30-50% crystallized relative to the same mantle-derived melt, and most of the crustal section is far more evolved than the upper crust. If we assume a traditional view of the accretion of the lower ocean crust, ~1200 meters of the lower crustal section has not been drilled. With the missing upper crust, this brings the minimum total crustal thickness to ~4.4 km, suggesting that the observed seismic Moho at Atlantis Bank is indeed the crust-mantle boundary rather than an alteration-front as suggested by Muller et al. (1997). We find that the olivine-gabbros have augites in disequilibrium with the coexisting plagioclase relative to crystallization at crustal pressures, and that this 10 mol% discrepancy may be accounted for by dissolution and reprecipitation within a crustal mush-zone due to the intrusions of melt from below.

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## Chapter 4

### Dissolution Kinetics of Oceanic Lower Crust

#### Abstract

We have carried out dissolution experiments using natural olivine, plagioclase and clinopyroxene, and equilibrium melting experiments on synthetic cumulate gabbro compositions to quantify the melting kinetics of oceanic gabbros. Dissolution experiments were carried out using olivine – plagioclase and clinopyroxene - plagioclase mineral pairs over the temperature range of 1330 - 1220 °C at the quartz-fayalite-magnetite buffer for durations of 0.25 to 24 hours. Isothermal experiments performed over the temperature range of 1230 – 1105 °C defined the equilibrium phase relations in the system. Dissolution mechanisms are similar to those found in other experimental studies that utilized end-member mineral compositions. The experimental results are used to assess the extent to which dissolution of cumulate igneous rocks from the oceanic lower crust will produce disequilibrium melts compositions. The efficiency of the reaction is dependent on the area of grain-boundaries between different minerals in a rock that is reheated. Coarse-grained rocks (>3mm) will disintegrate internally upon reheating, whereas finer grained rocks will melt from the outside in. If these disequilibrium melts are allowed to mix into the ascending magma, the new magma will have higher Na<sub>8</sub> and lower Fe<sub>8</sub> than the original. Lower crustal assimilation of gabbro will result in the destruction of troctolites in the lower crust, and will lead to the early onset of high-Mg# olivine-gabbros.

#### 1. Introduction

The oceanic lower crust is dominated by olivine gabbro, generally composed of olivine, plagioclase and augite (Coleman, 1977). It is commonly accepted that this lower crust is formed by crystallization of basaltic magmas that ascend from the upper mantle. However, there are two main lines of thought on the geodynamic method of accretion of

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the lower ocean crust. One is the so-called “gabbro glacier” model, whereas the other is the “many sills” model. The former end-member is based on observations in ophiolites (Sleep, 1975; Dewey and Kidd, 1977 and Quick and Denlinger, 1993), and developed further to account for geophysical observations at the East Pacific Rise (EPR) and other magmatically robust spreading segments (Henstock et al., 1993; Phipps Morgan and Chen, 1993). The gabbro-glacier model assumes that melt arrives directly from the upper mantle to a shallow melt sill, where it crystallizes gabbro cumulates that subsequently flows downwards in a near steady-state fashion. The many sills model (Pedersen, 1986; Bédard et al, 1988, 1991, 1993; Boudier et al., 1996; Kelemen et al., 1997b; Korenaga and Kelemen, 1997) postulates that the lower crust is accreted by the injections of individual sills that fractionate in-situ and eject their melt to higher levels. Thus, particularly for the fast-spreading ridges, the lower crust is accreted within the lower crust. Modifications of this latter model are generally favored for slow-spreading ridges.

In general, truth probably lies between the gabbro glacier and the many sills model (MacLennan et al., 2004). In either case, the extent to which the lower ocean crust interacts with subsequent ascending magmas is largely unknown. The rate of assimilation is naturally dependent on the heat available to incorporate and melt pre-existing cumulates into magma, and the residence time of the magma in the lower crust. The effect of the composition and grain-size of the assimilant, thus the compositional effects on the magma and lower crust, has not yet been established for the oceans.

The potential temperature of the melts generated in the upper mantle at mid ocean ridges could range from 1180°-1510°C (Klein and Langmuir, 1987; Kinzler and Grove, 1992), although other authors would disagree, arguing that the temperature is constant around 1250°C (e.g. Green et al, 2000; Presnall, 2002). The upper end of the temperature spectrum represents the fast spreading EPR. It has been documented that the EPR has a shallow axial magma chamber (AMC) (Detrick et al., 1987) that lies on a low velocity zone (Harding et al., 1989; Vera et al., 1990). Dunn et al. (2000) showed a 5-8 km wide low velocity zone under the spreading center of the EPR has as much as 20 % melt fractions within it. A second melt lens has been documented near the seismic Moho

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transition (Crawford and Webb, 2002); it has been argued that its stability is limited, as any crystallization in the chamber causes the surroundings to heat up and melt the surroundings in a manner that is not observed (Chen, 2001). The EPR therefore clearly has a heat-budget that is large enough to remelt lower crustal material. At intermediate-spreading ridges, the AMC reflector is found at deeper levels in the crust (Phipps Morgan and Chen, 1993b; Chen and Lin, in press). The existence of shallow steady-state magma-chambers at slower ridges is indeed debated (Sinton and Detrick, 1992), and very deep magma chambers below the conductive boundary-layer in the mantle (Mével et al., 2002) would not be possible to detect with the seismic techniques available today (Detrick and Reeves-Sohn, pers comm. 2004). Erupted MORB show extensive differentiation and mixing that require that magma chambers exist, even if only ephemerally. The evidence of magma-chamber type processes happening before eruption at all ridges, suggest that there should be sites that are hot enough to dissolve preexisting crustal material even if the magma-lenses evade us along the slower ocean ridges. Therefore, the magmas likely reside for long enough to interact with their surroundings in either steady state- or ephemeral magma chambers.

Ocean islands are also sites of elevated magma-temperatures (Watson and McKenzie, 1991). In fact, there are Hawaiian picrites that record crystallization temperatures in excess of 1400°C (Natland, pers comm. 2004). In Hawaii post-shield volcanism of very different composition develops after the earlier shield-stage lavas. Their isotopes are more like those of mid-ocean crust, and they are frequently very alkaline in composition (Yang et al., 2003). The post-shield lavas often contain xenoliths made out of the pre-existing, tholeiitic, lower-crust that had to be penetrated to get to the surface (Fodor and Galar, 1997; Neumann et al., 2000). This suggests that the post-shield stage magmas, which spend more time in the lower crust than the shield-stage magmas, may have digested some lower crustal material on ascent.

The lower crust is generally difficult to study at fast-spreading ridges and ocean islands due to the 2+ km of dikes and lavas that cover it and the lack of tectonic windows. However, the lower crust is often exposed in tectonic windows at slow-spreading ridges

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(Tucholke and Lin, 1998). The gabbroic outcrops on Atlantis Bank at the slow Southwest Indian Ridge have been studied extensively (Dick et al., 2000; Natland and Dick, 2000, and references therein). Even the most primitive, ophitic, gabbros sampled in the bank have high-Mg augite oikocrysts surrounding more evolved plagioclase chadacrysts with which they are not in equilibrium at magmatic temperatures (Chapter 3, this thesis). Nearly all these rocks have petrologic characteristics that provide strong evidence for complex magmatic histories that may involve several magmas and potential remelting.

The purpose of this study is to constrain on the processes of magma – cumulate interaction that occurs within the oceanic crust. We have therefore investigated the kinetics of melting between natural mineral-pairs representative of the those found in lower ocean-crustal rocks; finding rapid melting rates and unusual melt-compositions. Unlike previous dissolution-experiments, our mineral compositions are similar to those at mid-ocean ridges. Therefore, the melts produced in the experiments allow us to assess the chemical changes that incorporation of lower crustal material may cause in the ascending magmas.

## 2. Experimental Procedures

During the course of this study, we performed two types of experiments. The first type was dissolution-experiments on olivine-plagioclase and augite-plagioclase mineral pairs designed to provide information on the mineral melting kinetics in ocean gabbro cumulates. Second, in order to quantify the kinetic controls on melting, we also determined the solidus to near-liquidus phase relations of two synthetic CMASNF ( $\text{CaO} - \text{MgO} - \text{Al}_2\text{O}_3 - \text{SiO}_2 - \text{Na}_2\text{O} - \text{FeO}$ ) olivine – plagioclase - clinopyroxene cumulate-compositions that represented bulk compositions applicable to the mineral pairs used in the dissolution experiments.

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Table1: Theoretical starting composition  
isothermal experiment, in weight %

Label	Melt Fo	Melt AK3
$\text{SiO}_2$	53.1	50.6
$\text{Al}_2\text{O}_3$	17.7	14.7
$\text{FeO}$	3.2	6.7
$\text{MgO}$	7.0	12.0
$\text{CaO}$	16.1	13.4
$\text{Na}_2\text{O}$	2.9	2.4
<b>Total</b>	100	100
Mg#	79.7	76.1
Ca#	75.5	75.5

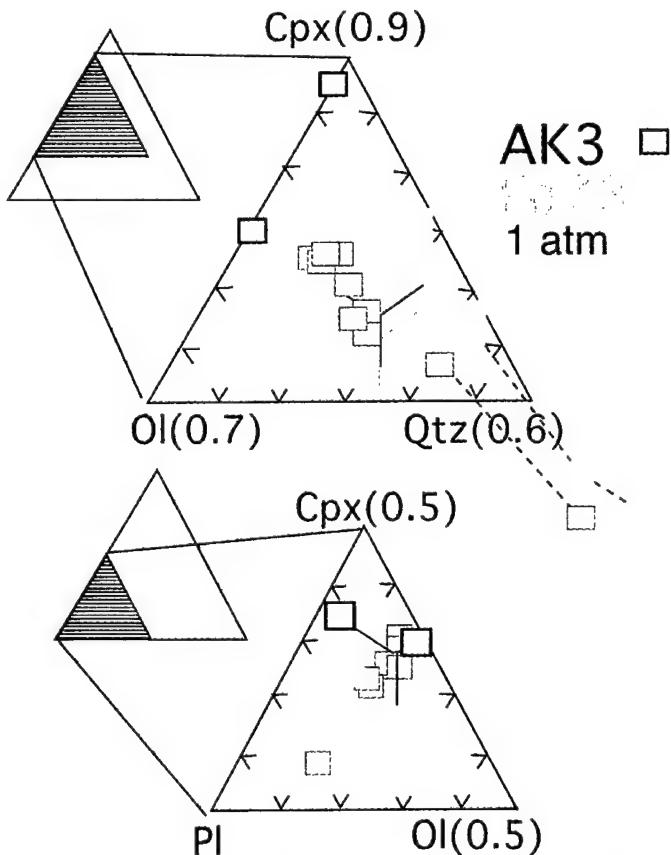


Figure 4-1: Pseudoternary projection-plots of the melt compositions of the isothermal phase-equilibrium experiments (projection from Grove et al., 1992). Cpx=augite, ol=olivine, qtz=quartz, pl=plagioclase. Open squares represent AK3, closed gray squares are Fo73. The inferred phase-saturation boundaries are indicated on the figures for reference. The Cpx-ol-qtz projection shows how the low-temperature experiments plot below the olivine-plagioclase-clinopyroxene-melt saturation-boundary for the basaltic system, towards more silicic compositions. Fo73 is has a very small olivine component, whereas AK3 generally follows the trends for basalts.

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### ***2.a. Phase equilibrium experiments***

**Starting material** - Phase equilibrium experiments were performed on two model olivine-gabbro compositions, “AK3” and “Fo73” (Table 4-1) where the former has the larger olivine component. Mechanical mixtures of each composition were prepared from high-purity oxides and silicates ground in an agate mortar under ethanol for 6 hours. Pressed pellets of AK3 or Fo73 powdered mix were sintered onto a 0.8 mm diameter Pt<sub>91</sub>-<sub>93</sub>-Fe<sub>7.9</sub> loop that was chosen to minimize Fe exchange with the pellet. The loop and pellet were suspended in the hotspot of a Deltech DT31VT vertical gas-mixing furnace. We controlled the oxygen fugacity by mixing CO<sub>2</sub> and H<sub>2</sub> gases and f<sub>O<sub>2</sub></sub> was monitored using a solid ZrO<sub>2</sub>-CaO electrolyte O sensor calibrated against the Fe-FeO, Ni-NiO, and Cu-Cu<sub>2</sub>O buffers. Temperature was measured using a Pt-Pt<sub>90</sub>Rh<sub>10</sub> thermocouple calibrated against the melting points of NaCl, Au and Pd on the IPTS 1968 temperature scale. The thermocouple junction was placed immediately adjacent to the experimental charge, and temperatures are believed to be accurate within ± 2°C. The experiments were carried out at 0.1MPa, with oxygen fugacity buffered at quartz – fayalite – magnetite (QFM) from near-solidus- to near-liquidus temperature. Experiments were run at constant temperature for 5 to 119 hrs, and were terminated by drop-quenching into a water bath. The results are listed in Table 4-2 (a, b and c).

Volatilization of sodium from the experimental charge, and mass transfer of iron between the experimental charge and wire loop can be significant in 0.1MPa experiments. We attempted to minimize the iron exchange by using a loop containing ~8 wt% iron, making it in equilibrium with tholeiite at QFM over the temperature interval of 1100°C to 1400°C (Grove, 1981). Sodium volatilization is related to sample size, gas flow rate, oxygen fugacity, experimental duration and experimental temperature (Tormey et al., 1987). As a result, we used larger sample sizes (100mg) and lower gas-flow rates (0.1ml/s) to reduce sodium volatilization.

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Table 2A: Results of isothermal Fo73 experiments.

Temperature°C	1190	1175	1165	1150	1135	1120	1105							
Glass	3 gl(13)	σ	1 gl(8)	σ	2 gl(8)	σ	4 gl(10)	σ	5 gl(11)	σ	6 gl(3)	σ		
SiO <sub>2</sub>	56.6	.15	59.8	.49	59.5	.41	60.4	.49	66.9	1.11	71.1	.73		
Al <sub>2</sub> O <sub>3</sub>	15.1	.09	14.1	.12	14.7	.33	13.8	.10	13.9	1.02	11.9	.44		
FeO	4.6	.15	5.6	.23	5.8	.34	6.7	.27	4.0	.30	2.5	.44		
MgO	6.9	.06	4.7	.18	5.4	.16	4.5	.08	3.1	.46	1.4	.40		
CaO	14.4	.11	11.0	.15	11.4	.27	10.9	.08	8.2	.92	6.1	.79		
Na <sub>2</sub> O	2.1	.10	2.8	.12	2.7	.08	2.6	.07	2.1	.40	2.2	.21		
Total	99.7		98.2		99.4		98.9		98.2		95.1			
Mg#	73.0		59.8		62.2		54.5		58.0		49.5			
Ca#	79.1		68.7		70.1		69.9		67.9		60.8			
Augite	3 px(15)	σ	1 px(10)	σ	2 px(15)	σ	4 px(14)	σ	5 px(9)	σ	6 px(15)	σ	7 px(13)	σ
SiO <sub>2</sub>	54.4	.94	54.2	.95	55.0	1.14	53.9	.94	53.5	.98	56.3	2.73	54.0	1.34
Al <sub>2</sub> O <sub>3</sub>	2.1	1.17	2.3	.85	2.5	1.37	1.8	.48	2.6	2.01	3.4	2.61	3.0	1.85
FeO	2.7	.20	3.6	.68	3.9	.34	5.0	.62	4.5	.61	4.5	1.21	6.5	1.23
MgO	17.5	.64	16.4	.64	17.2	1.00	16.0	.52	15.3	1.05	13.3	2.45	13.6	2.02
CaO	23.6	.53	22.4	1.25	21.6	1.30	22.8	1.16	22.4	1.26	21.9	2.57	23.5	1.39
Na <sub>2</sub> O	0.1	.04	0.3	.14	0.3	.24	0.2	.06	0.4	.23	0.5	.39	0.5	.35
Total	100.4		99.2		100.4		99.7		98.7		99.8		101.2	
Mg#	92.0		89.1		88.6		85.0		85.7		84.1		78.8	
Plagioclase	3 pl(12)	σ	1 pl(9)	σ	2 pl(10)	σ	4 pl(11)	σ	5 pl(6)	σ	6 pl(4)	σ	7 pl(11)	σ
SiO <sub>2</sub>	49.5	1.55	51.2	1.61	53.4	.96	53.9	.98	53.7	3.88	56.3	3.40	56.1	1.17
Al <sub>2</sub> O <sub>3</sub>	31.3	1.65	29.6	2.04	29.0	.92	28.5	.90	26.7	3.98	27.0	1.97	27.5	.91
FeO	0.6	.27	0.8	.34	0.5	.13	0.5	.13	1.1	.46	0.3	.04	0.3	.12
MgO	0.9	.49	0.8	.46	0.6	.12	0.4	.16	1.8	1.35	0.3	.03	0.2	.52
CaO	16.1	.81	14.5	.76	13.1	.47	13.0	.67	14.5	1.55	11.3	1.82	11.5	.88
Na <sub>2</sub> O	2.0	.49	2.8	.40	3.6	.19	3.6	.23	2.5	.54	4.3	.79	4.4	.29
Total	100.5		99.7		100.2		99.8		100.4		99.4		99.9	
Ca#	81.4		74.1		66.7		66.7		76.6		59.2		59.0	

Notes: gl=glass, px=plagioclase, pl=augite, ol=olivine, pig=pigeonite

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**Table 2B: Results of isothermal AK3 experiments**

Temperature°C	1230	1210	1197	1190	1190	1190	1170	1170	1150	1130	1110
Glass	4 gl(10)	2 gl(10)	7 gl(13)	1 gl(11)	3 gl(11)	5 gl(8)	10 gl(11)	8 gl(8)	6 gl(6)	9 gl(6)	9 gl(6)
SiO <sub>2</sub>	52.23	0.27	51.95	0.38	52.86	0.24	53.26	0.47	53.11	0.33	52.46
Al <sub>2</sub> O <sub>3</sub>	15.12	0.14	15.52	0.13	15.07	0.29	14.72	0.15	14.38	0.13	14.53
FeO	6.30	0.23	6.45	0.30	6.97	0.24	7.27	0.33	8.17	0.32	7.92
MgO	10.82	0.07	10.16	0.16	9.61	0.17	8.41	0.24	8.56	0.16	7.73
CaO	14.01	0.20	14.20	0.15	14.24	0.25	12.16	0.48	12.95	0.23	13.55
Na <sub>2</sub> O	1.74	0.15	1.57	0.11	1.64	0.09	1.84	0.13	1.81	0.19	1.95
Total	100.2	99.8	100.4	97.7	99.0	99.6	99.5	99.0	98.5	99.7	98.8
Mg#	75.4	73.7	71.1	67.4	78.5	79.8	65.1	67.3	63.0	60.4	55.2
Ca#	81.6	83.4	82.7	79.4	82.7	82.7	81.0	81.0	79.4	77.9	74.3
Augite				7 px(10)	1 px(14)	1 px(10)	3px(10)	5px(10)	10 px(9)	8 px(8)	6 px(5)
SiO <sub>2</sub>	55.1	0.48	51.9	0.77	53.7	0.01	52.0	1.01	53.6	0.88	54.5
Al <sub>2</sub> O <sub>3</sub>	2.0	0.56	4.1	0.85	3.4	0.05	4.3	1.56	3.4	0.72	3.4
FeO	3.6	0.10	4.8	0.39	4.7	0.01	4.5	0.84	4.6	0.70	5.9
MgO	20.2	0.59	18.1	1.09	19.4	18.3	19.5	0.97	18.5	1.07	18.5
CaO	20.0	0.61	20.7	1.43	19.5	1.26	21.2	20.1	19.4	0.80	19.1
Na <sub>2</sub> O	0.2	0.04	0.1	0.07	0.2	1.70	0.2	0.12	0.2	0.15	0.4
Total	101.1	99.8	100.9	100.5	100.5	100.5	100.5	100.3	101.4	100.9	100.7
Mg#				90.8	87.1	88.0	87.8	87.8	83.4	82.5	80.7
Plagioclase				2pl(10)	7 pl(10)	1 pl(13)	1 pl(10)	3pl(10)	5pl(7)	10 pl(9)	8 pl(7)
SiO <sub>2</sub>	47.0	0.48	47.8	0.79	47.5	1.56	48.2	1.23	47.0	0.75	49.0
Al <sub>2</sub> O <sub>3</sub>	33.5	0.14	31.8	0.19	31.3	0.29	30.6	0.57	31.1	0.25	30.6
FeO	0.7	0.19	0.9	0.29	1.3	0.39	1.4	0.67	1.7	0.64	0.9
MgO	0.5	0.00	1.0	0.02	1.4	0.02	1.3	1.2	0.00	1.0	0.00
CaO	16.9	0.46	16.7	0.56	16.2	0.47	16.2	1.65	0.84	15.0	0.66
Na <sub>2</sub> O	1.7	0.02	1.7	0.03	1.7	1.9	0.01	1.7	0.03	2.6	0.03
Total	100.5	99.9	99.3	99.5	99.5	99.2	99.1	99.8	100.9	99.3	101.5
Olivine				4ol(10)	2ol(10)	7 ol(10)	1 ol(15)	1ol(10)	3ol(10)	5ol(11)	10 ol(10)
SiO <sub>2</sub>	40.9	0.31	40.6	0.28	40.2	0.17	39.7	0.24	39.8	0.45	40.3
Al <sub>2</sub> O <sub>3</sub>	0.1	0.01	0.16	0.1	0.02	0.1	0.08	0.1	0.15	0.11	0.15
FeO	7.3	0.53	8.4	0.43	9.8	0.24	11.4	0.41	12.0	0.29	10.0
MgO	51.2	0.58	50.2	0.76	49.6	0.31	47.9	0.31	47.1	0.36	48.7
CaO	0.4	0.03	0.5	0.18	0.5	0.03	0.5	0.05	0.4	0.11	0.6
Na <sub>2</sub> O											
Total	100.0	99.8	100.2	99.6	99.4	99.6	99.6	99.2	100.0	101.3	100.9
Mg#	92.6	91.4	90.0	88.2	87.5	89.7	85.9	83.7	78.3	75.9	

Notes: gl=glass, px=plagioclase, ol=olivine, pig=pigeonite

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Table 2C: Results of isothermal phase-equilibrium experiments.

Experiment#	Temperature °C	Duration (hr)	Pressure (kbar)	Sum R <sup>2</sup>	Minerals	Kd <sup>Fe/Mg</sup>	Kd <sup>Ca/Na</sup>	Kd <sup>Fe/Mg</sup>	Mg# gl	Cat#	Mg# cpx	Fo ol	An pl
AK3-4	1230	5.5	0.1	0.75	g(94.9)ol(3.2) g(93.3)pl(0.5)ol(5)	0.25	1.10	0.26	75.4	81.7	73.7	83.4	92.6
AK3-2	1210	6	0.1	0.67	g(83.1)cpx(0.9)pl(6.9)ol(7.5)	0.25	1.14	0.27	71.1	82.8	80.8	91.4	84.6
AK3-7	1196	6	0.1	0.50	g(57.4)cpx(13.4)pl(19.7)ol(8.9)	0.25	1.21	0.27	65.1	79.9	88.0	87.5	82.8
AK3-1	1190	14	0.1	0.50	g(70.2)cpx(7.9)pl(13.6)ol(8.1)	0.29	1.38	0.24	67.3	79.4	87.8	89.7	84.2
AK3-3	1190	6.25	0.1	0.50	g(43.9)cpx(19.7)pl(25.1)ol(10.3)	0.24	1.16	0.28	63.0	73.5	87.8	85.9	76.3
AK3-5	1170	16.5	0.1	0.50	g(4)lcpx(33)pl(47)ol(13)	0.26	0.80	0.34	55.2	63.7	82.5	78.3	58.4
AK3-8	1150	15	0.1	0.50	g(3)lcpx(46)pl(46.4)	0.29	0.88	0.36	54.8	74.3	80.7	71.8	51.2
AK3-6	1130	6	0.1	0.65	g(-30)cpx(25)pl(71)pig(31)	0.28	0.93	0.36	52.9	52.9	80.1		
AK3-9	1110	20	0.1	0.49									

Experiment#	Temperature °C	Duration (hr)	P (kbar)	Sum R <sup>2</sup>	Minerals	Kd <sup>Fe/Mg</sup>	Kd <sup>Ca/Na</sup>	Kd <sup>Fe/Mg</sup>	Mg# gl	Cat#	Mg# cpx	Fo ol	An pl
F073-3	1190	70.5	0.1	1.45 (0.11)	g(53.3)cpx(15.5)pl(29.5)g(51.5)cp	0.24	1.16	1.16	73.0	79.1	92.0	81.4	
F073-1	1175	95	0.1	1.44 (0.57)	g(19.9)cpx(31.9)pl(47.1)g(19.1)cp	0.18	1.30	0.59	68.7	89.1	74.1		
F073-2	1165	69	0.1	1.58 (0.47)	g(8.9)cpx(36.1)pl(52.8)g(3.5)cpx(	0.21	0.85	0.86	62.2	70.1	88.6	66.7	
F073-4	1150	71	0.1	1.38	g(5.3)cpx(37)pl(56)	0.21	0.86	0.86	54.5	69.9	85.0	66.7	
F073-5	1135	119	0.1	2.74 (1.3)	g(5.5)cpx(33.3)pl(59)g(5.4)cpx(33	0.23	1.54	0.58	67.9	85.7	76.6		
F073-6	1120	45	0.1	2.57 (1.06)	g(-13.1)cpx(46.9)pl(64.1)g(-13.2)c	0.19	0.94	0.50	60.8	84.1	59.2		
F073-7	1105	72	0.1	3.09	cpx(40.5)pl(56.6)	0.28	0.93	0.36	52.9	52.9	80.1		

Notes: gl=glass, ol=olivine, cpx=olivine, pi=plagioclase, pig=pigeonite. All analyses done using natural mineral standards at the MIT electron microprobe, as mentioned in the text.

### **2.b. Mineral Dissolution Experiments**

Two sets of mineral pairs were used in these experiments (Table 4-3). The first consisted of clinopyroxene (Mg# 82-84), and plagioclase ( $An_{60}$  and  $An_{54}$ ). The second set of experiments used  $An_{60}$  plagioclase and  $Fo_{73}$  olivine; or  $An_{54}$  plagioclase with  $Fo_{82}$  olivine. The experiments were carried out over a the temperature range of 1240-1330 °C for the olivine – plagioclase pairs and 1180-1300 °C for the plagioclase – clinopyroxene pairs. The duration of the experiments varied from 15 minutes to 24 hours (Table 4-4).

The  $An_{62}$  plagioclase, the  $Fo_{73}$  olivine and the augite were sampled from olivine gabbros of homogeneous compositions from Atlantis Bank, Southwest Indian Ridge, ODP Hole 735B, at ~1400 m depth (Natland and Dick, 2002). Due to the limited availability and grain-size of the material from the Southwest Indian Ridge, we selected alternative sources of plagioclase and olivine, in order to run longer experiments. The  $An_{54}$  plagioclase was sampled from a single crystal from the layered Bjerkreim-Sokndal intrusion in Southwestern Norway. The  $Fo_{82}$  olivine is from a granular dunite from the Balmuccia Iherzolite massif, Italy (Rivalenti, 1980). Although the  $An_{54}$  plagioclase and the  $Fo_{82}$  olivine are not exactly the compositions expected from a gabbro from the ocean crust, they are desirable as they are natural materials with more appropriate compositions compared to synthetic endmembers of the solid solution series.

*Sample preparation and experimental procedures* Each mineral was cut with a diamond wafering-saw into parallelepipeds of roughly the same width. No attempt was made to select a specific crystallographic orientation for any of the minerals. One side of each piece were polished with  $Al_2O_3$ -paper and diamond pastes to  $5\mu m$  grit. The thickness of each mineral-piece was measured with a micrometer to a precision of  $\sim 10\mu m$  perpendicular to the polished surface. The mineral pairs were placed together with the polished surfaces in contact, and tied with a 0.8mm diameter Pt wire. Each experiment was held at a constant temperature at the QFM buffer for a specified time in the hotspot of a Deltech DT31VT vertical gas-mixing furnace. Each experimental assembly achieved thermal equilibrium within 5 minutes or less after being positioned in the hotspot. The beginning of dissolution was measured from that point on until the

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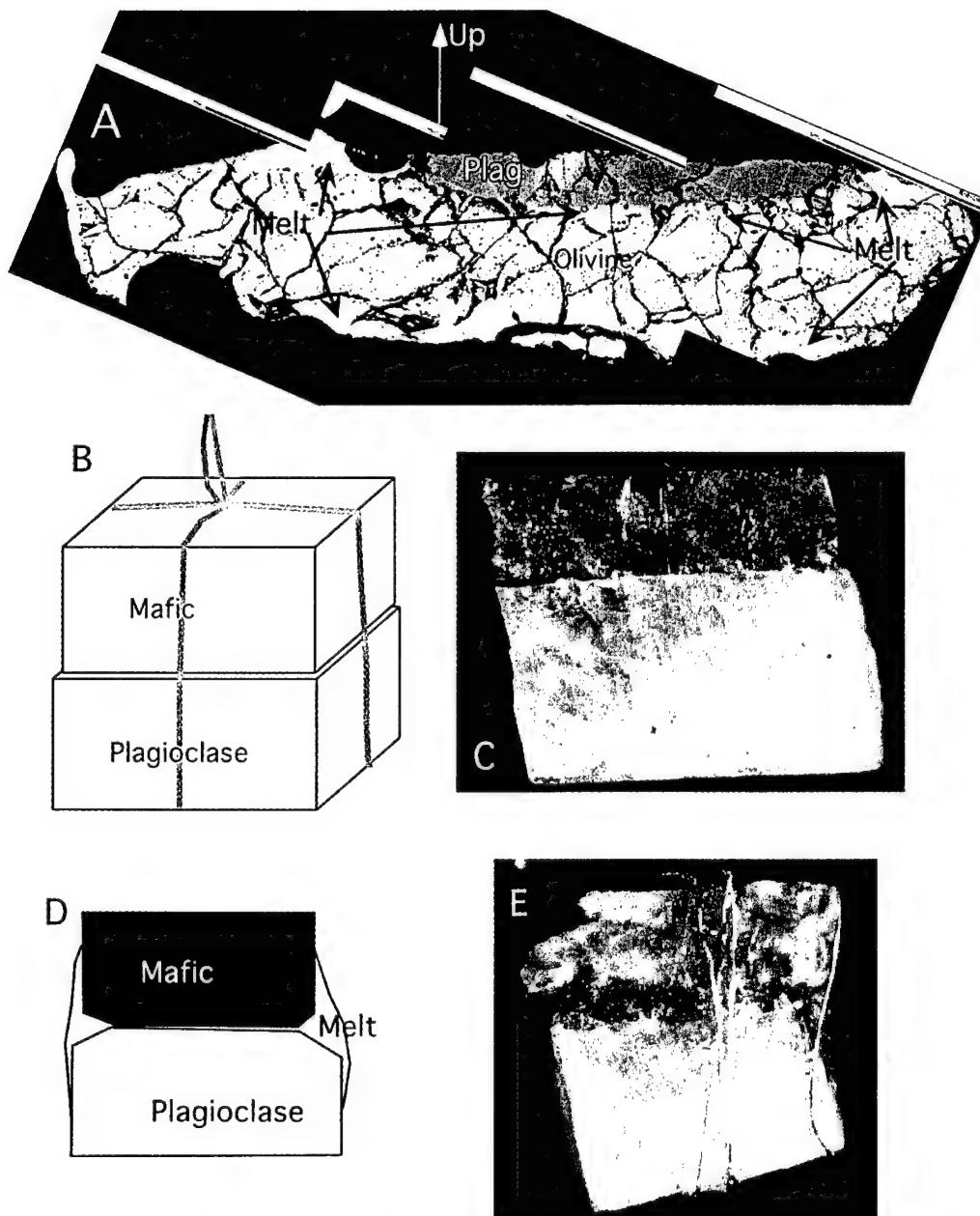


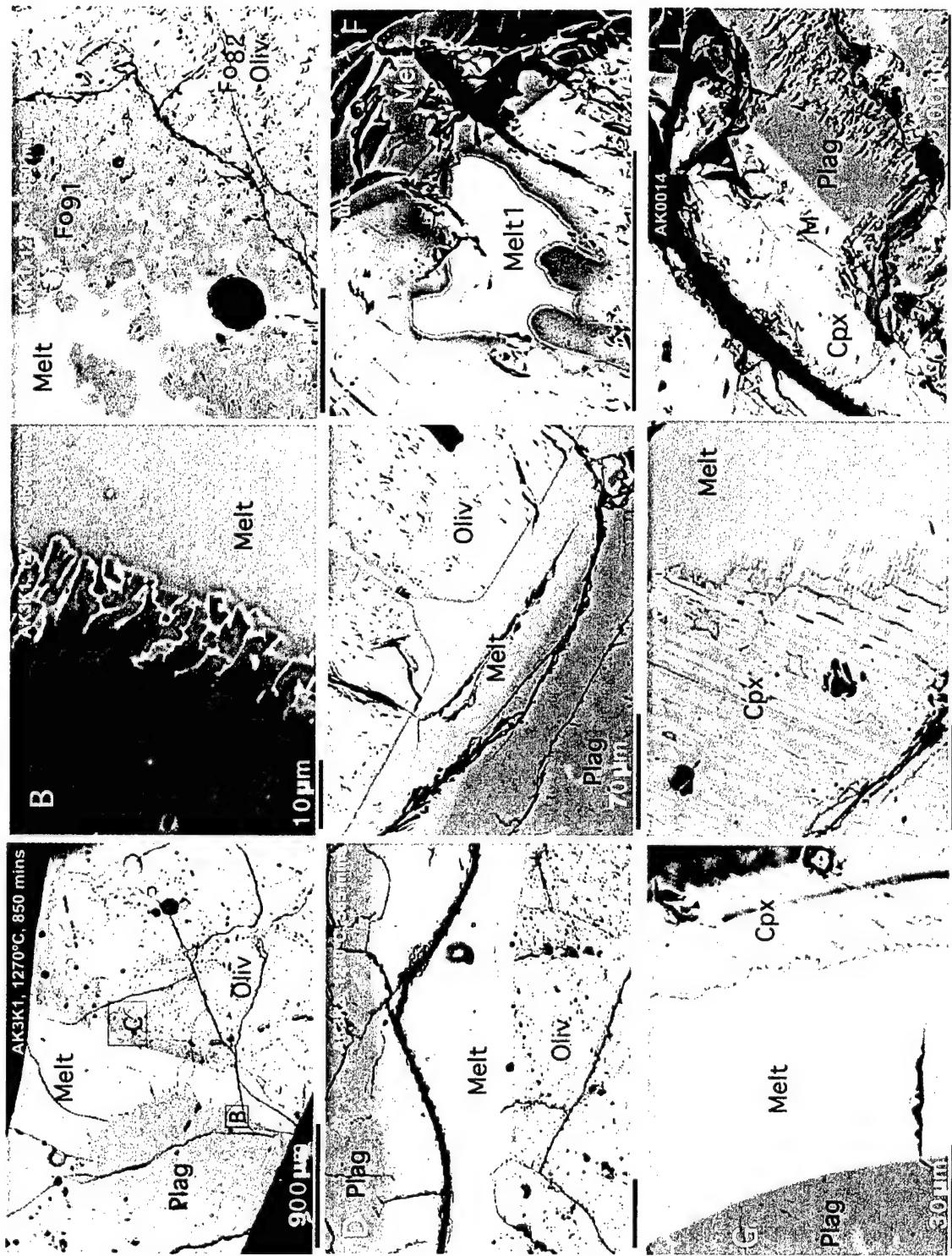
Figure 4-2: The experimental setup. A: Composite back-scatter electron image of an early plagioclase-olivine experiment. Early experiments were performed with the plagioclase on top of olivine, causing the interfacial melt to become thicker due to plagioclase floatation. The melt produced is pressed to the sides of the interface and wraps the crystals. B: A latter experiment performed with the mafic mineral on top of plagioclase, wrapped with a platinum quenching wire. C: The appearance of an experiment before melting. D: Schematic cross-section of the experiments after melting. The melt is produced at the interface between the minerals and at the immediate corners, and is pressed out to the sides. E: The appearance of an experiment after melting. The melt clearly wraps the outside of the mineral.

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**Table 3:** The composition of the starting material for the dissolution experiments.

Element	Whole-rock analysis, sample 1		Plag1a	Plag1b	Plag 2	Augite 1a	Augite 1b	Augite 2	Olivine 1a	Olivine 1b	Olivine 2
SiO <sub>2</sub>	52.1	SiO <sub>2</sub>	53.5	53.3	55.2	52.3	52.1	52.7	38.5	38.1	41.0
TiO <sub>2</sub>	0.4	TiO <sub>2</sub>				.57	.56	1.20			
Al <sub>2</sub> O <sub>3</sub>	15.4	Al <sub>2</sub> O <sub>3</sub>	29.4	30.0	28.9	2.88	3.28	2.70	.02	.01	.19
FeO(tit)	4.0	Cr <sub>2</sub> O <sub>3</sub>				.065	.304	.181	.004	.002	.078
Fe <sub>2</sub> O <sub>3</sub>	0.8										
FeO*	4.7	FeO	.33	.34	.19	6.5	6.5	5.5	24.0	23.7	15.7
MnO	0.1	MnO				.20	.16	.18	.37	.36	.20
MgO	9.5	MgO	.04	.02	.01	17.13	16.80	16.28	38.11	38.32	44.06
CaO	15.4	CaO	12.1	12.6	10.8	19.9	20.1	22.3	.07	.05	.12
Na <sub>2</sub> O	2.3	Na <sub>2</sub> O	4.82	4.57	4.93	.35	.37	.36			
K <sub>2</sub> O	0.0	K <sub>2</sub> O	.06	.06	.33				.01		
P <sub>2</sub> O <sub>5</sub>	0.0	NiO							.04	.03	.03
Total	100.0	Total	100.2	100.8	100.2	99.9	100.1	101.5	101.0	100.6	101.4
V (ppm)	220	An%	57.9	60.1	54.7						
Cr (ppm)	272	Mg#				82.4	82.2	84.1	73.9	74.2	83.3
Ni(ppm)	81	Wo				42	42	46			
Cu (ppm)	50	En				50	49	47			
Zn (ppm)	31	Fs				8	9	7			
Rb (ppm)	1	Notes: Plag=plagioclase. All minerals used are natural mineral separates cut from Hole 735B and other sites as mentioned in the text. The mineral compositions were determined on the MIT electron microprobe									
Sr (ppm)	150	An is Anorthite % of plagioclase, calculated from the molar Ca/(Ca+Na).									
Y (ppm)	14	Mg# is the molar Mg/(Mg+Fe) of clinopyroxene. Fo is forsterite content of olivine, and is calculated as Mg#.									
Zr (ppm)	26	All values are in wt% unless otherwise noted. The whole-rock analysis is from Coogan et al, 2000.									
Nb (ppm)	1.0	Minerals 1a, 1b and 2 refers to those used in the AK00- AK01- and AK3K- experiments.									

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples



## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Figure 4-3: Back-scatter electron images of cross-sections through dissolution experiments. A: A 1/4 cross-section through the experiment shown in 2E. The bottom edge is the center. The interface between the minerals is smooth, and the corners of the minerals are melted. Impurities inside the natural olivine have caused some melting in the olivine (light color), but this does not effect the general dissolution rate. B: Inset in A; the reaction surface between plagioclase and melt is very narrow, although jagged. C: inset in A; there is some recrystallized olivine at its corner, and this olivine has a very high forsterite content. D: The general appearance of the plagioclase-melt-olivine interface. Note the weakly undulating surfaces and the lack of Fe-Mg diffusion in the olivine indicated by a lack of significant color variations. E: Preferential dissolution occurs at an olivine grain boundary, though this does not affect the overall dissolution rate. F: Color variations adjacent to a melt pocket (Melt 1) inside olivine shows strong re-equilibration with the host, although the adjacent melt disk (Melt 2) is less than  $2\mu\text{m}$  away and in complete disequilibrium with the olivine. G: The interface melts between plagioclase and augite (cpx). The orthopyroxene exsolution lamellae in augite do not appear to affect the general dissolution at the interface. Note also how smooth the surface of plagioclase is, possibly because of the high Ca# of the melt. H: Higher temperature experiments see some melting inside the augite, although this does not affect the general interface. The perpendicular crystals are from quenching. I: The smooth interface to the melt-disk between plagioclase and augite is clear even when the lamellae are almost perpendicular to the surface.

experiment was terminated. The experiments were quenched in air, as it was discovered that very little quench-growth was produced compared to the experiments that were quenched in a water-bath. The sample assembly was cast in epoxy and cut perpendicular to the mineral-melt interface. The thickness of each mineral was measured using a microscope and the same micrometer as before in the center of each experiment after 5 $\mu\text{m}$  grit polishing. Finer polishing caused the plagioclase and glass to become almost indistinguishable in reflected light. In addition, the thickness of each mineral was measured in the electron microprobe using backscatter-electron imaging. Both measurement techniques produced similar results. The measurements of mineral thickness after each experiment probably have the largest errors, because the shapes of the partly dissolved minerals were not always uniform. Thus, we estimate the error in thickness and measurements of each side of the experimental charge based on repeated measurements to be ~5% relative. The absolute errors are smaller for the experiments that melted less, as the interfaces were smoother and the geometry simpler. No attempt was made to account for dissolution during heating or quenching, as this effect is small compared to the precision of the thickness measurements.

### ***2.c. Analytical techniques***

Phase compositions were analyzed with the MIT 4 and 5-spectrometer JEOL 733 Superprobes, using 15kV accelerating voltage and 10nA beam current. Beam-size is 10 $\mu\text{m}$  for glass analyses, except for some of the lower temperature experiments where only small glass-pockets were available, and 1 $\mu\text{m}$  for crystalline phases. In addition, the large mineral grains used in the dissolution experiments were analyzed with a 10 $\mu\text{m}$  beam. Counting times were 10-40 s depending on the elements. Data were reduced with the CITZAF correction package using the atomic number correction of Duncumb and Reed, the absorption coefficients of Heinrich and the fluorescence correction of Reed (Armstrong, 1995).

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Table 4: Results of dissolution experiment

Run #	Plag #	Oliv #	T (°C)	t (min)	Plagioclase thickness			Olivine thickness			Total melt (μm)	
					Before (μm)	After (μm)	Melt (μm)	Before (μm)	After (μm)	Melt (μm)		
AK0019	1a	1a	1270	24	940	851	89	4	965	930	35	2
AK0020	1a	1a	1270	60	2314	1899	415	21	2113	1824	290	14
AK0021	1a	1a	1270	60	1067	684	383	19	1245	950	295	15
AK0025	1a	1a	1272	240	1615	1161	454	23	864	679	184	9
AK0104	1b	1b	1300	30	2362	528	1834	92	1242	1001	241	12
AK0105	1b	1b	1300	240	2197	0	2197	110	841	508	333	17
AK0106	1b	1b	1300	60	2337	818	1519	76	1144	557	587	29
AK0106	1b	1b	1300	60	2337	234	2103	105	1144	557	587	29
AK3K1	2	2	1270	850	4216	3011	1205	60	2088	1804	284	14
AK3K2	2	2	1270	175	4191	4026	165	8	2129	2129		165
AK3K3	2	2	1270	565	2695	2225	470	23	2134	1905	229	11
AK3K4	2	2	1270	1320	3533	1975	1558	78	2306	2027	279	14
AK3K6	2	2	1247	1080	2631	2631			2383	2383		
AK3K7	2	2	1255	1451	2631	2032	599	30	2383	1617	766	38
AK3K8	2	2	1255	1118	2113	1473	640	32	2136	1588	549	27
AK3K10	2	2	1330	183	4012	1629	2383	119	2357	2066	291	15
AK3K11	2	2	1330	69	2431	1575	856	43	2418	2240	178	9
AK3K12	2	2	1330	30	2273	1778	495	25	2413	2151	262	13
AK3K18	2	2	1290	91	2332	1725	607	30	2421	2210	211	11
AK3K5	2	2	1290	60	3782	2455	1327	66	2395	1659	736	37

Notes: The errors are based on the errors in replicability of the measurements.  
 Plag=plagioclase, oliv=olivine, cpx=augite. t(min)=duration of the run in minutes.

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Table 4, continued: Results of dissolution experiment

Run #	Plag #	Cpx #	T (°C)	t (min)	Plagioclase thickness			Clinopyroxene thickness			Total melt (μm)
					Before (μm)	After (μm)	Melt (μm)	Before (μm)	After (μm)	Melt (μm)	
AK0018	1a	1a	1270	20	1676	1494	182	9	940	834	106
AK0024	1a	1a	1271	235	1476	825	651	33	1697	0	1697
AK0101	1b	1b	1240	60	597	448	149	7	991	802	189
AK0102	1b	1b	1240	120	643	494	149	7	978	738	240
AK0103	1b	1b	1240	840	1168	727	441	22	1168	426	742
AK3K14	2	2	1290	104	4493	3556	937	47	3005	508	2497
AK3K15	2	2	1280	30	2263	1956	307	15	2830	1778	1052
AK3K16	2	2	1260	60	2294	1981	312	16	2916	2769	147
AK3K17	2	2	1260	180	2314	1880	434	22	2992	2362	630
AK3K19	2	2	1210	670	4478	4069	409	20	2916	2819	97
AK3K21	2	2	1300	45	4036	3086	950	47	2974	0	2974
AK3K22	2	2	1300	15	3965	3429	536	27	2967	2253	714
AK3K23	2	2	1290	33	4018	3772	246	12	2817	0	2817
AK3K24	2	2	1240	293	3952	3193	759	38	2906	2187	719
AK3K26	2	2	1220	776	3828	2868	960	48	2819	1908	912
AK3K27	2	2	1220	420	3642	3183	460	23	2822	2517	305

Notes: The errors are based on the errors in replicability of the measurements.  
 Plag=plagioclase, oliv=olivine, cpx=augite. t(min) =duration of the run in minutes

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**Table 4 Cont: Results of dissolution experiment**

Dissolution rates (mm/s)

Errors

Run #	Plag	Olivine	Total	Dissolution rates mm/ $\sqrt{s}$			Dissolution rates mm/ $\sqrt{s}$		
				Plag	Oliv	Total	Plagio clase	Olivine	Total
AK0019	6.2E-6	2.4E-6	8.6E-6	3.1E-7	2.4E-6	4.3E-7	2.3E-4	9.3E-5	3.3E-4
AK0020	1.2E-5	8.0E-6	2.0E-5	5.8E-7	8.0E-6	9.8E-7	6.9E-4	4.8E-4	1.2E-3
AK0021	1.1E-5	8.2E-6	1.9E-5	5.3E-7	8.2E-6	9.4E-7	6.4E-4	4.9E-4	1.1E-3
AK0025	3.2E-6	1.3E-6	4.4E-6	1.6E-7	1.3E-6	2.2E-7	3.8E-4	1.5E-4	5.3E-4
AK0104	1.0E-4	1.3E-5	1.2E-4	5.1E-6	1.3E-5	5.8E-6	4.3E-3	5.7E-4	4.9E-3
AK0105	1.5E-5	2.3E-6	1.8E-5	7.6E-7	2.3E-6	8.8E-7	1.8E-3	2.8E-4	2.1E-3
AK0106	4.2E-5	1.6E-5	7.5E-5	2.1E-6	1.6E-5	2.9E-6	2.5E-3	9.8E-4	4.5E-3
AK0106	5.8E-5	1.6E-5	7.5E-5	2.9E-6	1.6E-5	3.7E-6	3.5E-3	9.8E-4	4.5E-3
AK3K1	2.4E-6	5.6E-7	2.9E-6	1.2E-7	5.6E-7	1.5E-7	5.3E-4	1.3E-4	6.6E-4
AK3K2	1.6E-6	1.6E-6	1.6E-6	7.9E-8	7.9E-8	1.6E-4	1.6E-4	1.6E-4	3.8E-4
AK3K3	1.4E-6	6.7E-7	2.1E-6	6.9E-8	6.7E-7	1.0E-7	2.6E-4	1.2E-4	6.5E-4
AK3K4	2.0E-6	3.5E-7	2.3E-6	9.8E-8	3.5E-7	1.2E-7	5.5E-4	9.9E-5	6.5E-4
AK3K6									
AK3K7	6.9E-7	8.8E-7	1.6E-6	3.4E-8	8.8E-7	7.8E-8	2.0E-4	2.6E-4	4.6E-4
AK3K8	9.5E-7	8.2E-7	1.8E-6	4.8E-8	8.2E-7	8.9E-8	2.5E-4	2.1E-4	4.6E-4
AK3K10	2.2E-5	2.7E-6	2.4E-5	1.1E-6	2.7E-6	1.2E-6	2.3E-3	2.8E-4	2.6E-3
AK3K11	2.1E-5	4.3E-6	2.5E-5	1.0E-6	4.3E-6	1.2E-6	1.3E-3	2.8E-4	1.6E-3
AK3K12	2.8E-5	1.5E-5	4.2E-5	1.4E-6	1.5E-5	2.1E-6	1.2E-3	6.2E-4	1.8E-3
AK3K18	1.1E-5	3.9E-6	1.5E-5	5.6E-7	3.9E-6	7.5E-7	8.2E-4	2.9E-4	1.1E-3
AK3K5	3.7E-5	2.0E-5	5.7E-5	1.8E-6	2.0E-5	2.9E-6	2.2E-3	1.2E-3	3.4E-3

Notes: The errors are based on the errors in replicability of the measurements.

Plag=plagioclase, oliv=olivine, cpx=augite. t(min) =duration of the run in minutes

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Table 4, continued: Results of dissolution experiment  
Dissolution rates (mm/s)

Run #	Errors						Dissolution rates mm/ <sup>1/s</sup>		
	Plag	Cpx	Total	Plag	Cpx	Total	Plag	Cpx	Total
AK0018	1.5E-5	8.8E-6	2.4E-5	7.6E-7	8.8E-6	1.2E-6	5.3E-4	3.1E-4	8.3E-4
AK0024	4.6E-6	1.2E-5	1.7E-5	2.3E-7	1.2E-5	8.3E-7	5.5E-4	1.4E-3	2.0E-3
AK0101	4.1E-6	5.2E-6	9.4E-6	2.1E-7	5.2E-6	4.7E-7	2.5E-4	3.1E-4	5.6E-4
AK0102	2.1E-6	3.3E-6	5.4E-6	1.0E-7	3.3E-6	2.7E-7	1.8E-4	2.8E-4	4.6E-4
AK0103	8.8E-7	1.5E-6	2.3E-6	4.4E-8	1.5E-6	1.2E-7	2.0E-4	3.3E-4	5.3E-4
AK3K14	1.5E-5	4.0E-5	5.5E-5	7.5E-7	4.0E-5	2.8E-6	1.2E-3	3.2E-3	4.3E-3
AK3K15	1.7E-5	5.8E-5	7.5E-5	8.5E-7	5.8E-5	3.8E-6	7.2E-4	2.5E-3	3.2E-3
AK3K16	8.7E-6	4.1E-6	1.3E-5	4.3E-7	4.1E-6	6.4E-7	5.2E-4	2.5E-4	7.7E-4
AK3K17	4.0E-6	5.8E-6	9.9E-6	2.0E-7	5.8E-6	4.9E-7	4.2E-4	6.1E-4	1.0E-3
AK3K19	1.0E-6	2.4E-7	1.3E-6	5.1E-8	2.4E-7	6.3E-8	2.0E-4	4.8E-5	2.5E-4
AK3K21	3.5E-5	1.1E-4	1.5E-4	1.8E-6	1.1E-4	7.3E-6	1.8E-3	5.7E-3	7.6E-3
AK3K22	6.0E-5	7.9E-5	1.4E-4	3.0E-6	7.9E-5	6.9E-6	1.8E-3	2.4E-3	4.2E-3
AK3K23	1.2E-5	1.4E-4	1.5E-4	6.2E-7	1.4E-4	7.7E-6	5.5E-4	6.3E-3	6.9E-3
AK3K24	4.3E-6	4.1E-6	8.4E-6	2.2E-7	4.1E-6	4.2E-7	5.7E-4	5.4E-4	1.1E-3
AK3K26	2.1E-6	2.0E-6	4.0E-6	1.0E-7	2.0E-6	2.0E-7	4.4E-4	4.2E-4	8.7E-4
AK3K27	1.8E-6	1.2E-6	3.0E-6	9.1E-8	1.2E-6	1.5E-7	2.9E-4	1.9E-4	4.8E-4

Notes: The errors are based on the replicability of the measurements.  
Plag=plagioclase, oliv=olivine, cpx=augite. t(min) =duration of the run in minutes.

### **2.d. Attainment of equilibrium**

Our phase equilibrium experiments were all made with synthetic compositions. Reversal experiments were not conducted and therefore equilibrium has not been demonstrated with respect to all exchange components. However, the Fe-Mg bearing crystalline phases show relatively consistent partitioning of magnesium and iron into the melt (Table 4-2B). The  $Kd_{cpx/melt}^{Fe/Mg}$  is  $0.27 \pm 0.02$ , the  $Kd_{oliv/melt}^{Fe/Mg}$  is  $0.27 \pm 0.03$  and  $Kd_{plag/melt}^{Ca/Na}$  is  $1.07 \pm 0.2$  measured in the AK3 experiments. For the Fo73 experiments  $Kd_{cpx/melt}^{Fe/Mg}$  was found to be  $0.21 \pm 0.02$  and  $Kd_{plag/melt}^{Ca/Na}$   $1.11 \pm 0.27$ . The partition coefficients are close to the expected values from other phase-equilibrium experiments, especially at these refractory compositions (Grove and Bence, 1977; Ulmer, 1989; Grove et al., 1992). Secondly, at least for the experiments 1170°C and higher, the mineral and melt-compositions are homogeneous within the analytical uncertainty for each analyzed element. To achieve very homogeneous phases for all experiments, we would have liked to run the experiments for longer times. However, this was not practical because we wanted to minimize Na-loss in these experiments. In fact, the Fo73 experiments were run approximately 10 times longer than AK3 experiments, resulting in better exchange Kd's but at the expense of much higher Na-loss. For the purpose of comparing the equilibrium melting behavior of these mineral pairs with the dissolution behavior, however, we feel that our experiments have approached equilibrium to a sufficient extent.

## **3. Discussion**

### **3.a. Experimental results**

*Phase Equilibrium Experiments:* AK3 has olivine on the liquidus in the highest temperature experiment (1230°C), and is accompanied by plagioclase at 1210°C, and augite at 1196°C. Below the olivine-plagioclase-augite saturation boundary, at 1150°C, we produced a second set of solid compositions; augite + plagioclase followed by pigeonite. The minerals in these lower temperature experiments (<1150°C) coexist with a very silicic melt and olivine was not found at these temperatures. The Fo73

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composition was not run at temperatures higher than 1190°C, as it was obvious to us that the composition had a very low olivine component, and we therefore chose to pursue the composition with a higher olivine component (e.g. AK3). At the highest temperatures of Fo73, investigated plagioclase and clinopyroxene coexist with a melt with ~56 wt% SiO<sub>2</sub>. The solidus was like in AK3 found at ~1150°C, although small fractions of unusually silicic melts were observed below this temperature. Biggar and Humphries (1981) showed that the temperature-difference between the olivine-clinopyroxene-plagioclase eutectic and the olivine-free clinopyroxene-plagioclase cotectic is only 4°C. We used a materials balance technique to assess whether the bulk composition of the experimental charged changed during the course of each experiments. In the case of the “Fo73” experiments the results of these calculations did not do a good job of recovering the starting compositions. Therefore, we also recalculated the modal proportions by excluding Na<sub>2</sub>O, as shown in italics in Table 4-2a, improving the errors. This improved the materials balanc results, and did not change the modes significantly, suggesting that the original modes are adequate. The compositions of the melts produced under equilibrium conditions are shown in Figure 4-1 in the Olivine – Clinopyroxene – Plagioclase pseudo-ternary projection scheme.

In these isothermal experiments, the saturation-boundary (~solidus) for the plagioclase-clinopyroxene-olivine-melt is 1150°C, and the saturation-boundary for augite-plagioclase-melt and olivine-plagioclase-melt is inferred to be ~5°C and ~40°C higher, respectively, based on comparison with previous studies (Biggar and Humphries, 1981) carried out in iron-free systems.

*Melting Kinetics Experiments:* In the earliest experiments (Fo<sub>73</sub>-An<sub>60</sub> (AK00- and AK01-), Mg#~83(Cpx) - An<sub>60</sub> (AK3K-)), we placed the less dense mineral, plagioclase, on top of the mafic mineral to prevent convection in the melt following the method used in previous investigations (e.g. Tsuchiyama, 1985, 1986). However, experiments that melted to a significant extent, and exceeded the generation of a melt film between the two, developed a complex geometry of melt distribution and mineral shape in the plagioclase (Fig. 4-2 a). We found that if we placed the denser mineral on top of the

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assembly (used for the latter half of the experiments;  $\text{Fo}_{82}$  -  $\text{An}_{54}$ , and  $\text{Mg\#}\sim83(\text{Cpx})$  -  $\text{An}_{54}$ ) the geometry of disintegration was easier to characterize (Fig. 4-2 b,c). With this geometry, the distance between the minerals stayed relatively constant and the melt was again pushed to the sides, but now it covered the minerals towards the top (Fig 4-2 d,e). Reaction only happened at the 2-phase interface with the melt. There was, however, excess melting towards the sides of each mineral slice, and along some phase-boundaries in the poly-crystals used for dissolution (Figure 4-3 a,e).

Olivine, the gabbroic mineral that has the fastest internal solid-state diffusion (Gaetani and Watson, 2000), developed only a very narrow (tens of  $\mu\text{ms}$ ) or no observable diffusion gradient along the actively melting interface (Fig. 4-3d, e, f), indicating that the melting rate is faster than solid-state diffusion for Fe/Mg in olivine. However, internally in olivine grains (Fig. 4-3f) melt-pockets originating around impurities that has melted completely re-equilibrate with the olivine. Some recrystallization occurred in the melt close to olivine, away from the most active melt interface (Fig. 4-3a, c). In addition, rare melt-pockets were found trapped inside olivine-grains away from the dissolving interface that have re-equilibrated with the host to form a lower Mg# melt (Fig. 4-3f and Table 4-5).

Plagioclase and clinopyroxene grains also melted without experiencing any internal diffusion of major elements. Plagioclase shows a very narrow ( $\sim10\mu\text{m}$ ) jagged reaction zone (Fig. 4-3b), but no significant anisotropy of melting is apparent at the melt-contact. There is no measurable chemical difference in plagioclase composition towards the melt-contact, within the resolution of the electron microprobe. The augite slab begins to melt internally in the highest temperature dissolution experiments (Fig. 4-3h), but does not show any sign of preferential melting of low-Ca pyroxene exsolution lamellae or preferential melting of different crystal faces (Fig. 4-3g-i) relative to the plagioclase interface.

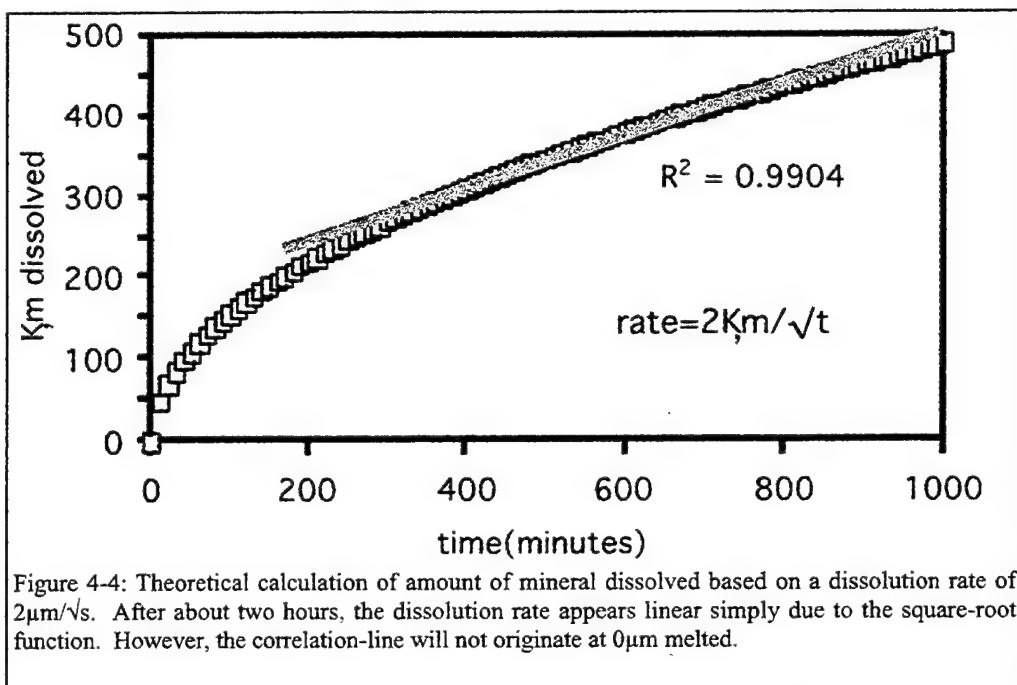
The traditional method of treating dissolution-experiments is to define a reaction function  $\xi = f(\sqrt{t})$  or  $\xi = f(t)$ . These types of experiments are usually designed with one large mineral-grain dissolving into a large volume of melt, both of known and constant

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far-field compositions. The interface between the mineral and the melt is approximately fixed, thus the dissolution rate of the mineral into the melt and the resulting diffusion of components through the melt can readily be measured. Our experiments, however, are very different. At the start of each run, no melt is present between the two mineral-disks. As soon as the minerals are heated above the melting temperature, a melt-film forms between them. When minerals precipitate, interface-kinetics between the melts and crystals may control the rate at which the crystals grow. However, Zhang et al. (1989) used the interface-reaction rates from Kuo and Kirkpatrick (1985) and demonstrated that the interface-melts are saturated with the bounding crystals after  $\sim 1$  second of dissolution, thus there is no interface-control of dissolution in our experiments. The melt-film has a near-constant thickness, as the weight of the overlying mineral squeezes the melt to the sides of the experiment. Each interface between the minerals and melt is therefore moving, and the center of the melt film experiences flow of melt from the minerals towards the center, in addition to flowing from the center of the run towards its sides. The reaction function cannot be solved analytically with these boundary conditions, and a thorough and extensive numerical model is beyond the scope of this paper. We found that the most meaningful way of approaching the reaction-function is to measure the linear loss of material from each mineral disk, in order to assess the rate of dissolution.

Our dissolution rates are summarized in Table 4-4. Marvin and Walker (1985) found that the melting rate at the onset of melting correlate with  $\sqrt{t}$  but as soon as a steady melt film is formed between the minerals (usually within 1-2 hrs), the newly formed melts are smeared out to the sides of the experiment, and the melting rate becomes linear with time. Unfortunately, they did not perform any experiments for longer than 200 minutes. We did perform experiments for as long as 24 hours, and these show good correlation with  $\sqrt{t}$  and linearly with time after a few hours. Thus, only the  $\sqrt{t}$  correlate with amount melted in a function that originates at the origin (Fig. 4-5a) as would be expected for diffusion-controlled dissolution. After a few hours, continued melting

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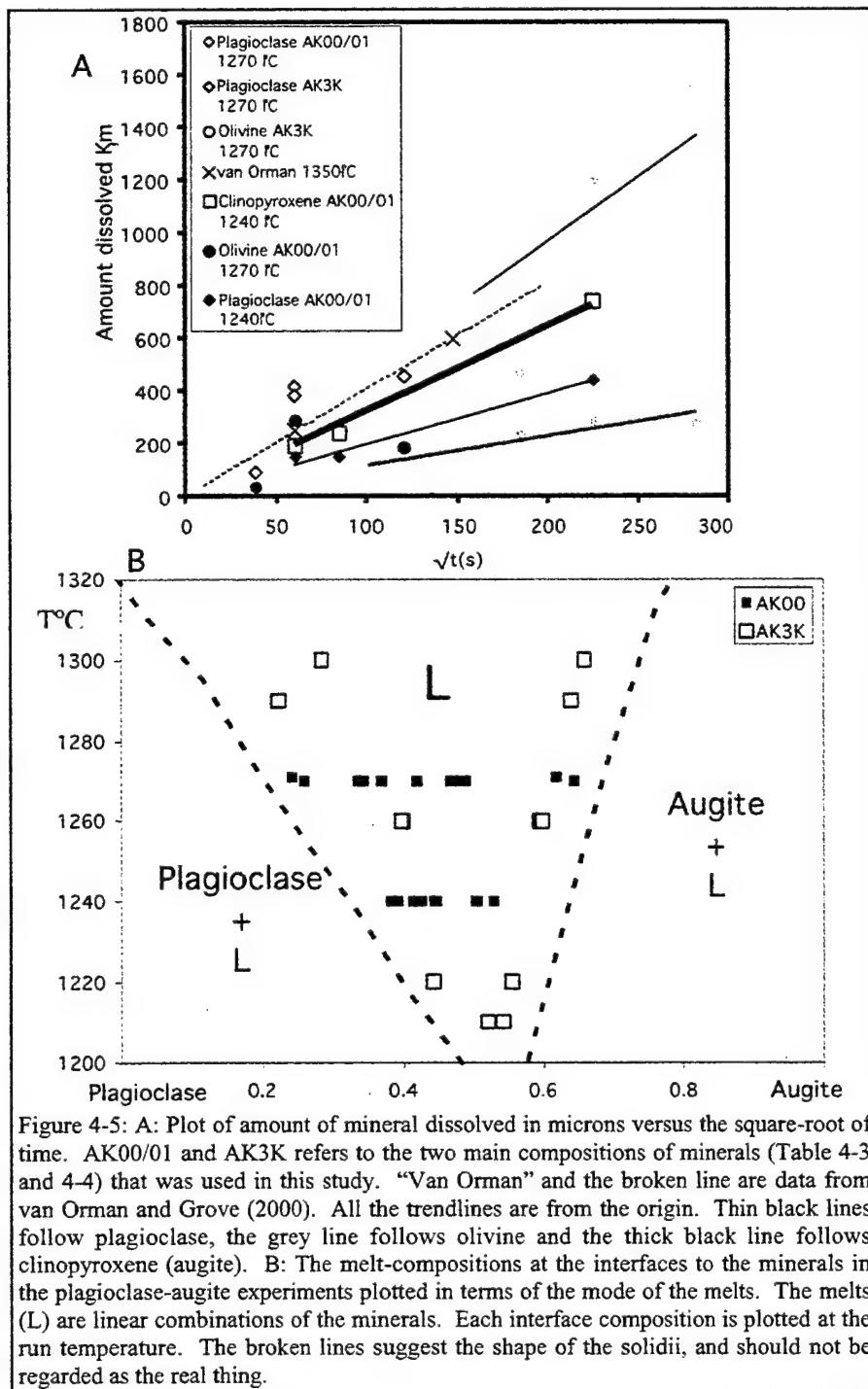


Figure 4-5: A: Plot of amount of mineral dissolved in microns versus the square-root of time. AK00/01 and AK3K refers to the two main compositions of minerals (Table 4-3 and 4-4) that was used in this study. “Van Orman” and the broken line are data from van Orman and Grove (2000). All the trendlines are from the origin. Thin black lines follow plagioclase, the grey line follows olivine and the thick black line follows clinopyroxene (augite). B: The melt-compositions at the interfaces to the minerals in the plagioclase-augite experiments plotted in terms of the mode of the melts. The melts (L) are linear combinations of the minerals. Each interface composition is plotted at the run temperature. The broken lines suggest the shape of the solidii, and should not be regarded as the real thing.

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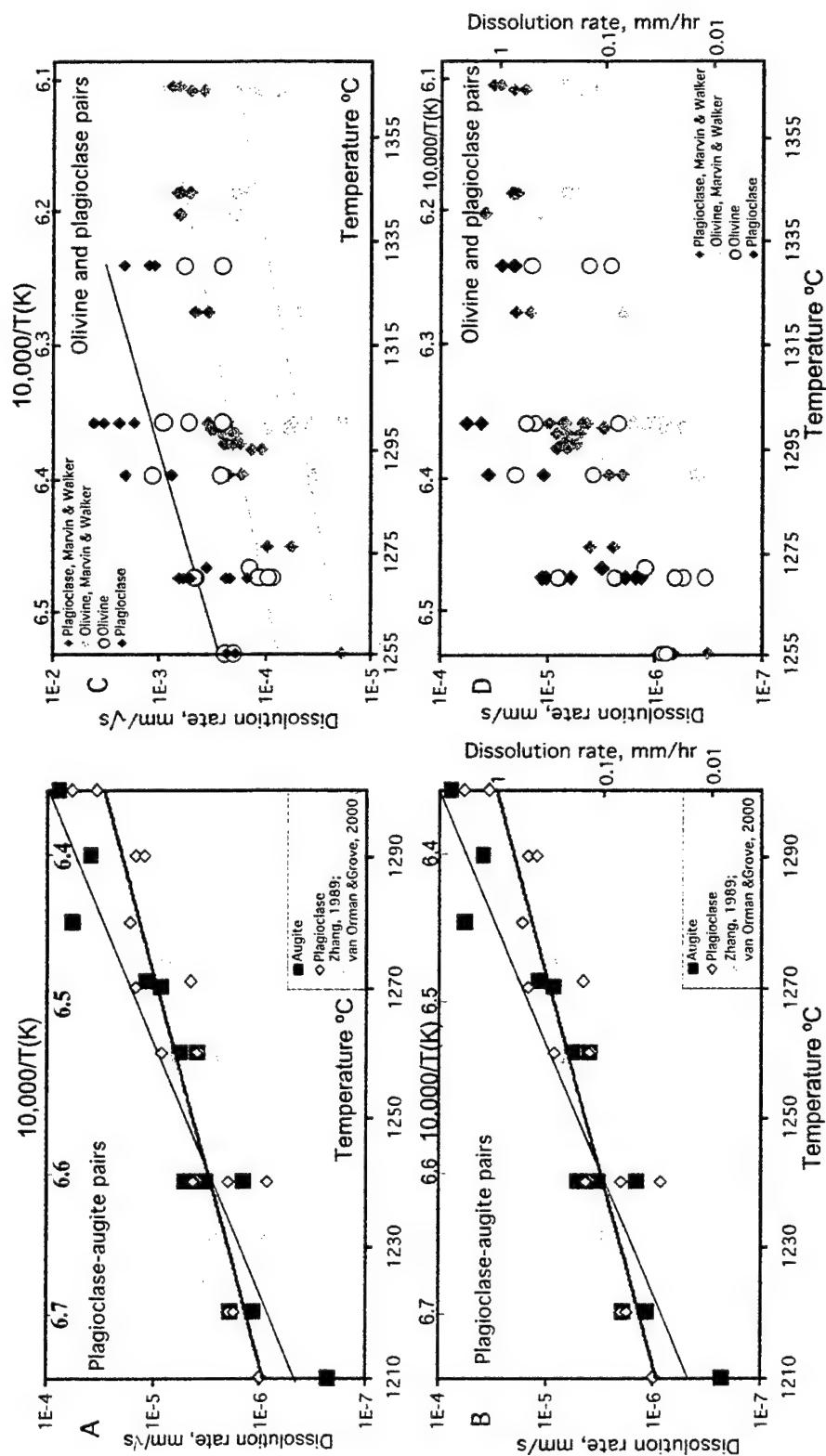


Figure 4-6: Dissolution rates for our experiments together with other studies (Zhang et al., 1989; van Orman and Grove, 2000).

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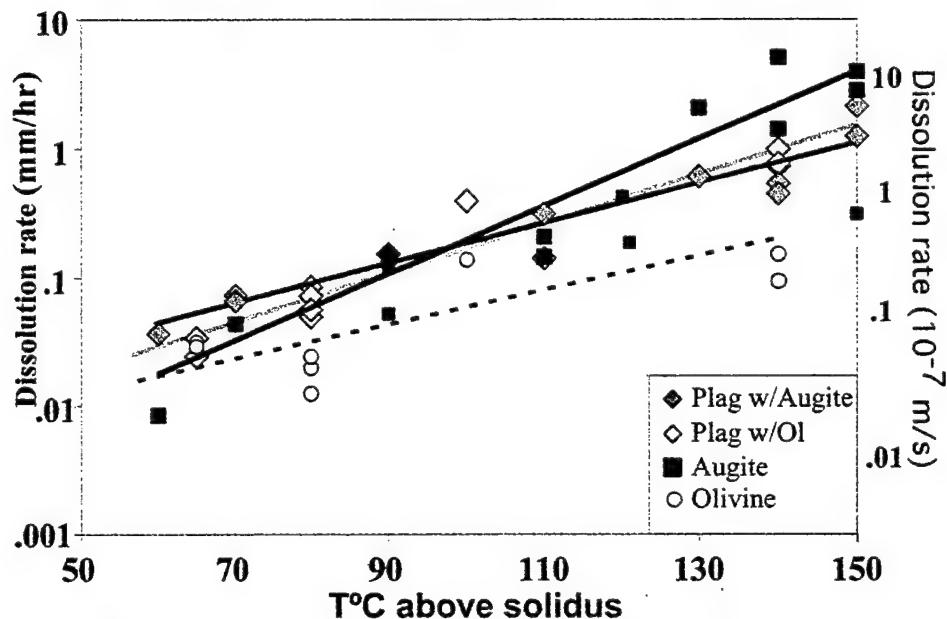


Figure 4-7: The dissolution rates for plagioclase dissolved with augite or olivine overlap when plotted vs.  $\Delta T^{\circ}\text{C}$  above solidus.

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appears linear as a simple consequence of the  $\sqrt{t}$  relationship (Fig. 4-4). The reaction functions for dissolution may therefore be approximated as:

$$\xi = 2.43 * 10^{(-26)} * e^{0.004 \sqrt{T(C)}} \text{ for augite melted with plagioclase;}$$

$$\xi = 3.47 * 10^{(-15)} * e^{0.002 \sqrt{T(C)}} \text{ for plagioclase melted with augite;}$$

$$\xi = 1.79 * 10^{(-21)} * e^{0.0032 \sqrt{T(C)}} \text{ for plagioclase melted with olivine}$$

The rates for olivine-dissolution have a poor correlation with  $\sqrt{t}$  (and t), and a reaction-function has not been calculated for olivine. For all these functions, the approximation with t is reasonable when working with 10's of mm's and times less than a week, and the dissolution-rates may be used in units of mm/hr.

Dissolution rates for the plagioclase - clinopyroxene and plagioclase - olivine pairs are plotted in Figure 4-6 along with other dissolution studies (Marvin and Walker, 1985; Zhang et al., 1989; Van Orman and Grove, 2000). Marvin and Walker (1985) investigated the melting rates of  $Fo_{90}-An_{54}$  at 0.1 MPa from 10 to 200 minutes, Zhang et al. (1989) dissolved diopside in an andesite, for 15 minutes to 23 hours at various pressures, whereas van Orman and Grove (2000) melted diopside with ilmenite at ~13 GPa for 1-6 hours. Our rates for plagioclase and augite are very similar to each other, and the other authors' data agree well. Note that the temperatures of their experiments have been translated to 1 atm conditions (by subtracting 9°C/kbar, on the basis of the algorithms from Grove et al. (1992)). At a given temperature, plagioclase melts slower when in contact with an olivine than when in contact with an augite. For instance,  $An_{54}$  plagioclase at 1290°C melts 0.54-mm/hr when in contact with augite, but only 0.17-mm/hr in contact with olivine. The minerals of the augite-plagioclase pairs, however, melt at comparable rates to the plagioclase of the plagioclase-olivine pair, relative to their solidus temperature (Figure 4-7). Thus, the intermediate (relative to the spectrum of gabbros at Atlantis Bank) gabbro-composition that our experiments represent, would react very effectively with a melt of a temperature of 1230°C or higher. Neither of the individual minerals was held significantly above their individual solidus ( $An_{54}$  has a solidus of ~1320°C).

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Table 5: Range of melts from the dissolution experiments using gabbroic raw material.

Sample #	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnO	Mgo	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	P <sub>2</sub> O <sub>5</sub>	Mg#	Ca#	What
<b>AK00 ol-pl</b>														
AK19	49.1	.03	21.0		7.7	.12	10.9	8.8	3.21	.04	.09	71.6	60.3	Rim to olivine
AK19	50.6	.00	23.1		5.1	.12	7.3	10.0	3.64	.05	.10	71.8	60.4	Rim to plagioclase
AK19	50.6	.03	20.6		8.1	.12	9.5	8.4	3.17	.04	.08	67.5	59.4	Cotectic melt
AK21	49.3	.01	21.5		7.1	.13	10.3	9.0	3.71	.05	.06	72.1	57.2	Rim to olivine
AK21	50.8	.01	22.5		6.7	.09	9.8	8.8	4.03	.04	.15	72.4	54.8	Rim to plagioclase
AK21	48.4	.20	19.1	.006	7.8	.16	11.2	11.5	2.29	.04	.12	72.0	73.4	Cotectic melt
AK25	48.7	.18	20.9		8.3	.16	10.2	9.6	3.20	.06	.11	68.5	62.4	Rim to olivine
AK25	49.6	.13	23.2		6.0	.13	8.6	8.9	4.02	.07	.11	72.1	54.9	Rim to plagioclase
AK25	49.1	.11	21.5	.002	7.3	.11	10.2	9.2	3.29	.04	.08	71.5	60.6	Cotectic melt
AK25	45.8	.32	20.1	.008	13.6	.19	8.1	10.4	2.38	.05	.10	51.5	70.8	Pocket inside olivine
AK20	49.7	.31	15.0	.038	7.4	.12	10.9	12.5	2.25	.05	.13	72.4	75.4	Rim to olivine
AK20	50.8	.25	19.4	.019	4.8	.10	8.2	13.0	3.20	.07	.20	75.2	69.2	Rim to plagioclase
AK20	51.7	.30	17.9	.027	6.7	.14	8.7	10.7	2.62	.05	.13	70.0	69.3	Cotectic melt
<b>AK00 cpx-pl</b>														
AK22-1	53.6	.37	11.8	.080	5.1	.13	10.8	15.0	1.85	.03	.17	79.3	81.7	Rim to clinopyroxene
AK22-1	52.3	.17	20.6		2.9	.10	6.5	15.0	3.01	.02	.14	79.9	73.3	Rim to plagioclase
AK22-1	50.4	.50	15.8	.009	7.9	.15	10.5	12.6	2.19	.04	.16	70.3	76.1	Cotectic melt
AK22-2	55.3	.33	12.9	.057	4.2	.15	11.3	16.3	2.48	.02	.30	82.6	78.3	Rim to clinopyroxene
AK22-2	56.8	.26	17.9	.013	3.7	.11	9.0	15.1	2.65	.04	.81	81.3	75.9	Rim to plagioclase
AK18-1	53.1	.41	13.2	.033	4.2	.16	10.4	17.3	2.02	.03	.20	81.4	82.6	Rim to clinopyroxene
AK18-1	53.0	.34	14.3	.030	4.5	.11	9.7	15.8	2.22	.03	.16	79.3	79.7	Rim to plagioclase
AK18-1	50.8	.52	16.1	.014	5.1	.15	10.0	16.4	2.10	.03	.19	77.8	81.1	Cotectic melt
AK18-2	53.7	.77	10.1		5.7	.16	11.9	17.2	1.79	.03	.19	78.7	84.2	Rim to clinopyroxene
AK18-2	52.9	.13	25.3		1.8	.06	3.1	13.3	4.28	.05	.10	75.4	63.1	Rim to plagioclase
AK24	52.1	.52	9.9	.107	5.5	.21	13.4	17.6	1.61	.03	.24	81.3	85.8	Rim to clinopyroxene
AK24	49.2	.31	20.4	.057	3.7	.17	8.0	16.5	2.08	.04	.20	79.3	81.4	Rim to plagioclase
<b>AK01 ol-pl</b>														
ak0104-1	47.8	.08	19.5	.000	8.3	.10	12.4	9.3	3.01	.05	.06	72.6	63.0	Rim to olivine
ak0104-1	51.6	.07	24.5	.021	4.7	.07	6.9	8.8	4.76	.08	.04	72.2	50.6	Rim to plagioclase
ak0104-2	48.5	.13	19.5	.010	9.8	.20	11.8	8.6	3.03	.03	.11	68.1	61.1	Rim to olivine
ak0104-2	50.1	.13	23.6	.006	5.7	.13	7.8	8.7	4.13	.05	.01	70.8	53.7	Rim to plagioclase
ak0106-1	49.2	.07	19.7	.037	9.0	.10	12.9	8.7	2.63	.03	.15	71.8	64.7	Rim to olivine
ak0106-1	50.3	.05	25.9	.017	4.8	.09	7.3	8.0	4.40	.08	.14	73.0	50.1	Rim to plagioclase
ak0106-2	48.9	.06	20.4	.010	7.3	.11	11.9	8.9	3.18	.01	.06	74.4	60.7	Rim to olivine
ak0106-2	51.3	.11	24.0	.000	4.6	.06	7.7	7.6	4.76	.07	.06	74.9	46.8	Rim to plagioclase
<b>AK01 cpx-pl</b>														
ak0101	53.2	.36	15.1	.051	4.3	.03	9.7	15.5	2.48	.04	.17	80.1	77.5	Rim to clinopyroxene
ak0101	52.6	.32	16.5	.090	3.6	.04	8.2	16.0	2.60	.05	.19	80.3	77.3	Rim to plagioclase
ak0101	53.8	.45	15.4	.083	3.4	.11	9.8	15.3	2.34	.04	.16	83.5	78.3	Cotectic melt
ak0102	53.7	.40	14.5	.061	4.5	.11	9.8	15.1	2.52	.03	.13	79.6	76.8	Rim to clinopyroxene
ak0102	52.9	.30	17.0	.086	4.0	.13	9.1	15.3	2.74	.05	.15	80.3	75.4	Rim to plagioclase
ak0102	53.4	.73	13.9	.052	3.8	.14	10.8	15.8	2.00	.03	.33	83.5	81.3	Cotectic melt
ak0103	53.4	.52	13.5	.038	4.5	.11	10.6	15.7	2.11	.03	.14	81.0	80.5	Rim to clinopyroxene
ak0103	52.5	.44	17.9	.048	3.9	.10	8.9	15.5	2.55	.02	.11	80.4	77.1	Rim to plagioclase
ak0103	53.8	.61	14.5	.056	3.8	.13	10.2	15.8	2.24	.03	.15	82.5	79.5	Cotectic melt
<b>AK3K ol-pl</b>														
AK3K1	51.6	.01	21.7	.06	4.9	.06	11.1	8.6	3.16	.25	.00	80.2	60.1	Rim to plagioclase
AK3K1	53.3	.00	24.1	.05	3.3	.05	7.9	7.6	4.43	.36	.00	80.8	48.5	Rim to olivine
AK3K3	50.8	.09	22.2	.00	4.6	.07	9.6	8.9	3.87	.25	.00	78.6	55.9	Rim to plagioclase
AK3K3	52.6	.00	23.8	.00	3.2	.00	7.8	8.1	4.42	.40	.15	81.4	50.3	Rim to olivine
AK3K4	50.8	.09	21.6	.00	4.5	.05	10.8	8.3	3.05	.25	.00	81.0	60.1	Rim to plagioclase
AK3K4	52.4	.06	23.7	.02	3.5	.06	8.1	7.6	4.58	.40	.06	80.6	48.0	Rim to olivine

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**Table 5: Range of melts from the dissolution experiments using gabbroic raw material.**

Sample #	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnO	Mgo	Cao	Na <sub>2</sub> O	K <sub>2</sub> O	P <sub>2</sub> O <sub>5</sub>	Mg#	Ca#	What
AK3K5	50.7	.00	20.5	.00	4.8	.04	12.2	8.6	3.06	.21	.00	81.8	60.9	Rim to plagioclase
AK3K5	53.0	.05	24.2	.02	2.7	.11	7.0	7.8	4.73	.40	.00	82.1	47.7	Rim to olivine
AK3K11	49.9	.00	19.4	.07	5.7	.01	14.5	8.3	3.14	.19	.00	82.0	59.2	Rim to plagioclase
AK3K11	53.3	.06	25.4	.00	2.1	.06	6.0	7.9	5.82	.41	.00	83.4	42.7	Rim to olivine
AK3K12	50.1	.05	19.3	.08	6.1	.02	14.8	8.4	3.23	.15	.00	81.2	59.0	Rim to plagioclase
AK3K12	53.6	.00	26.3	.04	2.0	.00	5.4	7.4	5.96	.44	.00	82.9	40.8	Rim to olivine
AK3K18	47.4	.00	16.7	.00	5.8	.06	19.3	6.4	3.39	.27	.00	85.6	50.9	Rim to plagioclase
AK3K18	52.2	.03	24.1	.00	2.4	.00	6.7	6.9	5.57	.63	.00	82.9	40.7	Rim to olivine
<b>AK3K cpx-pl</b>														
AK3K16	52.2	.39	11.9	.08	4.2	.01	10.9	15.3	2.45	.18	.00	82.2	77.5	Rim to augite
AK3K16	52.5	.25	17.6	.00	2.7	.00	7.6	14.3	3.17	.21	.00	83.5	71.4	Rim to plagioclase
AK3K17	52.3	.51	11.9	.09	3.8	.12	10.9	15.4	1.97	.17	.00	83.8	81.2	Rim to augite
AK3K17	51.8	.30	17.6	.11	2.5	.10	7.9	14.1	2.82	.26	.00	85.0	73.4	Rim to plagioclase
AK3K19	53.1	.27	15.3	.08	4.0	.13	8.5	14.3	2.95	.24	.00	79.1	72.8	Rim to augite
AK3K19	53.0	.32	16.1	.00	3.7	.08	8.3	14.2	2.92	.19	.00	80.2	72.9	Rim to plagioclase
AK3K22	52.6	.30	9.3	.10	4.5	.10	12.6	16.9	1.96	.12	.00	83.3	82.7	Rim to augite
AK3K22	52.1	.13	19.6	.03	2.2	.00	6.7	14.1	3.34	.29	.00	84.6	70.0	Rim to plagioclase
AK3K14	51.6	.56	7.2	.08	4.8	.10	14.1	16.8	1.42	.07	.00	83.9	86.7	Rim to augite
AK3K14	51.4	.19	18.4	.07	2.2	.07	7.0	13.9	3.24	.23	.00	85.0	70.3	Rim to plagioclase
AK3K26	53.6	.80	13.6	.15	5.8	.14	9.5	14.9	2.57	.17	.00	74.6	76.2	Rim to augite
AK3K26	52.7	.59	16.7	.08	5.0	.19	8.0	14.3	2.92	.16	.00	73.8	73.0	Rim to plagioclase

Notes: Ol=olivine, cpx=augite, pl=plagioclase. Cotectic melt=melt compositions on sides of runs

"Rim to" = compositions next to a mineral

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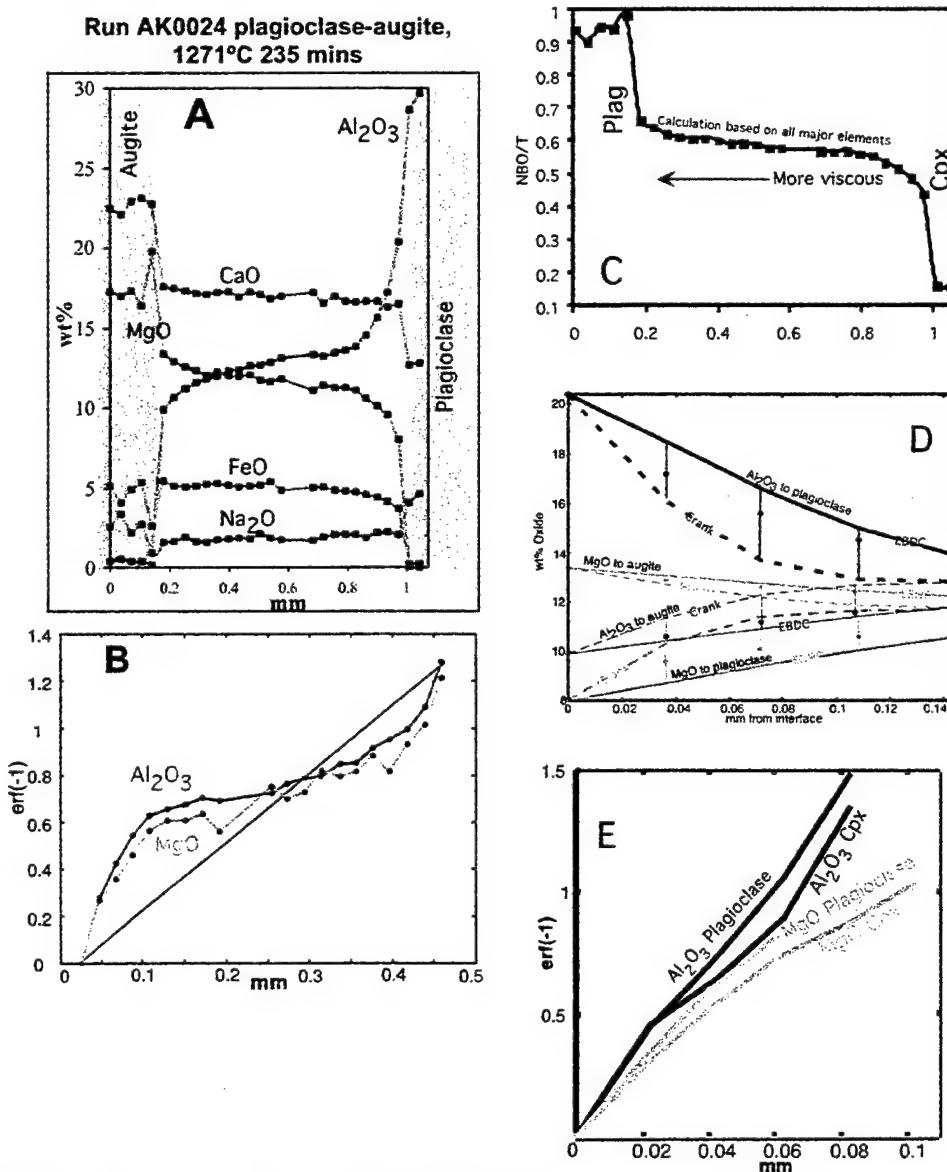


Figure 4-8: A: Chemical profiles across the augite-melt-plagioclase interface in sample AK0024. SiO<sub>2</sub> is not included as the concentrations are similar in both minerals. B: If the entire profile is diffusion controlled,  $\text{erf}^{-1}\left(\frac{x}{2\sqrt{Dt}}\right)$  should follow a straight line through the origin. C: The non-bridging oxygens per tetrahedrally coordinated cation (NBO/T) (Mysen, 1988) are inversely related to the polymerization of the melt and melts to plagioclase are more viscous than those to augite. D: Fits to the data, using diffusion-coefficients calculated from only the melts closest to the interface. Broken lines are calculated using the classical method from Crank (1975), the full lines are using the effective binary diffusion coefficients (EBDC) from Zhang et al., (1989). Thick black lines are Al<sub>2</sub>O<sub>3</sub> in the melt to plagioclase, thin black lines are Al<sub>2</sub>O<sub>3</sub> in the melt to augite, thick grey lines are MgO in the melt towards plagioclase, thin grey lines are MgO to augite. The data are dots in the respective colors, connected to the corresponding diffusion-profiles by vertical lines. E: The melts close to the mineral interfaces follow a near straight line for erf<sup>-1</sup>.

### 3.a.1. Melt compositions

We analyzed the melts in traverses perpendicular to the mineral surfaces (Fig. 4-8). The melt compositions are linear combinations of the mineral compositions (Figure 4-5b, the melt-compositions closest to the mineral-interfaces are recalculated as mineral-modes by simple inversion to the mineral compositions). In addition, in the experiments where we used the oceanic gabbros as starting material, we analyzed the melt that had been squeezed out on the sides of the minerals (labeled “Cotectic melt”). The results are listed in Table 4-5.

If the dissolution-rates are controlled by the rates of diffusion in the melt, the dissolution rates should be the same when we calculate them using the amount of material dissolved, similar to the above treatment, as if we calculate diffusion-coefficients for the melts near the mineral-melt interfaces to find the dissolution rates. In our crystal-pair dissolution-experiments, melt migration and convection during dissolution have modified the compositional profiles of the melt film. This approach will therefore be an approximation, at best. Due to the difficulties regarding the smeared melts, Tsuchiyama (1985) suggested a very different approach to calculating diffusion rates in the melt. However, we do not find our results using his techniques satisfactory.

The concentration profiles adjacent to the crystal interfaces for  $\text{Al}_2\text{O}_3$  and  $\text{MgO}$  (Fig. 4-8A), do resemble diffusion profiles. As noted in other diffusion studies (Watson, 1982) some elements (e.g.  $\text{Na}_2\text{O}$  and  $\text{CaO}$ ) undergo rapid initial diffusion and establish a steady state distribution. In addition, the chemical profile is strongly zoned in non-bridging oxygens per tetrahedrally coordinated cations (NBO/T, Fig. 4-8C) (Mysen, 1988) indicating that the melt is strongly polymerized near the plagioclase crystal making the melt more viscous. A quick analysis of the diffusion-profiles in sample AK0024 (clinopyroxene-plagioclase at  $1271^\circ\text{C}$ , ~4 hours) shows that if the diffusion-coefficient is modeled using the opposite mineral-melt-contact as the far-field composition, the profiles do not show diffusion behavior (Figure 4-8B). However, if only the profile next to the crystal is used (Figure 4-8E), and the composition at the middle of the melt film is used as the far-field melt, the diffusion coefficients may be calculated (Table 4-6).

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**Table 4-6: Method of calculating the dissolution-coefficients using diffusion coefficients in the melt**

**First: Calculate diffusion-coefficients in the melt-film at the interface to the minerals.**

A: Method from Crank (1975) for a semi-infinite volume:

Calculation	Result
$\frac{C(x,t) - C_0}{C_\infty - C_0} = \operatorname{erf}\left(\frac{x}{2\sqrt{Dt}}\right),$ <p><math>C_0</math> = composition at the interface to the mineral, <math>C_\infty</math> = the far-field melt composition, and <math>C(x,t)</math> = composition of the melt away from the interface</p>	Diffusion-coefficients ("D") for the melt-interface to: Plagioclase: $\text{Al}_2\text{O}_3$ : $5.9 \times 10^{-14} \text{ m}^2/\text{s}$ $\text{MgO}$ : $7.6 \times 10^{-14} \text{ m}^2/\text{s}$ Augite: $\text{Al}_2\text{O}_3$ : $1.6 \times 10^{-13} \text{ m}^2/\text{s}$ $\text{MgO}$ : $9.7 \times 10^{-14} \text{ m}^2/\text{s}$

B: Method from Zhang et al. (1989) diffusion-controlled dissolution in a semi-infinite melt reservoir.

The effective binary diffusion coefficient (EBDC) is obtained by:

$$\frac{C(x,t) - C_\infty}{C_0 - C_\infty} = \frac{\operatorname{erfc}\left(\frac{x}{2\sqrt{Dt}} - a\right)}{\operatorname{erfc}(-a)}$$

$$\text{where } a \text{ satisfies } \sqrt{\pi}ae^{a^2} \operatorname{erfc}(-a) = b; b = \frac{C_0 - C_\infty}{C_s - C_0},$$

$a$  is found by a numerical fit,  $C_0$  = composition at the interface to the mineral,  $C_\infty$  = the far-field melt composition,  $C_s$  = the composition of the solid, and  $C(x,t)$  = composition of the melt away from the interface.

EBDC's ("D") for the melt-interface to:

Plagioclase:  
 $\text{Al}_2\text{O}_3$ :  $3.6 \times 10^{-13} \text{ m}^2/\text{s}$   
 $\text{MgO}$ :  $5.2 \times 10^{-13} \text{ m}^2/\text{s}$

Augite:  
 $\text{Al}_2\text{O}_3$ :  $6.5 \times 10^{-13} \text{ m}^2/\text{s}$   
 $\text{MgO}$ :  $5.9 \times 10^{-13} \text{ m}^2/\text{s}$

**Second: The diffusion coefficients can be used to determine the dissolution-rates:**

$\text{rate} = a\rho\sqrt{D}$  $\rho$ = density of the minerals, $a$ is the same as above: $\sqrt{\pi}ae^{a^2} \operatorname{erfc}(-a) = b; b = \frac{C_0 - C_\infty}{C_s - C_0},$	<p><b>Dissolution-rates using diffusion-coefficients from A:</b>  Plagioclase: <math>7.4 \text{ gm}^{-2}\text{s}^{-1/2}</math> Augite: <math>23.5 \text{ gm}^{-2}\text{s}^{-1/2}</math></p> <p><b>Dissolution-rates using EBDC from B:</b>  Plagioclase: <math>0.403 \text{ gm}^{-2}\text{s}^{-1/2}</math> (<math>\text{Al}_2\text{O}_3</math>) <math>0.55 \text{ gm}^{-2}\text{s}^{-1/2}</math> (<math>\text{MgO}</math>)  Augite: <math>0.403 \text{ gm}^{-2}\text{s}^{-1/2}</math> (<math>\text{Al}_2\text{O}_3</math>) <math>0.33 \text{ gm}^{-2}\text{s}^{-1/2}</math> (<math>\text{MgO}</math>)</p> <p><b>Dissolution-rates using measured dissolution distance:</b>  Plagioclase: <math>\sim 1.4 \text{ gm}^{-2}\text{s}^{-1/2}</math> Augite: <math>2.5 \text{ gm}^{-2}\text{s}^{-1/2}</math></p>
--	---

Zhang et al. (1989) developed an approach for diffusion-controlled dissolution in a semi-infinite melt reservoir. This method has been used in other single crystal dissolution experiments (Van Orman and Grove, 2000) to quantify the dissolution process (Table 4-6). The fit to the chemical diffusion-profile using our calculated diffusion-coefficients is shown in figure 4-8D, and it is clear that the effective binary diffusion coefficient (EBDC) provide better fits than the traditional diffusion coefficients. Furthermore, the dissolution rates calculated from the diffusion-coefficients (Table 4-6) are smaller for Zhang et al. (1989) than the rates determined by volumetric dissolution (Table 4-6). This is probably because much of the melt has been pressed out of the interface to the sides of

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the experiment. The *measured* dissolution rate lies between the two determined from the method using the diffusion-coefficients. We therefore conclude that dissolution is controlled by the diffusion in the melt-disk between the minerals.

### **3.b. How efficient is dissolution?**

In his classical paper on assimilation, Bowen (1928) identifies three general mechanisms by which rock/magma chemical reactions may occur. These are: a) reaction of the melt with minerals that are unstable in it to produce new stable minerals; b) direct solution of unstable minerals into the melt; and c) fusion or partial fusion of country-rock or xenoliths and subsequent mixing with the magma by diffusion in the molten stage. All these processes consume energy that must be accommodated by crystallization in the melt if the melt is near or below its liquidus.

Liang (1999) states that almost all dissolution in magmatic silicates is controlled by the diffusion of mass in the liquid, as the diffusion in the liquid is much faster than in the solid-solutions. The chemical disequilibrium between a solid phase (crystal) and a liquid causes dissolution below the solidus of the crystal and require simultaneous diffusion of heat and mass (Kerr, 1995; Tsuchiyama, 1986; Woods, 1993).

### **3.c. Dissolution of gabbro in the lower ocean crust**

The lower ocean crust can interface with ascending magmas in several ways while it is conductively cooling. The efficacy of diffusion of heat into the country-rock depends on the geometry of the interface, and the size of the heat-reservoir. If the magma is a stagnant dike, the interface temperature between the dike and the wall-rock will never exceed half of the difference in temperature between them presuming that the heat diffusivity in the melt and solid are the same (Turcotte and Schubert, 1982). A flowing dike or convecting magma-chamber represents a more efficient heat-reservoir, since convection redistributes the heat, and the temperature of the wall-rock may rise towards that of the magma. In addition, new injection of magma will also increase the available heat. However, a xenolith stoping into a magma, completely encapsulated, would be the most efficient geometry to heat a wall rocks to magmatic temperatures.

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<b>Table 7: The thermodynamic properties of the minerals at 1250°C</b>						
	<b>Diopside</b>	<b>An60</b>	<b>An54</b>	<b>Fo90</b>	<b>Fo73</b>	<b>Fo82</b>
<b>ΔHf (J/mol)</b>	133228	94254	91556	117594	112716	107969
<b>Molecular weight (g/mol)</b>	216.6	271.8	270.9	147.0	147.0	147.0
<b>Cp solid (J/mol-K)</b>	268	350	349	192	196	194
<b>Cp liquid (J/mol-K)</b>	345	399	395	265	260	263
<b>Molecular volume (cm<sup>3</sup>)</b>	68.5	103.5	103.6	46.1	46.5	46.3
<b>T (K) of fusion</b>	1664	1650	1624	2096	1981	2042
<b>T (K) solidus</b>	1573 (Augite82)	1613	1593		1773	1873
<b>Density (g/cm<sup>3</sup>)</b>	3.16	2.63	2.62	3.19	3.16	3.18

Note: Source of data given in text

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The lower crust at fast-spreading ridges is thought to contain a mush-zone with  $\leq$  20% melt (Dunn et al., 2000); i.e. the system is above or close to the solidus. If a magma dissolves a xenolith of different isotopic and trace element signatures than itself, the effect on the hybrid magma composition will not be hidden by subsequent crystallization. Melts saturated with olivine only will show the effects of assimilation on its major-element compositions more than multiply-saturated melts. The addition of a primitive melt with only olivine as a liquidus phase into the lower mush-zone below the ridge may breakup previously generated networks of minerals (Philpotts and Dickinson, 2000; Jerram et al., 2003) and stope rocks into the magma. Bedard et al. (2000) reviewed the concept of syntaxis (the processes, mechanical and chemical, by which magmas react with and assimilate their host rocks). They stated: "it seems self evident that host-rocks and xenocrysts will dissolve into an invading melt only if they are exposed to it" as an argument to support porous melt-flow through pre-existing cumulates. However, if enclosed in hot magma, xenoliths with two- and three-phase mineral boundaries ("triplejunctions") with solidii significantly below the magma-temperature, the xenolith melt along the internal mineral-boundaries when exposed to the heat of the magma. The thermal instability of the phase-boundary requires it to melt, as this is the only way to cool the interface-melt towards a thermally stable condition. We will show that the cumulates only need to be exposed to this heat from the ascending melts, and that if the melts do not escape from these cumulates, but crystallize *in-situ*, they still have a profound effect on the mineral-compositions.

Our melting-rates allow disintegration of a xenolith in hours to days. If a gabbro (i.e. ~50/50 augite/plagioclase) becomes heated above the solidus, the boundaries between the minerals will start to melt. Figure 4-9 demonstrates theoretically how fast a coarser-grained rock will likely melt. At 1240 °C, the slide in Figure 4-9A will look like 4-9B and at 1270°C like 4-9C, after only one hour. However, as we have repeatedly shown in our experiments, plagioclase is buoyant in this melt, whereas augite (and olivine) is denser. If a mineral melts on all sides, it will settle as high or low as possible (Figure 4-9D). Therefore, a close contact with efficient melting will be maintained along at least one interface at all times.

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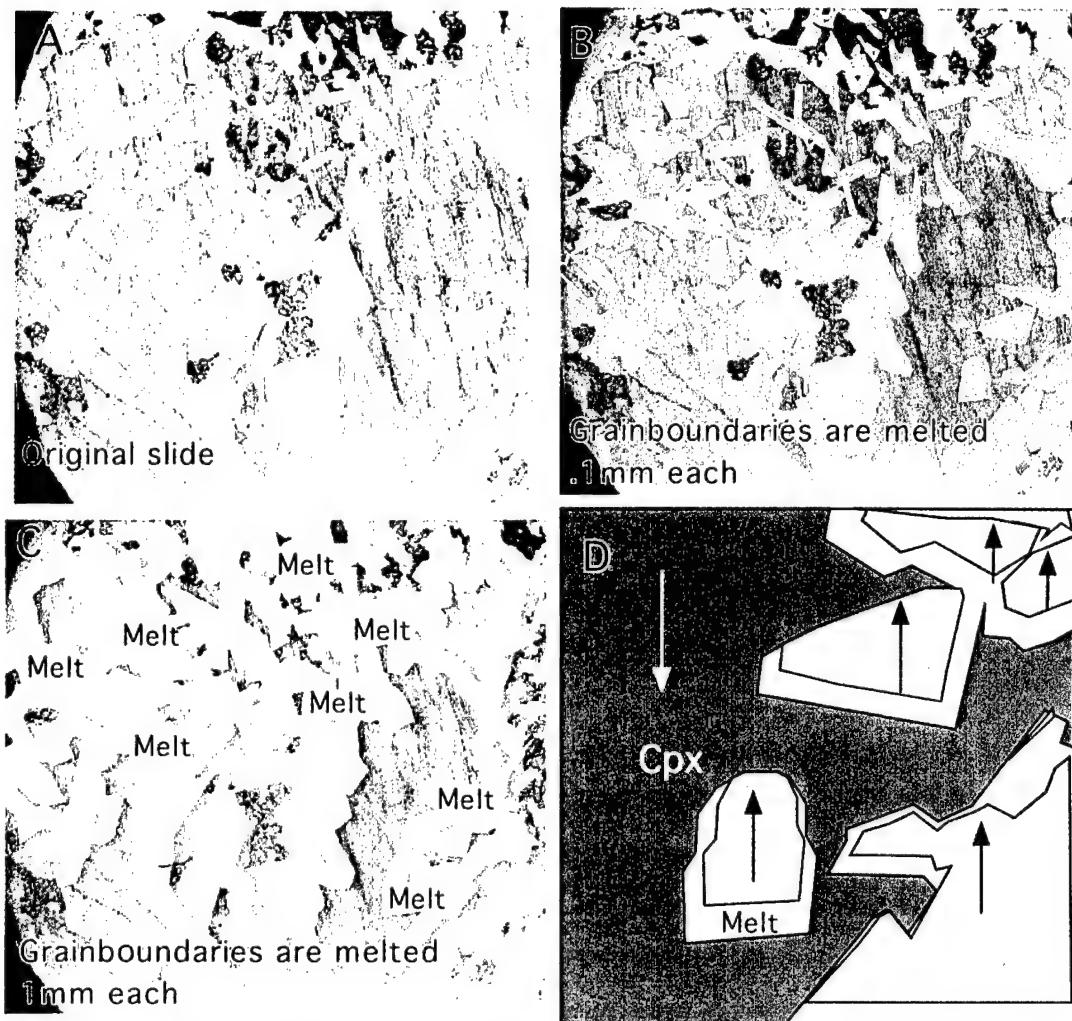


Figure 4-9: A: A thin-section in parallel light of olivine gabbro sample JR31-12-1 from Atlantis Bank. The ophitic minerals are augite, the long laths are plagioclase, and the other minerals are partially altered olivine. B: As a thought experiment, we draw the boundaries between plagioclase and augite and let the thickness of the lines represent amount dissolved on the mineral-boundary. Here, each mineral melted 0.1 mm each, as indicated in white. C: Each mineral has melted 1mm each. D: This sketch is of the lower right corner of the slide, with each grain melted 0.25mm. Here we let the minerals settle to their gravitationally stable level. Plagioclase will rise and clinopyroxene will sink relative to the melt. Thus, a tight connection is kept at least at one face of each mineral.

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### Thermodynamic considerations

As melting and dissolution generally are endothermic processes, “inclusions pass into solution by precipitating their heat equivalent of the phase with which the liquid is saturated” (Bowen, 1928). However, Bowen also stated that the reaction between a melt and an assemblage higher up in the reaction series thus would be exothermic. Addition of rocks from lower in the reaction-series is endothermic and causes the system to cool gradually (e.g. addition of a gabbro to a primitive, mantle-derived melt). It is not clear, however, that the addition of gabbroic xenoliths with the accompanying fractionation of olivine causes the melt volume to reduce significantly.

Our results indicate that the phase-boundaries of lower crust from ocean ridges will melt efficiently and out of equilibrium at a temperature above the plagioclase-clinopyroxene-melt saturation boundary at  $\sim 1150^{\circ}\text{C}$  (i.e. gabbro) and plagioclase-olivine-melt saturation boundary at  $\sim 1195^{\circ}\text{C}$  (i.e. troctolite) or higher. The heat of fusion ( $\Delta H_f$ ) and conductive heat-loss to the surroundings controls how effective the melting will be in infusing the ascending magmas with lower crustal material. The  $\Delta H_f$ 's of the minerals (calculated for  $1250^{\circ}\text{C}$ , Ghiorso and Carmichael, 1980) are listed in Table 4-7. The heat capacity ( $C_p$ ) used for the system is  $1.59 \text{ J/g}^{\circ}\text{C}$  (Robie et al., 1978). Stoping of xenoliths is the most efficient method of reheating wall-rock material by ascending melts. Other methods, including direct heating of the wall rock by a moving magma may be important, but will possibly not provide much new melt to the main magma-chamber as it will remain trapped in the wall rock where it is produced. We will consider two main scenarios: The stoping of gabbro-, or troctolite-xenoliths into a primitive magma (Table 4-8).

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**Table 4-8: Stoping of gabbro or troctolite-xenoliths into a primitive magma**

Thermal re-equilibration of a xenolith with ascending magma; the resulting temperature of the magma:

$$\Delta T = \frac{(Cp_{Augite} * M_{Augite} + Cp_{Plag} * M_{Plag}) * (T_{magma} - T_{Xenolith})}{Cp_{System} * M_{System}}$$

T=temperature; Cp = heat capacity of the mineral, M=mass (grams) of mineral or magma.

Therefore:

If the initial magma is: 1250°C

Add 10 wt% gabbro xenolith at 1150°C (i.e. solidus)

⇒ ~9°C drop in magma temperature

Add 10 wt% troctolite xenolith at 1190°C (i.e. solidus)

⇒ ~3°C drop in magma temperature

⇒ (The subsequent heat of fusion due to new crystallization is ignored.)

Thermal re-equilibration of a xenolith with ascending magma: how fast does it happen?

$$X = \sqrt{Dt}, D = 0.008 \text{ cm}^2/\text{s}; X = \text{distance, t = time in seconds}$$

5 cm cube: 16 minutes; 20 cm cube: 8 hours

Need to compensate assimilation of the xenolith by crystallization in the magma to prevent heat-loss/ -gain

$$\Delta H = M_a \Delta H_f^a - M_c \Delta H_f^c; \text{ Kelemen (1990)}$$

$\Delta H$ =heat of fusion, M=mass (grams) a=assimilated, f = “of fusion”, c=crystallized.

If 10 wt% xenoliths (relative to the magma) just below their solidus are emplaced into a primitive magma by stoping, assuming isenthalpic conditions, the temperature of the magma will only drop slightly (Table 4-8). If the magma is not above its liquidus, it will crystallize to compensate for the heat-loss. The rate of thermal re-equilibration is dependent on the size of the xenolith. When the temperature inside the xenolith oversteps the melting-temperature of the mineral-boundaries, the minerals starts to dissolve independently of the surrounding magma. The mono-mineralic phase-boundaries in contact with the surrounding melts have significantly higher temperature solidii, and will therefore dissolve more sluggishly. Assume the mineral-grains of our idealized xenolith are 1-cm cubes, and the xenolith is a 20-cm cube with the mineral-grains stacked randomly, almost all grains will touch a different mineral and start melting upon emplacement. The melting starts along the outer zone and progresses inward as the heat diffuses. Philpotts and Dickinson (2000) showed that plagioclase-frameworks produced during crystallization keep a rock coherent until 75% melting (by volume). Even still, our gabbro-xenolith will then completely disintegrate in less than 45 hours supplying the magma with a melt of similar viscosity and density as itself (Figure 4-10), together with 25 wt% xenocrysts of augite and plagioclase (relative to the initial

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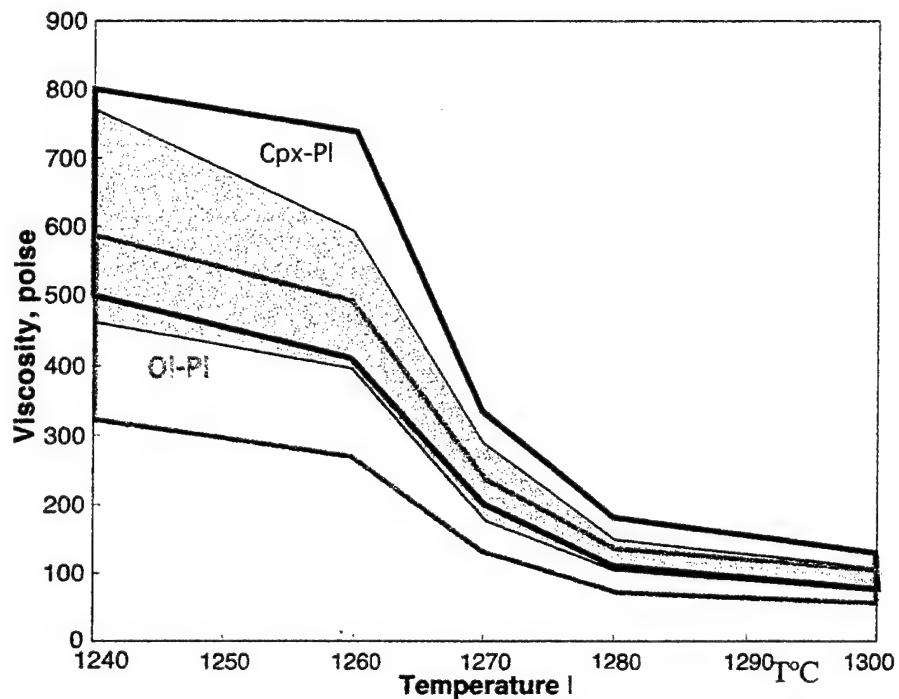


Figure 4-10: The viscosities of the melts produced in our dissolution experiments calculated using Bottinga and Weill (1972). The results are compared to viscosities calculated for lavas from the Southwest Indian Ridge.

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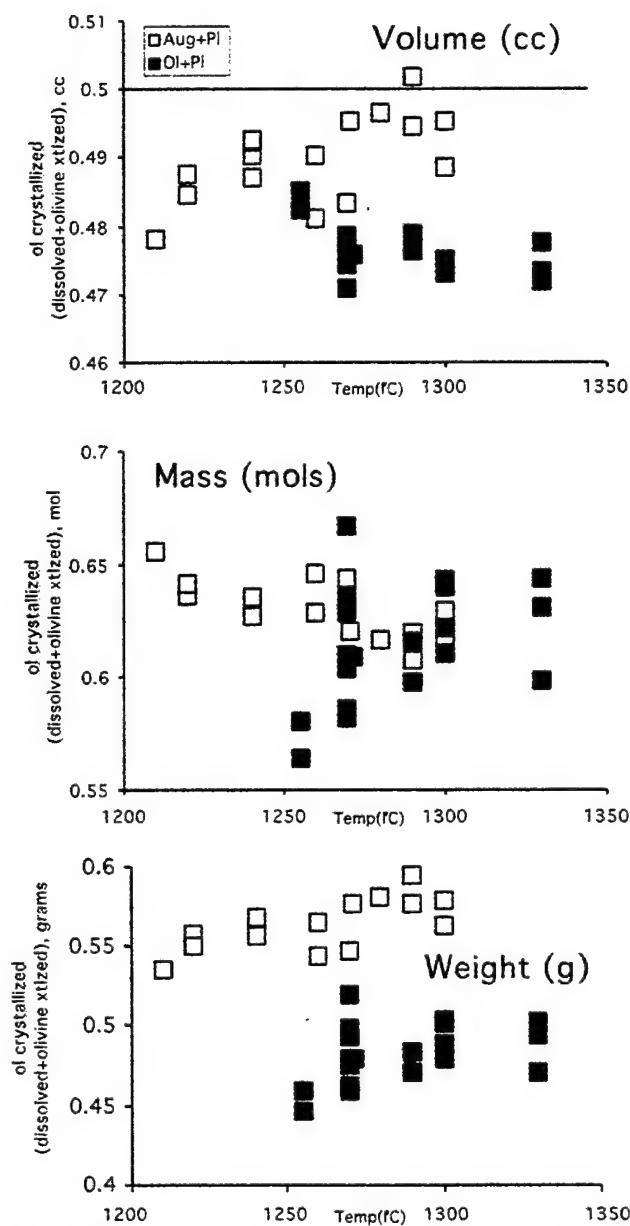


Figure 4-11: The relative fraction of olivine that must be crystallized in order to equal out the heat balance for dissolving gabbro or troctolite. Top: Volume; middle: molar; bottom: grams.

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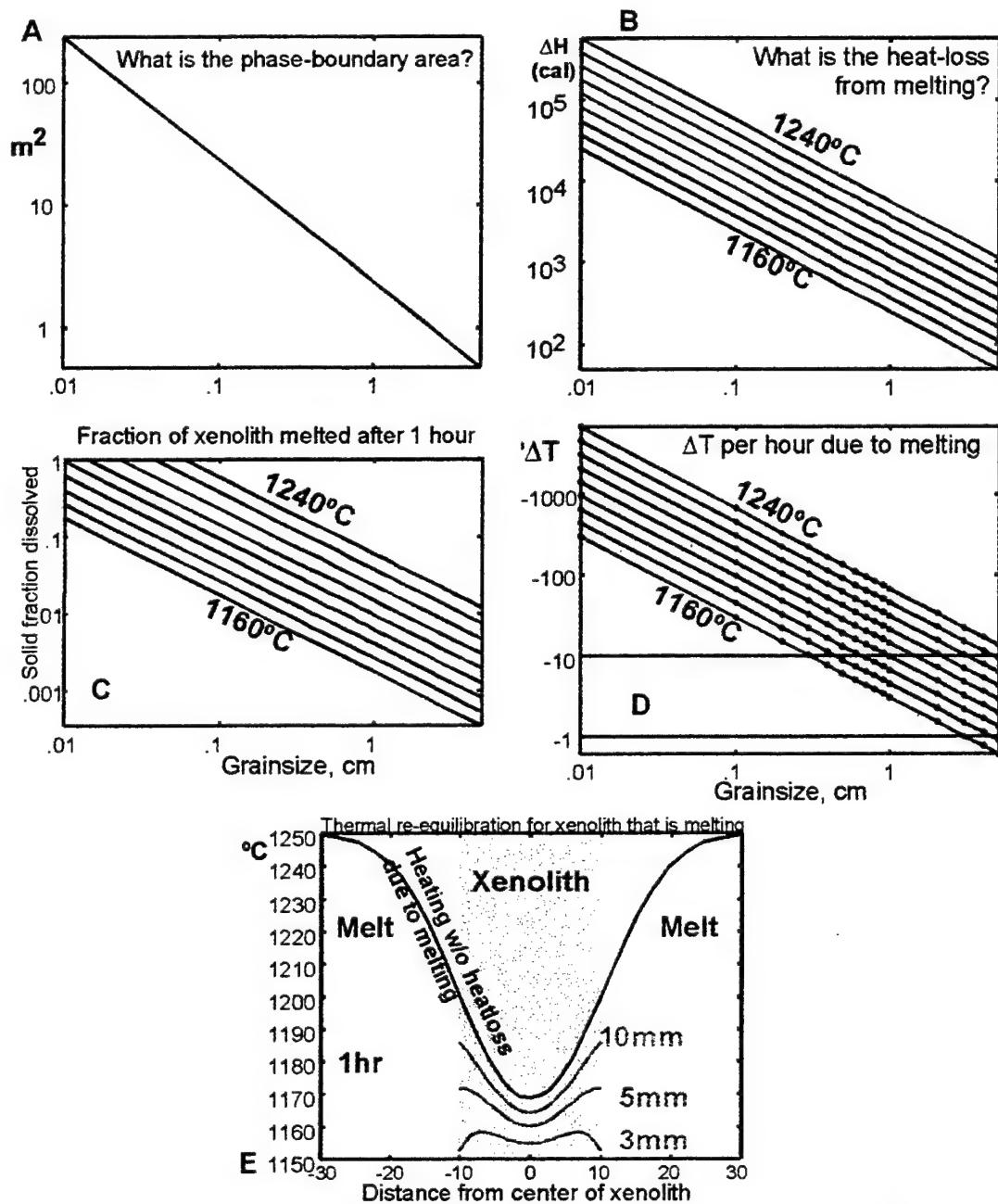


Figure 4-12: A: The geometric area of contact between the minerals of a 20cm bimimetic (50:50) cube for varying grainsizes. B: The heat-loss from melting after one hour at different temperatures and grainsizes. C: The fraction of melting relative to the volume of the original xenolith after one hour. D: the temperature drop calculated per hour due to melting on the boundaries. E: The thermal re-equilibration of the xenolith calculated with and without the heat-loss due to dissolution for varying grain sizes in one hour. The results for grain-sizes less than 3mm yield heat-losses so large that they prevent the heat from penetrating the xenolith.

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xenolith), possibly chained or clustered. A 20 cm xenolith-cube of troctolite will disintegrate in 8 days.

During the melting of the xenolith, the magma itself must compensate for the heat-loss due to the reequilibration and heat of fusion of the minerals. The olivine-saturated melt surrounding the xenolith will therefore start precipitating olivine as the temperature drops in the xenolith's immediate vicinity. In order to be isenthalpic (~no heat lost or gained),  $\Delta H$  has to be 0. The relative amount of olivine that has to be precipitated from the melt in order to accommodate the heat-loss due to the melting in our experiments is shown in Figure 4-11. Again, assuming isenthalpic conditions (i.e. no heat-loss to the surroundings), if the xenolith melts equal molar amounts of plagioclase and augite, the magma can compensate by crystallizing ~59% as many grams of  $\text{Fo}_{85}$ . This is both because the  $\Delta H_f$  for refractory olivine is high, and because plagioclase has a much smaller density than olivine (Table 4-7). In fact, by volume, balancing the heat by crystallization of olivine causes the volume of the melt to increase by 1.2%. A 20 cm xenolith-cube of troctolite cause crystallization of ~58% as many grams of  $\text{Fo}_{85}$  and the melt volume will increase by 0.4%. The volume-change depends mostly on the large density-difference between the mafic minerals and plagioclase.

### ***3.d. The effect of grain-size***

The previous calculations were performed as a simple two-step calculation, and thus ignored the effect of incipient melting upon the heating of the xenolith. It is clear that the grain-size of a given xenolith is important regarding the amount of heat that is lost due to melting on the phase-boundaries. During the melting-process, the solids must be super-heated relative to the solidus of the phase-boundaries (Wood, 1993). Once the minerals at the phase-boundaries are superheated, the phase-boundaries melt, as this is the more stable state. If the phase-boundaries are close to each other, the melting may cool the xenolith back down to the solidus-temperature. If the phase-boundaries are far enough apart, the heat-loss due to melting is not sufficient to bring them back to the saturation-boundary, and the coarser-grained xenolith-interior can be heated and melt. This provides a mechanism where xenoliths can provide super-solidus melts that are linear combinations of the mineral-phases, causing the entire xenolith to disintegrate.

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Figure 4-12A demonstrates the potential phase-boundary area for a 20cm xenolith-cube consisting of two phases (1:1) stacked randomly with varying grain-sizes. Since this resembles potential Hole 735B gabbro-xenoliths, we have used the thermodynamic data for typical Atlantis Bank plagioclase and diopside to model the melting. In basaltic systems, troctolite has more plagioclase than olivine, and therefore has a smaller area of two-phase boundaries. The possibility of superheating of the solids during melting of the phase-boundaries is therefore higher in troctolites than in the gabbros. Smaller grain-sizes provide more volume-loss on the phase-boundaries relative to the xenolith. Therefore, the heat supplied by the magma may not be able to penetrate the xenolith at a rate that is faster than the melting process can cool the rock.

A small amount of heating in a fine-grained rock would cause a large fraction of melting (Figure 4-12C), and quickly cause a substantial heat- and temperature-loss (Figure 4-12B, D). A fine-grained gabbro-xenolith would therefore be prevented from thermally re-equilibrating with the melt internally and would instead experience dissolution along its surface. However, if the grains are  $\geq 0.3\text{cm}$ , then the grain-boundary area is small enough, and the mineral-volume melted in one hour is less than 1% of the original xenolith. This results in less than  $10^\circ\text{C}$  temperature-drop due to the melting.

If we model our xenolith as infinite in Y and Z (i.e a one-dimensional model), this provides a simple calculation for the heating of the xenolith in the melt:

$$T(X, t) = T_c + \frac{1}{2} * (T_0 - T_c) * \left[ \operatorname{erf}\left(\frac{\xi + 1}{2\sqrt{\tau}}\right) - \operatorname{erf}\left(\frac{\xi - 1}{2\sqrt{\tau}}\right) \right],$$
$$\xi \equiv \frac{X}{a}, \quad \tau \equiv \frac{\kappa * t}{a^2},$$

where  $\kappa$  is the thermal diffusivity ( $= 0.008 \text{ cm}^2\text{sec}^{-1}$ ),  $T$  is the temperature at any given place in the X-direction,  $T_0$  is the original temperature at the intrusion time in the xenolith,  $T_c$  is the temperature of the melt,  $a$  is the half-width of the xenolith, and  $t$  is the heating time in seconds. We can then subtract the heat due to melting. In our one-hour model, we found  $\Delta T_{\text{melting}} = 9.495 * 10^{(-22)} * (1/\phi(m)) * e^{(3.87110^{02} * T^\circ\text{C})}$  ( $\phi$ =grain-size) by an exponential fit. Therefore  $T_{\text{total}} = T(X, t) - \Delta T_{\text{melting}}$ . Figure 4-12E shows how this affects the heating of the xenolith. Cooling of the melt next to the xenolith, as shown in Figure 4-12, would cause it to crystallize, though the thermal effect of this on the melt and xenolith is

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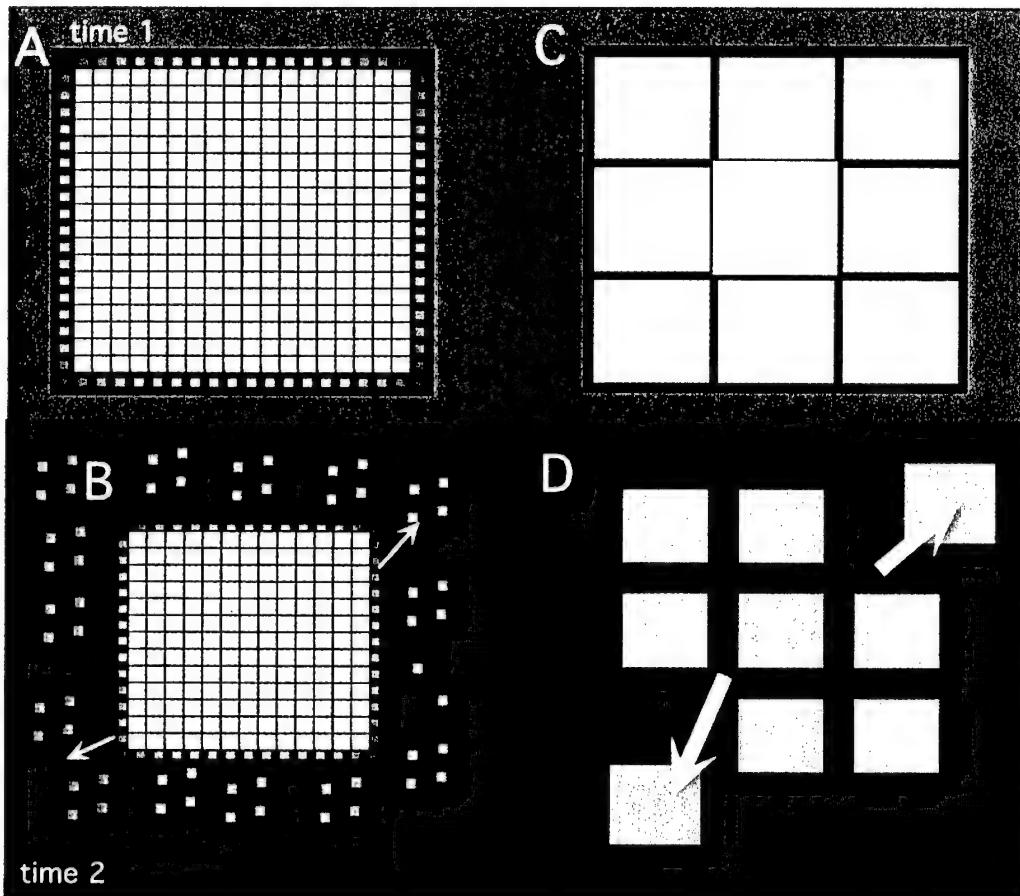


Figure 4-13: The mechanism of dissolution of a xenolith. A: Time 1: A fine-grained xenolith (gridded square) is penetrated by heat (orange) from the surrounding magma (black), causing the minerals to melt in the surroundings. Due to the volume of the melt produced, the dissolution consumes so much heat that none penetrates deep into the xenolith (white= $T < \text{solidus}$ ). B: Time 2: The same xenolith has generated a mush-zone (much like that described by McLeod and Sparks (1998)) around it, leaving the interior unheated. This creates xenocrysts smaller than in the initial xenolith. C: Time 1: A coarse-grained xenolith with the same amount of melting on the mineral-boundaries as in A has much less melt by volume. Thus, the heat of the magma is able to penetrate into the interior of the xenolith. D: Time 2: The outline of the xenolith is larger than that of B at the same amount of melting. The melt is released either by being compaction, or total disaggregation where (large) xenocrysts are released from the xenolith.

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not considered directly here. Again, only xenoliths with grain-sizes  $\geq 3\text{mm}$  will support superheat in the solids relative to the solidus of the phase-boundaries. Xenoliths with grain-sizes  $< 3\text{mm}$  will be effectively insulated by melting of the outer layer of the xenolith, effectively preventing heat-transfer into the xenolith (Figure 4-13). In fact, if the xenolith has grain-sizes  $\sim 0.1\text{mm}$ , the melting overwhelms the surface layer in less than 1 minute. Therefore, fine-grained rocks will disintegrate by melting on the outer surface only, whereas larger-grained rock may melt internally and disintegrate after a large fraction of disequilibrium melting.

Our results may explain why fine-grained diabase is frequently found as unmelted xenoliths in lavas, and why layer 2B (the sheeted dikes) is thermally stable and is hard to melt even with a stable magma-chamber lying right beneath it (Sinton and Detrick, 1992). A seismic layer 2B, composed of typical fine-grained sheeted dikes, will be insulated due to preferential melting along its base in contact with a magma-chamber. This effectively prevents the heat transport upward and the rocks are not easily disaggregated and assimilated into the melt. Likewise, it has been observed that fast ridges have finer-grained gabbros compared to slow ridges (observed by Dick et al., 2002, and quantified using crystal-size distributions by Coogan et al., 2002 (EPR); Meyer and Sapp, 1994 (Hole 735B)). In fact, the slow spreading ridges often have ultraphyric lavas containing large grains ( $>3\text{mm}$ ) of plagioclase (Mével et al., 2002), suggested to be xenocrysts (Coettze, 2001).

### ***3.e. The effect of mineral density***

We found that the melting rate slowed down in our early experiments as the minerals floated apart, and the melting-rates are thus faster if the minerals are kept in close contact. Plagioclase is buoyant ( $\rho \sim 2.7\text{g/cc}$ ) and clinopyroxene is dense ( $\rho \sim 3.2\text{g/cc}$ ) relative to the melt produced within the xenolith, allowing the minerals to separate if there is no rotation of the xenolith. If the olivine precipitating from the melt uses the xenolith to nucleate on, it might become completely capped by the olivine, further accommodating the separation of the minerals, since the release of the minerals to the magma will not happen independently of the others. If the enclosure is complete, pressure might start rising inside the xenolith, and failure of the olivine-layer may cause

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catastrophic mixing with the surrounding magma. The now isolated plagioclase crystals will rise and the clinopyroxene and olivine will sink.

### ***3.f. Geochemical consequences of syntaxis***

At least for major elements, xenoliths that melt out of equilibrium will not significantly equilibrate by solid-state diffusion with the magma until they are disaggregated into it as mono-mineralic clusters or are far enough away from other xenocrysts. The interstitial dissolution-derived melts formed within the xenoliths have similar viscosities as magma with the same temperature, and should therefore be able to readily mix with it following disaggregation (Figure 4-10).

If gabbro-xenoliths are assimilated into the ascending magma in this way, the liquid line of descent (LLD) will be disturbed (Figure 4-14, bottom). Normally, an olivine-saturated magma (a) will precipitate olivine along path 2 until the melt is of composition ‘b’ at the plagioclase-olivine-melt saturation-boundary and proceed towards the olivine-plagioclase-augite-melt saturation-boundary at ‘e’. However, ‘pl-ol’ and ‘cpx-pl’ represent typical melts produced in our melting-experiments. If melt ‘a’ assimilates a gabbro xenolith-suite, ‘cpx-pl’, and accommodates the heat-loss by crystallization of olivine, the melt will start to follow the curved path 1. If the melt assimilates troctolitic xenoliths (pl-ol), it will start to follow curved path 3. ‘e’ is the olivine-plagioclase-augite-melt saturation-boundary at 1atm, and the boundary moves downwards with pressure (white squares representing 2, 4 and 8kbar, respectively). If gabbro-assimilation happens at lower crustal pressures (~2-3kbar), then less troctolite will be initially produced before augite starts crystallizing. The melt that forms the initial olivine-gabbro assemblages at the 4-phase saturation boundary will therefore have higher Mg#’s and Ca#’s than the melt following path 2. Assimilation of troctolite, however, brings the liquid onto the plagioclase-olivine-melt saturation boundary higher up towards “pl-ol”, and causes more, high Mg#, troctolite to form. Since the mafic minerals dissolve without solid-state re-equilibration the Mg# of the consumed solid becomes the Mg# of the xenolith-derived melt causing elevated Mg#’s in the hybrid melts. Figure 4-14 (top) shows that gabbro-assimilation (the arrow with Cpx-pl) creates melts that have a larger clinopyroxene-component compared to the melt that only crystallized olivine (“Ol frac”)

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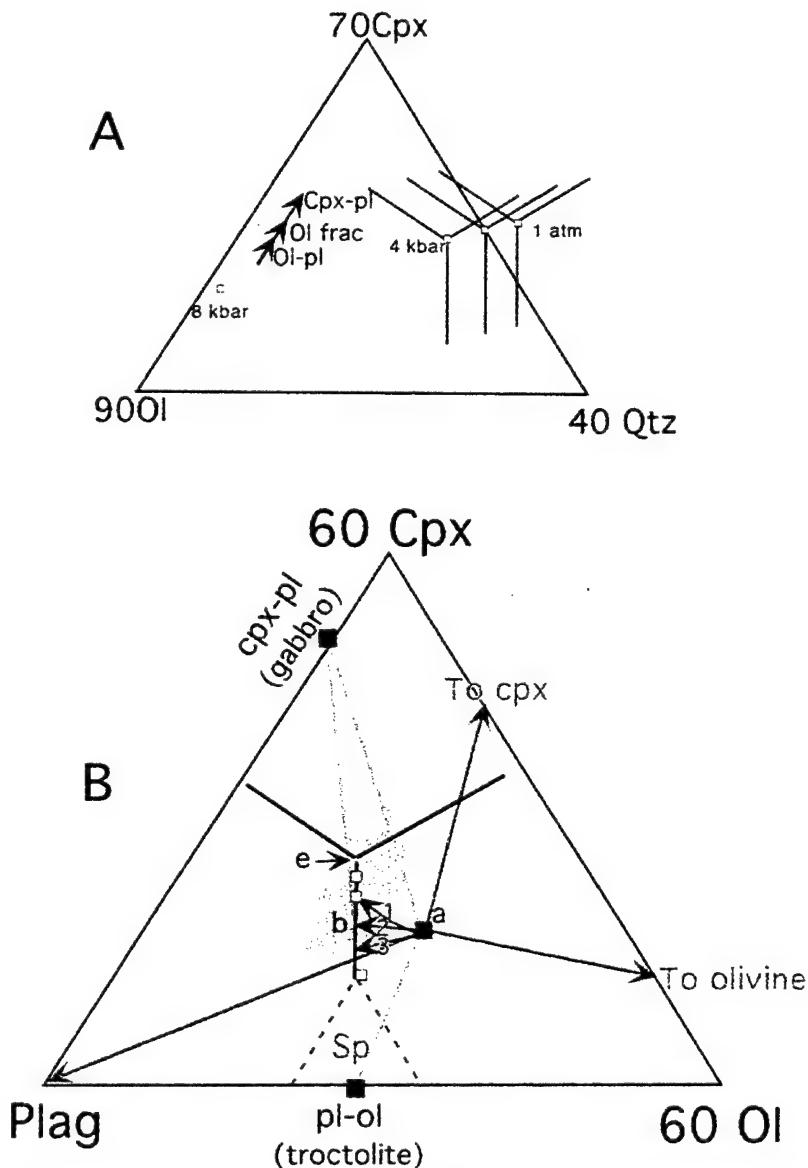


Figure 4-14: The pseudoternary projection (as in Figure 1) of the potential compositions of melts that have assimilated lower crustal material. A: The black arrow shows the LLD of a primary melt that has just crystallized olivine (Ol frac), has in addition assimilated gabbro (Cpx-pl) and has assimilated troctolite (Ol-pl). The experimentally determined saturation-boundaries are indicated (right to left) for 1 atm, 2 kbar, 4 kbar and 8 kbar. B: a=the mantle derived, primary melt (Kinzler and Grove, 1992); pl-ol=troctolite; cpx-pl=gabbro; e=4-phase saturation boundary; grey arrows point to mineral compositions; gray lines show the vectors from the melt to the lower crustal compositions; b=where 2 intersects the plagioclase-olivine cotectic; black lines=cotectics; broken black line is the projected field for spinel; Plag=100% plagioclase; gray squares are basalt glasses from the entire Southwest Indian Ridge. Curve 1 is fractionation of olivine only, curve 2 indicates gabbro-assimilation while fractionating olivine, and curve 3 indicates assimilation of troctolite while fractionating olivine. See text for additional explanation.

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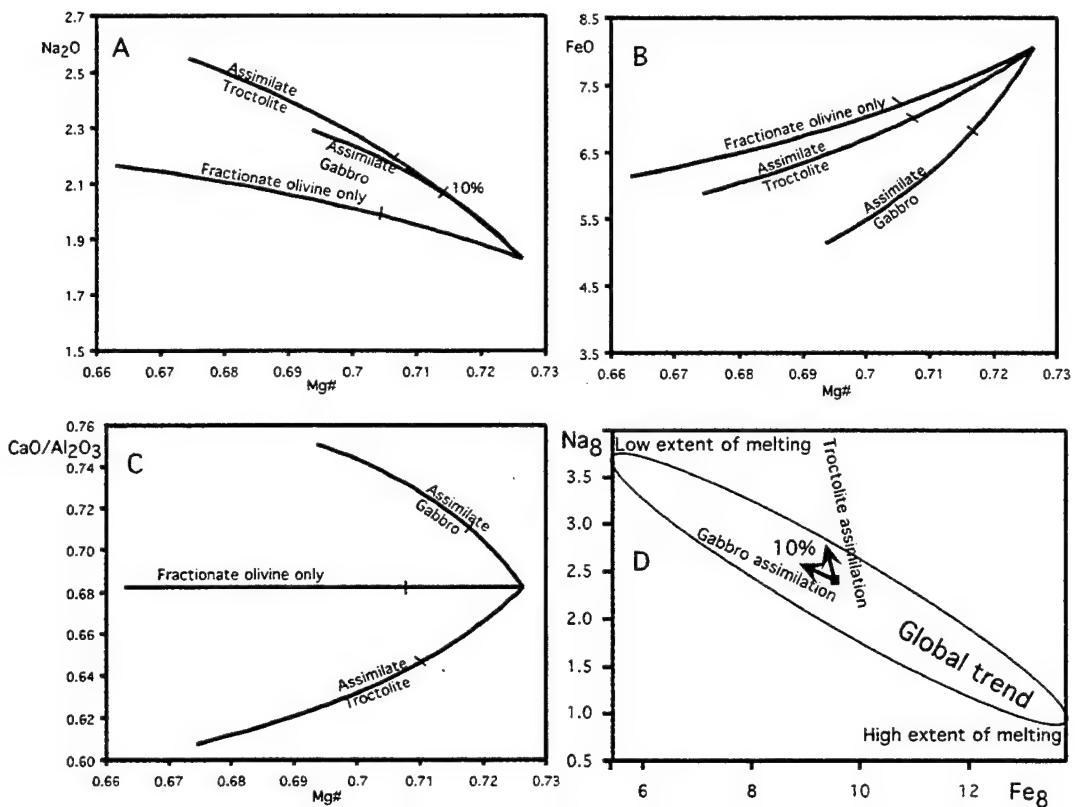


Figure 4-15: A-C: The chemical effect of assimilation of lower crustal material. The heat is balanced by fractionation of olivine. The tag on the lines indicates 10% assimilation or crystallization. Mg#=molar  $Mg/(Mg+Fe)$ . D: The effect of assimilation of 10wt% gabbros or troctolite on a primary melt composition. Klein and Langmuir (1987) suggested that the global trend of the basalt compositions emulate the potential temperature of the mantle. High temperatures cause higher extent of melting than lower temperature melting of identical sources.

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giving the appearance of originating at lower pressure. Troctolite-assimilation, on the other hand, makes the melts have less augite-component and more olivine component, causing them to appear as if they crystallized at a higher pressure, possibly causing the local trends described by Klein and Langmuir (1989).

In summary, melts produced by dissolution of cumulates will have higher Mg#’s than the liquids that originally precipitated the cumulate minerals. If these dissolution-produced liquids are added to a MORB melt, the resulting mixed melt would appear more primitive than the melt that was left from the initial crystallization of the cumulates. Mixing of xenolith-derived magmas into a mantle-derived magma as described above will increase the Na<sub>2</sub>O content of the magma, and decrease the FeO-content (Figure 4-15). However, the CaO/Al<sub>2</sub>O<sub>3</sub> ratio moves in opposite directions, gabbro assimilation increases the ratio, whereas troctolite assimilation decreases it. Therefore, with respect to the “Global trend” of Klein and Langmuir (1987), gabbro assimilation into primitive magma makes the hybrid melt appear as if was produced by a lower degree of melting, whereas assimilation of troctolite moves the melt along the local trend of Klein and Langmuir (1989).

## 4. Conclusions

We have shown that the lower ocean crust melts efficiently when exposed to ascending melts at ocean ridges. We find that average plagioclase-olivine and plagioclase-augite pairs from the lower crust at the Southwest Indian Ridge have solidii at 1195 and 1150°C, and melt fast and out of equilibrium at temperatures significantly above the solidii. Dissolution experiments performed using natural, average olivine – plagioclase and clinopyroxene - plagioclase mineral pairs, under conditions applicable to oceanic magmatism, show dissolution mechanisms similar to those found in other experimental studies that utilized endmember mineral compositions.

The experimental results indicate that melting of oceanic lower crust may produce disequilibrium melt compositions. However, lower crustal material can only be easily and pervasively superheated relative to the solidii of the phase-boundaries if the grain sizes are sufficiently large. Coarse-grained rocks (>3mm) will disintegrate internally upon reheating, whereas finer grained rocks will melt from the outside in. The resulting

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disequilibrium melts are linear mixes of the mineral-compositions. If these melts are allowed to mix into the ascending magma, the new magma will have higher Na<sub>8</sub> and lower Fe<sub>8</sub> than the original emulating the global trend of Klein and Langmuir (1987).

Assimilation of gabbro into magma with only olivine on the liquidus forces the liquid line of descent towards the olivine-plagioclase-augite-melt eutectic, effectively suppressing the formation of troctolite as it brings the liquid onto the plagioclase-olivine-melt saturation boundary closer to the olivine-plagioclase-augite-melt saturation boundary. Thus, crustal assimilation of gabbro will result in the less troctolites in the lower crust, and will lead to the early onset of high-Mg# olivine-gabbros. Troctolite assimilation, on the other hand, leads the liquid line of descent further up on the plagioclase-olivine saturation boundary, producing more, and more primitive-looking troctolites, causing the hybrid melts to define the local trends of Klein and Langmuir (1989).

Caution should be taken when using basalts and gabbros to deduce the extent and pressure of melting in the mantle and crystallization in the lower crust and estimating crustal thickness (e.g. Klein and Langmuir, 1987; McKenzie and Bickle, 1988). Instead, the entire crustal budget should be known, including lavas, dikes and gabbros, so that the effect of assimilation between the ascending magmas and the lower crust can be eliminated.

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## Chapter 5

### Mineral chemistry of gabbros from Atlantis Bank

### outside Hole 735B

#### Abstract

We have analyzed the major element compositions of the magmatic minerals of 99 gabbros from on and around Atlantis Bank, South West Indian Ridge, including the top carapace of the bank, the Western Wall, and the slopes to the north and the south, covering an area of almost 660-km<sup>2</sup>. Samples from the northern area have the most primitive magmatic minerals, extending to ferrogabbros with evolved mineral-compositions. The southern rocks are all evolved, including layered olivine-gabbros sampled at gabbro-peridotite contacts. Gabbros from the Western Wall and the top of the bank have intermediate compositions. Despite the large variations in mineral chemistry in our samples, the primitive and evolved ends of the ODP Hole 735B spectrum are not present in our data. No plagioclase with anorthite contents higher than An<sub>71.5</sub> has been recovered, anorthite higher than An<sub>64</sub> is rare, and no magmatic plagioclase has been found with anorthite below An<sub>32</sub>. Likewise, augite compositions generally range from 60 to 86.3 molar Mg/(Fe+Mg)\*100 or Mg#, where Hole 735B augite compositions range from 54 to 89. Although we have sampled the area extensively, we find that, like the gabbros from Hole 735B, the surface samples represent melts that have crystallized 50-90% relative to mantle-derived compositions. As discussed in Chapter 3, the overall variation of the gabbros is broadly consistent with control of the liquid-line of descent by fractional crystallization (modified by post-cumulus processes). If the lower crust was accreted by fractional crystallization, large amounts of troctolite should exist somewhere in the lowermost crust or intruded into residual mantle at depths below what is exposed on the sea-floor or drilled in Hole 735B.

As in Hole 735B, augite is more magnesian than expected for any coexisting plagioclase anorthite content, and plagioclase and augite are frequently reversely zoned. TiO<sub>2</sub> contents of augite vary inversely with Mg# at all levels of evolution, even after

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magnetite and ilmenite crystallized, suggesting extensive reaction, dissolution and reprecipitation within the cumulate pile. A simple fractional-crystallization model would therefore not be suitable. We propose an alternative model where some originally primitive cumulates may have been re-homogenized and transformed by later, interstitial melts.

### 1 Introduction

The nature of accretion of the lower crust at spreading-ridges has mainly been determined by the study of ophiolites as analogues. The origin of many ophiolites is unfortunately uncertain (e.g. Miyashiro, 1979), and the exact crystallization conditions and primary melt compositions may be different at modern ocean ridges.

Tectonic windows make the investigation of the lower ocean crust in the modern oceans possible, and current technology makes them accessible. Atlantis Bank is a well-known example of an paleo-inside corner high found at the ultraslow spreading Southwest Indian Ridge. It is located 11-Ma off axis, south of the present-day rift valleys on the eastern side of Atlantis II Fracture Zone (Dick *et al.*, 1987). Spreading rate is asymmetric, 8.5-mm/y to the South and 5.5-mm/y to the North. The 700-meter deep platform and its surrounding slopes represent the largest outcrop of gabbroic rocks known in the modern oceans. The lower crustal section is exposed along a low-angle detachment-fault surface that rooted in the dike/gabbro transition, or possibly in a shallow magma-chamber or crustal mush-zone (Dick *et al.*, 2001). The exhumation exposed lower crustal rocks for 4-My of seafloor spreading at the ridge, and we have mapped gabbroic outcrops over 660-km<sup>2</sup> (Figure 5-1) extending 39-km in the plate spreading direction (Dick *et al.*, 2001). ODP Hole 735B provides a 1500 meter deep vertical section, contributing to a 3D perspective of the platform. In addition, Hole 1105A was drilled about 1.3-km northeast of Hole 735B and penetrated 158-m (Shipboard Scientific Party, 1999).

The Western Wall of Atlantis Bank faces the Atlantis II Transform Valley. Altered mantle rocks are exposed on the trench wall, more than 2000-m below the bottom depth Hole 735B to the northwest and < 500-m to the southwest and at all depths

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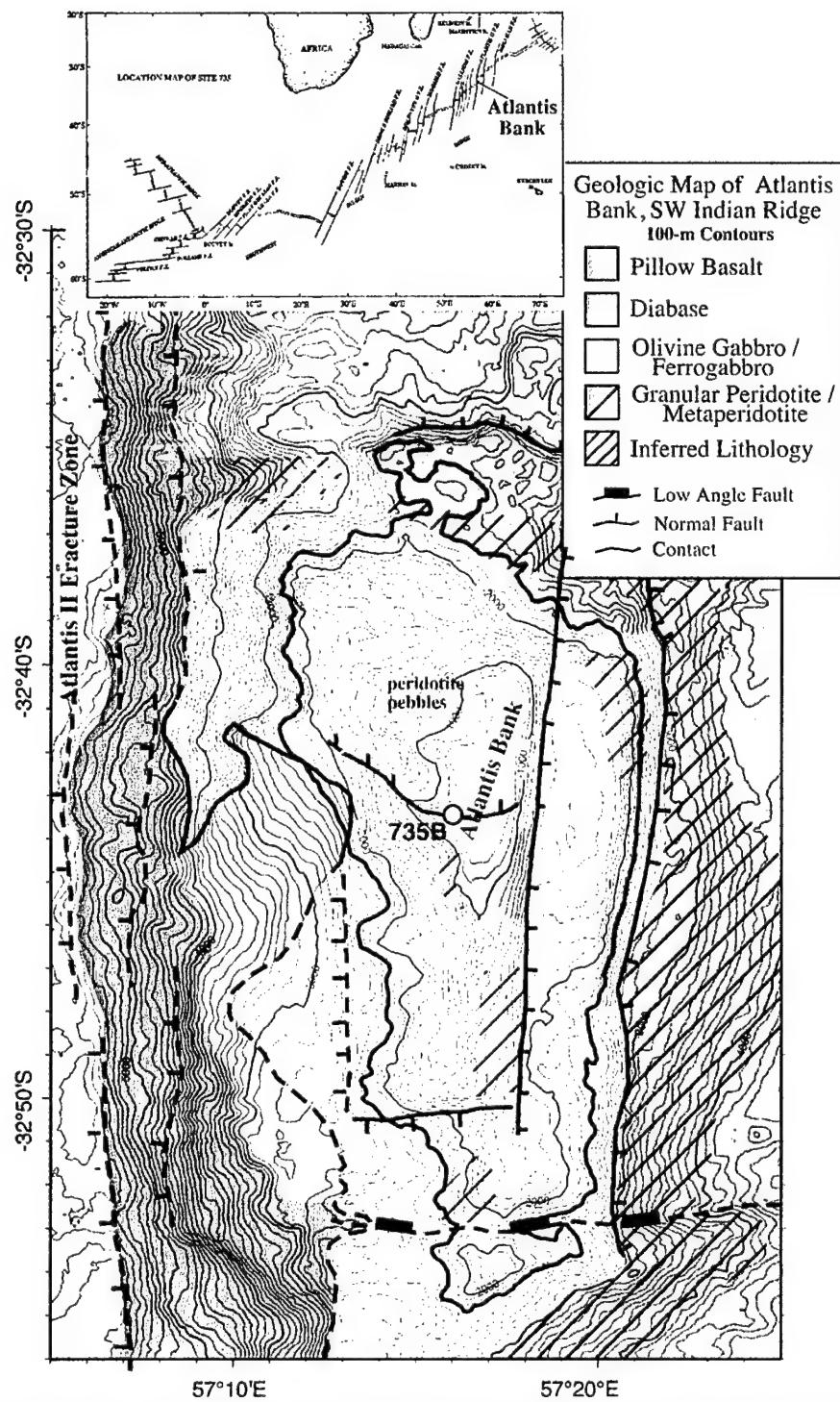


Figure 5-1: The geology of Atlantis Bank (Dick et al., in prep.). A thick contour-line outlines the bottom of Hole 735B.

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shallower than the Hole along the transform wall to the west (Fig. 5-1). The Northern Wall faces the termination of the detachment with pillow basalts exposed to the north. The Eastern wall has two down-faulted blocks that terminate in a volcanic terrain to the east. The detachment break-away zone to the south has not yet been found, although it is inferred morphologically from apparent constructional volcanic edifices that replace the gabbro-dike terrain south of 32° 52" S.

Bloomer et al. (1991), Dick et al., (2000, 2001) and Coogan et al., (2001) observed that the expected primitive section of cumulates is missing in Hole 735B. Coogan et al (op. cit.) argued that Hole 735B was drilled closer to Atlantis II Fracture Zone than to the magmatic segment center (AN-1) and the authors therefore proposed that the missing cumulates were either deeper in the section towards the crust-mantle boundary, or towards the eastern wall of the bank if not in the mantle itself. They also suggested that the evolved nature of the lower crustal rock is due to incorporation of large amounts of trapped melt. Serpentinites west of Hole 735B represent gouge intruded laterally along the detachment fault, and show that these locally overlie the gabbro massifs west of Hole 735B (Fig. 5-1) (Dick pers. comm.). Dick et al (1987) and Hosford et al (2003) showed that Atlantis Bank is the center of a small sub ridge-segment and Muller et al., (1997) found the seismic crustal thickness is  $5\pm1$ -km close to Hole 735B, thinning towards the sides of the platform. An extensive sampling-survey was performed on and around Atlantis Bank to allow us to constrain the regional distribution of lithologies and the magmatic mineral-compositions in both time and space.

The results of major element mineral analysis of gabbro samples from James Clark Ross 31 ("JR-31") Bridge ("BR)- and British Geological Survey ("BGS")-cores and dredges, and SHINKAI 6500 dive samples from the MODE 98 and ABCDE expeditions to Atlantis Bank are reported here. The purpose of this paper is to describe the 3-D distribution of mineral-compositions on and around Atlantis Bank and to find the primitive, missing cumulates, or explain why they are missing.

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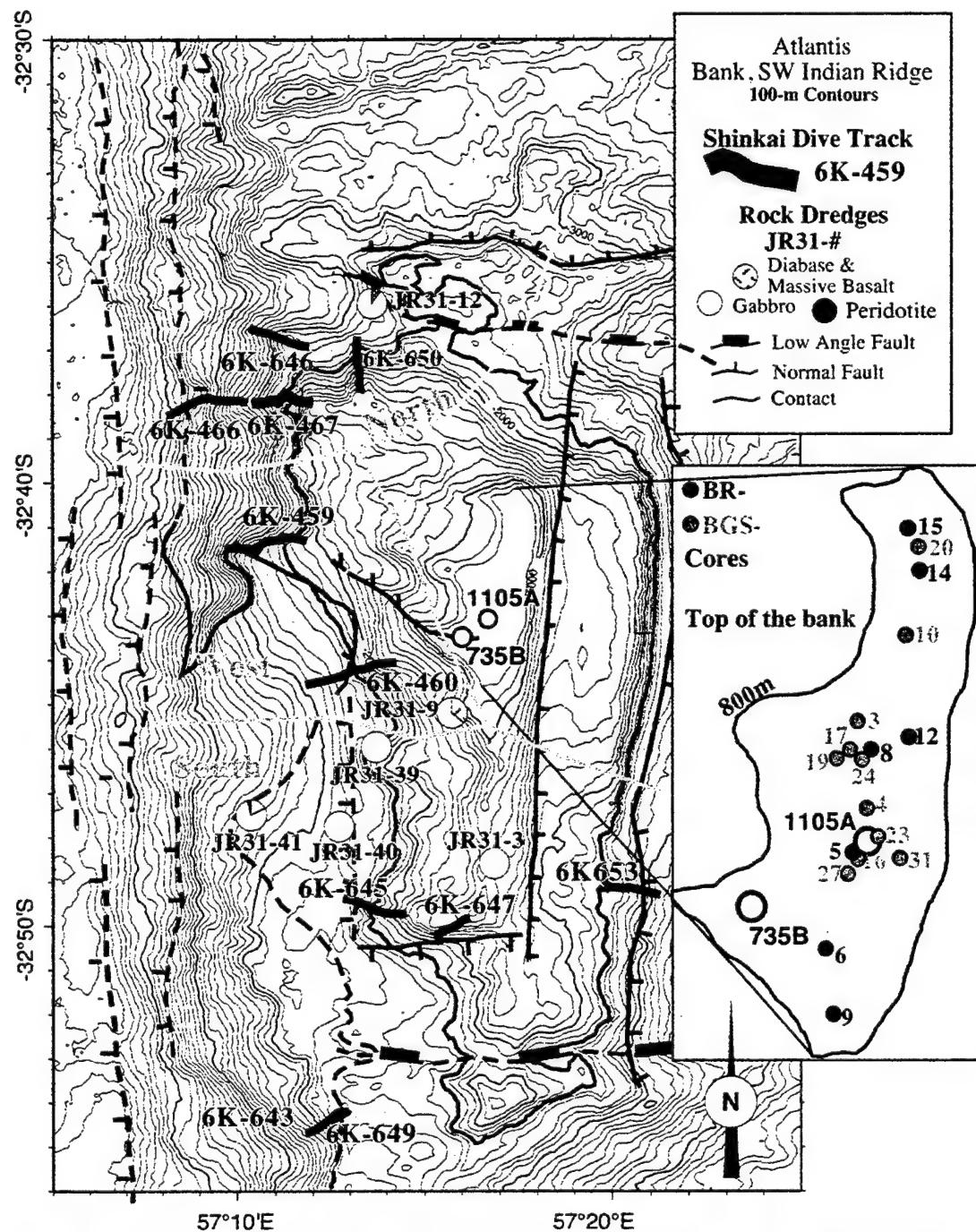


Figure 5-2: Sample locations for this study. The thick gray contour line indicates the termination of Hole 735B, and the thinner gray lines show the geographic sample divisions discussed in the text.

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### 2 Materials

The rocks from this study are sampled from the top and sides of Atlantis Bank. Site collections were chosen based on lack of alteration, number of useful samples in each collection and density of sampling in a given area. We attempted to analyze magmatic plagioclase and clinopyroxene ( $\pm$  olivine and orthopyroxene) for each sample, and if one of these two were completely altered to metamorphic minerals, the sample was not chosen for analysis. Some collections were picked even with a single sample, if it was the only one in the given area. Likewise, somewhat altered samples were not chosen in areas with a lot of coverage. The locations of the dives, dredges, and cores from which we chose samples are shown in Figure 5-2.

The JR-31 expedition sampled a number of shallow (1-2-m) rock drill-cores on the top of the bank, and 52 dredges around its sides. Six of these dredges (JR31-3, 9, -12, -39, -40, and -41) were chosen for this study as they contained several samples of fresh gabbro. Seven of the oriented BR-cores and 10 non-oriented BGS-cores from the 22-km<sup>2</sup> platform at the top of the bank were also selected for analysis.

The MODE 98 and ABCDE expeditions used SHINKAI 6500 to sample traverses up the sides of the bank. Mode 98 sampled the west wall of the bank, and dives 6K-459, 6K-460, 6K-466, and 6K-467 were selected for analysis. One sample, 6K-467-11, is a dike that contains plagioclase and clinopyroxene. Cruise ABCDE sampled all sides of the bank and seven dives were selected (6K-643, -645, -646, -647, -649, -650, and -653). The samples are listed in Table 5-1 and collection-locations are shown in Figure 5-2.

We have divided the bank into four study areas. The platform area ("Top of the bank") includes all the JR31-BR- and BGS-drill cores, and Hole 735B. The northern area ("North") includes JR31-12, 6K-466/467, 6K-646, and 6K-650, the western area ("West") JR31-9, and 6K-459 and 6K-460, and the southern area ("South") consists of JR31-3, -39, -40, and -41 and 6K-643, -645, -647, -649, and -653.

### 3 Methods

The magmatic minerals, olivine, clinopyroxene, orthopyroxene and plagioclase, were analyzed with the MIT 4 and 5-spectrometer JEOL 733 Superprobes, using 15kV accelerating voltage and 10nA beam current. When possible, continuous 10- $\mu\text{m}$  10-spot traverses in the core and the rim of adjacent mineral-grains of plagioclase and clinopyroxene were analyzed. Continuous 10- $\mu\text{m}$  6-spot traverses in the core and rim of orthopyroxene were analyzed, whereas olivines were analyzed in six spots throughout the crystal. A 10 nA beam current, 15 kV acceleration potential, and a 10- $\mu\text{m}$  beam were used for the analyses. Counting times were 10-40 seconds, depending on the elements. Standards used were DJ35, Marjalotti Olivine, Synthetic Fayalite, Alp7 orthopyroxene and Lake County Labradorite, and the major elements have a standard deviation of less than 2%. Data were reduced with the CITZAF correction package using the atomic number correction of Duncumb and Reed, the absorption coefficients of Heinrich and the fluorescence correction of Reed (Armstrong, 1995). The results of the mineral analyses from this study are listed in Table 5-2.

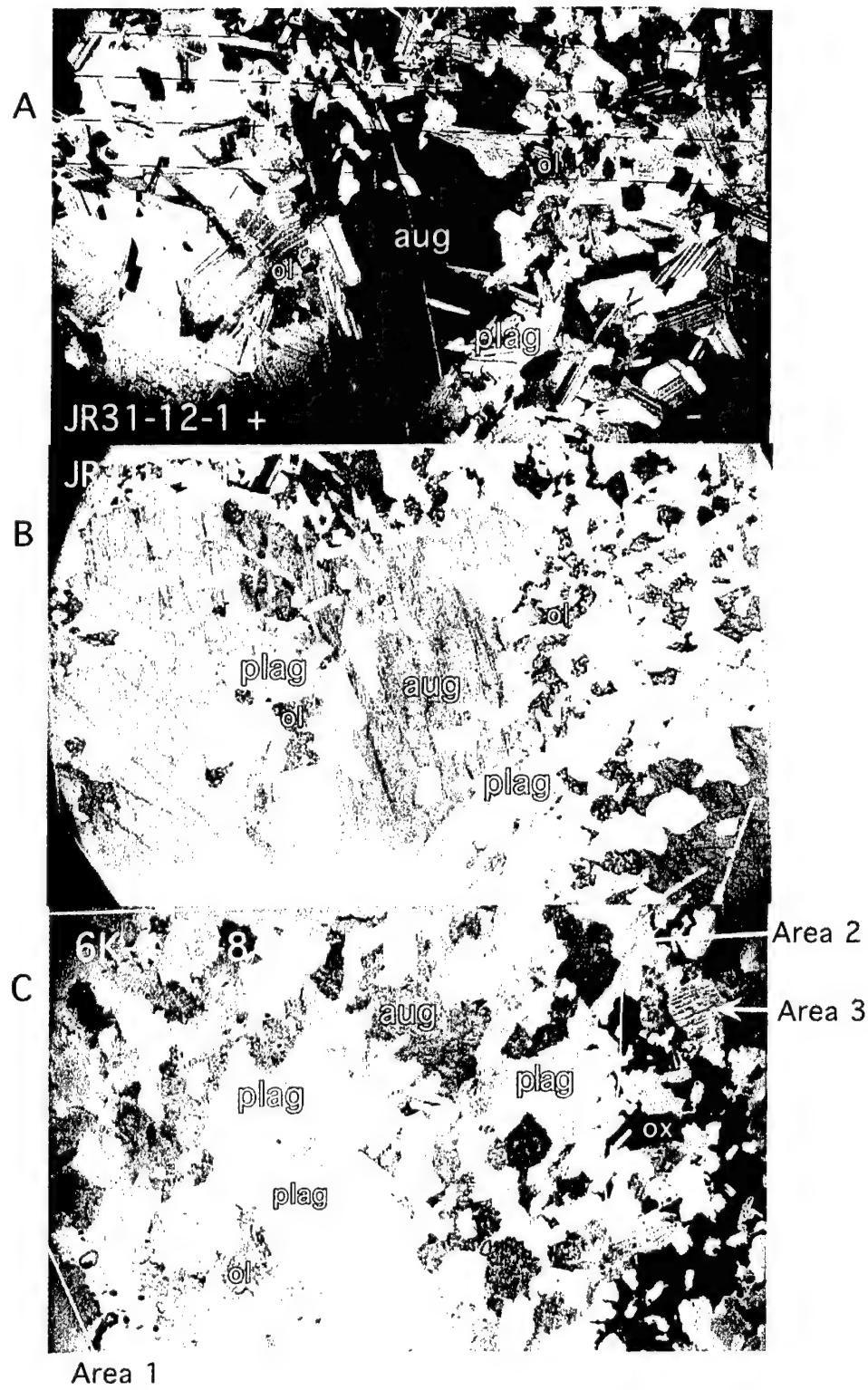
### 4 Results

#### 4.1 General mineral distribution and petrography

Plagioclase and clinopyroxene are ubiquitous throughout the sample-suite. The grain-sizes range from medium grained to very coarse-grained, and fine-grained micro-gabbros are only rarely present. In undeformed samples, fresh plagioclase appears as randomly oriented euhedral laths, surrounded by subophitic clinopyroxene (Fig. 5-3a,b). Olivine appears as anhedral grains, with a characteristic cracking-pattern, whereas orthopyroxenes are granular, or appear as rims around olivine (Fig. 5-3e). Oxide-minerals are common, and are usually interstitial to the other minerals (Fig. 5-3c). Ilmenite and magnetite occur in individual grains or as trellis-twin exsolution-lamellae after ulvöspinel.

High temperature deformation has caused kink-bands in plagioclase with anhedral and granular augite and plagioclase (Fig. 5-3d). Some samples have saussuritized

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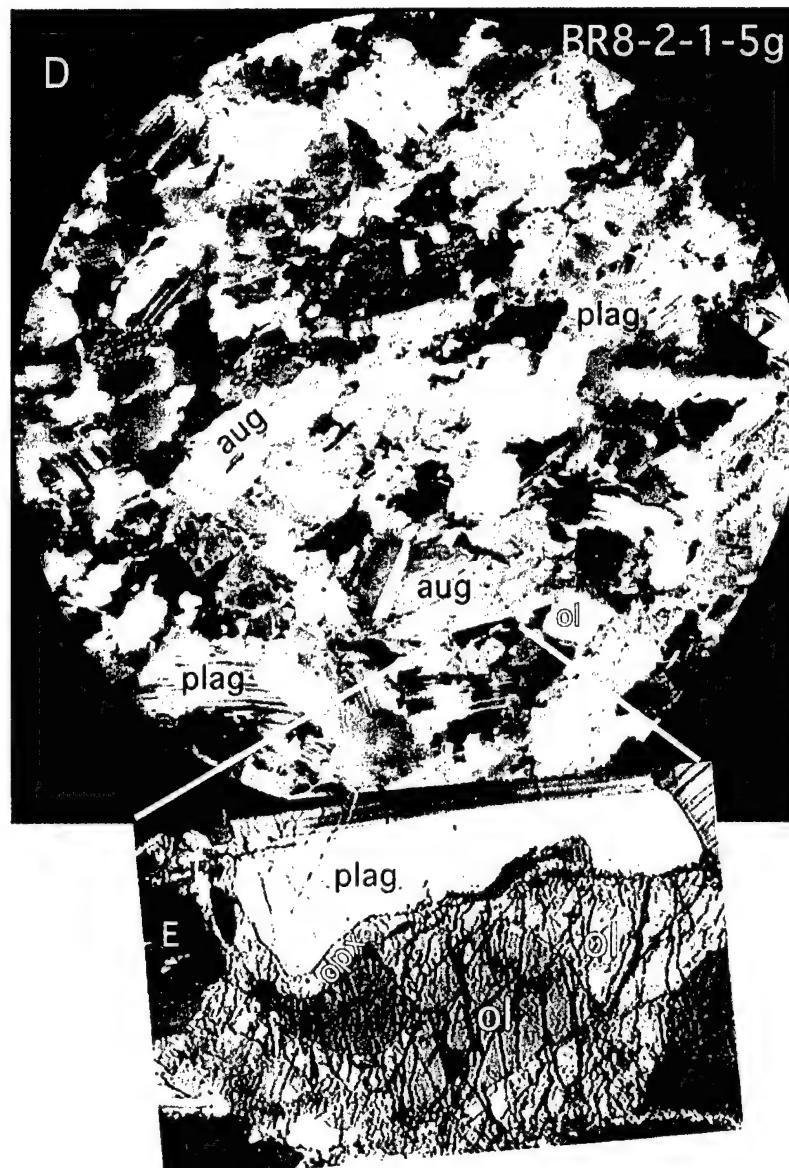


Figure 5-3: **A:** Sample # JR31-12-1: Sub-ophitic augite surrounds plagioclase-laths with interstitial olivine. Cross-polarized light. Plag=plagioclase, aug=augite, ol=olivine. Field of view is 3" from left to right. **B:** Same slide in parallel light. **C:** Sample 6K-466-8, field of view is 1" across. An olivine-gabbronorite-oxide-gabbronorite contact, as seen by the extensive interstitial oxide-crystals to the right. The boundaries between the lithologies are indicated by the white line. The picture is taken using a combination of reflected and parallel light to distinguish the different minerals. Three areas (indicated) have been analyzed. **D:** Sample # BR8-2-1-5g: The sample has been partially deformed at high temperatures, as seen in the kink-banded plagioclase and rounded augite. Cross-polarized light, the slide is 1" across. **E:** Detail of C; a thin orthopyroxene-grain lies between an olivine grain (with the characteristic cracking-pattern) and plagioclase

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plagioclase and clinopyroxene altered to green amphibole; these are signs of greenschist-facies alteration. These samples were therefore not chosen for mineral-analysis. Olivine is altered to iddingsite under the same conditions, and fresh material was not available representing most of the rocks that originally containing olivine, although we have found olivines from the entire compositional spectrum of gabbros. Orthopyroxene is also more susceptible to this low-grade alteration than augite, but small mineral-cores are often available for analysis.

Point-counting of thin-sections in order to obtain mineral norms has not been performed, as the rocks are generally so coarse-grained that even a large (2"x3") thin-sections does not adequately represent the mode of the rock (Fig. 5-3a). However, Coogan et al, 2000 analyzed the whole-rock compositions of the rocks around Atlantis Bank sampled during the JR-31 expedition. We have calculated norms based on the methods outlined in Grove et al. (1992) (Fig. 5-4), and compared them to MORB from the ridge to the north and transform walls, and to the average of Hole 735B (Dick, et al., 2000). If the gabbros were formed by pure fractional crystallization, the whole-rock compositional norms should start as pure olivine (dunite), continue along the plagioclase-olivine cotectic and end at the 4-phase saturation-boundaries outlined in Figure 5-4. The whole-rock compositions clearly span a wider range than this, with large excesses in plagioclase and augite, suggesting mechanical separation of the minerals or reactive modification of mineral proportions after cotectic crystallization. The presence of non-cotectic proportions of olivine, plagioclase and clinopyroxene is even more pronounced in the samples from Hole 735B. The Hole 735B samples are, however, very small in volume, so they may not be as representative for the bulk-rock composition as the grain-size is coarse.

### ***Mineral analyses***

#### **4.1.1 Plagioclase**

The average plagioclase has 49.5 % An, the median is 49.8%, and the highest is 71.5 % An. Only a few plagioclase grains have An below 32.3%, and those grains are possibly

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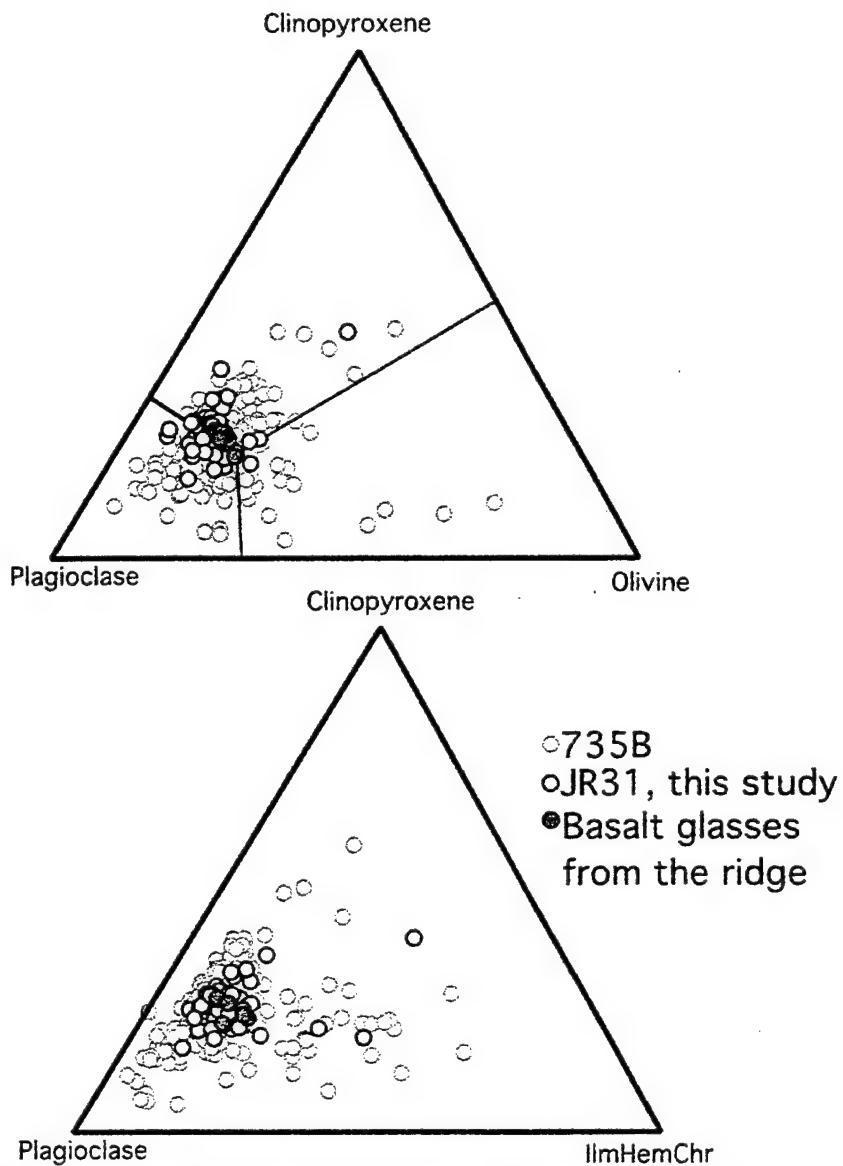


Figure 5-4: The mineral modes of Atlantis Bank gabbros and average Hole 735B compositions calculated from whole rock data using Grove *et al.*, (1992). The cotectic lines in the uppermost figure represents 1atm olivine-plagioclase-augite-melt saturation-boundary based on a basalt-glass from the spreading ridge, thus suggesting the cumulative nature of the gabbroic rocks.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

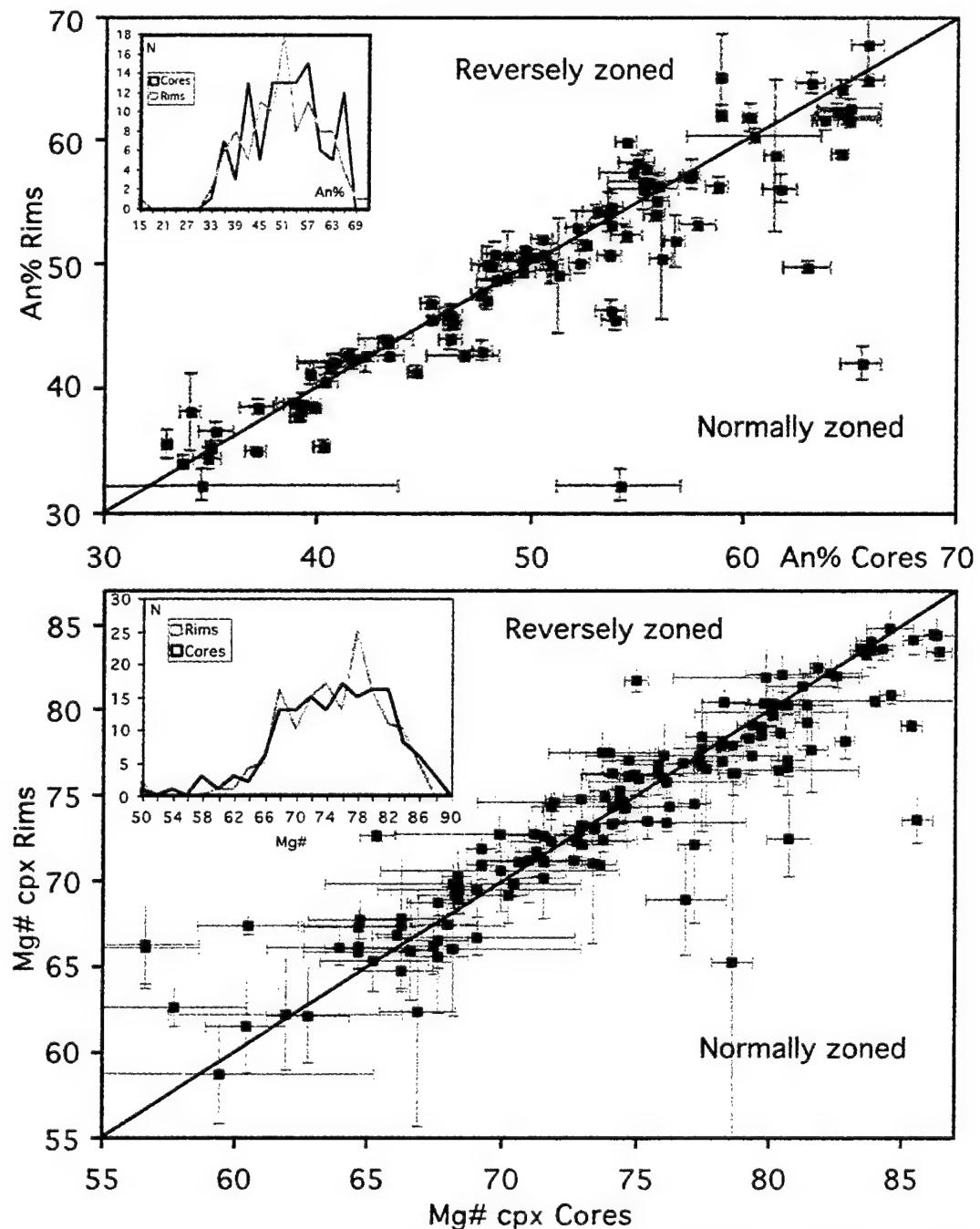


Figure 5-5A: Cores and rims of clinopyroxene and plagioclase from this study. Note how the distribution between cores and rims are indistinguishable (insert). Line indicate 1:1.

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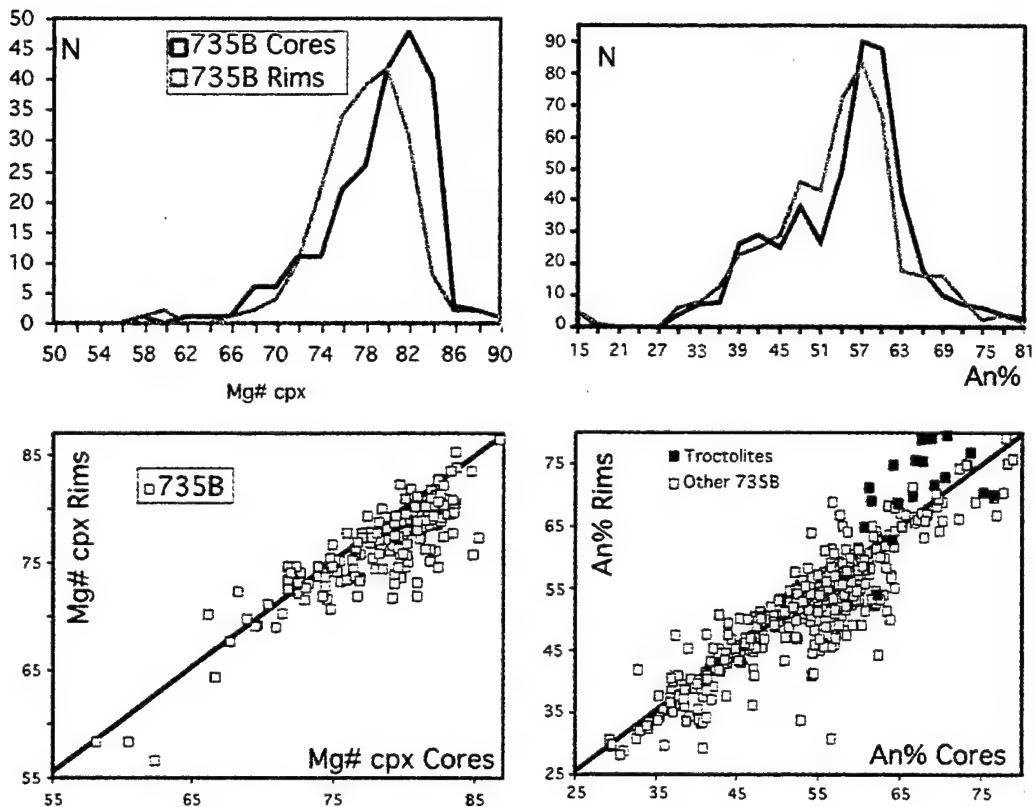


Figure 5-5B: The distribution of cores and rims of plagioclase and clinopyroxene from Hole 735B. The data are from Dick *et al.*, (2001) and Angeloni and Dick (1990) (troctolite-dikes from the upper 550m of the hole). The crossing line indicate 1:1. The core-rim information is not generally available for the upper 550 m of Hole 735B.

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altered. The orthoclase contents increase slightly as An decreases, although it rarely is higher than 1-mol%.

Zoning is common (Figure 5-5a). Plagioclase is reversely zoned in 50 out of 109 cases (i.e. rims are more An-rich than the cores), with respect to An. The core to rim difference in An contents is  $>1\sigma$  for microprobe uncertainty (and mineral inhomogeneity) in 23 of the 50 reversely zoned samples. 59 plagioclase grains are normally zoned, and the core to rim change in An content is  $> 1\sigma$  in 37 of the 59 normally zoned samples. Low anorthite plagioclase cores are also often overgrown by higher anorthite plagioclase with a different crystal orientation, suggesting that the low An-core acted as a nucleation site for the larger crystal (Fig. 5-7).

### 4.1.2 Clinopyroxene

The augites analyzed in this study ranges from Mg# 48.1 to 86.3 %, with an average of 74.5 % and a median of 74.9 % (Figure 5-5A and 5-6). Rims of the clinopyroxene are defined as the contacts with other minerals, whereas the cores are the respective areas furthest away from this contact. Cores and rims were analyzed in 158 clinopyroxenes (Figure 5-5A). Of these, 78 have reverse zoning ( $24 > 1\sigma$ ), whereas 80 are normally zoned ( $27 > 1\sigma$ ).

The minor elements also vary significantly. Figure 5-7A shows two large clinopyroxene crystals enclosing many large plagioclase laths. The cores and rims of these grains show a wide range in compositions. Both Cr<sub>2</sub>O<sub>3</sub> and TiO<sub>2</sub> show a range comparable with the range of clinopyroxene compositions analyzed in the entire lower half of Hole 735B (Dick et al., 2002). This large single-sample range for the minor elements is common throughout the suite (Figure 5-7B).

### 4.1.3 Orthopyroxene

Low-Ca pyroxene appears as orthopyroxene. The grains are susceptible to sea-floor alteration, and thus fewer have been analyzed than in the more protected and unaltered Hole 735B. Orthopyroxene compositions range from Mg# 56 to 78%, with an average of 67% (Figure 5-6). Chemical zoning is not pronounced in the orthopyroxenes.

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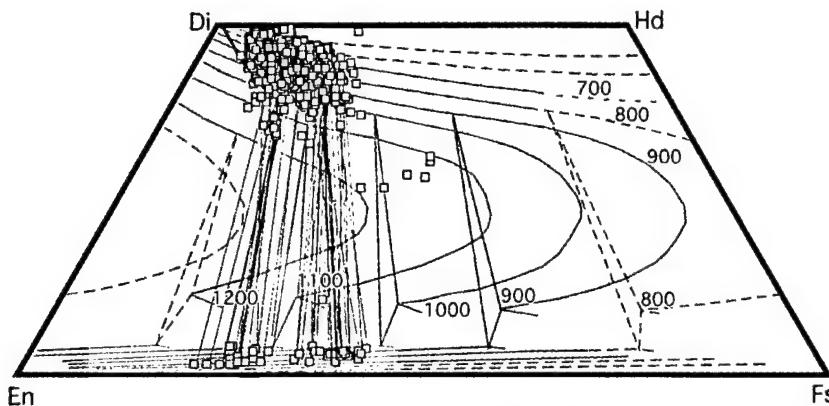


Figure 5-6: The tetrahedral components of the pyroxenes from this study calculated from Lindsley (1983). A thin black line connects coexisting minerals.



Figure 5-7: Sample #JR31-12-6. Each mineral analysis-point is shown with a white point and the analysis is shown with it. The key is to the upper left. The cross-sections through two large plagioclase crystals are shown. Two large augite-grains surround laths of plagioclase (their separation is marked by a yellow line). There is pronounced zoning in the minerals, and the Mg#, TiO<sub>2</sub> and Cr<sub>2</sub>O<sub>3</sub> span almost the entire range in composition seen otherwise in the suite (insert). Notice how the plagioclase-grains have near constant compositions throughout the thin-section. Field of view is ~4cm from left to right.

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### 4.1.4 Olivine

Seafloor alteration has also reduced the number of magmatic olivines available to be analyzed in this study. The forsterite content ranges from 46 to 80 %, with an average of 67%, and a median of 69%. No zoning is observed, and grains within a sample are generally uniform in composition.

## 4.2 *Geographical Variability Around Atlantis Bank*

Four areas on and around Atlantis Bank appear to have distinctive distributions of magmatic minerals. Figure 5-8 (a,b,c) shows the anorthite content of plagioclase vs. the Mg# of coexisting augite, and orthopyroxene, and the forsterite content of olivine respectively. The northern area clearly has the widest range in compositions, although they never reach the high Mg#'s in clinopyroxene and An contents in plagioclase of the primitive cross-cutting troctolites in the upper levels of Hole 735B. In fact, one of the more primitive analyses among the surface samples is from a dike crosscutting the Dive 6K-647 section. Dive 6K-466, directly below 6K-467, sampled gabbros just above a contact with the residual peridotites, and these rocks have plagioclase as low as An<sub>35</sub>. The samples from the top of the bank exhibit more intermediate compositions, where few plagioclases are higher than An<sub>61</sub> or lower than An<sub>40</sub>. Some of the individual cores are unusual, as shown by BGS-24, which has a composition trend opposite to the other samples (Fig. 5-8a). The Western Wall of the bank has an even narrower trend, and resembles the distribution for the lower 500 meters of Hole 735B. This is interesting, as traverse 6K-460 was sampled close to Hole 735B, at a water depth just below the total depth from sea-surface to the bottom of the Hole. The southern part of the bank is also evolved geochemically. The layered olivine gabbros from the crust-mantle transition in this area (Dive 6K-649 and 653) are surprisingly fractionated as well with An<sub>48-60</sub> and augite Mg#'s from 70 to 82 mol%. The sample areas do follow the composition fields for Hole 735B. By contrast, the composition field for Hole 1105A gabbros (Thy, 2003) is offset to consistently more anorthitic plagioclase (or less Mg-rich augite) compositions than Hole 735B. A comparison of the whole-rock data from Hole 735B (Natland and

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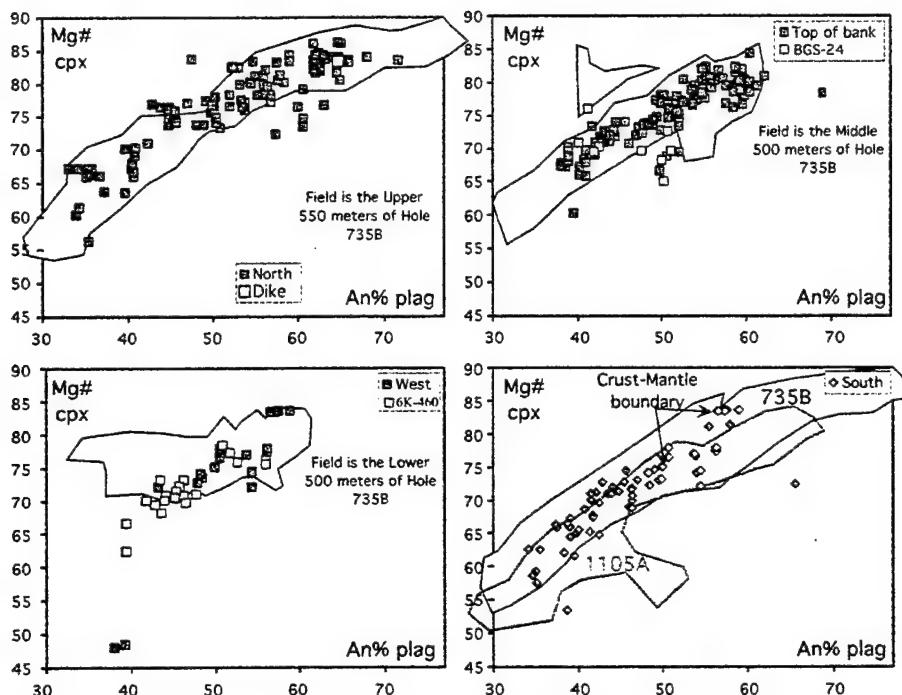


Figure 5-8: A: The composition of coexisting plagioclase (plag) and augite (cpx) for each of the areas from around Atlantis Bank. The fields represent the distribution of Hole 735B gabbros (Dick *et al.*, 2001) and the lower right diagram also shows Hole 1105A field (Thy *et al.*, 2003).

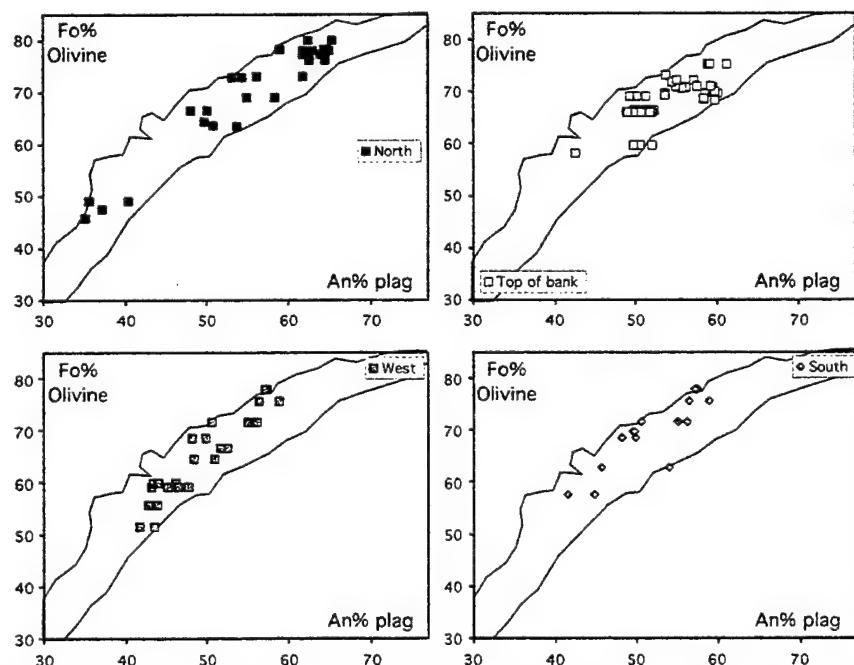


Figure 5-8: B: The composition of coexisting plagioclase (plag) and olivine (Fo%). The fields represent Hole 735B olivine-gabbros.

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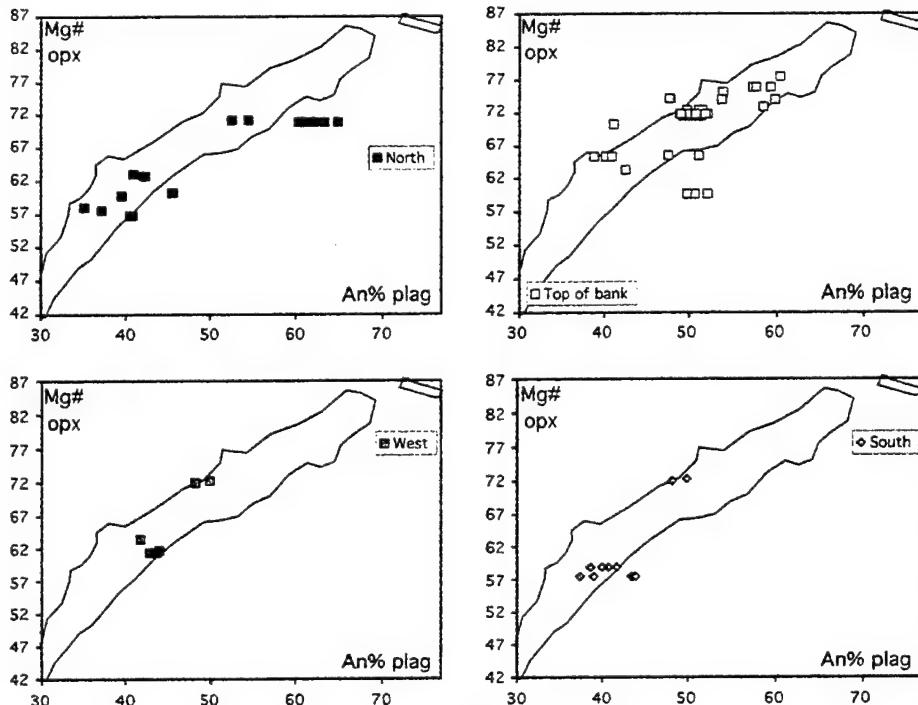


Figure 5-8: C: The composition of coexisting plagioclase (plag) and orthopyroxene (Mg#). The fields represent Hole 735B gabbronorites.

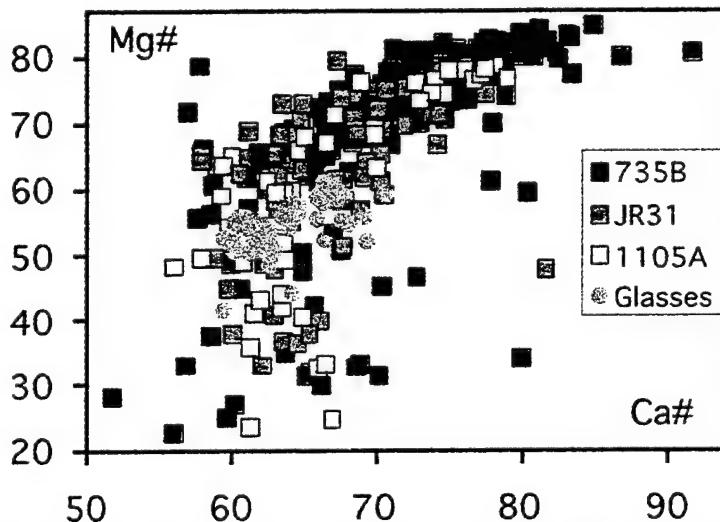


Figure 5-9: Whole-rock composition of gabbros from 735B (Dick et al., 2000), JR31 (Coogan et al. 2000), Hole 1105A (Casey et al., 2000) and unpublished glass-compositions from the fracture-zone and actively spreading ridge. It is clear that there is no distinctive difference in Mg#'s between 1105A and the rest of Atlantis Bank.

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Dick, 2002) to those obtained by JR-31 (Coogan et al., 2000) and Hole 1105A (Shipboard Scientific Party, 1999), however, show no discernable difference in Ca# vs. Mg#'s (Figure 5-9). Therefore, there may be analytical differences between the electron-probe methods used by Thy (2003) and the one in this study and by other investigators.

Olivine, as mentioned, is more rare than in Hole 735B due to iddingsite-alteration. The pause in olivine-crystallization, seen in many layered intrusions (Bowen and Schairer, 1935), is not expressed in these rocks or in Hole 735B, as olivine exists continuously from Fo<sub>80</sub> to Fo<sub>45</sub>. Orthopyroxene is even scarcer, but is again found at all the areas, with an intermediate range of compositions compared to Hole 735B.

### 5 Discussion

The rocks sampled on and around Atlantis Bank likely represent crystallization-products formed during 50 to almost 90% fractionation of a primitive magma in equilibrium with the mantle based on modeling their composition using Melts (Ghiorso and Sack, 1995) and the “SWIR PMORB” composition discussed in Chapter 3 (Fig. 5-10A, see Chapter 3 for extensive explanation of the fractionation model). The northern sample-collection has the widest compositional range, whereas the top of the bank has a narrower range, from ~55 to 82% crystallization. The rocks from the south are surprisingly evolved, representing 65-90% crystallization, including a well exposed layered olivine-gabbro with An<sub><60</sub> and Fo<sub><78</sub> at the gabbro-peridotite contact. The samples from the western wall are also evolved, and the rocks include mylonitized ferro-gabbros from the crust-mantle boundary. The most primitive sample along the wall is in fact a dike cross-cutting the section. Thus, we have not found the “missing cumulates” exposed anywhere around Atlantis Bank (i.e. hypothetical troctolites that should have been produced by the initial stages of crystallization from melts extracted from the mantle and emplaced in the Hole 735B section).

It is clear that the troctolites do not exist at or near the gabbro-peridotite contacts exposed or inferred on the transform wall, where numerous samples of evolved olivine-gabbros and ferrogabbros were recovered. Although the exact origin of this contact between serpentized peridotites and evolved gabbros is uncertain, it could be an

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intrusive contact between gabbroic rocks and residual peridotites, which was later altered and exhumed. Locally, it could be a faulted tectonic contact whose origin has been obscured by syn-and post-tectonic alteration, as believed to have been the case immediately to the west of Hole 735B. To the Southwest on that transform wall, it has been found in outcrop as a intrusive contact with gabbro overlying massive peridotite (Moroshita et al., in prep).

The northern suite has gabbros that are somewhat more primitive than found elsewhere over the bank, but they are not troctolites with primitive plagioclase and olivine. The most primitive troctolites found at Atlantis Bank are, in fact, the troctolite dikes that crosscut the upper 550-m of Hole 735B. The transform wall may be an unusual place to find primitive gabbros at very-slow spreading ridges if the mantle-derived magmas are emplaced at the segment centers and then intruded down-axis towards the transform (e.g. Whitehead et al., 1984; Dick, 1989). If we assume that melts in equilibrium with the mantle were indeed emplaced in the crust at the center of the segment, and that only evolved melts reached the transform-area (e.g. after more than 65% of crystallization), then it is possible to estimate a minimum crustal thickness below Hole 735B. Far to the southwest of Hole 735B the gabbro-peridotite contact is exposed 500 vertical meters below termination of the drill-hole. Vanko and Stakes (1991) found that Atlantis Bank has 1.65-km of missing (upper) crust based on fluid-inclusions. If the gabbro-peridotite boundary is that of the crust-mantle boundary, and it has no topography (unlikely), this would suggest an original crustal thickness on the order of 3.7-km at the ridge-transform intersection. Thus, as the crustal thickness suggested by fractional crystallization is 4.4-km (See Chapter 3), the center of the segment should be proportionally thicker. At the center of the segment, e.g. Hole 735B, the cross-cutting troctolites are the only cumulates in Fe-Mg and Ca-Na equilibrium with the spatially associated lavas. The latter are interpreted as liquids derived by 25-40% of crystallization from a primitive, mantle-derived parent (Fig. 5-10a and Chapter 3). In addition, troctolite dikes have sharp but somewhat irregular contacts with the coarse-grained olivine gabbros. This suggests that the troctolites in the upper part of Hole 735B

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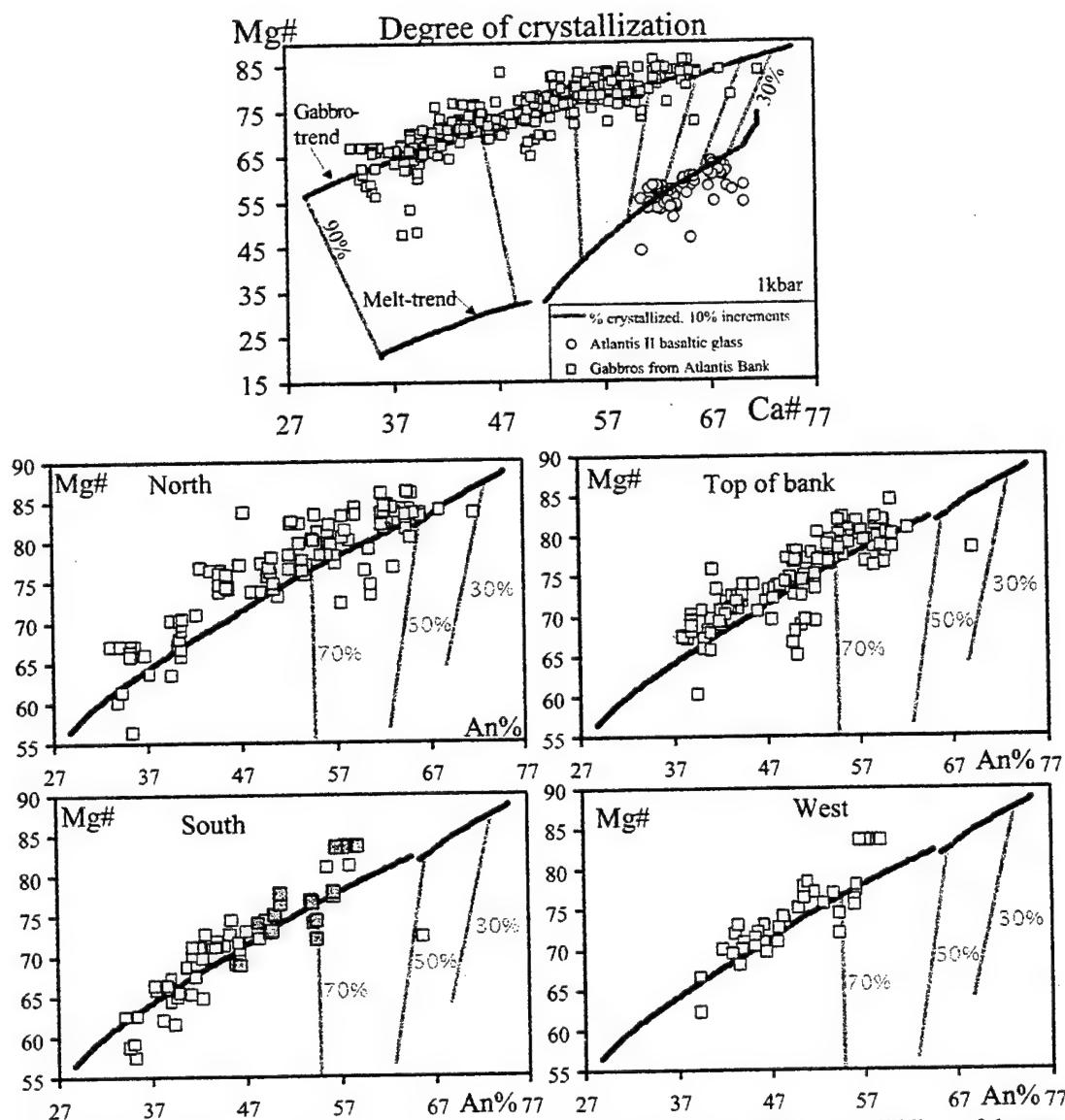


Figure 5-10: The distribution of mineral-compositions relative to modeled liquid and solid lines of descent. A: All the samples, compared to the 1kbar crystallization model presented in Chapter 2. The Gabbro-trend represents the anorthite of plagioclase with the Mg# of co-existing augite, whereas the Melt-trend represent the Mg# and Ca# of the evolving melt. The Atlantis II basaltic glasses are from Johnson et al. (1990). Each gray tie-line connects mineral-assemblage and a melt at the same stage of evolution in 10% increments. B: The Gabbro-trend compared to the area in the North. C: The Gabbro-trend compared to the top of Atlantis Bank. D: The Southern area. Note that the gray squares indicate the olivine-gabbros at the crust-mantle boundary. E: The Western area.

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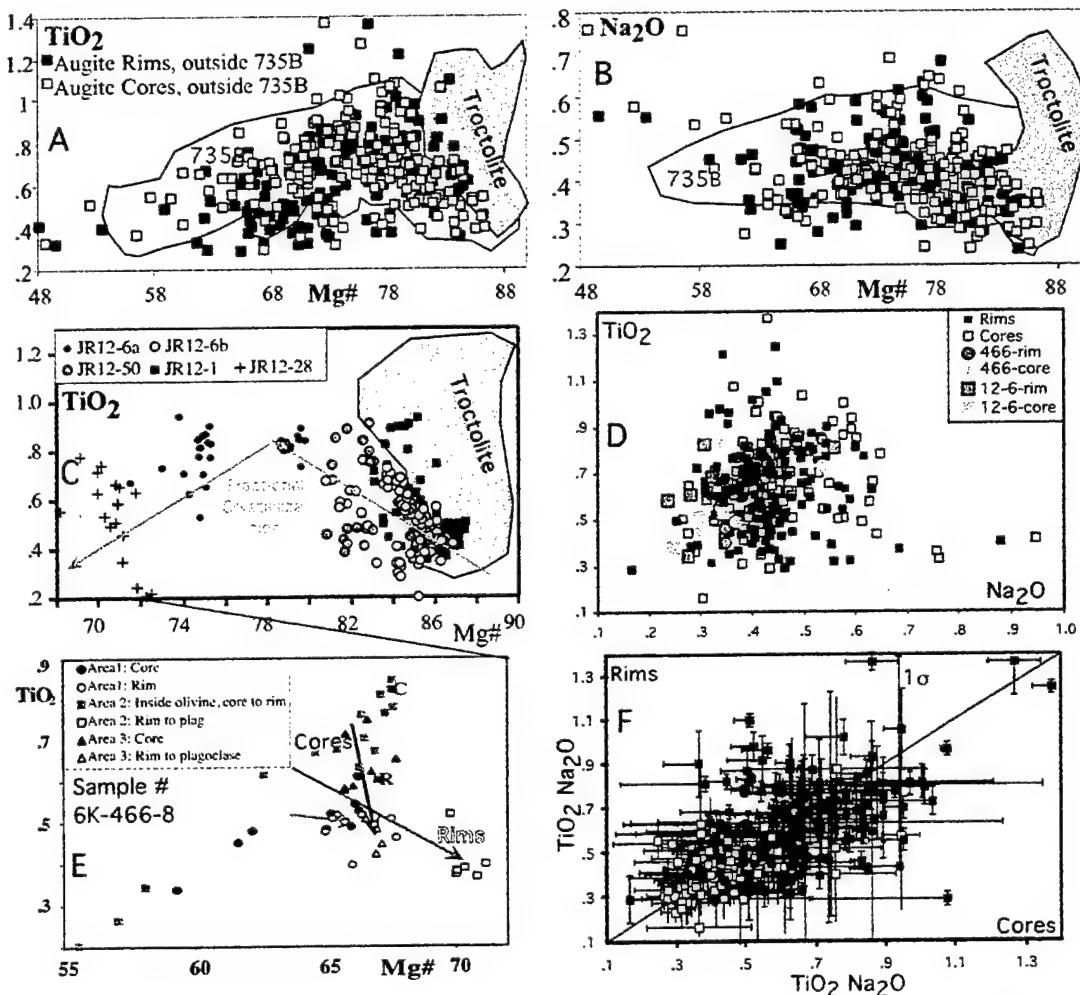


Figure 5-11: Minor element chemistry of augites analyzed in this study, compared to the upper 550-m of Hole 735B. A: Average Mg#-TiO<sub>2</sub> of augites from our study compared to the general outline of Hole 735B and the cross-cutting Hole 735B troctolites from the upper 550 meters of the core. B: Average Mg#'s vs. Na<sub>2</sub>O for the same rocks. C: Rocks in a single dredge in the northern showing wide ranges of compositions in augite (individual analyses). In addition, single samples can have dramatically different augite compositions between grains (JR12-6a and 6b). The more evolved JR12-28 sample plot perpendicular to a normal fractionation trend. D: There is no direct linear relationship between Na<sub>2</sub>O and TiO<sub>2</sub> in augite. E: Augites from a single, evolved oxide-olivine gabbro has more refractory cores than rims, plotting perpendicular to the fractional crystallization trend in C. F: There are as many reversely zoned augites as there are normally zoned ones: true both for Na<sub>2</sub>O (open squares with bars indicating 1 $\sigma$  variation) and for TiO<sub>2</sub> (closed squares with bars indicating 1 $\sigma$  variation).

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have an intrusive/reactive relationship to the wall-rocks and represent melt-channels for magma potentially erupted through the lower crust.

Figure 5-11 shows the distribution of minor-elements for the surface samples (this study) and for Hole 735B. The high Ti-Na troctolites from Hole 735B are clearly different from the rest of the core. Meyer et al. (1989) and Bloomer et al. (1989) both noted the Mg/Fe ratios of the mafic minerals in Southwest Indian Ridge gabbros are high, compared to the anorthite content of the coexisting plagioclase. Sample JR31-12-6 (Fig 5-7, 5-11E) has a large chemical range, while the more evolved rock from the same dredge-haul (12-28) plots perpendicular to the fractionation-trend. Sample 6K-466-8 is an olivine-gabbronorite – oxide-gabbronorite contact (Fig. 5-3C and 5-11G) and has augite-rims with more primitive compositions than the respective cores. In addition, the oxide-rich part has higher Mg# augites than the olivine-rich ones in the same sample. Figure 5-11F shows how the minor-element distribution is indistinguishable between the primitive and evolved gabbros, thus and that there are as many reversely zoned as normal zoned augites (Fig. 5-11H).

Meyer et al., (1989) argued that the interstitial melts in 7°E Southwest Indian Ridge gabbros reacted and equilibrated Fe and Mg with olivine, causing interstitial melt to maintain a higher Mg# during fractionation. Therefore, the augite that precipitated from this melt appears refractory relative to the melt that produced it. Minor, incompatible oxides like  $\text{TiO}_2$  increase in a fractionating melt, explaining why the high Mg# clinopyroxene has high incompatible element concentrations. If the melts that produced the entire crustal section reacted in this way with olivine, individual dredge suites or even individual samples would show a positive correlation between Mg# and  $\text{TiO}_2$  and Mg# and  $\text{Na}_2\text{O}$ . Inspection of Figure 5-11 indicates that the opposite is the case. When the fine-grained troctolites are excluded from the upper 550-m of Hole 735B, its hard to say that there are any systematics in the scatter at all. In our surface-sample data-set, individual samples and dredge-hauls have unusually large variability.

Meyer et al.'s (1989) plagioclase and clinopyroxene analyses also commonly showed reverse zoning. Angeloni and Dick (1990) demonstrated that the cross-cutting troctolites

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of Hole 735B consistently have reversely-zoned plagioclase, suggesting a process of melt percolation and overgrowth on plagioclase by later, more calcic melts with higher levels of incompatible elements. In addition, Bloomer et al. (1991) found that Unit V of Hole 735B characteristically has reversely zoned plagioclase. They argued that the high Mg# clinopyroxene and the plagioclase originated from separate magma sources, and that the crystallization happened in a mushy boundary layer. The interstitial melts are therefore a mix between main magma chamber and the interstitial equilibrium melt. Langmuir (1989) developed such a model for *in situ* crystallization in magma chambers. In this model, “melts from magma-chamber margins” mix with comparatively primitive melts within “magma-chamber melts”, and magmas then evolve along a chemical path different from that of fractional crystallization. However, the *in-situ* crystallization model only returns the interstitial melts to the magma-chamber, it does not allow for fresh, primitive melts to enter the mush. Therefore, Langmuir’s model would not make the intercumulus minerals of the most primitive cumulates appear more refractory.

Three processes may lower the TiO<sub>2</sub>-content of melt along a given liquid-line of descent: Mixing with a primitive melt; crystallization of ilmenite; and melting of, and reaction with, pre-existing cumulates by interstitial melts. Mixing of evolved and primitive melt will increase the melt Cr<sub>2</sub>O<sub>3</sub> and decrease the Na<sub>2</sub>O relative to the end-members. Cr<sub>2</sub>O<sub>3</sub> is generally low in the more evolved rocks, thus extensive mixing with primitive melts are unlikely. Crystallization of ilmenite will not increase the Mg# of the melt significantly. It was shown in Chapter 4 that dissolution of gabbroic cumulates produce melts that are low in TiO<sub>2</sub>, high in Na<sub>2</sub>O and have high Mg#’s and Ca#’s (Ca/(Ca+Na)). In fact, if plagioclase-augite phase-boundaries melt, the new melt will have ~10-15 mol% higher Mg#’s and 10% of the TiO<sub>2</sub> relative to the melt that produced the gabbros in the first place. Reprecipitation consequently produce the reverse zoning seen in the evolved cumulates.

Our results from gabbroic surface samples from on and around Atlantis Bank suggest that, with respect to chemistry, the rocks are more like the lower 1000-m of Hole 735B than the upper 550-m. In terms of petrography, the presence of numerous massive

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ferrogabbros is more like the characteristics of the upper 500-m of Hole 735B. In Chapter 3, the olivine-gabbros from the lower 1000-m were shown to be out of equilibrium. It is therefore possible that reaction and dissolution with migrating melts has modified the entire section. This process evidently may have eradicated both the very primitive, and the highly evolved compositions around the Bank (except at the most extreme end of the spreading segment). There, termination of detachment fault, which is believed to have rooted in a shallow crystal mush zone below the sheeted dikes (Dick pers. comm.), may have halted the process of melt migration before the primitive character of the gabbros was lost. Thus, even more primitive cumulates could at some point have existed, but may have been largely modified by later ascending melts. This assimilation-model is explored further in the next chapter, using trace-element compositions for the gabbroic rocks on and around Atlantis Bank.

### 6 Conclusion

The gabbros on and around Atlantis Bank, Southwest Indian Ridge are likely the products of ~50-90% fractional crystallization from a mantle-derived magma, and do not represent a complete cumulate section. The area to the north of Hole 735B has the most primitive magmatic minerals, whereas the southern gabbros are all evolved, including layered olivine-gabbros sampled at gabbro-peridotite contacts. Gabbros from the Western Wall and the top of the bank have intermediate compositions. The primitive and evolved ends of the ODP Hole 735B spectrum are not recovered in our data.

Augite is more magnesian than expected for any coexisting plagioclase anorthite content, and plagioclase and augite are frequently reversely zoned.  $TiO_2$  of augite varies inversely with Mg# at all levels of evolution, even after magnetite and ilmenite crystallization. We propose an alternative model to simple fractional crystallization, as there is evidence for extensive melt-rock reaction, possibly with simultaneous reaction and dissolution in the cumulate pile. Assimilation of pre-existing cumulates during crystallization will change the magma-budget and make our interpretation of the amount of crystallization at Atlantis Bank very different. Thus, more primitive cumulates may

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have existed at an early magmatic stage, but have been completely reconstituted by later ascending melts. This alternative model will be investigated further in the next chapter.

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## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 1: The samples analysed in this study

Tool	Core	Piece	interval	Tool	Core	Piece	interval	Tool	Dive	Piece	Tool	Dive	Piece	Tool	Dredge	Piece								
BGS	2-1	7c	37-45	BGS	27-1	1a	3	SHINKAI	459	3	SHINKAI	643	8	JR31	3	51								
BGS	3-1	2a	0-5	BGS	27-1	1b	8.5	SHINKAI	459	9	SHINKAI	645	18	JR31	9	2								
BGS	4-1	4c	43-48	BGS	27-1	1b	19.5	SHINKAI	460	6	SHINKAI	645	17	JR31	9	4								
BGS	4-2	4d	0	BGS	27-1	1b	25.5-31	SHINKAI	460	11	SHINKAI	646	1	JR31	12	1								
BGS	4-2	4d	17	BGS	27-1	1d	56-62	SHINKAI	460	13	SHINKAI	646	4	JR31	12	6								
BGS	4-2	4d	31	BGS	31-1	1c	34.5-39.5	SHINKAI	460	14	SHINKAI	646	6	JR31	12	8								
BGS	10-1	1a	31-36	BGS	31-1	1a	0-5	SHINKAI	460	15	SHINKAI	647	1	JR31	12	10								
BGS	16-1	1	7-12	BR	5-1	1b	4.5-8.5	SHINKAI	460	16	SHINKAI	647	10	JR31	12	16								
BGS	17-1	1a	25-30	BR	5-1	1h	74	SHINKAI	466	4	SHINKAI	647	13	JR31	12	28								
BGS	17-1	1b	65-67	BR	6-1	1g	20.5-24.5	SHINKAI	466	6	SHINKAI	647	9	JR31	12	39								
BGS	19-2	12a	10-11.5	BR	8-1	5d	34-36	SHINKAI	466	7	SHINKAI	649	8	JR31	12	42								
BGS	19-2	12a	38.5	BR	8-1	5g	69-71.5	SHINKAI	466	9	SHINKAI	649	8	JR31	12	45								
BGS	20-1	3b	28.5-32	BR	8-2	5h	3	SHINKAI	466	11	SHINKAI	649	12	JR31	12	46								
BGS	20-1	6a	64-67	BR	8-2	5m	18.5-23	SHINKAI	467	1	SHINKAI	649	14	JR31	12	47								
BGS	20-2	6b	0-5.5	BR	9-1	1d	26-28.5	SHINKAI	467	2	SHINKAI	650	4	JR31	12	50								
BGS	20-2	6b	12	BR	12-1	3b	33-35.5	SHINKAI	467	3	SHINKAI	650	6	JR31	12	68								
BGS	23-1	2b	23-28	BR	14-1	7a	35-37.5	SHINKAI	467	4	SHINKAI	650	9	JR31	12	8								
BGS	23-1	2b	29-34	BR	15-1	3b	17.5-20	SHINKAI	467	6	SHINKAI	653	2	JR31	12	3								
BGS	23-1	2c	44.5-48.5	BR=Oriented drill cores from the top of Atlantis Bank																				
BGS	23-1	2e	6-10	12 thinsections																				
BGS	24-1	3a	21-24	BGS=non-oriented drill cores from top of Atlantis Bank																				
BGS	24-1	3c	24	35 thinsections																				
BGS	24-1	3d	47	Mode 98: 25 thinsections from 4 dives																				
BGS	24-1	3h	51.5	from 7 dives							ABCDEF: 18 thinsections from 7 dives						41							
BGS	24-1	3i	60	JR31: 27 thinsections from 7 dredges																				
BGS	24-1	3k	77	JR31: 40																				

Total: 99

JR31: 27 thinsections from 7 dredges

Table 2a: The average composition of plagioclase from on and around Atlantis Bank.

Sample# and placement	Averages										Standard deviations									
	C	R	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%
6K-459-3 pl(c6)	x	x	67.28	22.22	.02	2.62	10.69	.03	102.9	1.9	.55	.13	.03	.29	.18	.01	.40	1.32		
6K-459-3 pl(r6)	x	x	68.00	21.35	.15	1.66	11.38	.03	102.6	7.5	.93	.72	.08	.03	.25	.44	.01	.94	3.36	
6K-459-9 pl(c6)	x	x	59.33	27.03	.11	8.43	7.25	.11	102.3	39.1	.25	.13	.02	.06	.11	.01	.25	.36		
6K-459-9 pl(r5)	x	x	59.58	26.92	.16	.01	8.17	7.43	.09	102.4	37.8	.18	.08	.03	.05	.11	.01	.47		
6K-460-6 pl c-r(8)	x	x	54.70	27.57	.65	.13	10.44	5.61	.07	99.2	50.7	1.07	.59	.02	.01	.07	.10	.04	.11	
6K-460-6 pl c-r(7)	x	x	53.50	28.45	.60	.19	11.44	4.99	.09	99.2	55.9	1.29	1.04	.27	.02	.66	.39	.02	.45	
6K-460-11 pl(outside cpx)c(6)	x	x	56.82	27.31	.02	.00	8.94	6.48	.02	99.6	43.3	.44	.23	.01	.18	1.02	.53	.34	4.86	
6K-460-11 pl(outside cpx)r(6)	x	x	56.97	27.47	.12	.00	9.04	6.39	.02	100.0	43.9	.20	.26	.06	.00	.24	.17	.01	.39	
6K-460-11 pl(c6)	x	x	56.14	27.44	.19	.02	9.36	6.06	.09	99.3	46.0	.31	.11	.06	.00	.12	.09	.01	.41	
6K-460-11 pl(c-r(3)	x	x	56.73	27.67	.26	.02	9.43	6.13	.08	100.3	46.0	.49	.22	.07	.01	.05	.10	.01	.49	
6K-460-13 pl(c6)	x	x	55.13	28.19	.10	.01	10.85	5.43	.09	99.8	52.5	.27	.13	.01	.01	.08	.09	.00	.65	
6K-460-13 pl(c-r(6)	x	x	55.37	28.02	.26	.08	10.68	5.55	.08	100.0	51.5	.28	.17	.25	.01	.07	.01	.07	.29	
6K-460-14 pl(c-r(5)	x	x	57.66	25.47	.21	.01	8.73	6.77	.04	98.9	41.6	.48	.51	.09	.15	.11	.08	.02	.36	
6K-460-14 pl(c-r(3)	x	x	56.88	25.29	.19	.35	8.97	6.38	.08	99.1	43.7	.98	.57	.42	.01	.06	.07	.01	.55	
6K-460-14 pl2-c-r(3)	x	x	57.10	25.88	.06	.00	9.09	6.56	.06	98.7	43.4	.33	.42	.02	.60	.22	.20	.05	.42	
6K-460-14 pl3-c(6)	x	x	57.59	25.33	.14	.01	8.90	6.61	.07	98.7	42.6	.41	.31	.03	.00	.14	.09	.00	.52	
6K-460-14 pl3(r(6)	x	x	58.12	26.44	.12	.00	8.11	6.95	.15	99.9	39.2	.50	.36	.03	.02	.12	.07	.01	.56	
6K-460-15pl(c12)	x	x	56.23	25.71	.41	.15	9.85	6.01	.10	98.5	47.5	.34	.28	.34	.01	.28	.15	.01	.76	
6K-460-16 pl3-c-r(6)	x	x	56.81	25.48	.21	.01	9.65	6.19	.10	98.5	46.3	.22	.55	.19	.21	.13	.06	.01	.26	
6K-460-16 pl1(c5)	x	x	56.97	25.45	.14	.00	9.41	6.24	.10	98.5	45.4	.09	.51	.02	.00	.04	.09	.01	.23	
6K-460-16 pl1r1(5)	x	x	56.38	24.64	.47	.14	9.34	6.28	.10	97.4	45.1	.36	.48	.29	.00	.09	.07	.01	.39	
6K-460-16 pl1r2(6)	x	x	56.52	24.82	.15	.01	9.41	6.36	.10	97.4	45.0	.33	.55	.01	.19	.09	.01	.33	.69	
6K-460-16 pl2-c-(5)	x	x	56.21	27.85	.11	.01	9.71	6.10	.09	100.1	46.8	.42	.40	.04	.01	.21	.17	.01	.45	
6K-466-4 pl(c6)	x	x	57.06	27.16	.13	.01	8.88	6.60	.06	99.9	42.7	.33	.12	.02	.01	.35	.20	.03	.67	
6K-466-4 pl(r6)	x	x	56.64	27.50	.12	.01	9.30	6.35	.07	100.0	44.7	.37	.26	.03	.01	.05	.10	.03	.38	
6K-466-4 pl(r(6)	x	x	56.85	27.33	.13	.01	9.09	6.47	.07	99.9	43.7	.35	.19	.02	.01	.20	.15	.03	.50	
6K-466-4 pl2-c-(5)	x	x	54.69	28.81	.18	.02	10.67	5.48	.05	99.9	51.8	.71	.54	.03	.01	.13	.12	.03	.45	
6K-466-6 pl(c6)	x	x	53.38	29.46	.16	.01	11.63	4.91	.06	99.6	56.7	.26	.07	.02	.41	.25	.01	.69	2.10	
6K-466-6 pl(c-r(6)	x	x	58.02	26.88	.10	.01	8.37	6.86	.10	100.4	40.3	.33	.25	.02	.00	.14	.08	.00	.36	
6K-466-6 pl(r(6)	x	x	58.81	25.83	.13	.01	7.40	7.45	.05	99.7	35.4	.45	.16	.05	.01	.09	.07	.01	.58	
6K-466-7 pl(r(6)	x	x	55.56	27.61	.17	.02	10.42	5.61	.09	99.5	50.7	.27	.22	.01	.05	.29	.15	.01	.46	
6K-466-8 alpl(c6)	x	x	54.49	28.32	.24	.02	10.97	5.25	.16	99.4	53.6	.46	.29	.14	.01	.07	.05	.01	.30	
6K-466-8 alpl(r(6)	x	x	55.62	27.32	.23	.02	7.82	7.32	.13	99.8	37.1	.14	.09	.03	.01	.20	.10	.01	.50	
6K-466-9 pl(c6)	x	x	58.43	26.32	.23	.02	7.45	7.64	.14	100.2	35.0	.15	.16	.03	.00	.06	.12	.03	.45	
6K-466-9 pl(c-r(6)	x	x	57.84	27.05	.20	.03	8.28	7.02	.12	100.5	39.5	.54	.16	.04	.00	.05	.07	.04	.22	
6K-466-8 a2pl-c-(9)	x	x	57.28	27.26	.25	.04	8.52	6.92	.09	100.4	40.5	.40	.31	.13	.01	.16	.18	.03	.47	
6K-466-8 a3pl(c10)	x	x	55.56	27.61	.17	.02	10.42	5.61	.09	99.5	50.7	.27	.22	.01	.05	.29	.15	.04	1.32	
6K-466-7 pl(r(6)	x	x	54.49	28.32	.24	.02	10.97	5.25	.16	99.4	53.6	.46	.29	.14	.01	.07	.05	.01	.30	
6K-466-8 alpl(c6)	x	x	55.62	27.32	.23	.02	7.82	7.32	.13	99.8	37.1	.14	.09	.03	.01	.20	.10	.01	.50	
6K-466-9 pl2-r-(6)	x	x	58.13	26.03	.12	.01	8.39	6.85	.13	99.6	40.4	.13	.27	.05	.18	.27	.18	.01	.66	
6K-466-1 pl(c6)	x	x	58.46	26.20	.15	.02	8.43	6.81	.12	100.2	40.6	.45	.24	.05	.01	.04	.14	.01	.36	
6K-466-7 pl(c6)	x	x	54.28	29.21	.11	.03	11.29	5.14	.04	100.1	54.9	.16	.11	.02	.03	.07	.09	.01	.65	
6K-466-9 pl(r(6)	x	x	53.53	29.83	.18	.04	11.88	4.72	.04	100.2	58.2	.28	.13	.04	.01	.10	.13	.01	.54	
6K-467-1 pl(cpx(6)	x	x	59.93	25.54	.14	.01	6.95	7.54	.14	100.2	33.7	.14	.15	.04	.04	.15	.07	.01	.23	
6K-467-1 pl(c-r(6)	x	x	59.78	25.62	.13	.00	6.99	7.46	.13	100.1	34.1	.23	.15	.03	.01	.05	.10	.01	.17	
6K-467-2 pl(r(6)	x	x	60.18	25.04	.20	.00	6.80	7.65	.12	100.0	32.9	.42	.12	.19	.01	.09	.11	.01	.37	
6K-467-3 pl(c6)	x	x													.01	.05	.09		.46	

Each analysis is indexed with the mineral (ol) and core (c), rim (r-c) (core to rim traverse) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample# and placement	Averages										Standard deviations									
	C	R	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%
6K-467-3 pl1(r6)	x	x	59.25	25.56	.24	.04	7.33	7.33	.10	99.8	35.6	.27	.35	.25	.09	.24	.15	.07	.18	1.10
6K-467-3 pl2 c-r(12)	x	x	59.63	25.74	.24	.00	7.27	7.42	.10	100.4	35.1	.48	.12	.16	.00	.17	.15	.01	.34	.90
6K-467-3 pl3 c-r(10)	x	x	59.70	24.46	.16	.00	7.06	7.57	.17	99.1	34.0	.35	.20	.05	.01	.16	.08	.01	.49	.39
6K-467-4a pl(c6)	x	x	59.74	25.63	.13	.02	7.27	7.45	.15	100.4	35.0	.46	.19	.05	.01	.16	.13	.02	.60	.73
6K-467-6 pl c-(9)	x	x	57.15	26.76	.11	.01	8.87	6.06	.75	99.7	44.6	1.23	1.26	0.4	.01	1.48	.63	1.39	.51	5.47
6K-467-9 pl(c6)	x	x	55.76	28.00	.16	.05	10.01	5.80	.06	99.8	48.8	.19	.28	.09	.07	.11	.05	.03	.13	.36
6K-467-9 pl1r(3)	x	x	55.61	28.07	.58	.21	10.08	5.81	.03	100.4	49.0	.17	.13	.19	.17	.11	.09	.00	.13	.54
6K-467-9 pl2c-r(4)	x	x	55.66	28.62	.21	.04	10.48	5.75	.05	100.8	50.2	.35	.11	.07	.06	.14	.00	.42	.52	
6K-467-10 pl(c6)	x	x	56.43	28.07	.09	.01	9.87	5.94	.07	100.5	47.9	.18	.18	.03	.01	.12	.12	.01	.21	.70
6K-467-10 pl(cpx)(6)	x	x	55.66	28.54	.17	.06	10.29	5.70	.06	100.5	50.0	.38	.21	.23	.12	.25	.20	.01	.18	1.36
6K-467-11 pl1r(4)	x	x	55.37	25.95	2.19	1.00	9.01	5.82	.07	99.4	47.3	6.07	2.51	3.57	1.72	3.17	2.53	.05	2.01	18.19
6K-467-11 pl2c(c4)	x	x	54.62	28.04	.15	.02	10.69	5.44	.07	99.0	52.0	1.56	1.07	.07	.02	.33	.08	.03	2.93	.51
6K-467-11 pl2c(6)	x	x	55.00	28.40	.29	.18	10.88	5.34	.07	100.7	52.9	.30	.52	.40	.47	.08	.01	.31	1.26	
6K-467-11 pl3r(12)	x	x	52.61	29.70	.59	.30	13.16	3.98	.04	100.4	64.6	.49	.34	.32	.23	.39	.14	.01	.38	1.37
6K-467-13 pl(c5)	x	x	56.28	27.26	.16	.03	9.36	6.24	.08	99.4	45.3	.17	.10	.04	.01	.07	.04	.01	.25	.11
6K-467-13 pl(c6)	x	x	56.62	27.44	.16	.03	9.41	6.23	.08	100.0	45.5	.45	.15	.02	.01	.07	.04	.01	.47	.23
6K-467-13 pl(c6)	x	x	56.45	27.35	.16	.03	9.39	6.24	.08	99.7	45.4	.31	.13	.03	.01	.07	.04	.01	.36	.17
6K-467-13 pl(c5)	x	x	54.24	29.16	.26	.03	11.49	5.02	.05	100.2	55.9	.37	.16	.03	.02	.04	.09	.01	.42	.37
6K-467-15 pl(c5)	x	x	53.94	29.36	.26	.02	11.57	4.97	.05	100.2	56.3	.33	.06	.04	.02	.01	.17	.00	.24	.85
6K-467-15 pl(c6)	x	x	55.00	28.95	.21	.01	11.02	5.26	.06	100.5	53.6	.34	.26	.02	.01	.19	.12	.01	.40	.97
6K-467-17 pl(c6)	x	x	54.94	28.77	.20	.01	10.96	5.34	.07	100.3	53.2	.28	.10	.03	.01	.09	.10	.01	.26	.56
6K-467-17 pl(c7)	x	x	65.33	21.93	.02	.00	2.95	10.17	.00	100.4	13.8	.58	.23	.03	.00	.36	.27	.01	.49	1.77
6K-467-17 pl2 c-r(10)	x	x	52.17	30.79	.22	.03	13.07	4.11	.02	100.4	63.7	.10	.17	.02	.01	.10	.08	.00	.20	.58
6K-467-20 pl(c6)	x	x	52.77	30.62	.20	.01	12.64	4.34	.02	100.6	61.7	.27	.22	.05	.01	.15	.11	.01	.37	.81
Ler ABCDE			57.49	27.58	.00	.00	9.21	6.67		101.0	43.3	.36	.14	.01	.01	.04	.15		.47	.64
6K-645-r18 pl1(10)	x	x	57.88	27.40	.14	.01	9.13	6.59	.10	101.3	43.3	.18	.19	.12	.01	.10	.10	.01	.38	.49
6K-645-r18 pl(c6)	x	x	57.86	27.65	.07	.00	9.20	6.53	.07	101.4	43.8	.18	.11	.04	.01	.05	.06	.01	.24	.30
6K-645-r18 pl(r6)	x	x	59.95	26.02	.23	.00	7.30	7.39	.05	101.0	35.3	.22	.18	.25	.00	.07	.10	.02	.24	.39
6K-645-r17 pl(r5)	x	x	59.80	25.96	.09	.00	7.22	7.40	.18	100.7	35.0	.24	.28	.02	.00	.07	.03	.01	.12	.21
6K-645-r17 pl(c5)	x	x	56.30	27.95	.21	.00	9.62	6.03	.04	100.1	46.8	.14	.19	.02	.00	.12	.02	.11	.46	
6K-647-1 pl(c5)	x	x	56.63	27.45	.41	.02	9.28	6.20	.08	100.1	45.3	.36	.18	.41	.01	.10	.06	.01	.20	.42
6K-647-1 pl10 pl(r5)	x	x	59.53	25.77	.27	.01	7.22	7.58	.05	100.4	34.5	.15	.18	.04	.02	.14	.16	.03	.25	.82
6K-647-10 pl(c5)	x	x	59.63	25.73	.17	.00	7.23	7.47	.11	100.3	34.9	.44	.07	.01	.00	.13	.02	.40	.64	
6K-647-11 pl(c6)	x	x	55.98	28.27	.24	.00	10.01	5.93	.02	100.4	48.3	.32	.24	.16	.01	.33	.17	.03	.18	1.48
6K-647-1 pl(c5)	x	x	55.83	28.31	.24	.01	10.11	5.85	.04	100.4	48.8	.33	.19	.05	.01	.16	.08	.02	.56	.44
6K-647-1 pl(c5)	x	x	56.53	27.83	.23	.01	9.49	6.15	.05	100.3	46.0	.36	.23	.04	.01	.09	.16	.02	.19	.83
6K-647-1 pl(c5)	x	x	52.33	31.02	.19	.01	13.28	4.08	.08	99.5	65.7	.32	.11	.12	.03	.12	.13	.00	.17	.86
6K-646-r18 pl(c6)	x	x	52.66	30.73	.16	.01	12.77	4.26	.03	100.6	62.4	.38	.49	.03	.00	.21	.09	.01	.16	.65
6K-646-r17 pl(c6)	x	x	51.87	30.45	.25	.04	13.07	3.76	.08	99.5	65.7	.32	.31	.02	.02	.28	.09	.03	.57	.77
6K-646-r17 pl(c6)	x	x	51.11	31.06	.24	.03	13.64	3.58	.08	99.7	67.8	.77	.57	.03	.01	.65	.39	.05	.38	3.35
6K-646-r17 pl(c6)	x	x	50.50	31.27	.23	.04	14.19	3.13	.05	99.4	71.5	.33	.45	.04	.02	.13	.03	.04	.28	.38
6K-646-r17 pl(c6)	x	x	51.87	30.45	.25	.04	13.07	3.76	.08	99.5	65.7	.32	.31	.02	.02	.28	.09	.03	.57	.77
6K-646-r10 pl(c6)	x	x	51.79	30.40	.27	.04	13.06	3.90	.13	99.6	64.9	.21	.44	.01	.01	.13	.12	.05	.18	.50
6K-646-r06 pl(c6)	x	x	52.67	30.93	.22	.02	12.85	4.19	.04	100.9	62.9	.87	.44	.01	.01	.17	.16	.01	.66	1.12
6K-646-r06 pl(c6)	x	x	56.19	28.65	.19	.00	10.33	5.77	.03	101.2	49.7	.28	.15	.06	.00	.10	.08	.02	.18	.41

Each analysis is indexed with the mineral (oh) and core (c), rim (r) or rim traverse) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample# and Placement	Averages										Standard deviations									
	C	R	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	Total	An%	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%	
6K-649-r06 plc(6)	x	x	56.80	27.99	.26	.04	9.98	5.96	.08	101.1	48.1	.23	.21	.05	.00	.07	.10	.01	.44	.52
6K-649-r06 plt(6)	x	x	56.36	28.33	.22	.02	10.37	5.78	.07	101.1	49.8	.12	.17	.06	.01	.06	.10	.00	.22	.45
6K-643-r08 plt(6)	x	x	56.69	28.32	.13	.02	9.96	5.65	.34	101.1	49.3	.32	.15	.02	.01	.10	.11	.03	.43	.49
6K-643-r08 plc(6)	x	x	56.58	28.22	.19	.05	9.96	5.60	.19	100.8	49.6	.10	.21	.07	.03	.05	.13	.11	.28	.63
6K-649-r08 pl2c(6)	x	x	53.99	29.93	.11	.03	11.97	4.92	.07	101.0	57.3	.31	.26	.03	.01	.05	.09	.01	.43	.41
6K-649-r08 pl2t(6)	x	x	54.00	29.85	.19	.04	11.93	4.97	.07	101.0	57.0	.34	.13	.14	.03	.07	.08	.01	.57	.33
6K-649-r08 plc(6)	x	x	54.40	29.57	.33	.24	11.87	5.08	.07	101.6	56.3	.31	.25	.48	.41	.30	.12	.02	.14	.63
6K-649-r08 plc(6)	x	x	54.02	29.98	.07	.06	12.32	4.79	.06	101.3	58.7	.18	.07	.04	.01	.16	.08	.01	.30	.51
6K-649-r12 plc(6)	x	x	55.35	28.55	.25	.05	10.72	5.14	.29	100.4	53.5	.32	.13	.06	.01	.08	.11	.07	.19	.66
6K-649-r12 plt(6)	x	x	55.18	29.16	.20	.03	10.89	5.09	.08	100.6	54.2	.38	.28	.02	.02	.29	.21	.06	.48	1.59
6K-649-r14 plc(6)	x	x	54.50	29.37	.34	.02	11.60	5.03	.07	100.9	56.0	.48	.08	.13	.01	.11	.15	.01	.62	.54
6K-649-r14 plc(5)	x	x	56.07	28.96	.20	.01	10.62	5.79	.07	101.7	50.4	.124	.75	.04	.01	.87	.72	.01	.32	4.98
6K-649-r14 pl2c(5)	x	x	54.87	29.65	.26	.01	11.50	5.20	.06	101.6	55.0	.44	.20	.10	.01	.19	.11	.01	.28	.83
6K-650-r04 plc(5)	x	x	53.22	30.99	.15	.01	13.11	4.25	.03	101.8	53.0	.18	.13	.02	.01	.03	.12	.01	.26	.65
6K-650-r04 plt(6)	x	x	52.19	31.70	.23	.02	13.25	3.99	.04	101.4	64.7	.136	.72	.24	.03	.14	.18	.05	.71	.85
6K-650-r06 plc(2)	x	x	53.70	29.99	.12	.00	12.03	4.65	.03	100.5	58.9	.71	.47	.04	.00	.37	.18	.00	1.75	.18
6K-650-r06 plc(cpx)(6)	x	x	53.25	30.87	.19	.01	12.80	4.29	.00	101.4	62.2	.20	.15	.02	.01	.16	.08	.01	.12	.70
6K-650-r06 plc(ol)(6)	x	x	52.38	31.27	.25	.03	13.46	3.97	.01	101.4	65.2	.94	.68	.08	.05	.77	.39	.01	.23	3.51
6K-650-r16 plt(6)	x	x	55.12	29.16	.22	.19	11.21	5.50	.05	101.4	55.0	.55	.44	.24	.40	.22	.14	.01	.60	.66
6K-650-r19 plc(5)	x	x	54.97	29.59	.11	.00	11.48	5.35	.04	101.6	54.2	.22	.14	.01	.01	.12	.08	.01	.21	.56
6K-650-r09 plt(6)	x	x	51.36	27.36	.23	.23	8.45	5.40	.06	97.0	46.2	8.46	1.88	4.35	4.87	2.37	1.42	.03	4.92	.82
6K-653-r2 plc(5)	x	x	54.59	28.90	.21	.01	10.80	5.14	.08	99.7	53.7	.30	.24	.03	.01	.15	.08	.02	.35	.70
<b>BGS-Cores</b>																				
BGS-2-17c-37.41 plc(5)	x	x	58.00	26.41	.22	.05	8.98	6.91	.12	99.8	39.3	.17	.21	.05	.09	.07	.11	.02	.25	.40
BGS-2-17c-37.41 plt(5)	x	x	58.04	26.16	.20	.06	7.85	6.99	.11	99.4	38.3	.25	.11	.08	.13	.19	.10	.02	.24	.29
BGS-3-1-2a 0-5 plc(6)	x	x	53.62	29.45	.06	.00	11.55	5.13	.04	99.8	55.4	.14	.26	.02	.00	.10	.07	.04	.31	.44
BGS-3-1-2a 0-5 plc(4)	x	x	53.46	29.74	.07	.00	11.85	5.02	.02	100.2	56.6	.14	.27	.03	.00	.07	.11	.01	.26	.69
BGS-4-1-4c 43-48 plc(6)	x	x	52.84	30.03	.17	.01	12.19	4.61	.07	100.5	59.4	.71	.52	.06	.01	.45	.01	.45	.02	2.41
BGS-4-2-0-4 plc(6)	x	x	52.85	30.02	.14	.03	12.48	4.56	.15	100.2	60.2	1.89	.97	.06	.01	1.16	.68	.04	.61	5.75
BGS-4-2-0-4 plc(6)	x	x	52.52	30.66	.09	.06	12.94	4.39	.07	100.7	62.0	.54	.28	.04	.04	.16	.15	.03	.51	1.06
BGS-4-2-17-dpl(6)	x	x	53.62	29.66	.14	.03	12.09	4.75	.15	100.4	58.4	1.24	.87	.04	.02	.97	.51	.03	.35	4.54
BGS-4-2-17-dpl(26)	x	x	53.07	29.90	.17	.03	12.44	4.59	.16	100.4	60.0	.63	.49	.05	.02	.40	.17	.03	.58	1.62
BGS-4-2-4d plc(ext)-r-c(10)	x	x	54.35	29.56	.15	.09	11.44	4.95	.05	100.6	56.1	.41	.44	.12	.20	.38	.13	.02	.59	1.31
BGS-4-2-4d plc(int)-r-(8)	x	x	54.72	29.35	.19	.03	11.34	5.11	.06	100.8	55.1	.33	.17	.04	.01	.08	.12	.02	.46	.60
BGS-4-2-4d plc(int)-r-(8)	x	x	54.79	29.44	.11	.03	11.43	5.04	.04	100.9	55.6	.53	.26	.05	.03	.31	.20	.02	.32	1.54
BGS-10-1-1a 31-36 plc(10)	x	x	55.22	28.25	.20	.00	10.28	5.76	.08	99.8	49.7	.27	.13	.01	.00	.10	.08	.00	.19	2.0
BGS-10-1-1a 31-36 plc(6)	x	x	55.16	28.47	.20	.00	10.48	5.57	.04	99.9	51.0	.10	.13	.02	.00	.07	.08	.02	.25	.37
BGS-16-1-17-12 plc(6)	x	x	55.13	28.53	.06	.00	10.38	5.63	.08	99.8	50.5	.16	.21	.02	.00	.06	.10	.00	.21	.47
BGS-16-1-17-12 plc(6)	x	x	54.77	28.74	.07	.00	10.72	5.48	.08	99.9	52.0	.15	.25	.01	.00	.07	.05	.01	.28	.29
BGS-16-1-17-12 plc(6)	x	x	55.26	28.31	.13	.00	10.20	5.72	.08	99.7	49.6	.27	.26	.03	.00	.14	.11	.01	.23	.73
BGS-17-1-1A 12 pl(10)	x	x	55.55	30.36	.28	.00	11.97	4.91	.05	103.1	57.4	.90	.54	.01	.01	.65	.35	.02	.41	3.07
BGS-17-1-1A 12 pl(5)	x	x	54.90	29.41	.06	.01	11.46	5.37	.01	101.2	54.1	.94	.45	.02	.01	.56	.37	.01	.55	2.93
BGS-17-1-17-1b 25-30 plc(3)	x	x	60.20	25.48	.13	.00	6.82	7.89	.04	100.6	32.3	.88	.42	.06	.00	.22	.19	.04	1.21	1.24
BGS-17-1f-65-67 plc(4)	x	x	58.80	25.34	.38	.06	7.08	7.33	.03	99.0	34.7	2.58	1.75	.41	.05	1.99	.87	.02	.21	9.23
BGS-19-2-10-11.5 pl(10)	x	x	53.44	28.99	.26	.04	12.28	4.75	.06	99.8	58.8	.46	.47	.10	.00	.12	.12	.01	.73	.78
BGS-19-2-12a 38.5 pl(6)	x	x	54.25	30.12	.15	.01	12.33	4.82	.04	101.7	58.6	.21	.29	.03	.01	.15	.06	.01	.36	.42

Each analysis is indexed with the mineral (ol) and core (c), rim (r) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample# and placement	Averages										Standard deviations									
	C	R	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%
BGS-19-2-12a 38.5 pl2(15)	x	x	54.07	29.68	.28	.05	12.42	4.72	.06	101.3	59.2	.55	.41	.28	.14	.38	.24	.01	.55	1.82
BGS-20-1-3B 28.5-32 plc(9)	x	x	55.24	27.62	.36	.24	10.10	5.65	.05	99.3	49.7	.39	.73	.33	.37	.45	.35	.04	.84	2.31
BGS-20-1-6a pl inside cpx(5)	x	x	55.52	28.81	.38	.04	10.72	5.72	.05	101.2	50.9	.30	.13	.05	.01	.11	.08	.01	.24	.52
BGS-20-1-6a pl(r0)-c(10)	x	x	55.76	29.09	.23	.03	10.92	5.76	.06	101.8	51.1	.20	.18	.03	.01	.27	.14	.01	.31	1.12
BGS-20-1-6a pl(r0x)-r(cpx)(10)	x	x	55.76	28.86	.26	.11	10.40	5.75	.04	101.4	50.0	.77	.27	.27	.24	.35	.23	.01	.68	1.58
BGS-20-1-6a 64-67	x	x	55.34	28.59	.27	.17	10.58	5.37	.03	100.3	52.1									
BGS-20-1-6a 64-67 plc(6)	x	x	55.85	27.59	.19	.11	10.19	5.74	.47	99.7	49.5	.47	.64	.10	.15	.34	.16	.01	.72	1.43
BGS-20-1-6a 64-67 plt(9)	x	x	55.52	27.65	.21	.08	10.40	5.72	.08	100.4	52.0	.65	.33	.05	.01	.45	.28	.01	.42	.78
BGS-20-1-6a 64-67 plc2(6)	x	x	55.40	27.78	.30	.13	10.50	5.52	.04	100.3	49.9	.45	.18	.04	.01	.17	.11	.01	.53	
BGS-20-1-6a 64-67 pl(cpx)(10)	x	x	55.83	27.33	.25	.14	10.07	5.85	.04	99.5	49.1	.10	.07	.16	.24	.05	.11	.09	.02	.46
BGS-20-2-6b al pl(undeformed) r-(10)	x	x	55.96	28.05	.14	.04	9.67	5.63	.04	99.5	48.7	.24	.24	.05	.05	.11	.09	.02	.42	.36
BGS-20-2-6b al pl(r(cpx)-c(10))	x	x	55.71	28.24	.09	.00	9.90	5.64	.03	99.6	49.2	.28	.23	.04	.00	.16	.12	.03	.41	.78
BGS-20-2-6b al2 pl(r(cpx)-c(8))	x	x	55.53	28.75	.22	.01	10.48	5.35	.01	100.4	52.0	.39	.33	.05	.01	.45	.28	.01	.42	.32
BGS-20-2-6b al2 pl(r(c(10)))	x	x	56.12	28.31	.11	.01	10.08	5.60	.04	100.3	49.9	.45	.18	.04	.01	.17	.11	.01	.42	.78
BGS-20-2-6b 0.5-8 plc(6)	x	x	55.27	28.27	.19	.00	10.14	5.87	.02	99.8	48.8	.35	.23	.02	.00	.09	.10	.01	.34	.45
BGS-20-2-6b 0.5-8 pl(r7)	x	x	54.25	27.73	.109	.60	10.25	5.53	.02	99.5	50.6	.83	1.14	.34	.88	.35	.36	.01	.22	1.98
BGS-20-2-6b 0.5-8 pl(6)	x	x	54.55	27.46	.89	.36	10.40	5.39	.04	99.1	51.6	.11	.96	.52	.63	.42	.65	.02	.64	.15
BGS-23-1-2b 23-28 al pl(inside cpx)-r-(10)	x	x	57.91	27.22	.33	.04	8.72	6.52	.08	100.8	42.5	.39	.33	.12	.07	.17	.17	.01	.57	1.00
BGS-23-1-2b 23-28 al pl(outside cpx)-r-(10)	x	x	58.23	27.31	.14	.01	8.60	6.59	.08	101.0	41.9	.43	.24	.05	.01	.23	.17	.02	.35	1.06
BGS-23-1-2b 23-28 al pl(inside) r-(5)	x	x	58.22	27.44	.27	.06	8.77	6.45	.11	101.3	42.9	.28	.29	.13	.11	.12	.14	.01	.54	.35
BGS-23-1-2b 23-28 a2 pl(r-c(2))	x	x	58.23	26.71	.52	.28	8.89	6.51	.10	101.2	43.0	.23	.99	.26	.35	.25	.02	.00	.36	.79
BGS-23-1-2b 23-28 a2 pl(outside) r-c(5)	x	x	57.51	27.54	.65	.13	9.01	6.28	.10	101.2	44.2	.91	.20	.28	.10	.29	.02	.28	1.02	
BGS-23-1-2b 23-28 a3 pl(r-t9)	x	x	57.84	27.76	.17	.00	8.96	6.54	.05	101.3	43.1	.42	.17	.05	.00	.17	.19	.02	.41	1.10
BGS-23-1-2b 23-28 a3 pl(r-c(6))	x	x	57.67	27.60	.29	.10	8.91	6.36	.06	101.0	43.6	.66	.14	.32	.18	.11	.19	.02	.49	.84
BGS-23-1-2b 23-28 a4 pl(r-c(10))	x	x	59.57	26.66	.19	.03	7.78	7.08	.11	101.4	37.8	.64	.48	.09	.03	.49	.31	.02	.50	2.45
BGS-23-1-2b 29-34 a1 pl(r-(4))	x	x	56.61	27.67	.34	.09	9.44	6.16	.08	100.4	45.9	.29	.13	.13	.11	.09	.07	.02	.26	.23
BGS-23-1-2b 29-34 a4 pl(p2 r-c(t0))	x	x	57.37	27.31	.19	.07	8.97	6.57	.09	100.6	43.0	.20	.20	.09	.11	.11	.08	.01	.33	.43
BGS-23-1-2b 29-34 a2 pl(r-t9)	x	x	57.72	26.99	.14	.00	8.63	6.74	.10	100.3	41.5	.43	.20	.02	.00	.06	.11	.00	.53	.29
BGS-23-1-2b 29-34 a3 pl(r-c(6))	x	x	57.31	27.27	.16	.01	8.86	6.60	.09	100.3	42.6	.22	.17	.03	.01	.09	.09	.00	.55	.50
BGS-23-1-2b 29-34 a3 pl(r-t9)	x	x	56.42	27.69	.56	.17	9.22	6.45	.05	100.5	44.1	.89	.39	.82	.46	.33	.24	.00	.55	.60
BGS-23-1-2b 29-34 a4 pl(outside cpx) r-(5)	x	x	56.14	27.02	.21	.04	9.02	6.28	.07	98.8	44.3	.67	.24	.06	.04	.09	.11	.01	.83	.61
BGS-23-1-2b 29-34 a4 pl(p2 r-c(t8))	x	x	56.53	27.65	.28	.06	9.30	6.17	.10	100.1	45.4	.83	.25	.30	.14	.42	.23	.02	.62	1.93
BGS-23-1-2b 29-34 a5 pl(r-t8)	x	x	56.52	26.93	.65	.21	8.94	6.40	.10	99.7	43.6	.14	.60	.96	.37	.20	.26	.03	.73	.85
BGS-23-1-2b 29-34 a5 pl(r-t7)	x	x	56.14	25.57	1.26	.94	9.60	6.04	.14	99.7	46.7	.81	2.70	1.15	.71	1.36	.62	.03	.71	6.01
BGS-23-1-2c 44.5-48.5 pl(r-t9)	x	x	57.63	26.44	.18	.01	8.27	6.83	.11	99.5	40.1	.48	.22	.02	.01	.28	.22	.01	.32	1.52
BGS-23-1-2c 44.5-48.5 pl(r(t0))	x	x	58.34	27.01	.18	.03	8.28	6.72	.13	100.7	40.5	.60	.27	.12	.05	.36	.22	.03	.14	1.85
BGS-23-1-2e a1 pl(r-c(8))	x	x	58.33	26.58	.15	.06	7.89	6.92	.11	100.0	38.6	.62	.23	.04	.01	.21	.13	.02	.57	1.04
BGS-23-1-2e a3 pl(r-c(10))	x	x	58.19	26.82	.16	.02	8.19	6.77	.12	100.3	40.1	.44	.14	.08	.03	.15	.11	.02	.39	.71
BGS-23-1-2e a4 pl(r-c(10))	x	x	58.25	26.92	.19	.03	8.43	6.77	.11	100.7	40.8	.41	.23	.07	.14	.11	.10	.03	.51	.58
BGS-24-1-3a 6-10 plc(6)	x	x	55.65	28.50	.17	.02	10.34	5.71	.09	100.5	50.0	.37	.08	.02	.01	.11	.13	.01	.45	.70
BGS-24-1-3a 6-10 pl(t6)	x	x	55.18	28.27	.22	.04	10.36	5.59	.09	99.8	50.6	.25	.14	.10	.07	.08	.10	.01	.25	.57
BGS-24-1-3c 21-24 pl(c-r(6))	x	x	57.61	27.20	.20	.01	8.80	6.63	.08	100.5	42.3	.45	.27	.05	.01	.25	.19	.04	.57	1.18
BGS-24-1-3e a3 pl(r-c(10))	x	x	60.48	27.19	.64	.22	8.19	6.82	.30	103.8	39.9	.74	1.11	.128	.70	.60	.34	.27	.53	1.21
BGS-24-1-3h 47 pl(6)	x	x	58.04	26.46	.20	.02	8.78	6.78	.15	100.4	41.7	.16	.49	.03	.04	.10	.03	.07	.71	1.33
BGS-24-1-3h 51.5 pl(3)	x	x	57.35	28.02	.19	.00	9.88	6.07	.09	101.6	47.4	.36	.24	.03	.00	.09	.10	.01	.30	.15
BGS-24-1-3h 51.5 pl(6)	x	x	56.52	28.40	.20	.00	10.59	5.65	.10	101.5	50.9	.52	.36	.02	.01	.51	.30	.01	.39	2.49

Each analysis is indexed with the mineral (ol) and core (c), rim (r)-c (core to rim traverse) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample# and placement	Averages										Standard deviations										
	C	R	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%	
BGS-24-1-3i 60 pl(6)	x	x	56.16	27.89	.18	.00	10.39	5.78	.09	100.5	49.8	1.01	.36	.03	.64	.36	.01	.41	3.08		
BGS-24-1-3k 77 pl(9)	x	x	59.88	27.53	.26	.09	8.48	6.74	.12	103.1	41.0	.66	.41	.25	.27	.05	.02	.68	.72		
BGS-27-1-1A 3 pl(9)	x	x	56.37	29.73	.18	.00	10.89	5.21	.04	102.4	53.6	.31	.15	.03	.01	.14	.08	.02	.30	.59	
BGS-27-1-1B 8.5 pl(6)	x	x	55.63	29.29	.18	.00	11.30	5.14	.04	101.6	54.8	.24	.25	.03	.00	.16	.01	.33	.93		
BGS-27-1-1B 8.5 pl(8)	x	x	55.02	29.37	.22	.05	11.26	5.22	.02	101.2	54.4	.28	.29	.09	.13	.18	.21	.01	.34	1.39	
BGS-27-1-1B 19.5 pl(15)	x	x	54.19	29.05	.44	.18	11.39	5.15	.04	100.4	55.0	.63	.56	.93	.54	.68	.21	.02	.76	1.92	
BGS-27-1-1b 25.5-31 pl(5)	x	x	53.56	28.98	.13	.00	11.17	5.14	.05	99.0	54.6	.18	.11	.05	.00	.17	.09	.01	.21	.52	
BGS-27-1-1d 36-62 pl-r-c(10)	x	x	51.09	31.72	.40	.33	13.56	3.37	.03	100.5	69.0	.88	.79	.75	.02	.09	.01	.22	.02	.99	1.53
BGS-31-1-1c 34.5-39.5 pl(c6)	x	x	55.95	27.83	.19	.01	9.88	5.95	.10	99.9	47.8	.24	.25	.07	.00	.04	.01	.04	.01	.35	.26
BGS-31-1-1c 34.5-39.5 pl(r6)	x	x	56.33	27.93	.14	.01	9.73	6.05	.05	100.2	47.0	.28	.15	.03	.01	.09	.12	.02	.41	.69	
BGS-31-1-a 0-5 pl(c12)	x	x	54.83	28.73	.20	.02	10.77	5.51	.08	100.1	52.0	.31	.22	.03	.01	.10	.13	.01	.51	.70	
BR-Drills																					
BR-5-1-1B 4.5-8.5 pl(c6)	x	x	54.16	29.42	.12	.01	11.40	5.01	.05	100.2	55.7	.19	.09	.04	.01	.06	.09	.00	.20	.48	
BR-5-1-1B 4.5-8.5 pl(r6)	x	x	54.57	29.09	.11	.08	11.09	5.22	.05	100.2	54.0	.43	.20	.11	.15	.14	.10	.00	.25	.57	
BR-5-1-1H 74-78 pl2(c6)	x	x	54.62	28.55	.21	.11	11.07	5.27	.09	99.9	53.7	.36	.18	.11	.18	.07	.12	.01	.34	.53	
BR-5-1-1H 74-78 pl2(t6)	x	x	54.35	28.88	.11	.01	11.26	5.16	.07	99.8	54.7	.28	.19	.01	.01	.21	.13	.01	.25	1.02	
BR-5-1-1H 74-78 pl-c-t(12)	x	x	54.60	28.65	.16	.03	11.04	5.33	.07	99.9	53.4	.26	.18	.07	.04	.09	.08	.01	.31	.49	
BR-5-1-2e 48-51.5 plc	x	x	53.73	29.00	.20	.01	10.85	5.46	.05	99.3	52.4	.29	.20	.01	.02	.07	.08	.01	.31	.47	
BR-6-1-1g 20.5-24.5 pl(c6)	x	x	55.74	27.87	.13	.03	10.01	6.08	.11	100.0	47.6	.28	.20	.01	.02	.07	.08	.01	.31	.47	
BR-6-1-1g 20.5-24.5 pl(r6)	x	x	55.90	27.92	.13	.01	10.00	6.10	.09	100.2	47.5	.39	.09	.05	.01	.03	.11	.01	.45	.44	
BR-8-1-5D 34-36 pl1-r-t(11)	x	x	54.42	27.00	.22	.03	10.95	5.26	.07	98.0	53.5	.38	.39	.09	.04	.19	.09	.01	.55	.67	
BR-8-1-5D 34-36 pl2-c-t(8)	x	x	52.96	28.17	.23	.05	12.10	4.52	.04	98.1	59.6	.29	.38	.08	.08	.23	.23	.01	.39	.76	
BR-8-1-5G 69-71.5 pl-r-t(15)	x	x	53.02	27.78	.25	.02	11.86	4.69	.07	97.7	53.3	.65	.21	.02	.01	.28	.17	.02	.51	1.47	
BR-8-2-5h 3 plH(10)	x	x	53.40	28.73	.26	.05	11.81	4.92	.07	99.2	57.0	.81	.50	.04	.09	.49	.26	.02	.57	2.28	
BR-8-2-5h 3 pl2(6)	x	x	53.83	29.37	.20	.00	11.99	4.92	.04	100.3	57.4	.59	.44	.04	.00	.28	.08	.01	.76	.85	
BR-8-2-5h 3 pl3(4)	x	x	52.89	29.25	.31	.01	12.29	4.69	.07	99.5	59.1	.62	.47	.18	.01	.26	.14	.05	.99	1.25	
BR-8-2-5m 18.5-23 pl-c-t(8)	x	x	52.17	29.19	.45	.20	11.80	4.30	.04	98.2	60.3	.56	.31	.27	.26	.21	.10	.02	.19	.56	
BR-9-1-1D 26-28.5 pl(c5)	x	x	55.13	27.59	.76	.03	9.80	5.80	.05	99.2	48.3	.89	.36	.41	.02	.24	.12	.01	.37	.82	
BR-9-1-1D 26-28.5 pl(r6)	x	x	55.06	27.93	.15	.01	9.99	5.81	.06	99.0	48.8	.41	.17	.05	.01	.05	.16	.07	.56	.62	
BR-14-1-7a 35-37.5 pl(t10)	x	x	54.47	28.12	.16	.01	10.73	5.55	.04	99.1	51.7	.22	.19	.06	.04	.11	.07	.01	.30	.37	
BR-14-1-7a 35-37.5 pl2(c6)	x	x	54.68	28.19	.25	.12	10.40	5.63	.04	99.3	50.5	.35	.25	.19	.24	.23	.09	.01	.24	.78	
BR-14-1-7a 35-37.5 pl2(r6)	x	x	54.55	28.13	.33	.20	10.37	5.57	.04	99.2	50.7	.47	.39	.25	.29	.23	.20	.01	.42	1.27	
BR-15-3B 17.5-20 pl(c-t10)	x	x	52.68	30.00	.23	.04	12.70	4.46	.05	100.2	61.1	.31	.21	.06	.01	.13	.07	.01	.39	.48	
BR-15-3B 17.5-20 pl2-c-t(10)	x	x	53.54	29.67	.17	.05	12.20	4.68	.03	100.4	59.0	.39	.24	.07	.09	.10	.09	.01	.45	.51	
JCR31-Dredges																					
JCR31-3-5l pl(c6)	x	x	52.33	30.99	.56	.17	13.70	3.99	.03	101.8	65.5	.07	.35	.03	.02	.15	.13	.01	.38	.94	
JCR31-3-5l pl(r6)	x	x	57.98	26.67	.81	.11	8.98	6.86	.09	101.5	42.0	.38	.32	.07	.04	.31	.16	.01	.37	1.33	
JCR31-9-2 pl(c6)	x	x	56.69	28.69	.23	.03	10.13	6.01	.07	101.8	48.2	.12	.15	.03	.01	.16	.08	.01	.26	.58	
JCR31-9-2 pl(t6)	x	x	56.16	29.10	.27	.02	10.77	5.76	.06	102.1	50.8	.30	.09	.05	.01	.14	.14	.01	.24	.90	
JCR31-9-4 pl(c6)	x	x	57.05	28.35	.26	.06	10.13	6.14	.09	102.1	47.7	.24	.15	.03	.01	.08	.08	.00	.30	.44	
JCR31-9-4 pl(t5)	x	x	58.53	27.32	.31	.16	8.95	6.55	.10	101.9	43.0	.30	.12	.17	.14	.13	.01	.25	.80		
JCR31-12-1 ai pl(cpx)(6)	x	x	52.89	30.53	.20	.03	12.67	4.35	.03	100.7	61.7	.34	.23	.03	.01	.13	.07	.01	.22	.51	
JCR31-12-1 ai pl(c6)	x	x	51.94	31.00	.20	.05	13.05	3.91	.03	100.2	64.9	.09	.09	.04	.02	.04	.09	.01	1.00	1.51	
JCR31-12-1 ai pl(r6)(6)	x	x	52.75	30.65	.27	.05	12.93	4.24	.03	100.9	62.7	.24	.24	.02	.02	.22	.07	.01	.41	.64	
JCR31-12-1 a2 pl(cpx)(6)	x	x	53.09	29.81	.29	.24	12.06	4.65	.04	100.2	58.9	.21	.32	.05	.42	.16	.11	.02	.28	.33	
JCR31-12-1 a2 pl(c6)	x	x	52.46	30.92	.23	.06	13.16	4.01	.05	100.9	64.4	.29	.13	.01	.04	.19	.07	.02	.54	.40	

Each analysis is indexed with the mineral (ol) and core (c), rim (r) or traverse (t) and are accompanied by standard deviations (s).

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample# and placement	C	R	Averages										Standard deviations									
			SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%		
JR31-12-1 a2 pl(r)(0)(6)	x	x	52.39	30.97	.04	13.16	4.06	.04	100.8	64.2	.24	.17	.03	.01	.16	.09	.02	.16	.73	.16	.37	3.17
JR31-12-6 pl(c4)(6)	x	x	52.79	30.16	.19	03	12.30	4.45	.04	100.0	60.4	.90	.61	.05	.02	.29	.11	.29	.11	.01	.44	1.12
JR31-12-6 pl(r8)(6)	x	x	53.12	30.44	.23	03	12.70	4.31	.03	100.9	61.9	.11	.29	.02	.01	.29	.11	.29	.11	.01	.44	1.12
JR31-12-6 pl(r2)(6)	x	x	52.95	30.06	.13	00	12.38	4.49	.03	100.0	60.4	.32	.20	.03	.01	.08	.15	.14	.05	.01	.53	.59
JR31-12-6 pl(r1)(6)	x	x	52.20	30.24	.16	04	12.77	4.34	.04	99.8	61.9	.08	.13	.02	.02	.23	.06	.01	.23	.06	.01	.31
JR31-12-6 pl(c3)(6)	x	x	51.65	30.84	.19	03	13.25	3.99	.03	100.0	64.7	.39	.29	.03	.01	.30	.18	.01	.19	.19	.49	1.49
JR31-12-6 pl(c7)(6)	x	x	53.31	30.04	.38	06	12.36	4.52	.06	100.7	60.2	.45	.38	.05	.03	.09	.04	.02	.02	.46	.31	
JR31-12-6 pl(r-t)(30)	x	x	53.21	30.17	.20	02	12.64	4.46	.04	100.7	61.0	.48	.35	.05	.02	.40	.23	.02	.43	1.84		
JR31-12-6 pl(r-t5)(50)	x	x	52.35	30.54	.22	03	12.84	4.14	.04	100.1	63.2	.56	.39	.07	.02	.44	.24	.02	.34	.20		
JR31-12-8 pl(c6)	x	x	54.76	29.42	.11	13	11.42	5.24	.02	101.1	54.7	.46	.38	.04	.17	.38	.16	.01	.19	1.52		
JR31-12-8 pl(t6)	x	x	54.11	30.14	.20	03	11.98	4.92	.01	101.4	57.4	.36	.15	.03	.01	.16	.12	.01	.22	.75		
JR31-12-8 pl(toutside ol)(4)	x	x	53.45	30.20	.27	18	12.40	4.76	.02	101.2	58.8	1.30	1.26	.16	.26	1.29	.72	.01	.42	6.21		
JR31-12-8 pl(coutside ol)(4)	x	x	52.98	30.57	.14	03	12.77	4.44	.03	100.9	61.4	.22	.15	.07	.01	.07	.09	.01	.32	.46		
JR31-12-8 pl(inside ol)c-r(4)	x	x	53.25	30.48	.11	03	12.38	4.53	.03	101.0	60.5	.44	.16	.01	.02	.44	.24	.02	.34	.20		
JR31-12-10 pl(c6)	x	x	54.13	29.35	.27	04	11.67	4.77	.05	100.3	57.6	.18	.19	.08	.02	.13	.09	.01	.36	.43		
JR31-12-10 pl(cnp)(3)	x	x	53.79	29.67	.17	03	11.62	4.80	.04	100.1	57.3	.42	.05	.02	.05	.07	.04	.04	.50	1.22		
JR31-12-16 pl(cnp)(5)	x	x	54.23	29.47	.14	00	11.66	5.06	.02	100.6	56.1	.25	.20	.04	.00	.21	.12	.01	.60	.60		
JR31-12-16 pl(cpx)(2)	x	x	54.77	29.52	.15	03	11.97	5.07	.02	101.5	56.6	.52	.02	.03	.02	.16	.40	.01	.78	2.26		
JR31-12-16 pl(cpx)(4)	x	x	54.62	29.02	.26	19	12.27	5.06	.05	100.5	55.2	.58	.45	.14	.28	.38	.19	.01	.19	1.69		
JR31-12-16 pl(cpx)(6)	x	x	58.16	26.84	.18	08	8.46	6.81	.06	100.6	40.7	.63	.44	.18	.18	.30	.26	.09	.48	1.60		
JR31-12-28 pl(cpx)(6)	x	x	58.23	27.13	.19	00	8.82	6.72	.07	101.2	42.0	.14	.08	.07	.00	.08	.16	.03	.29	.74		
JR31-12-28 pl(cpx)(6)	x	x	57.99	26.89	.19	04	8.78	6.66	.07	100.6	42.2	.36	.21	.10	.04	.17	.04	.06	.65	.50		
JR31-12-30 pl(cpx)(6)	x	x	59.67	26.01	.31	10	7.16	7.26	.05	100.6	35.3	.45	.51	.49	.16	.21	.19	.07	.59	.81		
JR31-12-39 pl(cpx)(5)	x	x	59.67	26.46	.20	00	7.67	7.34	.02	101.4	36.6	.21	.18	.04	.00	.08	.14	.01	.41	.68		
JR31-12-42 pl(c6)	x	x	55.69	29.13	.15	04	10.92	5.53	.05	101.5	52.2	.34	.20	.02	.01	.09	.08	.01	.49	.48		
JR31-12-42 pl(r6)	x	x	55.70	28.79	.22	03	10.57	5.82	.02	101.2	50.1	.30	.19	.15	.01	.08	.17	.02	.25	.88		
JR31-12-45 pl(c6)	x	x	54.85	29.63	.15	02	11.46	5.31	.01	101.4	54.4	.14	.17	.02	.01	.14	.14	.00	.35	.75		
JR31-12-45 pl(r6)	x	x	55.06	29.35	.16	04	11.09	5.58	.02	101.3	52.3	.18	.26	.03	.01	.11	.09	.01	.29	.57		
JR31-12-46 pl(c5)	x	x	54.63	29.00	.28	16	11.24	5.20	.05	100.6	54.4	.97	.75	.29	.25	.08	.07	.06	1.16	.47		
JR31-12-46 pl(cpx)(1)	x	x	54.27	27.98	.11	00	11.80	4.37	.02	99.6	59.9	.33	.17	.05	.01	.30	.11	.01	.46	.89		
JR31-12-47 pl(c6)	x	x	54.29	29.66	.18	01	11.68	5.12	.07	101.0	55.8	.14	.09	.02	.01	.12	.11	.02	.39	.49		
JR31-12-47 pl(t6)	x	x	54.41	29.62	.16	01	11.62	5.23	.05	101.1	55.1	.18	.20	.02	.01	.08	.11	.02	.16	.62		
JR31-12-50 pl(c6)	x	x	51.83	29.90	.28	06	12.69	4.36	.04	99.2	61.6	.48	.30	.14	.01	.18	.11	.01	.84	.79		
JR31-12-50 pl(c5)	x	x	53.19	29.05	.19	05	11.47	4.96	.04	99.0	56.1	.31	.30	.03	.02	.10	.18	.01	.41	1.10		
JR31-12-68 pl(c6)	x	x	53.91	29.85	.36	05	11.85	4.79	.07	100.9	57.8	.33	.17	.05	.01	.30	.11	.01	.46	.89		
JR31-12-68 pl(r6)	x	x	55.13	29.33	.21	05	10.89	5.30	.03	100.9	53.2	.21	.11	.03	.03	.14	.07	.01	.39	.49		
JR31-39-1 pl(c6)	x	x	54.09	28.99	.27	06	11.04	5.23	.05	99.7	53.9	.18	.18	.19	.02	.11	.12	.02	.16	.62		
JR31-39-1 pl(r6)	x	x	56.52	27.55	.23	05	9.26	6.14	.08	99.8	45.4	.25	.15	.01	.01	.20	.11	.02	.46	.81		
JR31-39-3 pl(c6)	x	x	57.05	27.67	.16	01	9.21	6.32	.08	100.5	44.6	.14	.23	.13	.01	.18	.06	.01	.35	.34		
JR31-39-3 pl(r6)	x	x	57.90	27.18	.33	01	8.50	6.67	.11	100.7	41.3	.15	.14	.05	.01	.14	.10	.02	.30	.40		
JR31-39-4 pl(c6)	x	x	58.86	26.77	.19	03	8.20	7.11	.12	101.3	38.9	.23	.06	.02	.01	.10	.15	.00	.29	.79		
JR31-39-4 pl(r6)	x	x	59.22	26.44	.28	03	7.88	7.32	.11	101.3	37.3	.27	.36	.05	.00	.38	.31	.01	.30	2.14		
JR31-39-4 pl(inclusion)(4)	x	x	59.03	26.73	.20	03	8.24	7.14	.09	101.5	38.9	.25	.15	.01	.01	.04	.08	.02	.29	.18		
JR31-39-4 pl(r6)	x	x	60.38	26.11	.13	01	7.27	7.79	.14	101.8	34.1	.30	.17	.04	.01	.03	.15	.02	.53	.48		
JR31-39-7 pl(c6)	x	x	58.77	26.34	.47	10	7.82	7.00	.32	100.8	38.1	.95	.10	.42	.16	.88	.53	.52	1.73	3.03		
JR31-39-7 pl(r6)	x	x	56.70	27.82	.25	00	9.40	6.04	.09	100.3	46.2	.11	.15	.06	.01	.13	.10	.01	.31	.50		

Each analysis is indexed with the mineral (c), rim (r) or core (c), rim (r-c) (core to rim traverse) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample# and placement	C	R	Averages										Standard deviations									
			SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Total	An%		
JR31-39-8 pl1(2)	x	x	56.68	27.81	.26	.00	9.30	6.06	.08	100.2	45.9	.22	.11	.03	.01	.29	.09	.01	.27	.79		
JR31-39-8 pl1(6)	x	x	56.94	27.68	.27	.01	8.96	6.32	.08	100.3	43.9	.31	.27	.03	.02	.29	.10	.02	.71	.75		
JR31-39-10 pl1c(6)	x	x	58.45	26.68	.18	.00	8.13	6.94	.09	100.5	39.3	.19	.13	.07	.01	.17	.16	.01	.34	.83		
JR31-39-10 pl2(6)	x	x	58.41	26.60	.37	.01	8.06	7.04	.08	100.6	38.7	.31	.28	.08	.01	.24	.07	.03	.23	.86		
JR31-39-10 plc(6)	x	x	57.37	26.85	.28	.01	8.59	6.66	.07	99.8	41.6	.24	.18	.03	.01	.15	.06	.04	.33	.47		
JR31-39-10 plt(6)	x	x	57.36	27.03	.34	.02	8.73	6.59	.07	100.1	42.3	.24	.20	.05	.01	.18	.10	.02	.44	.78		
JR31-39-11 plc(6)	x	x	59.10	26.69	.18	.01	8.14	6.85	.12	101.1	39.7	.19	.16	.03	.01	.05	.06	.01	.16	.15		
JR31-39-11 plt(6)	x	x	58.69	27.19	.24	.00	8.58	6.77	.07	101.5	41.2	.25	.27	.03	.01	.23	.11	.03	.48	.84		
JR31-39-13 plc(6)	x	x	59.59	26.51	.23	.03	7.91	7.38	.14	101.8	37.2	.19	.12	.02	.01	.12	.15	.02	.16	.81		
JR31-39-13 plf(6)	x	x	59.13	26.95	.29	.03	8.26	7.28	.13	102.1	38.5	.30	.23	.02	.01	.10	.11	.02	.34	.59		
JR31-39-13 plt(6)	x	x	58.03	26.77	.17	.03	8.24	6.87	.10	100.2	39.8	.27	.18	.02	.01	.14	.06	.02	.26	.38		
JR31-39-17 pl1c(6)	x	x	58.24	26.57	.18	.02	8.03	7.08	.08	100.2	38.5	.14	.22	.03	.01	.11	.10	.03	.37	.41		
JR31-39-17 pl2(6)	x	x	57.66	27.18	.36	.04	8.63	6.70	.10	100.7	41.6	.27	.14	.03	.03	.10	.07	.02	.17	.40		
JR31-39-17 pl2c(5)	x	x	57.88	27.00	.22	.08	8.30	6.71	.10	100.3	40.6	.40	.29	.10	.09	.25	.10	.02	.59	.71		
JR31-40-1 plc(6)	x	x	54.86	29.87	.21	.02	11.81	5.27	.04	102.1	55.3	.84	.35	.04	.01	.04	.08	.01	1.07	.38		
JR31-40-1 plt(6)	x	x	54.79	30.45	.28	.02	12.33	4.98	.04	102.9	57.8	.34	.36	.05	.01	.37	.15	.01	.44	1.45		
JR31-41-4 plc(6)	x	x	58.69	27.94	.07	.01	9.09	6.86	.07	102.7	42.3	.28	.12	.02	.00	.15	.06	.02	.26	.57		
JR31-41-4 plt(6)	x	x	58.67	27.97	.20	.04	9.15	6.80	.04	102.9	42.6	.42	.45	.12	.06	.31	.15	.02	.24	1.33		

Each analysis is indexed with the mineral (ol) and core (c), rim (r) (core to rim traverse) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 2B: The average composition of augite from on and around Atlantis Bank.

Sample # and placement	Averages										Standard deviations														
	C	R	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnO	MgO	CaO	Na <sub>2</sub> O	Total	Mg#	SiO <sub>2</sub> , TiO <sub>2</sub> , Al <sub>2</sub> O <sub>3</sub> , Cr <sub>2</sub> O <sub>3</sub> , FeO	MnO	MgO	CaO	Na <sub>2</sub> O	Total Mg#						
Mode 98	x	x	51.61	.51	3.58	.01	18.67	.21	11.59	12.37	.58	99.1	52.5	.49	.13	.49	.01	1.09	.05	.51	.17	.15	.55	.244	
6K-459-3 cpxc(8)	x	x	51.14	.32	4.13	.03	19.96	.16	11.01	12.23	.56	99.5	49.6												
6K-459-3 cpxc(1)	x	x	50.22	.33	3.88	.02	20.13	.21	10.71	13.37	.76	99.6	48.7	.61	.13	1.60	.02	.92	.04	.60	1.59	.32	.58	.201	
6K-459-9 cpxc(7)	x	x	49.41	.41	4.75	.03	19.89	.18	10.31	13.43	.88	99.3	48.1	.20	.11	1.35	.01	1.84	.04	.35	2.69	.25	.70	.308	
6K-459-9 cpxc(7)	x	x	51.81	.86	2.14	.16	8.10	.25	16.67	18.96	.00	99.0	78.6	.35	.04	.16	.02	.18	.01	.13	.20	.00	.73	.50	
6K-460-6 cpxc(6)	x	x	49.76	1.37	3.72	.22	8.18	.21	14.90	20.13	.00	98.5	76.4	.25	.07	.37	.10	.48	.01	.28	.25	.00	.41	1.34	
6K-460-6 cpx2c(6)	x	x	50.77	1.27	3.13	.16	8.65	.21	15.15	20.28	.00	99.6	75.7	.56	.15	.10	.06	.41	.02	.56	.27	.00	.53	1.53	
6K-460-11 cpxc(8)	x	x	52.10	.49	2.19	.00	9.03	.28	13.91	21.57	.44	100.0	73.3	.39	.06	.21	.01	.15	.03	.16	.25	.05	.57	.28	
6K-460-11 cpxc(2)	x	x	51.50	.33	1.84	.01	9.70	.31	13.33	21.07	.59	100.0	71.1	.74	.08	1.14	.01	1.85	.01	.47	1.65	.32	.56	4.65	
6K-460-13 cpxc(10)	x	x	52.39	.72	2.59	.02	10.18	.24	17.48	16.37	.41	100.4	76.0	.79	.23	.62	.01	3.78	.06	3.39	7.31	.21	.37	2.95	
6K-460-13 cpxc(10)	x	x	52.43	.70	2.41	.02	8.17	.22	15.56	20.45	.38	100.3	77.4	.50	.12	.18	.01	1.60	.04	1.20	2.80	.06	.55	1.84	
6K-460-14 cpx3 c-(10)	x	x	52.33	.68	1.85	.00	10.43	.30	13.86	20.33	.42	100.2	70.3	.28	.07	.07	.01	.34	.03	.17	.41	.03	.30	.63	
6K-460-14 cpx2c(10)	x	x	51.94	.61	1.80	.01	12.11	.32	14.45	18.65	.37	100.3	68.3	.30	.07	.24	.01	2.53	.07	1.18	3.56	.07	.44	2.50	
6K-460-14 cpx2c(6)	x	x	52.21	.66	1.82	.01	11.03	.30	14.17	20.02	.39	100.6	69.6	.40	.08	.14	.01	.94	.05	.28	1.24	.02	.75	1.51	
6K-460-15 cpxc(6)	x	x	52.30	.17	.95	.00	12.23	.30	13.82	17.97	.30	98.1	66.8	.39	.11	.60	.01	1.42	.04	1.46	4.86	.13	1.19	1.40	
6K-460-15 cpxc(2)	x	x	51.90	.29	2.92	.00	13.82	.26	12.72	16.92	.46	98.2	62.5	.77	.29	1.84	.00	3.95	.03	.01	7.18	.28	2.61	6.74	
6K-460-16 cpxc(8)	x	x	52.00	.68	2.02	.01	11.08	.30	15.00	18.94	.40	100.4	71.2	.44	.15	.36	.01	2.77	.06	1.60	4.19	.12	.47	2.40	
6K-460-16 cpxc(10)	x	x	52.22	.67	2.00	.01	10.08	.27	14.73	20.22	.41	100.6	72.4	.34	.08	.12	.01	1.35	.03	.84	2.26	.05	.45	1.34	
6K-460-16 cpxc(8)	x	x	51.89	.66	1.99	.00	13.05	.32	16.13	15.70	.36	100.1	69.9	.48	.26	.74	.01	5.85	.14	3.28	8.55	.23	.20	4.43	
6K-460-16 cpxc(6)	x	x	51.79	.75	2.17	.00	10.21	.28	14.46	20.20	.46	100.3	71.7	.57	.13	.33	.00	1.51	.02	1.00	2.37	.14	.73	1.50	
6K-460-16 cpxc(10)	x	x	52.10	.65	1.91	.01	10.98	.31	14.68	19.63	.38	100.7	70.7	.36	.13	.23	.01	2.35	.03	1.29	3.58	.07	.49	2.31	
6K-466-4 cpxc(9)	x	x	51.67	.78	2.82	.01	7.54	.21	14.36	21.69	.65	99.7	77.2	.28	.04	.15	.01	2.4	.02	.14	.29	.10	.40	.54	
6K-466-4 cpxc(10)	x	x	52.35	.55	2.11	.01	7.67	.24	14.46	21.91	.54	99.9	77.1	.38	.05	.15	.01	2.1	.01	.15	.24	.07	.59	.66	
6K-466-4 cpxc-t(5)	x	x	52.22	.63	2.28	.01	7.92	.26	14.55	21.63	.48	100.0	76.6	.51	.14	.43	.01	.39	.02	.28	.25	.11	.30	1.20	
6K-466-6 cpxc(6)	x	x	51.39	.81	2.97	.06	7.59	.27	14.16	21.35	.60	99.2	76.8	.55	.24	.58	.02	1.15	.13	.91	.44	.20	.58	3.84	
6K-466-6 cpxc(10)	x	x	51.19	.97	2.96	.06	10.98	.31	14.68	19.63	.38	100.7	70.7	.36	.13	.23	.01	2.35	.03	.14	.18	.05	.50	.60	
6K-466-6 cpxc(9)	x	x	51.72	.81	2.71	.07	7.47	.22	15.31	20.27	.44	99.0	72.7	.44	.38	.44	.01	.55	.03	.46	.96	.10	.40	.54	
6K-466-7 cpxc(7)	x	x	51.56	.70	2.16	.03	11.16	.32	12.93	20.91	.44	100.2	67.4	.43	.04	.15	.02	.43	.01	.21	.01	.22	.02	.19	.33
6K-466-7 cpxc(7)	x	x	51.82	.63	2.30	.01	12.24	.33	13.57	17.87	.58	100.0	66.3	.88	.53	.78	.01	.77	.03	.35	3.10	.19	.99	1.67	
6K-466-8 al cpxc(8)	x	x	52.17	.49	1.62	.00	14.12	.43	14.00	18.36	.37	101.6	63.9	.28	.08	.23	.00	2.86	.08	.07	3.54	.10	.36	2.70	
6K-466-6 cpxc(9)	x	x	52.01	.50	1.68	.00	11.76	.40	12.95	20.83	.37	100.5	66.2	.56	.02	.04	.01	.49	.03	.12	.42	.03	.69	1.04	
6K-466-8 a2 cpxc(inside ol) c-(15)	x	x	51.81	.63	1.79	.01	14.48	.44	14.32	16.87	.34	100.7	63.8	.23	.20	.51	.01	4.94	.12	1.71	6.13	.14	.45	4.19	
6K-466-7 cpxc(tl)(6)	x	x	52.66	.40	1.62	.00	13.30	.33	13.72	21.87	.35	101.3	70.4	.29	.06	.07	.01	.22	.04	.08	.27	.04	.47	.52	
6K-466-8-a3 cpxc(9)	x	x	51.69	.74	2.42	.01	12.11	.39	13.20	20.26	.53	101.3	66.0	.28	.77	.01	.61	.04	.36	.95	.24	.60	.93		
6K-466-8-a3 cpxc(9)	x	x	52.09	.46	1.90	.01	11.62	.36	13.19	21.27	.35	101.3	66.9	.36	.04	.04	.01	.23	.03	.12	.39	.03	.58	.26	
6K-466-9 cpxr(10)	x	x	51.78	.70	2.61	.04	9.73	.26	15.14	19.52	.41	100.2	73.5	.21	.03	.07	.02	.40	.02	.31	.59	.04	.42	.47	
6K-466-9 cpxc(9)	x	x	51.91	.66	2.54	.03	9.56	.25	16.42	17.98	.37	99.7	76.1	.60	.18	.58	.02	4.18	.06	3.69	7.64	.17	.48	2.96	
6K-466-9 cpxc3(10)	x	x	51.52	.71	2.46	.03	8.24	.22	14.56	21.54	.40	99.7	75.9	.46	.09	.23	.01	.33	.02	.20	.40	.03	.55	.93	
6K-466-11 cpx2c(0)	x	x	51.63	.67	2.04	.05	11.66	.32	13.65	20.42	.42	99.9	68.2	.67	.05	.02	.01	.69	.03	.20	.82	.04	1.05	1.38	
6K-466-11 cpx2r(10)	x	x	52.43	.48	1.82	.04	10.80	.32	13.65	20.42	.42	100.4	69.3	.27	.06	.18	.01	.35	.02	.10	.81	.04	.50	.78	
6K-466-11 cpxc(10)	x	x	52.01	.56	2.01	.05	10.12	.31	13.50	21.33	.44	100.3	70.4	.40	.08	.23	.01	.18	.04	.12	.33	.04	.30	.53	
6K-466-11 cpxc3(10)	x	x	52.33	.48	1.66	.03	10.31	.30	13.44	21.45	.37	100.4	69.9	.33	.06	.17	.02	.40	.04	.09	.46	.02	.33	.77	
6K-467-1 cpxc(9)	x	x	52.57	.51	2.76	.11	6.70	.19	16.34	20.67	.26	100.1	81.3	.27	.06	.29	.01	.56	.02	.41	1.08	.06	.33	.94	
6K-467-1 cpxc(p)(9)	x	x	52.20	.63	2.53	.10	6.60	.22	15.15	22.30	.31	100.0	80.4	.55	.08	.28	.01	.24	.03	.27	.33	.05	.48	.60	

Each analysis is indexed with the mineral (ol) and core (c), rim (r) +c (particular notes), and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample # and placement	Averages										Standard deviations														
	C	R	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnC	MgO	CaO	Na <sub>2</sub> O	Total	Mg#	SiO <sub>2</sub>	TiO <sub>2</sub> , Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnO	MgO	CaO	Na <sub>2</sub> O	Total Mg#			
6K-467-2 cpx(c9)	x	49.76	.66	2.45	.00	13.83	.40	11.81	19.48	.55	98.9	60.4	.40	.03	.19	.13	.40	.01	.68	.05	.19	.54	.02	.41	1.50
6K-467-2 cpx(t3)	x	50.74	.33	1.79	.00	13.70	.36	12.30	18.96	.45	98.6	61.6	.82	.06	.45	.00	1.29	.04	.24	1.81	.06	.77	2.70		
6K-467-3 cpx(c10)	x	52.04	.29	1.25	.02	11.20	.33	12.95	20.97	.43	99.5	67.3	.83	.11	.27	.01	.40	.04	.28	.38	.10	.80	1.02		
6K-467-4 cpx(c9)	x	51.55	.51	1.90	.03	11.50	.36	12.83	20.80	.48	99.9	66.5	.43	.04	.18	.01	.37	.03	.11	.15	.04	.50	.84		
6K-467-4 cpx(t10)	x	51.82	.37	1.59	.03	11.96	.32	13.03	20.49	.44	100.1	66.1	.05	.03	.17	.01	1.32	.03	.34	.80	.04	.73	2.92		
6K-467-4 cpx(t10)	x	51.42	1.02	2.84	.03	8.84	.26	14.02	21.12	.69	100.2	73.9	1.14	.64	.97	.01	.79	.02	.19	1.34	.26	.62	1.87		
6K-467-6 cpx-c-(t16)	x	50.93	.94	3.25	.03	8.93	.25	14.92	19.72	.59	99.5	74.9	.85	.77	.78	.05	.46	.12	.59	1.24	.12	.59	1.84		
6K-467-8 cpx(c10)	x	52.79	.43	1.82	.03	8.00	.26	14.46	22.11	.46	100.4	76.3	.77	.26	.90	.01	.47	.07	.15	1.18	.22	.40	1.22		
6K-467-8 cpx(t9)	x	52.14	.83	2.60	.03	9.76	.24	15.44	18.91	.42	100.4	73.9	.34	.08	.22	.01	1.49	.05	1.13	2.56	.06	.32	1.50		
6K-467-9 cpx(c10)	x	52.23	.70	2.44	.04	7.50	.24	14.55	21.88	.43	100.0	77.6	.49	.06	.19	.01	.27	.03	.13	.48	.04	.45	.51		
6K-467-9 cpx Ir(t10)	x	51.68	.86	2.80	.04	9.19	.26	14.86	20.12	.53	100.3	74.3	.44	.11	.29	.02	.97	.02	.70	1.57	.09	.48	1.07		
6K-467-9 cpx2c(t10)	x	51.81	.80	2.51	.04	8.88	.25	14.94	20.46	.49	100.2	75.0	.23	.16	.34	.01	.90	.03	.44	1.63	.06	.48	1.40		
6K-467-9 cpx2c(t10)	x	51.99	.84	2.61	.03	9.70	.28	15.45	18.98	.44	100.3	74.0	.20	.05	.12	.01	.98	.02	.71	1.48	.05	.28	1.98		
6K-467-10 cpx(c8)	x	52.39	.61	2.23	.03	8.02	.23	14.52	21.83	.41	100.3	76.4	.21	.08	.18	.01	.49	.03	.21	.71	.03	.23	.93		
6K-467-10 cpx(c10)	x	51.75	.66	3.37	.11	6.05	.15	17.61	18.81	.00	99.5	83.8	.96	.97	.22	.58	.02	1.01	.74	.00	.36	.51			
6K-467-11 cpx Ir(t12)	x	50.36	1.06	4.57	.03	6.15	.14	16.29	19.68	.00	99.3	82.5	.39	.17	.24	.19	.35	.03	.38	.61	.00	.44	.85		
6K-467-11 cpx2c(t11)	x	52.21	.66	2.95	.02	5.87	.16	17.63	19.24	.00	99.5	84.2	.85	.24	1.02	.17	.61	.02	.73	.55	.00	.33	1.60		
6K-467-11 cpx3r-t(t8)	x	51.07	.85	2.58	.03	10.10	.27	15.02	18.89	.48	99.3	72.7	.98	.07	.15	.01	1.17	.03	.84	2.05	.05	1.26	1.90		
6K-467-12 cpx(c10)	x	51.77	.85	2.54	.03	10.51	.29	15.40	18.27	.48	100.1	72.4	.31	.07	.17	.01	1.17	.04	.81	1.93	.07	.45	1.28		
6K-467-12 cpx Ir(t10)	x	51.63	.89	2.64	.03	10.67	.28	15.18	18.50	.46	100.3	71.7	.07	.13	.02	.03	.56	.13	.04	.28	.71				
6K-467-12 cpx2c(t10)	x	51.95	.75	2.46	.03	8.92	.24	14.56	20.87	.46	100.3	74.4	.48	.06	.16	.01	.53	.03	.48	.91	.04	.51	.78		
6K-467-12 cpx2r(t10)	x	52.06	.63	2.38	.02	7.96	.26	14.27	21.96	.48	100.0	76.2	.54	.11	.31	.01	.36	.03	.22	.38	.04	.44	1.04		
6K-467-13 cpx(c9)	x	51.35	.88	2.80	.01	8.72	.24	14.25	21.04	.51	99.8	74.5	.53	.13	.19	.01	.44	.02	.21	.52	.02	.55	1.01		
6K-467-13 cpx(c10)	x	51.26	1.01	2.86	.01	8.48	.22	13.93	21.40	.58	99.7	74.5	.28	.08	.05	.01	.27	.02	.12	.24	.06	.44	.48		
6K-467-13 cpx(t10)	x	51.94	.79	2.41	.02	8.63	.31	14.33	21.59	.50	100.5	74.7	.50	.13	.39	.01	.50	.03	.65	.75	.06	.72	.91		
6K-467-13 cpx Ir(t1)	x	51.42	.82	2.65	.03	8.65	.24	14.05	21.33	.51	99.7	74.3	.50	.13	.39	.01	.50	.03	.65	.75	.06	.72	.91		
6K-467-13 cpx(c9)	x	51.62	.70	2.50	.01	8.64	.26	14.15	21.35	.45	99.7	74.5	.36	.12	.23	.01	.72	.02	.11	.61	.04	.69	1.67		
6K-467-13 cpx-c(t9)	x	51.18	.62	4.00	.08	6.60	.19	14.90	20.74	.51	98.8	80.1	.87	.03	.44	.01	.26	.04	.68	.69	.06	.66	.97		
6K-467-13 cpx(c7)	x	51.42	.80	2.95	.06	7.04	.21	15.19	21.28	.46	99.4	79.4	.79	.07	.42	.01	.62	.02	.16	.97	.03	.90	1.37		
6K-467-13 cpx(t7)	x	50.58	.59	3.52	.07	6.56	.16	16.02	20.75	.45	98.7	81.3	.90	.06	.30	.01	.22	.02	.21	.40	.08	.97	.44		
6K-467-13 cpx(t9)	x	50.96	.66	3.55	.07	6.96	.18	15.41	20.75	.49	99.0	79.8	.02	.04	.26	.01	.26	.03	.12	.81	.04	1.24	.59		
6K-467-13 cpx-c(t9)	x	52.09	.60	2.71	.05	8.55	.20	16.37	18.95	.44	100.0	77.4	.36	.05	.11	.01	1.19	.02	1.28	2.36	.05	.57	1.03		
6K-467-17 cpx(c10)	x	51.88	.80	2.65	.06	7.62	.19	14.97	21.02	.52	99.7	77.8	.38	.04	.16	.02	.40	.02	.37	.74	.04	.47	.62		
6K-467-17 cpx(t10)	x	51.99	.70	2.68	.05	8.09	.19	15.67	19.99	.48	99.8	77.6	.37	.05	.14	.01	.80	.02	.82	1.55	.04	.52	.83		
6K-467-17 cpx(c10)	x	52.37	.50	3.43	.51	5.66	.15	16.91	20.39	.33	100.2	84.2	.14	.02	.05	.02	.22	.02	.30	.40	.03	.08	.37		
6K-467-20 cpx(c9)	x	52.03	.78	3.03	.29	5.52	.18	15.91	21.91	.38	100.0	83.7	.58	.04	.55	.02	.26	.01	.14	.59	.10	.37	.71		
<b>Leg ABCDE</b>																									
6K-645-r18 cpxr(8)	x	51.78	.83	2.56	.01	10.35	.10	14.16	20.87	.54	101.2	71.2	.14	.40	.01	3.00	.10	.56	1.18	.11	.61	6.02			
6K-647-r17 cpxr(p)(9)	x	51.68	.52	1.89	.01	13.14	.40	12.41	20.15	.43	100.6	62.8	.46	.05	.64	.01	.85	.04	.34	1.49	.05	.56	1.10		
6K-647-r17 cpxc(9)	x	48.94	.55	1.90	.02	15.44	.43	11.78	19.86	.54	99.5	57.7	.17	.06	.13	.02	1.37	.07	.24	.41	.02	.47	2.74		
6K-647-r17 cpxr(p)(10)	x	51.68	.65	2.70	.01	9.01	.24	13.86	21.10	.57	99.8	73.3	.64	.17	.89	.01	.41	.02	.18	1.47	.17	.59	.76		
6K-647-r01 cpxc(10)	x	51.63	.87	2.74	.01	9.31	.26	14.04	20.74	.57	100.2	72.9	.09	.02	.04	.01	.38	.01	.16	.41	.05	.17	.71		
6K-647-r01 cpxr(p)(5)	x	49.48	.49	2.24	.03	14.65	.42	11.74	19.81	.45	99.3	58.9	.19	.04	.78	.02	1.24	.04	.45	.87	.03	.75	2.90		
6K-647-r10 cpxc(6)	x	49.80	.54	2.05	.03	14.64	.43	11.83	20.02	.43	99.8	59.4	.27	.12	.01	3.30	.04	.57	1.45	.03	1.08	5.85			
6K-647-r13 cpx-c(14)	x	52.35	.43	1.96	.15	9.56	.26	14.07	21.18	.38	100.3	72.4	.29	.08	.20	.03	.38	.03	.17	.44	.05	.39	.86		
6K-647-r09 cpxr(p)(7)	x	52.19	.79	2.75	.01	8.97	.24	14.90	19.40	.46	99.7	74.7	.76	.21	.15	.01	.60	.03	1.08	2.86	.06	.78	.54		
6K-647-r09 cpxc(8)	x	51.26	1.01	2.89	.02	10.66	.28	15.19	18.65	.41	100.4	71.9	.59	.08	.20	.01	1.74	.03	.44	1.25	.04	.37	2.88		

Each analysis is indexed with the mineral (ol) and core (c), rim (h-c) (core to rim traverse) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample # and placement	Averages										Standard deviations											
	C	R	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	CaO	MnO	MgO	Total	Mg#	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	CaO	MnO	MgO	Total	Mg#		
6K-646-r01 cpxc(5)	x	51.15	.54	3.75	.63	6.60	.13	17.24	18.91	.39	99.3	82.4	.32	.05	1.22	.03	1.65	2.78	.09	.36	1.22	
6K-646-r01 cpxr(p)(5)	x	50.36	.78	3.23	.61	5.91	.14	15.17	22.23	.42	98.8	82.1	.68	.05	.34	.03	.35	.03	.28	.49	.03	.51
6K-646-r01-cpx2(r4)	x	50.10	.63	3.58	.67	6.27	.14	15.77	20.71	.47	98.3	81.7	.37	.03	.12	.02	.12	.02	.47	.81	.03	.23
6K-646-r01-cpx2(r4)	x	50.28	.91	3.71	.59	5.62	.12	14.98	21.70	.44	98.3	82.6	.21	.04	.38	.02	.34	.04	.25	.42	.03	.65
6K-646-R04 cpxc(5)	x	50.98	.41	3.41	.37	5.73	.16	16.51	21.09	.38	99.0	83.7	.61	.03	.32	.01	.35	.01	.62	1.04	.03	.58
6K-646-R04 cpxr1(7)	x	51.91	.58	2.57	.29	5.40	.16	16.05	22.62	.34	99.9	84.1	.10	.04	.08	.02	.12	.02	.24	.44	.04	.33
6K-646-R04 cpxr2(7)	x	51.59	.52	2.55	.32	5.58	.17	16.98	22.52	.33	99.7	83.7	.48	.05	.25	.02	.19	.03	.28	.43	.04	.49
6K-646-R04 cpxc2(6)	x	51.04	.56	3.73	.46	5.47	.18	15.35	22.76	.41	99.9	83.4	.33	.03	.12	.02	.07	.03	.25	.22	.06	.51
6K-646-r06 cpxc(2)	x	51.40	.58	3.41	.11	8.99	.21	16.97	16.10	.63	98.4	76.9	.04	.28	.55	.04	.28	.05	.236	5.01	.13	1.48
6K-649-r06 cpx1(c8)	x	50.83	.86	3.12	.07	9.80	.29	15.86	18.37	.49	99.7	74.3	.67	.07	.22	.01	1.13	.03	.95	2.04	.06	.79
6K-649-r06 cpx1(r8)	x	50.64	.93	3.25	.07	8.96	.29	15.29	19.64	.46	99.5	75.3	.50	.06	.38	.02	1.29	.04	1.30	2.66	.09	.57
6K-649-r08 a2 cpxc(9)	x	53.33	.71	2.83	.20	5.78	.00	16.69	21.45	.47	101.5	83.7	.26	.05	.06	.02	.47	.00	.39	.79	.03	.26
6K-649-r08 a2 cpxr(10)	x	53.47	.69	2.73	.19	5.85	.00	16.75	21.40	.47	101.5	83.6	.22	.03	.07	.02	.63	.00	.60	.84	.04	.33
6K-649-r12 cpxc(8)	x	51.78	.69	2.65	.02	8.15	.25	15.42	20.59	.46	100.0	77.1	.55	.06	.13	.01	.36	.02	.83	1.66	.06	.73
6K-649-r12 cpxr1(6)	x	51.43	.71	3.63	.01	10.41	.28	15.11	16.96	.50	99.0	72.3	.29	.12	1.01	.01	2.16	.02	.59	2.73	.12	1.05
6K-649-r12 cpxr2(8)	x	51.22	.87	2.72	.02	8.96	.25	14.76	20.91	.43	100.2	74.6	.56	.04	.19	.01	1.00	.04	2.4	1.06	.07	.70
6K-649-r14 cpx1(c9)	x	50.19	1.04	3.68	.02	7.68	.23	14.89	21.21	.49	99.4	77.6	.43	.06	1.00	.01	.24	.04	.28	.30	.02	.32
6K-649-r14 cpx1(r4)	x	50.71	.73	3.54	.02	9.83	.20	17.31	16.67	.29	99.4	76.6	.30	.45	.02	.49	.09	.50	1.12	.20	.27	3.28
6K-649-r14 cpx2(c9)	x	51.07	.66	3.01	.07	7.95	.24	15.92	20.04	.45	99.5	78.1	.33	.02	.22	.01	.69	.04	.71	1.16	.06	.54
6K-649-r14 cpx2(r3)	x	51.07	.75	2.89	.01	7.53	.23	15.00	21.56	.39	99.4	78.0	.94	.04	.64	.01	.32	.01	.38	.43	.05	.60
6K-649-r14 cpxc(5)	x	52.54	.37	1.93	.24	5.56	.12	16.18	22.39	.29	99.6	83.9	.20	.15	.54	.11	1.27	.03	.40	.50	.05	.30
6K-650-r04 cpxr(p)(3)	x	51.14	.90	3.44	.12	6.37	.14	14.90	21.55	.54	99.1	80.7	.67	.10	.68	.00	.12	.03	.27	.66	.17	.85
6K-650-r06 cpxc(8)	x	51.70	.63	3.48	.17	5.47	.12	16.62	20.79	.39	99.4	84.4	.43	.03	.44	.02	.50	.03	.65	1.19	.03	.54
6K-650-r06 cpxr(p)(6)	x	51.29	.62	3.05	.19	4.95	.13	15.58	22.85	.35	99.0	84.9	.12	.16	1.48	.02	.44	.02	.55	1.01	.10	.62
6K-650-r09 cpxc(8)	x	53.22	.82	2.50	.03	7.18	.10	16.16	21.34	.48	101.8	80.1	.35	.04	.01	.47	.11	.36	.52	.03	.41	.86
6K-650-r09 cpxr(7)	x	53.14	.83	2.45	.02	6.86	.07	15.84	22.12	.43	101.8	80.5	.33	.07	.18	.01	.44	.04	.13	.46	.02	.39
6K-653-r2 cpxr(p)(7)	x	51.25	.51	3.02	.01	10.68	.29	13.28	19.71	.46	99.2	69.0	.72	.15	1.40	.00	1.74	.05	.33	.29	.18	.75
6K-653-r2 cpxc(7)	x	51.87	.67	3.29	.02	9.25	.24	17.06	17.06	.35	99.8	76.8	.67	.11	.82	.01	1.68	.04	1.66	3.17	.08	.22
BGS-Cores																						
BGS-2-1-7c-37-41 cpx1c(5)	x	50.77	.42	3.68	.05	16.14	.28	13.86	11.63	.95	97.8	60.5	.68	.14	1.28	.01	.61	.01	.60	.34	.33	.54
BGS-2-1-7c-37-41 cpx2r(4)	x	49.09	.55	2.10	.03	10.84	.35	12.60	20.24	.57	96.4	67.5	.75	.02	.10	.01	.36	.02	.21	.37	.06	2.31
BGS-3-1-2a 0-5 cpxr(5)	x	51.54	.49	2.74	.16	7.24	.20	15.94	20.40	.61	99.3	79.8	.16	.09	1.04	.04	1.57	.02	1.06	3.34	.30	.68
BGS-3-1-2a 0-5 cpxr(7)	x	52.15	.46	2.43	.20	6.18	.19	15.73	22.02	.44	99.8	82.0	.56	.10	.42	.03	.69	.03	.21	.73	.10	.28
BGS-4-1-4c 43-48 cpx2(4)	x	50.03	.98	3.16	.11	6.90	.18	15.06	21.00	.47	98.9	79.5	.26	.07	.16	.01	.14	.01	.06	.38	.11	.36
BGS-4-1-4c 43-48 cpx3(5)	x	51.66	.51	3.22	.17	6.33	.18	16.12	21.01	.33	99.5	81.9	.50	.02	.04	.02	.22	.02	.12	.21	.02	.68
BGS-4-1-4c 43-48 cpx4(4)	x	50.71	.81	2.71	.11	7.89	.22	16.51	17.04	.33	96.3	78.9	.24	.05	.07	.02	.20	.01	.39	.88	.03	.56
BGS-4-1-4c 43-48 cpx5(5)	x	52.45	.10	2.96	.11	8.01	.24	16.14	19.56	.31	99.3	78.2	.45	.03	.09	.02	.78	.01	.61	1.25	.04	1.06
BGS-4-1-4c 43-48 cpx5(5)	x	51.44	.96	2.85	.10	8.36	.25	15.58	19.56	.24	99.5	76.9	.18	.04	.04	.01	.13	.04	.16	2.29	.05	.26
BGS-4-1-4c 43-48 cpx5(5)	x	51.77	.22	3.18	.11	7.27	.23	15.40	21.10	.34	100.6	79.0	.30	.05	.06	.01	.17	.03	.12	.33	.04	.11
BGS-4-2-0-4d cpx1c(10)	x	52.11	.45	2.99	.16	5.74	.17	15.59	20.16	.27	99.7	84.5	.41	.02	.08	.01	.33	.02	.51	.84	.03	.32
BGS-4-2-0-4d cpx1r(8)	x	52.51	.52	2.20	.17	6.50	.19	15.53	21.72	.30	99.6	81.0	.30	.05	.13	.01	.18	.03	.16	.36	.08	.46
BGS-4-2-17-4d cpx1c(10)	x	51.67	.56	3.13	.27	7.42	.19	16.23	19.53	.34	99.3	79.6	.30	.03	.07	.02	.51	.02	.54	1.07	.05	.32
BGS-4-2-17-4d cpx1r(8)	x	51.43	.96	2.97	.12	7.13	.22	15.18	20.98	.32	99.3	79.1	.29	.07	.09	.02	.36	.04	.22	.49	.06	.46
BGS-4-2-4d al cpxc(9)	x	52.86	.53	2.57	.17	7.14	.19	16.63	20.79	.34	101.2	80.6	.44	.06	.13	.02	1.16	.02	1.12	2.20	.04	.61
BGS-4-2-4d al cpx c-r(p)(15)	x	52.72	.66	2.54	.17	6.71	.19	16.12	21.53	.34	101.0	81.1	.39	.14	.23	.02	.54	.03	.51	.98	.06	.47
BGS-4-2-4d cpxr(o)(6)	x	51.90	.98	2.84	.16	6.80	.19	15.62	22.33	.34	101.1	80.4	.35	.06	.09	.02	.21	.03	.06	.27	.04	.44
BGS-4-2-4d al cpx r(o)-r(p)	x	52.72	.71	4.04	.21	15.88	.23	16.23	23.34	.32	101.8	82.4	.77	.05	.13	.01	.23	.08	.23	.18	.05	.84

Each analysis is indexed with the mineral (o) and core (c), rim (r-c) (core to rim traverse) or particular notes (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample # and placement	Averages										Standard deviations													
	C	R	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnC	MgO	CaO	Na <sub>2</sub> O	Total	Mg#	SiO <sub>2</sub> , TiO <sub>2</sub> , Al <sub>2</sub> O <sub>3</sub> , Cr <sub>2</sub> O <sub>3</sub> , FeO	MnO	CaO	Na <sub>2</sub> O	Total Mg#						
BGS-4-2-4d al cpx r(expl)-r(intpl)	x	x	52.18	.62	2.83	.19	6.52	.19	15.47	22.50	.37	100.9	80.9	.81	.20	.58	.03	.34	.02	.33	1.0	.09	.77	.79
BGS-4-2-4d a2 cpxc(8)	x		52.44	.55	2.72	.14	7.05	.18	15.87	21.66	.33	100.9	80.0	.23	.11	.42	.03	.52	.02	.39	2.48	.08	1.17	2.86
BGS-4-2-4d a2-cpxr(p)(5)	x	x	52.00	.92	2.76	.12	6.71	.19	15.02	22.16	.35	100.2	80.0	.45	.10	.27	.01	.11	.01	.18	1.12	.05	1.19	.42
BGS-10-1-a 31-36 cpxc(5)	x	x	49.61	.86	3.46	.06	9.48	.22	14.27	20.99	.60	98.7	72.9	.57	.02	.12	.01	.50	.02	.23	.32	.02	.50	1.34
BGS-10-1-a 31-36 cpxc(6)	x	x	51.58	.78	2.67	.05	8.90	.26	14.88	20.04	.54	99.7	74.9	.71	.11	.43	.01	.69	.02	.129	.05	.42	1.48	
BGS-16-1-1-7-12 cpxc(10)	x	x	50.78	.89	3.10	.06	12.07	.35	14.66	17.13	.59	99.6	69.0	.77	.19	.56	.02	3.72	.06	1.71	.574	.17	.39	3.66
BGS-16-1-1-7-12 cpxr(9)	x	x	51.56	.73	2.25	.05	10.88	.35	13.91	19.98	.43	100.1	69.6	.28	.09	.10	.01	1.28	.05	.62	1.70	.03	.52	1.62
BGS-16-1-1-7-12 cpxr(28)	x	x	51.42	.66	1.98	.04	11.78	.37	13.31	19.66	.50	99.7	66.8	.46	.06	.17	.01	.67	.02	.33	.64	.06	.50	1.06
BGS-17-1-1-A 12-cpx1(9)	x	x	53.56	.44	2.39	.03	8.35	.20	15.41	20.05	.32	100.8	70.0	.25	.24	.02	.03	2.34	.03	.50	3.70	.17	1.66	4.64
BGS-19-2-10-11-5-cpx(5)	x	x	51.79	.94	2.74	.11	6.80	.17	15.27	22.32	.29	100.4	80.0	.64	.10	.08	.01	.16	.03	.09	.24	.04	.55	.43
BGS-19-2-10-11-5-cpx(16)	x	x	52.96	.53	2.95	.12	6.67	.18	16.46	21.03	.28	101.2	81.5	.54	.04	.05	.01	.59	.02	.77	1.26	.07	.72	.75
BGS-19-2-12a 38.5-cpx(14)	x	x	54.21	.58	2.18	.13	6.32	.19	15.76	22.36	.30	102.0	81.6	.26	.12	.35	.02	.16	.03	.11	.24	.06	.56	.46
BGS-19-2-12a 38.5-cpx(16)	x	x	52.36	.53	2.48	.17	6.10	.18	15.69	22.61	.33	100.4	82.1	.69	.08	.22	.02	.10	.02	.10	.39	.05	.62	.26
BGS-17-1-1-A 12-cpx1(9)	x	x	52.81	.65	2.79	.15	6.01	.17	15.54	22.60	.34	101.0	82.2	.54	.08	.18	.02	.13	.01	.16	.21	.02	.32	.40
BGS-19-2-12a 38.5-cpx(25)	x	x	53.18	.70	3.44	.12	5.92	.16	15.57	21.87	.46	101.4	82.4	.54	.15	.95	.05	.11	.03	.12	.81	.26	.27	.32
BGS-19-2-12a 38.5-cpx(36)	x	x	52.95	.38	1.66	.01	7.81	.21	14.71	22.21	.43	100.4	77.0	.42	.10	.30	.01	.26	.03	.17	.32	.04	.31	.70
BGS-20-1-3B-28.5-32 cpxr(7)	x	x	51.31	.44	1.76	.08	6.96	.25	13.99	21.17	.64	96.6	78.2	.61	.06	.16	.01	.11	.02	.32	.10	.06	1.66	.65
BGS-20-1-3B-28.5-32 cpx1c(2)	x	x	50.22	.37	1.43	.06	7.37	.25	14.92	20.98	.68	96.3	78.3	.08	.02	.04	.01	.37	.02	.16	.28	.05	.64	.67
BGS-20-1-3B-28.5-32 cpx1cr(2)	x	x	52.47	.84	2.51	.04	9.43	.28	15.89	19.07	.38	100.9	75.0	.43	.08	.16	.02	.91	.02	.80	.25	.10	.60	.82
BGS-20-1-6a al cpxc(7)	x	x	51.53	.71	2.50	.03	8.33	.24	14.85	21.00	.43	99.6	76.1	.88	.07	.47	.02	.64	.04	.31	.79	.11	.61	1.01
BGS-20-1-6a al cpxr(6)	x	x	52.06	.87	2.53	.04	9.16	.28	15.53	19.80	.41	100.7	75.2	.39	.07	.14	.02	.84	.04	.70	1.28	.08	.66	1.07
BGS-20-1-6a al r-cpxx(20)	x	x	52.00	.72	2.30	.03	7.87	.27	15.02	22.23	.43	101.1	77.3	.27	.07	.10	.02	.22	.01	.07	.33	.08	.36	.50
BGS-20-1-6a-67 cpx1c-r(15)	x	x	52.36	.94	2.50	.10	7.95	.27	15.26	20.48	.00	99.9	77.4	.36	.04	.20	.01	.57	.03	.53	1.06	.00	.54	.67
BGS-20-1-6a-67 cpx2h(5)	x	x	52.74	.46	2.17	.02	7.30	.20	14.67	21.91	.41	99.9	78.2	.28	.11	.12	.01	.14	.07	.09	.18	.02	.31	.33
BGS-20-1-6a a3 cpxr(9)	x	x	52.07	.70	2.29	.03	7.83	.25	15.07	21.94	.40	100.6	77.4	.49	.09	.22	.01	.27	.04	.13	.26	.06	.65	.55
BGS-20-1-6a a3 cpxc(14)	x	x	51.31	.88	2.48	.00	9.41	.21	14.83	20.20	.48	99.8	73.8	.81	.21	.27	.01	.19	.04	.56	2.53	.05	.69	1.59
BGS-20-2-6b al cpxc(10)	x	x	50.87	.77	2.41	.00	8.43	.21	14.21	21.67	.49	99.1	75.0	.10	.07	.20	.01	.30	.02	.57	.42	.11	1.94	.58
BGS-20-2-6b al cpxr(10)	x	x	51.38	.95	2.38	.00	9.33	.21	15.40	20.09	.44	100.2	74.6	.28	.05	.13	.00	.87	.04	.66	1.43	.04	.28	.97
BGS-20-2-6b a2 cpxc(10)	x	x	52.10	.55	2.60	.01	7.94	.18	15.05	22.62	.40	100.9	77.2	.24	.10	.17	.01	.39	.02	.20	.53	.04	.35	.72
BGS-20-2-6b a2 cpxr(8)	x	x	52.06	.70	2.23	.00	8.26	.19	14.82	22.49	.44	101.2	76.2	.29	.05	.07	.01	.22	.03	.17	.60	.02	.20	.34
BGS-20-2-6b a2 cpxr(p)(8)	x	x	52.38	.41	1.75	.02	7.89	.22	15.31	22.37	.47	100.8	77.6	.48	.13	.29	.01	.40	.01	.22	.56	.05	.47	1.05
BGS-20-2-6b 0-5 cpxc(8)	x	x	51.37	.93	2.45	.03	9.03	.24	14.72	19.53	.55	98.8	74.4	.53	.03	.14	.01	.42	.02	.46	.66	.03	.68	.49
BGS-20-2-6b 0-5 cpxr(8)	x	x	51.29	.78	2.43	.03	8.55	.23	14.18	21.31	.61	99.4	74.7	.61	.11	.25	.01	.35	.03	.40	.43	.10	.63	1.17
BGS-23-2b 23-28 a1 cpx1(8)	x	x	52.44	.46	1.74	.00	10.07	.32	14.49	20.38	.40	100.3	71.9	.23	.03	.11	.00	.60	.02	.28	.02	.08	.19	.03
BGS-23-2b 23-28 a1 cpxr(p)(8)	x	x	51.69	.83	2.43	.00	11.83	.35	14.86	18.09	.46	100.6	69.1	.29	.04	.09	.00	.43	.02	.16	.61	.05	.32	.55
BGS-23-2b 23-28 a1 cpxc(7)	x	x	52.22	.61	2.03	.00	10.44	.34	14.35	19.71	.44	100.1	71.0	.42	.08	.14	.00	.56	.03	.21	.94	.05	.78	.82
BGS-23-2b 23-28 a1 cpxr2(p)(5)	x	x	52.16	.61	2.02	.01	10.41	.32	14.55	21.64	.43	101.0	71.1	.54	.07	.09	.01	.47	.01	.35	1.17	.03	.54	1.10
BGS-23-1-2b a2 cpxc(8)	x	x	53.08	.37	1.54	.02	9.70	.32	14.55	21.52	.42	101.5	72.8	.16	.06	.10	.02	.28	.02	.08	.19	.03	.34	.49
BGS-23-1-2b a2 cpxr(6)	x	x	52.62	.49	1.78	.02	10.05	.32	14.46	21.08	.42	101.2	72.0	.35	.06	.10	.01	.54	.02	.22	.68	.03	.44	.80
BGS-23-1-2b 23-28 a3 cpxr(p)(6)	x	x	51.95	.73	2.38	.00	10.31	.26	14.39	20.36	.41	100.8	71.3	.45	.06	.26	.01	.26	.03	.11	.36	.04	.29	.38
BGS-23-1-2b 23-28 a3 cpxc(9)	x	x	52.28	.65	2.28	.00	10.08	.26	14.17	20.70	.41	100.8	71.5	.59	.09	.33	.01	1.00	.04	.44	1.74	.06	.64	1.32
BGS-23-1-2b 23-28 a3 cpxr(o)(6)	x	x	52.69	.52	2.05	.01	9.42	.24	14.10	21.42	.38	100.8	72.7	.53	.08	.25	.01	.47	.03	.21	.69	.03	.56	.76
BGS-23-1-2b 23-28 a3 cpx-c-r(10)	x	x	51.08	.67	2.34	.00	10.36	.26	14.37	20.35	.40	100.8	71.2	.43	.10	.21	.01	.32	.02	.08	.19	.03	.34	.49
BGS-23-1-2b 23-28 a4 cpxc(12)	x	x	51.61	.70	2.07	.02	11.26	.35	13.35	19.92	.53	99.8	67.9	.67	.06	.08	.02	.37	.02	.43	.73	.04	.41	.80
BGS-23-1-2b 23-28 a4 cpxr(p)(8)	x	x	51.27	.60	2.23	.02	11.44	.36	13.33	20.07	.47	99.8	67.6	.31	.08	.55	.01	1.14	.03	.22	1.13	.03	.44	2.20
BGS-23-1-2b 23-28 a4 cpxr(p)(8)	x	x	51.33	.73	2.26	.02	10.37	.31	14.10	20.37	.43	99.9	70.8	.76	.06	.14	.01	.79	.02	.36	.40	.04	.84	1.78
BGS-23-1-2b 23-28 a4 cpxr(p)(9)	x	x	51.36	.63	2.41	.02	10.26	.30	14.20	20.53	.42	100.1	71.2	.17	.01	.03	.03	.41	.03	.41	.93	.04	.12	2.55

Each analysis is indexed with the mineral (o) and core (c), lim (l) - core to rim (r-c), and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample # and placement	Averages										Standard deviations												
	C	R	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnC	MgO	CaO	Na <sub>2</sub> O	Total	Mg#	SiO <sub>2</sub> , TiO <sub>2</sub> , Al <sub>2</sub> O <sub>3</sub> , Cr <sub>2</sub> O <sub>3</sub>	FeO	MnO	MgO	CaO	Na <sub>2</sub> O	Total Mg#			
BGS-23-1-2b 29-34 a2 cpxct(10)	x	53.06	.31	1.33	.01	9.15	.28	14.29	21.69	.38	100.5	73.6	.40	.05	.16	.01	.27	.02	.17	.25	.04	.67	.64
BGS-23-1-2b 29-34 a2 cpxt(10)	x	52.07	.31	1.33	.01	10.14	.32	13.96	21.02	.44	100.3	71.1	.59	.06	.21	.00	.31	.03	.19	.34	.04	.78	.65
BGS-23-1-2b 29-34 a3 cpxct(10)	x	52.76	.48	1.51	.02	9.56	.30	14.36	21.46	.36	100.8	72.8	.51	.27	.14	.01	.24	.04	.15	.24	.02	.66	.52
BGS-23-1-2b 29-34 a3 cpxt(10)	x	52.76	.35	1.50	.01	9.47	.28	14.34	21.19	.36	100.3	73.0	.43	.07	.22	.01	.24	.03	.11	.38	.04	.47	.57
BGS-23-1-2b 29-34 a4 cpxt(pl)-c(1)	x	52.76	.35	1.50	.01	8.69	.24	13.97	20.89	.56	97.5	74.1	.65	.07	.25	.01	.36	.04	.17	.45	.07	1.26	.61
BGS-23-1-2b 29-34 a5 cpxc-t(10)	x	52.17	.39	1.57	.01	9.67	.31	14.04	21.22	.41	100.3	72.1	.41	.10	.20	.01	.39	.02	.13	.26	.04	.74	.89
BGS-23-1-2b 29-34 a5 cpxc-t(p)	x	52.15	.54	1.97	.02	11.30	.35	13.57	20.79	.44	100.8	68.5	.64	.06	.16	.00	.42	.03	.21	.21	.02	.50	1.10
BGS-23-1-2e cpx a1 (froxide)-t(p)	x	52.27	.50	1.73	.00	11.03	.34	13.30	20.48	.47	100.3	68.2	.62	.12	.14	.00	.45	.03	.43	1.37	.05	.83	.52
BGS-23-1-2e a2 cpx cpxt(7)	x	51.81	.63	2.23	.00	10.19	.32	13.59	21.93	.32	100.7	70.4	.57	.22	.71	.00	.63	.02	.20	1.11	.14	.51	1.48
BGS-23-1-2e a2 cpx cpxt(p)(8)	x	52.78	.32	1.24	.00	10.68	.33	13.86	21.23	.37	101.3	69.8	.63	.04	.10	.00	.64	.01	.26	.31	.02	.43	1.65
BGS-23-1-2e a2 cpx cpxt(lm)(3)	x	52.99	.37	1.46	.00	10.96	.33	13.70	20.72	.43	100.7	69.0	.26	.03	.08	.00	.27	.04	.22	.27	.04	.42	.79
BGS-23-1-2e a2 cpx cpxt(mb)(5)	x	52.60	.44	1.52	.00	10.59	.34	13.81	21.20	.41	101.1	69.9	.43	.02	.14	.00	.45	.02	.13	.31	.04	.31	.87
BGS-23-1-2e a3 cpxt(pl)(8)	x	51.56	.55	1.73	.00	11.40	.34	13.32	21.06	.42	100.4	67.5	.210	.07	.29	.00	.1.91	.04	.20	2.46	.07	.94	3.79
BGS-23-1-2e a3 cpxc(8)	x	51.68	.61	2.41	.00	11.96	.31	13.17	19.49	.47	100.1	66.2	.68	.17	.29	.00	.2.09	.04	.23	.27	.04	.86	3.87
BGS-23-1-2e a4 cpxt(pl)(8)	x	51.19	.76	2.13	.00	12.02	.36	13.18	20.49	.52	100.6	61.9	.06	.14	.00	1.83	.03	.40	.68	.03	.86	3.87	
BGS-23-1-2e a4 cpxc(9)	x	51.45	.67	2.38	.00	10.90	.31	13.07	21.32	.63	100.7	68.1	.1.86	.41	.46	.00	.1.68	.01	.89	1.94	.49	.93	4.77
BGS-23-1-2e a4 cpxt(mb)(10)	x	52.60	.51	1.68	.00	10.59	.34	13.81	21.20	.41	101.1	69.9	.32	.02	.14	.00	.45	.02	.21	.38	.04	.37	1.13
BGS-24-1-3a 6-10 cpxc(3)	x	49.04	.85	2.75	.01	12.28	.37	12.94	19.78	.53	98.6	65.3	.32	.02	.07	.01	.39	.04	.17	.02	.52	.66	
BGS-24-1-3a 6-10 cpxt(10)	x	52.54	.42	1.66	.00	9.44	.28	14.13	21.16	.40	100.0	72.7	.47	.05	.16	.00	.15	.01	.13	.28	.02	.75	.32
BGS-24-1-3c21-24 cpxc(3)	x	52.11	.62	2.07	.00	10.21	.28	13.63	20.71	.45	100.1	70.4	.89	.02	.13	.00	.30	.03	.23	.51	.04	1.53	.62
BGS-24-1-3d 21-24 cpx(9)	x	52.76	.78	1.72	.01	10.72	.28	14.55	19.31	.46	101.6	71.0	.04	.25	.63	.01	.274	.05	.96	4.66	.10	.98	2.68
BGS-24-1-3h 47 cpx(4)	x	52.50	.72	2.50	.01	10.53	.30	13.52	21.10	.50	101.7	69.6	.36	.04	.10	.01	.12	.04	.13	.10	.03	.36	.27
BGs-24-1-3h 51.5 cpx(14)	x	53.35	.65	2.66	.02	11.46	.30	14.80	19.03	.40	102.7	69.8	.15	.19	.42	.01	.1.05	.05	.77	.2.10	.08	1.78	1.18
BGS-24-1-3i 60 cpx(6)	x	53.31	.61	2.46	.01	11.03	.32	13.37	21.21	.47	102.8	68.4	.00	.05	.08	.01	.89	.09	.30	.62	.04	.93	2.19
BGS-24-1-3k 77 cpx(4)	x	51.75	.71	2.74	.01	8.10	.24	14.49	20.45	.36	98.8	76.2	.50	.16	.57	.01	.91	.03	.14	.23	.14	.67	2.12
BGS-24-1-3l 21-24 cpx(6)	x	54.16	.70	2.86	.03	7.04	.23	15.36	21.50	.36	102.2	79.5	.31	.03	.12	.01	.17	.03	.32	.31	.04	.56	.38
BGS-24-1-3n 47 cpx(4)	x	54.02	.76	2.50	.03	9.22	.23	18.02	21.08	.31	102.2	77.9	.49	.08	.30	.02	.2.09	.06	.244	4.74	.08	.57	1.88
BGS-27-1-1A 3 cpx(2)	x	53.67	.81	2.69	.05	8.59	.23	17.47	21.00	.32	100.8	78.6	.47	.13	.27	.01	.2.24	.05	.2.32	3.16	.21	3.06	2.36
BGS-27-1-1B 8.5 cpx(9)	x	53.01	.69	2.57	.04	6.89	.19	15.33	22.00	.29	101.0	79.9	.60	.05	.18	.01	.25	.02	.10	.36	.02	.55	.60
BGS-27-1-1B 19.5 cpx(7)	x	52.55	.90	2.77	.06	7.89	.22	15.48	20.57	.35	100.8	77.8	.44	.13	.26	.01	.59	.02	.40	1.29	.05	.94	1.02
BGS-27-1-1b 25.5-31 a1 cpxt(mb) x	x	51.58	.67	2.21	.08	6.62	.18	15.24	21.73	.31	98.6	80.4	.1.62	.12	.62	.02	.60	.03	.51	1.10	.03	1.97	1.96
BGS-27-1-1b 25.5-31 a1 cpxc(10) x	x	52.20	.59	2.66	.13	7.64	.19	17.57	18.69	.31	100.0	80.4	.42	.08	.20	.02	.1.66	.05	.207	3.89	.07	.34	1.39
BGS-27-1-1b 25.5-31 a1 cpxt(pl)interr x	x	51.48	.68	2.51	.10	6.13	.15	15.85	21.60	.35	98.8	82.2	.06	.58	.02	.48	.02	.23	.83	.06	.92	1.06	
BGS-27-1-1b 25.5-31 a2 cpxt(10) x	x	52.28	.60	2.50	.10	7.96	.23	16.46	19.35	.42	99.7	79.3	.30	.08	.01	.52	.03	.54	.95	.04	.31	.59	
BGS-27-1-1b 25.5-31 a2 cpxt(9) x	x	51.96	.78	2.66	.10	7.70	.23	16.46	19.35	.42	99.7	79.2	.25	.02	.07	.01	.29	.03	.28	.54	.03	.35	.45
BGS-27-1-1b 25.5-31 a2 cpxt(9) x	x	50.87	.95	2.34	.10	7.13	.22	15.21	21.58	.45	98.8	79.2	.03	.19	.28	.03	.66	.02	.14	.25	.04	1.65	3.1
BGS-27-1-1b 25.5-31 a3 cpxt(6) x	x	51.44	1.06	2.78	.09	7.22	.25	14.74	21.50	.43	99.5	78.5	.32	.05	.06	.01	.30	.03	.24	.42	.03	.62	.65
BGS-27-1-1d 56-62 cpxc(8) x	x	51.92	1.08	2.67	.10	7.33	.17	15.07	22.84	.36	101.5	78.6	.29	.03	.04	.02	.33	.03	.11	.48	.04	.77	.76
BGS-27-1-1d 56-62 cpxt(9) x	x	51.87	.97	2.38	.08	7.29	.18	14.50	23.41	.36	101.0	78.0	.55	.02	.08	.02	.1.7	.02	.17	.19	.04	.71	.44
BGS-27-1-1d 56-62 cpxt(different)(3) x	x	53.20	.29	.84	.05	10.88	.11	11.64	23.60	.16	101.8	65.4	.37	.40	.84	.01	.3.27	.81	2.33	.94	.15	.26	11.2
BGS-31-1-1c 34.5-39 cpxt(10) x	x	51.97	.63	2.32	.01	8.88	.25	14.18	21.31	.44	100.0	74.0	.82	.08	.41	.01	.24	.03	.23	.92	.06	1.05	.60
BGS-31-1-1c 34.5-39 cpxt(9) x	x	52.40	.44	1.89	.00	8.96	.27	13.85	22.08	.38	100.3	73.4	.58	.09	.36	.01	.20	.03	.16	.39	.04	.40	.55
BGS-31-1a 0.5 cpxt(10) x	x	52.22	.65	2.30	.02	8.27	.26	14.19	22.07	.38	100.3	75.4	.46	.16	.42	.02	.31	.04	.17	.44	.07	.30	.70
BGS-31-1a 0.5 cpxt(9) x	x	51.96	.57	2.31	.00	9.06	.26	14.14	20.94	.49	99.7	73.6	.75	.22	.58	.01	.41	.04	.24	1.29	.08	1.09	1.03

BR-Cores

BR-5-1-B4.5-8.5 cpxt(10)

BR-5-1-B-4.5-8.5 cpxt(10)

BR-5-1-H 74-78 cpxt(10)

Each analysis is indexed with the mineral (o) and core (c), rim (r-c) (core to rim traverse) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample # and Placement	Averages										Standard deviations										
	C	R	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	Fe <sub>2</sub> O <sub>3</sub>	MnC	MgO	CaO	Na <sub>2</sub> O	Total	Mg# <sup>a</sup>	SiO <sub>2</sub> , TiO <sub>2</sub> , Al <sub>2</sub> O <sub>3</sub> , Cr <sub>2</sub> O <sub>3</sub> , Fe <sub>2</sub> O <sub>3</sub> , MnO, MgO	CaO	Na <sub>2</sub> O	Total Mg#				
BR-5-1-1H 74-78 cpxr(10)	x	51.79	.78	2.72	.03	.721	.19	14.82	21.75	.41	99.7	78.6	.41	.13	.30	.01	.22	.02	.14	.38	.05
BR-5-1-1H 74-78 cpx	x	51.82	.76	2.90	.02	.705	.20	14.96	21.59	.42	99.7	79.1	.34	.08	.22	.01	.24	.02	.17	.45	.04
BR-5-1-5e 48-51.5 a cpxc(10)	x	51.74	.71	2.82	.03	.646	.16	15.08	22.30	.37	99.7	80.6	.73	.20	.62	.01	.30	.03	.13	1.28	.25
BR-5-1-5e 48-51.5 b cpxr(10)	x	51.44	.63	2.54	.03	.79.5	.20	14.71	20.47	.63	98.6	76.7	.87	.26	.99	.01	.63	.02	.37	1.61	.22
BR-5-1-5e 48-51.5 cpx c-r(8)	x	51.87	.84	2.74	.05	.706	.23	14.55	22.04	.41	99.8	78.6	.34	.02	.09	.01	.25	.02	.14	.21	.03
BR-5-1-5e 48-51.5 cpx	x	51.74	.71	2.82	.03	.646	.16	15.08	22.30	.37	99.7	80.6	.73	.20	.62	.01	.30	.03	.13	1.28	.25
BR-5-1-5e 48-51.5 cpxr	x	51.66	.59	2.34	.03	.782	.20	14.79	20.70	.59	98.7	77.1	.55	.24	.80	.01	.50	.02	.16	1.52	.19
BR-6-1-1G 20-5-24.5 cpxc(10)	x	52.35	.71	2.43	.00	.952	.22	14.98	19.02	.47	99.7	77.7	.1.00	.21	.40	.01	.52	.03	.1.16	2.72	.05
BR-6-1-1G 20-5-24.5 cpxr(10)	x	52.06	.58	2.13	.07	.645	.16	15.01	21.95	.41	98.8	80.6	.66	.14	.55	.02	.29	.01	.15	.84	.16
BR-8-1-5D 34-36 cpxc(10)	x	52.13	.50	2.85	.10	.968	.25	14.27	20.42	.50	100.2	74.4	.32	.11	.28	.01	.29	.05	.24	.69	.07
BR-8-1-5D 34-36 cpxr(10)	x	51.71	.72	2.50	.06	.831	.19	15.59	19.73	.38	99.1	77.0	.49	.04	.09	.01	.89	.03	.104	1.87	.04
BR-8-1-5D 34-36 cpx1(10)	x	51.73	.71	2.39	.06	.831	.19	15.59	19.73	.38	99.1	77.0	.49	.04	.09	.01	.77	.04	.77	1.61	.04
BR-8-1-5D 34-36 cpx2c(10)	x	52.26	.63	2.51	.07	.800	.19	16.09	19.28	.36	99.4	78.2	.47	.05	.09	.01	.77	.04	.55	1.30	.05
BR-8-1-5D 34-36 cpx2r(8)	x	52.06	.58	2.13	.07	.645	.16	15.01	21.95	.41	98.8	80.6	.66	.14	.55	.02	.29	.01	.15	.84	.16
BR-8-1-5G 69-71.5 cpxc(10)	x	52.13	.50	2.85	.10	.750	.17	15.92	19.59	.39	99.4	78.7	.30	.04	.15	.01	.57	.02	.84	1.33	.05
BR-8-1-5G 69-71.5 cpxr(10)	x	51.71	.72	2.50	.06	.831	.20	15.32	19.90	.40	99.1	76.7	.50	.03	.11	.02	.59	.02	.61	1.24	.04
BR-8-2-2.5h 3 cpx1(1)	x	54.62	.39	2.92	.04	.672	.17	15.86	22.49	.24	102.4	80.8	.04	.20	.01	.62	.02	.63	1.32	.07	
BR-8-2-2.5h 3 cpx1(10)	x	51.98	.57	2.92	.10	.749	.20	16.35	20.27	.50	100.4	79.6	.06	.20	.71	.02	.85	.02	.70	1.77	.27
BR-8-2-2.5h 3 cpx2(6)	x	52.44	.66	2.98	.09	.775	.19	16.33	20.19	.43	101.0	79.0	.77	.13	.54	.02	.78	.02	.91	1.43	.18
BR-8-2-2.5m 18.5-23 cpxc(10)	x	51.40	.78	2.57	.06	.659	.16	15.10	21.82	.41	98.9	80.3	.59	.08	.23	.01	.16	.02	.20	.25	.03
BR-8-2-2.5m 18.5-23 cpxr(10)	x	51.26	1.02	2.83	.06	.827	.18	15.77	19.92	.40	99.0	78.8	.31	.11	.14	.01	.59	.04	.47	1.09	.05
BR-9-1-1D 26-28.5 cpxc(10)	x	51.07	.86	2.99	.01	.887	.23	14.12	19.88	.49	98.5	74.5	.1.28	.01	.87	.01	.26	.03	.16	1.91	.20
BR-9-1-1D 26-28.5 cpxr(10)	x	51.59	.60	2.29	.01	.842	.24	13.79	21.35	.38	98.7	74.5	.31	.05	.14	.01	.26	.03	.14	.43	.02
BR-12-1-3B 33-35.5 cpx1r(9)	x	51.19	.54	1.53	.04	.965	.24	13.78	21.68	.45	99.1	71.8	.61	.04	.11	.01	.73	.03	.27	.24	.03
BR-12-1-3B 33-35.5 cpx1c(9)	x	51.55	.63	1.64	.05	.984	.24	13.64	21.65	.45	99.7	71.2	.75	.05	.10	.01	.29	.02	.57	.33	.07
BR-12-1-3B 33-35.5 cpx2c(5)	x	52.25	.35	.75	.02	.1036	.19	13.86	22.02	.42	100.2	70.5	.68	.04	.04	.01	.19	.02	.19	.10	.04
BR-14-1-7a 35-37.5 cpxc1(10)	x	51.72	.76	2.63	.02	.910	.24	15.82	19.27	.43	100.0	75.7	.54	.09	.24	.01	.156	.03	.152	3.20	.07
BR-14-1-7a 35-37.5 cpxr1(10)	x	51.57	.74	2.22	.03	.812	.21	15.12	20.49	.41	98.9	76.1	.26	.09	.19	.02	.75	.03	.78	2.36	.06
BR-14-1-7a 35-37.5 cpx2r(10)	x	51.64	.70	2.21	.03	.845	.24	15.10	20.93	.40	99.7	76.2	.76	.09	.19	.01	.93	.03	.54	1.75	.06
BR-15-3B 17.5-20 cpxc(7)	x	50.82	.74	1.68	.21	7.21	.14	15.87	20.86	.44	98.0	79.7	.80	.51	.10	.02	.67	.02	.39	.72	.03
BR-15-3B 17.5-20 cpxr(1)	x	52.06	.72	1.25	.16	6.70	.11	15.50	22.41	.38	99.3	80.5									
<b>JR31-Dredges</b>																					
JR31-3-51 cpxc(10)	x	50.72	1.37	2.96	.02	10.41	.20	15.48	19.25	.43	100.8	72.6	.14	.03	.03	.02	.45	.03	.20	.65	.04
JR31-3-51 cpxr(10)	x	50.85	1.25	2.85	.02	11.24	.23	15.64	18.36	.45	100.9	71.3	.21	.05	.01	.17	.02	.18	.16	.03	
JR31-9-2 cpxc(10)	x	52.04	.79	2.62	.00	10.27	.24	15.87	19.67	.45	102.0	73.6	.39	.11	.25	.01	.251	.05	.199	3.53	.09
JR31-9-2 cpxr(10)	x	52.17	.66	2.52	.01	.773	.17	15.00	22.24	.39	100.9	77.6	.29	.04	.16	.01	.55	.01	.27	.59	.03
JR31-9-4 cpxc(11)	x	51.74	.80	2.60	.01	.946	.28	14.28	21.70	.51	101.4	72.9	.28	.06	.17	.01	.64	.03	.22	.63	.05
JR31-9-4 cpxr(10)	x	51.88	.56	2.36	.00	9.69	.30	14.15	21.48	.44	100.9	72.2	.52	.09	.59	.00	.39	.04	.17	.56	.06
JR31-12-1 a1 cpx r-r(20)	x	52.55	.54	3.36	.50	4.70	.12	16.46	22.36	.36	100.9	86.2	.28	.13	.15	.08	.35	.03	.52	.80	.03
JR31-12-1 a1 cpxc(12)	x	52.86	.45	3.23	.41	4.90	.11	17.24	21.10	.35	100.7	86.2	.48	.03	.14	.03	.82	.02	1.37	2.20	.04
JR31-12-1 a1 cpxr(12)	x	52.20	.69	3.24	.45	5.42	.13	16.52	21.74	.35	100.7	84.5	.44	.14	.13	.03	.71	.02	.93	1.81	.04
JR31-12-1 a2 cpxr(10)	x	52.47	.64	3.15	.50	5.93	.14	16.85	21.16	.36	101.2	83.5	.24	.15	.10	.04	.47	.02	.74	1.18	.05
JR31-12-1 a2 cpxc(10)	x	52.77	.40	3.25	.74	4.91	.12	17.37	21.19	.34	101.1	86.3	.18	.02	.08	.03	.40	.02	.79	1.29	.03
JR31-12-6 cpx1r(10)	x	52.57	.59	2.93	.33	5.44	.12	16.70	21.78	.23	100.7	84.6	.28	.02	.21	.02	.10	.02	.20	.30	.04
JR31-12-6 cpx2c(10)	x	51.85	.53	3.85	.61	5.30	.11	17.26	20.74	.34	100.6	85.3	.34	.04	.10	.02	.28	.02	.49	.83	.06
JR31-12-6 cpx3r(10)	x	53.22	.35	1.86	.28	5.33	.11	15.89	23.75	.27	101.1	84.2	.37	.09	.45	.04	.34	.02	.21	.93	.11
JR31-12-6 cpx4r(10)	x	51.93	.83	2.87	.21	7.78	.16	16.64	20.04	.31	100.8	79.2	.28	.04	.09	.02	.25	.03	.20	.53	.04
JR31-12-6 cpx5c(10)	x	52.40	.52	3.40	.46	5.39	.12	17.50	20.59	.29	100.6	85.3	.18	.03	.05	.02	.20	.02	.28	.47	.05
JR31-12-6 cpx5r(10)	x																				

Each analysis is indexed with the mineral (c), rim (r) or core (c), and the number of analyses (n) and particular notes, and the ratio to rim (traverse) or particular notes, and the number of analyses (n) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample # and placement	C	R	Averages										Standard deviations											
			SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnC	Mg#	CaO	Na <sub>2</sub> O	Total	Mg#	SiO <sub>2</sub> , TiO <sub>2</sub> , Al <sub>2</sub> O <sub>3</sub> , Cr <sub>2</sub> O <sub>3</sub> , FeO	MnO	CaO	Na <sub>2</sub> O	Total	Mg#					
JR31-12-6 cpx6r(10)	x	52.35	.61	2.54	.23	6.24	.13	15.78	22.55	.28	100.7	81.8	.36	.09	.41	.02	.25	.01	.17	.44	.07	.30	.69	
JR31-12-6 cpx7c(10)	x	51.33	.84	2.84	.14	8.65	.21	14.48	21.88	.33	100.7	74.9	.68	.08	.41	.02	.21	.03	.19	.82	.10	.91	.45	
JR31-12-6 cpx8r(10)	x	51.07	.81	3.36	.10	9.10	.20	14.23	21.20	.44	100.6	73.6	.78	.25	1.00	.01	.64	.05	.26	1.39	.23	.58	1.28	
JR31-12-6 cpx9c(9)	x	53.76	.39	2.70	.48	5.36	.15	17.73	20.44	.25	101.3	85.5	.19	.03	.10	.03	.54	.02	.94	1.44	.08	.59	.60	
JR31-12-6 cpx9r(10)	x	52.49	.69	2.93	.26	5.61	.13	16.33	22.04	.32	100.8	83.8	.28	.03	.11	.01	.19	.02	.30	.69	.06	.50	.29	
JR31-12-6 cpx10c(10)	x	53.10	.45	3.29	.58	4.93	.13	17.13	21.34	.30	101.2	86.1	.18	.02	.07	.02	.20	.02	.29	.68	.04	.44	.37	
JR31-12-8 cpxc(8)	x	52.15	.51	3.24	.44	5.87	.09	16.76	20.94	.39	100.4	83.6	.22	.03	.05	.02	.27	.04	.34	.61	.03	.30	.49	
JR31-12-8 cpxr(p)(10)	x	51.42	1.0	3.29	.23	5.70	.11	16.04	21.74	.45	100.1	83.4	.27	.04	.11	.02	.14	.04	.14	.51	.05	.27	.35	
JR31-12-10 cpxr(p)(10)	x	52.08	.53	2.79	.04	6.86	.17	16.19	20.04	.38	80.7	73.7	.33	.02	.10	.02	.53	.03	.54	.95	.08	.45	.83	
JR31-12-10 cpxr(p)(8)	x	52.55	.45	2.64	.01	10.41	.19	15.40	15.41	.48	97.5	72.6	.14	.44	.20	.75	.01	.51	.02	.90	.82	.15	1.18	2.25
JR31-12-16 cpxc(8)	x	52.44	.66	2.58	.05	7.32	.19	14.97	22.65	.41	101.3	78.5	.63	.07	.20	.01	.26	.02	.24	.27	.06	1.21	.51	
JR31-12-28 cpxc(10)	x	51.63	.65	2.22	.00	10.23	.28	13.59	20.73	.44	99.8	70.5	.32	.09	.29	.00	.60	.02	.20	1.45	.09	.77	1.07	
JR31-12-28 cpxr(p)(8)	x	52.46	.40	1.37	.00	9.98	.30	13.82	20.76	.29	99.4	71.2	.58	.16	.40	.02	.42	.02	.11	.80	.07	.50	.90	
JR31-12-39 cpxr(lm)(7)	x	52.04	.40	1.46	.00	11.93	.32	13.18	19.77	.40	99.5	66.4	.62	.15	.61	.01	1.11	.02	.48	2.31	.17	.85	2.26	
JR31-12-39 cpxc(8)	x	51.35	.36	3.01	.00	17.54	.24	12.82	11.61	.76	97.7	56.6	.64	.04	.05	.95	.00	.66	.05	.64	.57	.21	.58	2.04
JR31-12-39 cpxr(p)(7)	x	51.46	.43	1.77	.00	11.78	.32	12.93	19.54	.52	97.8	66.2	.50	.10	.43	.00	1.23	.06	.34	1.73	.13	.70	2.39	
JR31-12-42 cpxc(10)	x	52.24	.50	2.65	.04	6.37	.13	17.16	18.97	.33	98.4	82.8	.31	.02	.04	.01	.28	.02	.42	.61	.03	.43	.39	
JR31-12-42 cpxr(p)(10)	x	51.88	.63	2.36	.03	8.10	.20	16.30	18.58	.35	98.4	78.3	.42	.05	.13	.02	1.04	.04	1.16	2.30	.07	.48	.99	
JR31-12-46 cpxc(5)	x	51.47	.51	2.68	.03	7.58	.16	17.37	17.74	.56	98.1	80.3	.71	.10	.63	.02	1.09	.05	1.00	1.41	.17	1.45	3.01	
JR31-12-46 cpxr(p)(7)	x	51.26	.87	2.78	.02	8.01	.19	14.73	20.19	.43	99.4	76.5	.67	.08	.19	.02	.25	.03	.25	.94	.14	1.26	1.00	
JR31-12-47 cpxl(10)	x	51.48	.73	2.48	.21	7.22	.17	14.81	21.90	.44	99.4	78.5	.37	.05	.09	.02	.25	.02	.19	.50	.06	.50	.66	
JR31-12-50 cpxc(10)	x	51.62	.45	2.85	.27	6.42	.17	16.63	19.75	.42	98.6	82.2	.36	.03	.26	.01	.65	.02	.75	1.15	.13	.62	1.14	
JR31-12-50 cpxr(p)(8)	x	51.16	.82	2.84	.10	6.19	.16	16.09	21.04	.39	98.8	82.3	.40	.06	.18	.03	.44	.03	.47	.84	.07	.40	.79	
JR31-12-68-cpx(large)c(10)	x	52.14	.52	2.87	.15	6.78	.16	16.73	19.96	.40	99.7	81.5	.34	.03	.18	.02	.27	.03	.22	.69	.03	.62	.57	
JR31-12-68-cpx(small)c(10)	x	51.18	.83	2.99	.03	8.03	.19	14.56	21.31	.48	99.6	76.6	.22	.08	.09	.01	.27	.02	.17	.35	.05	.35	.42	
JR31-12-68 cpxr(p)(10)	x	51.86	.60	2.90	.08	7.90	.17	15.25	20.23	.51	99.5	77.7	.38	.20	.79	.01	1.21	.04	.55	2.74	.10	.90	2.46	
JR31-12-71 cpxc(8)	x	52.24	.71	2.59	.08	9.36	.24	15.19	20.63	.40	101.4	74.3	.35	.05	.07	.02	.45	.03	.45	.83	.05	.40	.61	
JR31-12-71 cpxr(p)(8)	x	52.23	.77	2.95	.05	9.13	.23	15.02	20.88	.36	101.0	74.7	.53	.15	.28	.01	1.61	.05	1.09	2.42	.08	.57	1.71	
JR31-12-73 cpxc(10)	x	52.31	.79	2.31	.02	10.05	.28	14.13	21.56	.41	101.9	71.5	.24	.04	.10	.01	.58	.03	.26	.77	.06	.24	.81	
JR31-12-73 cpxr(p)(8)	x	52.32	.74	2.03	.01	10.61	.30	14.10	20.99	.33	101.4	70.3	.17	.05	.14	.01	.27	.03	.32	.56	.07	.30	.40	
JR31-12-73 cpxr(29)	x	52.20	.77	2.13	.01	10.50	.27	14.62	20.17	.39	101.1	71.0	.07	.30	.90	.01	1.21	.04	.81	.89	.07	.55	3.33	
JR31-12-75 cpxc(10)	x	51.68	.55	1.67	.00	13.38	.34	13.61	18.49	.36	100.1	64.6	.22	.07	.15	.01	.95	.03	.79	2.57	.07	.27	2.08	
JR31-12-75 cpxr(p)(8)	x	51.87	.48	1.58	.00	12.34	.31	13.42	19.84	.33	100.2	66.0	.33	.09	.13	.00	.72	.03	.18	.93	.03	.31	1.06	
JR31-12-76 cpxr(p)(10)	x	51.88	.50	1.72	.00	12.06	.34	13.24	19.87	.35	100.0	66.2	.23	.06	.13	.00	.54	.04	.13	.80	.03	.55	1.14	
JR31-12-76 cpxr(6)	x	51.92	.32	2.03	.00	11.07	.31	14.64	20.50	.41	101.2	72.1	.23	.03	.05	.01	.16	.01	.16	.43	.07	.37	.89	
JR31-12-77 cpxc(10)	x	51.45	.57	1.92	.00	13.43	.32	12.66	20.50	.43	101.3	62.7	.11	.08	.37	.00	.91	.02	.39	.77	.06	.48	1.55	
JR31-12-77 cpxr(p)(9)	x	51.96	.45	1.42	.00	14.68	.33	13.45	18.51	.35	101.2	62.3	.21	.06	.12	.00	2.55	.05	.67	.89	.06	.55	2.74	
JR31-12-78 cpxr(7)	x	52.27	.77	2.23	.00	10.96	.32	14.41	19.99	.37	101.3	70.1	.22	.05	.07	.00	.79	.02	.44	1.06	.09	.23	.89	
JR31-12-78 cpxr(p)(10)	x	52.20	.82	1.20	.01	12.01	.34	15.09	18.57	.31	101.6	69.2	.42	.17	.14	.01	1.40	.04	.73	2.11	.04	.91	1.56	
JR31-12-78 cpxr(6)	x	52.04	.90	2.26	.03	10.10	.31	14.64	20.50	.41	101.2	72.1	.23	.03	.05	.01	.16	.01	.16	.43	.07	.37	.88	
JR31-12-79-8 cpxr(6)	x	52.18	.91	2.35	.02	10.11	.28	14.50	20.24	.41	101.0	71.9	.26	.09	.24	.01	.43	.04	.14	.63	.03	.45	.86	
JR31-12-79-8 cpxr(p)(10)	x	52.06	.82	2.29	.02	11.24	.31	15.18	18.79	.39	101.1	70.9	.45	.13	.31	.01	2.63	.05	.91	4.33	.14	.52	1.94	
JR31-12-79-8 cpxr(4)	x	52.42	.69	2.07	.02	11.49	.34	15.59	18.15	.41	101.2	71.3	.39	.11	.32	.01	4.32	.08	3.24	7.11	.13	.59	2.46	
JR31-12-79-8 cpxr(p)(10)	x	51.16	.65	1.86	.01	15.43	.38	13.55	17.53	.28	100.9	61.9	.55	.17	.47	.01	5.56	.08	1.72	6.37	.15	.72	4.41	
JR31-12-79-8 cpxr(6)	x	51.62	.63	2.06	.05	12.40	.34	13.57	19.76	.31	100.7	66.2	.18	.06	.15	.01	1.33	.04	.52	1.40	.08	.47	1.41	
JR31-12-79-8 cpxr(10)	x	52.22	.39	1.63	.04	11.53	.31	13.66	21.12	.28	101.2	67.9	.38	.05	.25	.01	.49	.02	.16	.37	.05	.24	1.14	
JR31-12-79-8 cpxr(p)(10)	x	51.34	.67	1.88	.00	14.53	.35	13.36	18.53	.33	101.0	62.3	.06	.08	.26	.00	2.70	.05	.92	3.10	.10	.96	3.22	

Each analysis is indexed with the mineral (of) and core (c), rim (r) (core to rim traverse) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Sample # and placement	C	R	Averages	Mg#												Standard deviations								
				SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnC	MgO	CaO	Na <sub>2</sub> O	Total	SiO <sub>2</sub> , TiO <sub>2</sub> , Al <sub>2</sub> O <sub>3</sub> , Cr <sub>2</sub> O <sub>3</sub> , FeO	MnO	MgO	CaO	Na <sub>2</sub> O	Total	Mg#				
JR31-39-10 cpx(r)(p)(6)	x	x	51.96	.50	1.64	.00	12.84	.35	13.28	20.28	.25	101.1	64.9	.15	.04	.00	1.08	.03	.44	1.29	.04	.43	1.08	
JR31-39-11 cpx(c)(10)	x	x	51.17	.71	1.96	.02	14.13	.38	13.85	17.68	.33	100.2	65.2	.67	.12	.20	.01	4.25	.07	1.12	4.85	.14	.57	2.01
JR31-39-11 cpx(r)(10)	x	x	51.85	.61	1.79	.00	12.92	.35	13.69	19.35	.32	100.9	65.4	.26	.05	.11	.01	1.29	.03	.53	1.63	.04	.39	1.82
JR31-39-11 cpx(inside mt)r(10)	x	x	51.90	.45	1.74	.00	12.20	.34	13.81	19.71	.36	100.5	67.8	.42	.10	.14	.00	2.05	.04	.49	3.58	.06	.76	1.10
JR31-39-11 cpx(inside mt)c(10)	x	x	51.50	.64	1.97	.00	13.49	.38	13.89	18.23	.40	100.5	64.7	.20	.05	.19	.00	1.80	.04	.71	2.29	.05	.27	1.96
JR31-39-13 cpx(c)(10)	x	x	51.61	.71	2.01	.03	12.03	.31	13.40	20.39	.45	100.9	66.5	.08	.06	.10	.00	.82	.02	.31	.57	.01	.50	.99
JR31-39-13-cpx(r)(10)	x	x	50.56	.40	3.47	.03	18.55	.22	12.01	12.87	.55	98.7	53.6	.16	.10	.16	.00	1.11	.00	.56	1.64	.13	.76	2.65
JR31-39-17 cpx1(inside mt)r(15)	x	x	52.20	.63	1.79	.00	13.25	.39	13.94	18.91	.37	101.5	65.7	.27	.13	.32	.01	3.86	.09	1.28	4.92	.10	.45	3.36
JR31-39-17 cpx2(outside mt)r(p)(10)	x	x	51.92	.61	2.00	.01	12.23	.36	13.65	20.19	.34	101.3	66.6	.39	.05	.11	.01	1.21	.03	.36	1.54	.05	.47	1.62
JR31-39-17 cpx3c(8)	x	x	52.47	.43	1.58	.01	11.64	.35	13.58	21.05	.35	101.5	67.5	.22	.04	.05	.01	.46	.04	.17	.69	.06	.47	1.11
JR31-39-17 cpx3r(9)	x	x	52.48	.41	1.46	.00	11.00	.33	13.61	20.84	.43	100.6	68.8	.41	.05	.11	.01	.44	.02	.13	.48	.08	1.05	.76
JR31-40-1 cpx(c)(10)	x	x	53.11	.58	2.96	.04	7.57	.14	18.17	19.08	.41	102.1	81.2	.37	.06	.21	.01	1.58	.03	2.07	3.67	.10	.31	1.34
JR31-40-1 cpx(r)(10)	x	x	53.19	.64	2.38	.04	6.66	.14	16.41	21.99	.42	101.9	81.5	.20	.04	.07	.01	.44	.04	.30	.75	.04	.31	.73
JR31-41-4 cpx(c)(10)	x	x	52.39	.58	2.25	.01	12.36	.27	15.44	18.29	.46	102.0	69.8	.62	.20	.66	.01	5.07	.12	2.83	7.84	.20	.41	3.86
JR31-41-4 cpx(r)(10)	x	x	52.24	.54	2.26	.00	9.48	.23	14.26	22.34	.46	101.8	72.8	.57	.01	.06	.00	.45	.01	.26	.43	.03	.37	1.02

Each analysis is indexed with the mineral (c), core (c), rim (r) (core to rim traverse) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 2C: The average composition of low-Ca pyroxene from on and around Atlantis Bank.

Sample # and placement	Averages										Standard deviations											
	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnO	MgO	CaO	Na <sub>2</sub> O	Total	Mg#	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnO	MgO	CaO	Na <sub>2</sub> C	Total	Mg#
Mode 98	52.20	.42	1.08	.02	20.88	.50	19.75	5.24	.10	100.2	63.6	.77	.12	.40	.01	5.93	.13	3.01	8.42	.19	.66	4.64
6K-460-14 opx(c)(10)	52.80	.31	.84	.01	23.17	.55	21.19	1.49	.02	100.4	62.0	.57	.06	.04	.01	.36	.02	.40	.11	.03	.88	.66
6K-460-14 opx c-(8)	52.89	.36	.84	.02	23.31	.54	21.01	1.55	.01	100.5	61.6	.44	.03	.04	.01	.18	.02	.17	.08	.01	.63	.23
6K-460-14 opx(r)(10)	51.40	.32	.83	.00	25.34	.67	19.47	1.12	.01	99.2	57.8	.17	.11	.02	.00	.11	.04	.30	.12	.01	.21	.47
6K-466-8 al opx c-(8)	51.65	.35	.78	.02	25.14	.64	19.66	1.41	.17	99.8	58.2	.17	.03	.07	.01	.37	.04	.18	.20	.06	.40	.44
6K-466-8 al opx (inside cpx) c-(7)	52.59	.22	.73	.01	24.89	.66	20.86	1.04	.01	101.0	59.9	.49	.04	.02	.02	.23	.03	.22	.11	.02	.60	.12
6K-466-8 a2 opx (inside cpx) c-(7)	51.19	.30	.67	.03	24.86	.60	18.51	1.85	.18	99.2	57.0	1.48	.08	1.32	.01	1.04	.06	1.54	.137	.27	1.66	1.13
6K-466-11 opx(4)	52.54	.35	.89	.02	23.33	.58	20.65	1.57	.00	99.9	61.2	.29	.02	.03	.01	.47	.02	.20	.39	.00	.72	.28
6K-466-11 opx(c)(6)	52.14	.32	.82	.02	23.88	.60	20.37	1.31	.00	99.5	60.3	.18	.03	.03	.01	.54	.02	.16	.12	.00	.77	.38
6K-467-12 opx r-c-(pl)	53.23	.45	1.35	.00	18.72	.42	23.86	1.84	.04	99.9	69.4	.36	.09	.07	.01	.25	.03	.26	.39	.02	.47	.25
Leg ABCDE																						
6K-645-R18 opx(5)	52.44	.43	.78	.02	25.77	.40	19.73	1.81	.01	101.4	57.7											
6K-649-R06-2 opx(4)	52.58	.44	2.41	.03	16.95	.34	24.98	1.08	.01	98.8	72.4											
6K-643-r08 opx(r)(7)	53.82	.47	1.69	.00	16.73	.35	25.62	2.26	.05	101.0	73.2	.64	.04	.04	.01	.17	.02	.43	.12	.05	1.08	.36
6K-643-r08 opx(c)(9)	54.00	.43	1.70	.00	16.73	.39	25.85	2.13	.03	101.3	73.4	.22	.03	.05	.00	.30	.04	.34	.58	.03	.42	.21
BGS Cores																						
BGS-16-1-1 7-12 opxc	52.37	.36	1.02	.01	23.79	.59	20.36	1.61	.05	100.2	60.4	.36	.08	.08	.01	.45	.03	.18	.40	.01	.75	.42
BGS-16-1-1 7-12 opxr	52.39	.34	.92	.02	23.76	.61	20.36	1.66	.04	100.1	60.4	.30	.04	.05	.01	.36	.06	.13	.19	.01	.49	.41
BGS-16-1-1 7-12 opx c-(r)(10)	52.30	.33	.75	.00	24.08	.67	20.10	1.63	.00	99.9	59.8	.24	.03	.05	.00	.34	.02	.16	.18	.01	.39	.39
BGS-20-1-6a a2 opx r(cpx)-r(pl)(6)	54.26	.36	1.15	.03	17.68	.42	26.35	1.24	.00	101.5	72.7	.22	.11	.08	.01	.17	.04	.34	.26	.00	.30	.32
BGS-20-1-6a a3 opx r(ol)-r(pl)(12)	53.94	.40	1.24	.03	18.23	.45	25.76	1.40	.01	101.5	71.6	.22	.07	.04	.01	.30	.03	.14	.24	.02	.44	.31
BGS-20-2-6b a2 opx (inside ol)(6)	53.94	.13	1.02	.03	18.03	.37	25.91	.80	.09	100.3	71.9	.28	.02	.06	.02	.14	.03	.11	.07	.04	.21	.16
BGS-23-1-2e opx(9)	52.85	.19	.77	.00	21.25	.49	22.65	1.28	.03	99.5	65.5	.39	.03	.05	.00	.39	.04	.15	.07	.03	.53	.40
BGS-24-1-3c 21-24 opx(11)	52.79	.26	.88	.00	22.67	.56	22.08	.98	.03	100.2	63.4	.20	.02	.02	.00	.21	.05	.06	.01	.02	.25	.23
BGS-24-1-3c 21-24 opx(r)(7)	52.77	.26	.86	.00	22.67	.55	22.10	.98	.02	100.2	63.5	.99	.04	.11	.00	.28	.05	.11	.07	.02	.92	.38
BGS-24-1-3h 51.5 opx(6)	52.69	.39	1.12	.00	20.73	.52	22.27	1.58	.00	99.3	65.7	.32	.03	.01	.40	.03	.20	.30	.01	.50	.34	
BGS-24-1-3k 77 opx(6)	55.17	.29	.94	.00	18.92	.45	25.14	1.12	.00	102.0	70.3	.47	.10	.11	.01	.36	.03	.19	.44	.00	.96	.48
BGS-27-1-1A 3 opx(6)	56.67	.28	1.13	.01	16.12	.37	27.55	.88	.00	103.0	75.3	.54	.06	.11	.01	.25	.01	.35	.11	.00	.75	.37
BR Cores																						
BR-6-1-1g 20.5-24.5 opx c-(r)(7)	51.80	.29	1.52	.00	9.18	.20	14.85	21.67	.02	99.5	74.3	1.04	.05	.13	.00	1.47	.03	.17	.50	.01	.48	3.00
BR-8-1-5d 34-36 opx c-r	54.68	.16	.97	.00	16.33	.37	26.25	.83	.01	99.6	74.1	.42	.03	.12	.00	.30	.02	.21	.04	.01	.16	.44
BR-8-1-5g 69-71.5 opx r-r	53.66	.29	1.27	.00	16.74	.37	25.49	1.07	.04	98.9	73.1											
BR-8-2-5h opx(2)	54.50	.22	1.14	.00	15.51	.32	27.58	.95	.00	100.2	76.0	.11	.02	.01	.00	.58	.02	.08	.01	.00	.33	.44
BR-8-2-5m 18.5-23 opx c-(r)(8)	53.80	.27	1.28	.01	14.29	.31	27.89	.84	.01	98.7	77.7	.11	.05	.15	.01	.42	.03	.91	.10	.01	1.14	.71
JR31 Dredges																						
JR31-12-6 opx c-(r)(10)	53.93	.51	1.47	.05	18.01	.39	24.88	1.96	.00	101.2	71.1	1.31	.13	.06	.02	1.43	.03	.38	.32	0	0.43	0.74
JR31-12-28 opx(c)(10)	52.39	.39	.94	.00	22.25	.51	21.51	2.05	.05	100.1	63.3	.19	.12	.08	0	.37	.04	.31	.58	.04	.37	.4
JR31-12-28 opx(r)(9)	52.23	.39	1.00	.00	22.50	.52	21.37	1.67	.06	99.8	62.9	.41	.03	.05	.01	.62	.03	.26	.33	.04	.76	.71
JR31-12-28 opx(5)	52.45	.38	.93	.00	22.43	.50	21.31	1.70	.04	99.7	62.9	.37	.05	.07	0	.55	.03	.26	.3	.04	0.79	0.48
JR31-12-45 opx(6)	53.51	.35	1.33	.00	17.35	.37	24.41	1.78	.04	99.1	71.5	.48	.08	.01	0	.6	.05	.33	.34	.03	0.2	0.83
JR31-39-4 opx(6)	51.88	.32	.81	.01	25.39	.63	19.42	1.54	.00	100.0	57.7	.13	.04	.03	.01	.12	.03	.1	.15	0	0.17	0.18
JR31-39-17 opx c-(r)(6)	51.28	.33	.74	.01	24.48	.61	19.79	1.68	.02	98.9	59.0	.19	.05	.04	.01	.46	.03	.14	.19	.01	0.44	0.49

Each analysis is indexed with the mineral (ol) and core (c), rim (r) r-c (core to rim traverse) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

Table 2D: The average composition of olivine from on and around Atlantis Bank.

Label	Averages										Standard deviations											
	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnO	MgO	CaO	NiO	Total	Fo	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnO	MgO	CaO	NiO	Total	Fo
Mode 98																						
6K-460-R11 ol-r(4)	36.04	.00	.01	.00	34.55	.64	29.10	.05	.06	100.4	60.0	.21	.00	.01	.00	.66	.03	.57	.02	.03	1.07	.64
6K-460-R13 olc(6)	37.17	.04	.00	.02	29.60	.44	33.27	.05	.00	100.6	66.7	.08	.01	.00	.00	.42	.03	.15	.01	.00	.54	.26
6K-460-R13 ol(9)	37.14	.03	.00	.02	29.50	.45	33.26	.04	.00	100.4	66.8	.19	.01	.00	.01	.39	.04	.44	.01	.00	.85	.23
6K-460-R14 ol1 c-r(5)	34.87	.00	.02	.02	40.36	.60	24.11	.03	.00	100.0	51.6	.11	.00	.02	.00	.42	.01	.16	.01	.00	.37	.35
6K-460-R14 ol2 c-r(9)	35.65	.00	.01	.00	37.33	.60	26.59	.01	.00	100.2	55.9	.22	.00	.00	.00	.31	.03	.18	.01	.00	.37	.26
6K-460-R16 olc(7)	35.77	.00	.00	.00	34.80	.56	28.58	.03	.00	99.7	59.4	.12	.00	.00	.01	.57	.03	.23	.00	.00	.68	.31
6K-460-R16 ol(r7)	35.95	.00	.00	.00	35.13	.56	28.83	.04	.00	100.5	59.4	.37	.00	.00	.00	.38	.04	.09	.01	.00	.47	.27
6K-466-R7 ol(r4)	34.44	.00	.02	.00	41.79	.83	22.68	.06	.11	99.9	49.2	.08	.01	.01	.01	.39	.05	.16	.01	.07	.30	.36
6K-466-R8 a2 ol-r(r10)	33.99	.05	.00	.07	42.86	.75	21.89	.09	.18	99.9	47.6	.30	.01	.01	.02	.29	.04	.31	.02	.04	.60	.37
6K-466-R8 a2 r(p1)-r(mb)(10)	34.03	.07	.00	.07	44.38	.74	21.18	.06	.21	100.7	46.0	.42	.05	.00	.01	.27	.03	.43	.02	.03	.73	.59
6K-466-R9 olc(6)	36.57	.05	.00	.03	31.69	.51	31.36	.06	.00	100.3	63.8	.17	.01	.00	.01	.28	.04	.26	.00	.00	.36	.19
6K-466-R9 ol(r4)	36.72	.04	.00	.03	31.68	.47	31.16	.06	.00	100.2	63.7	.16	.02	.00	.01	.16	.04	.46	.01	.00	.52	.37
6K-466-R9 ol(r4)	36.72	.00	.00	.01	31.01	.48	31.79	.04	.02	100.1	64.6	.48	.00	.00	.00	.78	.03	.71	.01	.02	.64	1.05
6K-467-R9 olc-r(10)	37.46	.01	.01	.00	27.63	.42	34.79	.04	.07	100.4	69.2	.18	.02	.01	.01	.20	.04	.12	.01	.01	.10	.19
6K-467-R1 ol(6)	36.56	.00	.00	.00	29.54	.41	33.35	.05	.05	100.0	66.8	.64	.00	.00	.00	.39	.01	.20	.01	.02	.95	.27
6K-467-R10 ol(6)	38.64	.02	.01	.00	20.82	.30	40.44	.05	.08	100.4	77.6	.07	.00	.01	.01	.19	.02	.62	.01	.04	.48	.40
<b>Leg ABCDE</b>																						
6K-646-R01 ol(6)	37.62	.01	.38	.01	21.74	.36	39.44	.02	.11	99.7	76.4	.15	.01	.45	.01	.34	.04	.38	.01	.02	.44	.27
6K-646-R01 ol2 c-r(6)	37.49	.01	.07	.00	21.59	.33	39.70	.05	.03	99.3	76.6	.30	.01	.07	.00	.47	.02	.21	.01	.03	.76	.35
6K-649-R06 ol(2)	36.61	.00	.63	.00	27.48	.42	33.87	.07	.04	99.1	68.7	.04	.00	.72	.00	.44	.06	.59	.03	.02	.71	.72
6K-649-R08 ol-r(c6)	37.33	.02	.00	.01	27.69	.43	35.95	.04	.16	101.6	69.8	.13	.02	.00	.01	.41	.07	.15	.03	.03	.55	.28
6K-649-R08 ol(26)	38.59	.01	.03	.00	20.47	.42	40.89	.03	.01	100.4	78.1	.52	.01	.01	.00	.24	.03	.62	.01	.02	.67	.28
6K-649-R14 ol(6)	36.75	.01	.28	.00	25.48	.37	36.19	.04	.04	99.2	71.7	.35	.01	.45	.01	.33	.04	.27	.03	.02	.49	.33
6K-650-R06 ol c-r(7)	38.20	.00	.49	.00	18.25	.31	41.63	.03	.04	99.0	80.3	.57	.00	.40	.01	.30	.03	.63	.01	.03	.74	.42
6K-650-R09 ol(9)	37.80	.01	.00	.02	25.22	.56	38.56	.03	.00	102.2	73.2	.34	.01	.01	.02	.33	.04	.24	.01	.00	.27	.24
<b>BGS Cores</b>																						
BGS-4-1-4c 43.48 ol(6)	37.26	.03	.01	.02	26.54	.36	36.44	.02	.10	100.8	71.0	.25	.01	.01	.01	.12	.03	.18	.02	.02	.21	.18
BGS-4-2-17 ol(6)	37.47	.00	.02	.00	27.18	.42	35.37	.05	.12	100.4	69.9	.17	.01	.01	.00	.31	.02	.26	.01	.02	.48	.16
BGS-4-2-4d al1 (pl int)-r(pl ext)(10)	37.47	.00	.00	.00	26.29	.35	35.97	.03	.02	100.1	70.9	.28	.00	.00	.00	.94	.06	.18	.01	.02	1.00	.30
BGS-4-2-4d a2 al1 (pl)-r(lim)(10)	37.33	.00	.00	.00	26.39	.36	35.96	.04	.01	100.1	70.8	.13	.00	.00	.00	.33	.04	.17	.01	.01	.30	.30
BGS-16-1-17-12 ol	32.30	.33	.75	.00	24.08	.67	20.10	.63	.00	99.9	59.8	.24	.03	.05	.00	.34	.02	.16	.01	.01	.39	.39
BGS-19-2-10-11-15 ol(6)	38.38	.00	.00	.00	22.47	.31	38.67	.02	.07	99.9	75.4	.24	.00	.01	.00	.36	.03	.32	.01	.02	.73	.22
BGS-20-1-6a 3 ol r-r(6)	36.82	.00	.03	.00	29.88	.44	33.47	.02	.09	100.8	66.6	.20	.00	.00	.02	.19	.04	.31	.01	.08	.42	.13
BGS-20-1-6a 64-67 ol2 c-r(7)	38.23	.09	.08	.10	27.42	.45	34.60	.08	.00	101.0	69.2	.26	.00	.01	.01	.32	.02	.18	.01	.00	.46	.27
BGS-20-2-6b a2 ol r-c(10)	36.69	.00	.00	.00	29.84	.46	32.72	.03	.06	99.8	66.1	.19	.01	.00	.00	.40	.05	.37	.02	.03	.48	.49
BGS-24-1-3c 21 ol c-r(6)	35.71	.04	.00	.07	35.94	.59	38.22	.07	.00	100.6	58.3	.19	.01	.00	.01	.19	.02	.14	.01	.00	.31	.14
BGS-27-1-1a 3 ol(2)	36.71	.01	.00	.03	24.21	.34	37.26	.00	.00	98.5	73.3	.18	.01	.00	.01	.37	.01	.15	.00	.00	.37	.38
BGS-27-1-1b 8.5 ol1(4)	36.86	.02	.01	.03	25.60	.39	35.58	.03	.00	98.5	71.2	.30	.01	.01	.02	.28	.03	.25	.01	.00	.33	.35
BGS-27-1-1b 8.5 ol2(10)	37.32	.01	.00	.02	25.23	.38	36.44	.04	.00	99.4	72.0	.46	.01	.00	.01	.25	.03	.31	.01	.00	.57	.25
BGS-27-1-1b 19.5 ol(6)	37.51	.01	.03	.03	25.14	.38	36.85	.02	.00	99.9	72.3	.30	.01	.01	.02	.27	.02	.63	.01	.00	.78	.46
<b>BR Cores</b>																						
BR-8-1-5D 34-36 olc(5)	37.51	.00	.00	.00	26.64	.35	34.57	.03	.00	99.1	69.8	.06	.00	.00	.00	.13	.02	.14	.01	.00	.23	.12
BR-8-1-5D 34-36 olr(6)	37.51	.00	.00	.00	26.97	.36	34.47	.03	.00	99.3	69.5	.42	.00	.00	.00	.69	.02	.23	.01	.00	.63	.58
BR-8-1-5D 34-36 ol2(3)	37.13	.00	.00	.00	27.56	.40	33.71	.00	.00	98.8	68.6	.17	.00	.00	.00	.08	.04	.19	.00	.00	.32	.10
BR-8-1-5D 34-36 ol2c(5)	37.36	.00	.00	.00	26.25	.36	34.72	.03	.00	98.7	70.2	.08	.00	.00	.00	.40	.01	.13	.01	.00	.43	.29

Each analysis is indexed with the mineral (ol) and core (c), rim (r) or traverse (tr) of particular notes, and the number of analyses (#) and are accompanied by standard deviations.

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Label	Averages										Standard deviations											
	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnO	MgO	CaO	NiO	Total	Fo	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Cr <sub>2</sub> O <sub>3</sub>	FeO	MnO	MgO	CaO	NiO	Total	Fo
BR-8-1-5g olc(5)	37.13	.00	27.65	.37	34.15	.03	.00	99.3	.26	.00	.41	.03	.13	.03	.00	.62	.00	.33	.00	.00	.33	
BR-8-1-5g olc(5)	36.88	.00	27.44	.39	34.05	.01	.00	98.8	.24	.00	.43	.04	.39	.01	.00	.40	.00	.00	.40	.00	.00	.56
BR-8-2-5h ol(3)	36.23	.00	.01	24.97	.35	36.78	.03	.06	98.4	.24	.00	.25	.05	.39	.01	.01	.14	.01	.01	.14	.01	.02
BR-8-2-5h ol(24)	36.75	.02	.01	25.83	.43	35.97	.02	.06	99.1	.24	.02	.01	.00	.17	.05	.28	.01	.01	.33	.01	.01	.29
BR-15-3b 17.5-20 ol c-r(7)	38.79	.07	.22	.09	22.16	.36	38.15	.06	.00	99.9	.14	.00	.39	.01	.64	.02	.93	.01	.00	.00	.00	.42
JR31 Dredges																						
JR31-9-2 olc-r(6)	36.52	.01	.01	31.48	.41	32.32	.07	.09	100.9	64.7	.34	.01	.00	.66	.04	.79	.02	.02	.02	.02	.02	.98
JR31-9-4 olc(r5)	35.76	.01	.01	35.33	.46	28.97	.06	.02	100.6	59.4	.15	.01	.01	.65	.05	.20	.01	.01	.70	.01	.01	.44
JR31-12-1 al ol(10)	38.41	.02	.00	20.11	.25	40.43	.05	.07	99.3	78.2	.24	.01	.00	.41	.03	.22	.04	.02	.02	.02	.02	.30
JR31-12-1 a2 ol(10)	38.83	.02	.00	20.05	.25	40.78	.02	.07	100.0	78.4	.22	.02	.00	.30	.02	.34	.01	.02	.02	.02	.02	.25
JR31-12-50 ol c-r(10)	37.59	.00	.00	02	24.27	.37	37.24	.06	.08	99.6	73.2	.15	.00	.01	.63	.02	.26	.02	.03	.03	.02	.51
JR31-39-1 ol(6)	36.10	.02	.00	.00	32.55	.48	31.02	.07	.04	100.3	62.9	.11	.02	.00	.34	.02	.21	.02	.02	.02	.02	.34
JR31-39-3 ol(8)	35.99	.00	.00	.01	36.33	.53	27.91	.07	.00	100.9	57.8	.11	.01	.01	.48	.04	.49	.01	.00	.00	.00	.54

Each analysis is indexed with the mineral (ol) and core (c), rim (r) r-c (core to rim traverse) or particular notes, and the number of analyses (#) and are accompanied by standard deviations.

## Chapter 6

### Trace element mineral-chemistry of gabbros from Atlantis Bank: Evidence for melt-rock interaction

#### Abstract

Trace elements in augite from gabbros on and around Atlantis Bank reflect melts that crystallized 50-95% relative to a parental magma in equilibrium with the mantle. This agrees well with our previous results using major-elements in Chapter 3. The gabbros cannot have formed by simple fractional crystallization alone, as the rocks have increasing LREE/HREE ratios with fractionation. The gabbros may instead have crystallized by in-situ crystallization, or by significant assimilation of pre-existing olivine gabbros. Even if the magmas have assimilated host-rocks upon ascent, they still must have crystallized ~50% elsewhere before emplacement in the crust. Our previous estimate for the original crustal thickness at Atlantis Bank of 4.4-km is therefore valid.

Augites from a diabase-dike represent much more depleted melts than the melts that produced the gabbros. No known mineral-fractionation can produce the refractory pattern seen in this rock and we suggest that this dike represents depleted melts that did not mix with the melts found on and around Hole 735B.

#### 1 Introduction

Lower ocean crust gabbros are believed to form from the same basaltic melts that are parental to the upper crust. Trace- and rare-earth element analyses are frequently used to interpret the origin and evolution of such oceanic lavas. The lower crust, however, is more rarely analyzed for REE and trace element variability due both to limited accessibility and because it consists of cumulus and inter-cumulus minerals with potentially very TE- and REE-rich accessory-minerals. For the Atlantis Bank gabbros, an extensive major and trace element data-set now exists in compilations of whole-rock analyses from Hole 735B and Atlantis Bank area (Hart et al., 1999; Natland and Dick, 2002; and Coogan et al., 2001).

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The method of accretion of the lower ocean-ridge crust has long been debated, though fractional crystallization is considered the main process for the evolution of the upper crust. *In-situ* crystallization (Langmuir, 1989), assimilation-fractionation-crystallization (AFC, e.g. DePaolo 1981 and discussion in Chapter 3 and 4), and magma-mixing may also be important in the evolution of the lower crust. *In-situ* mineral-analyses of oceanic gabbros may help in interpreting the origin and evolution of the melts that formed lower crustal rocks. Coogan et al. (2000) showed that very primitive Mid-Atlantic Ridge olivine gabbros have plagioclase-grains with cores that have a different origin than that of the overgrowing plagioclase rims and interstitial clinopyroxene, and interpreted this as the result of magma mixing within the lower crust. Dixon et al. (1986) investigated an unusual aphyric ferrobasalt containing evolved gabbro-xenoliths from the Juan de Fuca Ridge, and concluded that the basalt represents hybrid magma formed by highly evolved magma, mixed with more primitive melts. The evolved melt-component produced the oxide-gabbros by filter pressing along a partially crystallized mush boundary. Coogan et al. (2002) demonstrated that East Pacific Rise gabbros exposed at Hess Deep represent a complex mixture of melts, some of which are mixed in the mantle, some of which are not. They suggested that all the melts added to the crust do not mix within a central axial magma chamber, and some partially crystallize in isolation within the lower crust. The homogeneity within the upper ocean crust is therefore surprising, as the existing knowledge of accretion of lower ocean-crust suggests a very heterogeneous process of accretion.

We have analyzed augites in eight Atlantis Bank gabbros for trace- and rare-earth elements. The sample set includes the entire compositional range of gabbros recovered from Atlantis Bank, from very primitive troctolites in Hole 735B to ferrogabbros recovered along the side of the platform, and one coarser-grained diabase sample from the northern wall (see Chapter 5). This data-set allows us to re-evaluate the magma-budget for Atlantis Bank when assimilation is taken into account.

## 2 Analytical methods

Unaltered cores and rims of clinopyroxene from eight samples were analyzed on the CAMECA IMS 3f ion probe at Woods Hole Oceanographic Institution using the

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**Table 6-1: REE and Trace elements content of augites in ppm**

	Ti	V	Cr	Sr	Y	Zr	La	Ce	Nd	Sm	Eu	Dy	Er	Yb
ODP-735B-83-7-139 rim to plagioclase	4282	370	3000	7.7	21.7	39.3	.45	3.27	4.80	2.94	.69	4.36	2.57	2.71
ODP-735B-83-7-139 core	1590	230	3764	4.7	6.5	3.5	.23	1.10	1.53	1.09	.41	1.90	1.14	1.24
ODP-735B-83-7-139 rim to olivine														
JR31-12-6 grain 1 rim	2593	340	939	10	11.9	7.8	.57	3.29	4.54	2.70	.55	4.53	2.58	2.65
JR31-12-6 grain 1 core	2424	315	1739	12.3	13.9	10.2	.24	1.25	2.21	1.34	.46	2.33	1.35	1.24
JR31-12-6 grain 1 in between	2212	301	1476	10.6	12	8.3	.44	1.88	2.30	1.53	.64	2.77	1.70	1.62
JR31-12-6 grain 2 rim	3056	338	1776	12.2	19.1	22.1	.24	1.09	1.58	.96	.45	1.81	1.03	.93
JR31-12-6 grain 2 core	1686	264	787	6.4	7.5	4.6	.023	.128	.292	.222	.072	.540	.675	.945
JR31-12-28 rim	2547	198	112	0.9	5.7	4.6	.036	.234	.507	.423	.134	1.025	.811	1.150
JR31-12-28 core	4187	200	109	1.9	5.4	5.3	.030	.171	.387	.274	.072	.635	.632	.844
JR31-12-28 repeat	3207	323	1227	7.7	20.7	44	.30	1.65	2.86	1.86	.51	3.41	1.99	1.98
JR31-12-1 rim	1876	267	1848	8.3	9.1	7.1	.19	1.06	1.67	1.26	.41	1.90	1.13	1.21
JR31-12-1 core	2732	369	175	8.9	61.5	107.4	.190	9.20	12.38	5.82	1.28	9.82	5.71	5.89
6K-466-R11 rim	3307	452	253	9.5	76.3	135.1	1.84	10.46	13.56	6.32	1.31	11.29	6.42	6.75
6K-466-R11 core	4196	429	147	10.1	28.4	25.9	.40	2.52	3.76	1.85	.65	3.37	2.00	2.10
6K-467-R09 rim	3027	395	145	9.5	26	28.8	.43	2.79	4.72	2.37	.83	4.57	2.62	2.51
6K-467-R09 core	3086	477	98	9.6	58.1	52.9	1.79	9.83	12.59	5.98	1.65	11.03	6.65	6.56
6K-466-R08 rim1	5009	560	133	12.2	68.1	54.1	1.78	8.48	11.06	5.43	1.66	9.81	5.54	5.99
6K-466-R08 core	3820	525	112	10.5	61.9	58	1.43	7.33	9.73	4.85	1.36	8.59	5.42	5.44
6K-466-R08 rim2	3208	428	82	10.5	69.5	65	1.52	7.74	10.99	5.92	1.49	10.92	6.77	6.42
6K-466-R08 rim3	4246	398	99	9.9	54.3	94.7	1.62	7.85	9.79	4.88	1.26	8.37	4.64	4.66
JR31-39-8 rim	3430	389	80	10.3	52.9	113.9	1.85	8.27	10.40	5.55	1.41	8.40	4.71	5.09
JR31-39-8 core	3706	396	138	9.4	34.9	52.4	.78	3.96	5.23	2.71	.67	4.46	2.55	2.85
6K-460-R13 rim	2705	303	116	8.2	23.9	31.2	.52	2.27	2.65	1.61	.46	2.31	1.51	1.20
6K-460-R13 core1														
6K-460-R13 core2														

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methods of Shimizu and Hart (1982). A primary beam of O<sup>-</sup> ions was focused to ~20μm for REE (La, Ce, Nd, Sm, Eu, Dy, Er, Yb) and ~10μm for other trace elements (Ti, V, Cr, Sr, Y, Zr). Molecular interferences were eliminated by energy-filtering and a secondary voltage offset of -30 to -60V for the REE, and -90 for the other trace elements. Uncertainties based on counting statistics were 5-10% (1σ) for REE and 1-5% (1σ) for the other trace elements. The data is presented in Table 6-1.

### 3 Results

The spider-diagrams and REE-plots for the augites are shown in Figure 6-1 and 6-2 respectively. It is clear that there are several populations. The diabase (JR31-12-28) has a very steep, refractory pattern, different from the rest. Almost all of the augites have negative Eu-anomalies, indicating that the augite crystallized after substantial volume of plagioclase had been fractionated. The REE-concentrations have been inverted using the partition-coefficients from Table 6-2 in order to present the data as the hypothetical equilibrium melts for the augites (Bédard, 1984)(Figure 6-4B).

### 4 Discussion

Except for the refractory diabase pattern, melt produced by 5.5 % by batch equilibrium melting in the spinel facies of depleted upper mantle (DMM) provides an adequate, sub-parallel, fit to the melts compositions calculated from the most primitive gabbroic augites (Figure 6-2B and Table 6-2). We used the DMM-composition and the partition-coefficients of Workman (in prep). Robinson et al. (2001) showed that basalts representing the upper crust formed contemporaneously to Atlantis Bank likely formed by fractional melting of a mantle with a potential temperature as high as that of the Mid-Atlantic Ridge at 23°N or Juan de Fuca Ridge at 45°N. They found that the difference in trace-element compositions and melt-production is due to a 20-km conductive lid on the mantle in this region, causing a larger fraction of melting to occur in the garnet stability-field below the dry solidus. Thus, the melts that are produced have slightly higher La/Sm-values than normal MORB even if the isotopes are those of a normal MORB (Robinson et al., 2001; Snow, 1993). The calculated crustal thickness as inverted from trace-elements corresponds to 3±1km of crust (Robinson et al., 2001), shallower than

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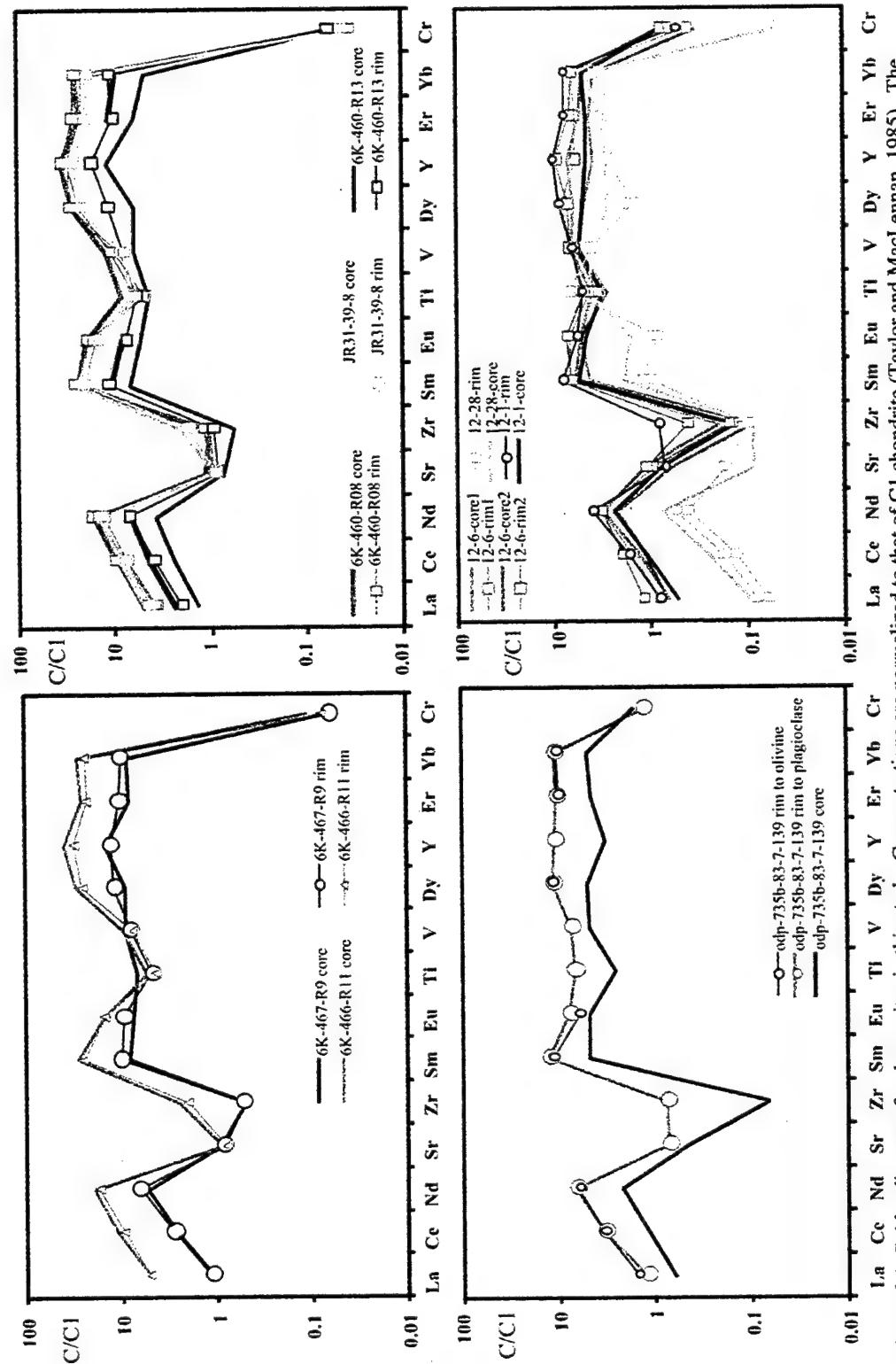


Figure 6-1: Spider diagrams for the augites in this study. Concentrations are normalized to that of C1 chondrite (Taylor and MacLennan, 1985). The sample-areas are described in Chapter 5.

Evolution of Oceanic Gabbros: In-situ and Ancient Examples

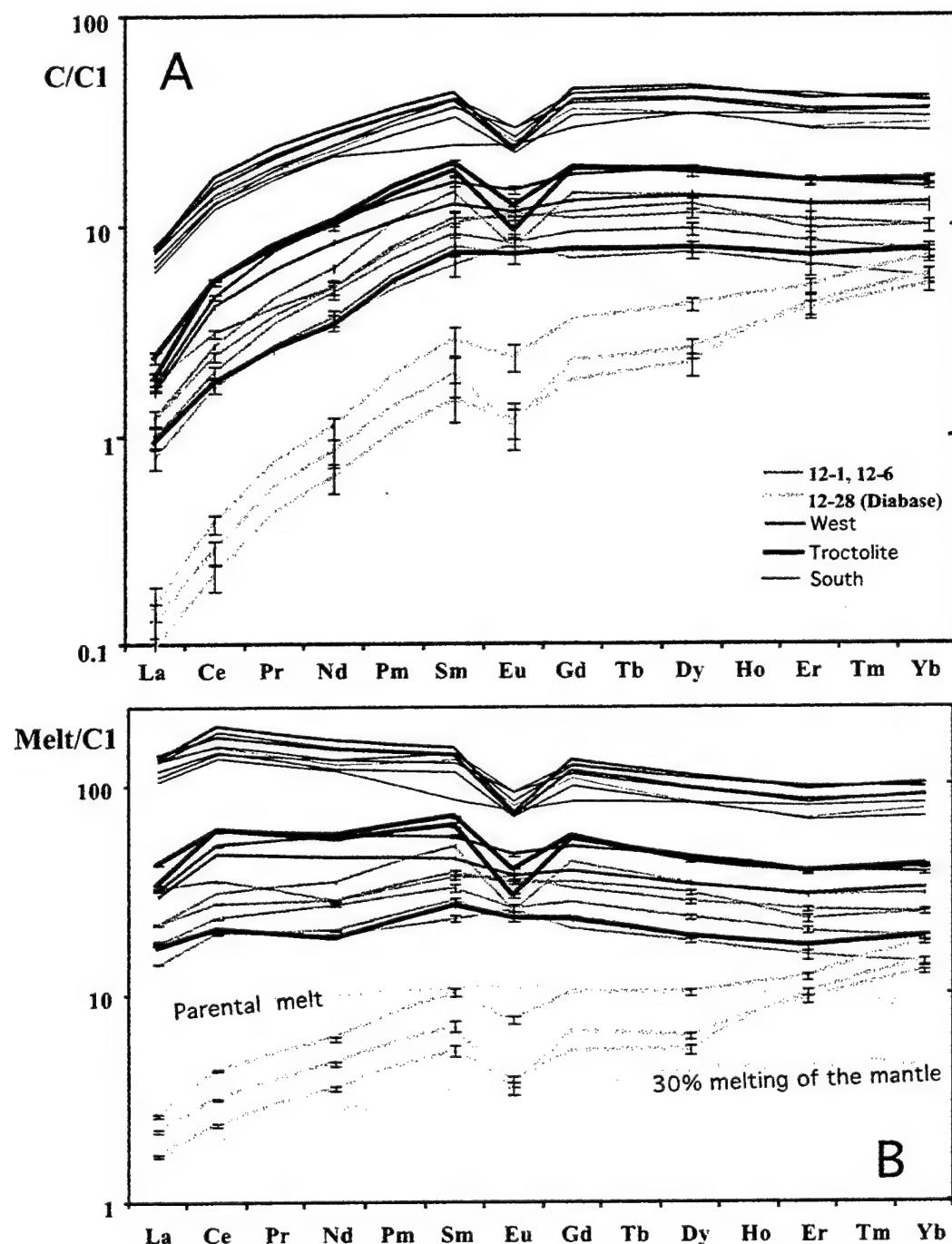


Figure 6-2: A: REE-plots for the augites in this study. The error-bars indicate  $1\sigma$ . B: The hypothetical melts in equilibrium with the gabbro augite-crystals. Two parental melts, made from 5.5 and 30% batch equilibrium melting of the depleted upper mantle (Workman, in prep) are also shown. The partition coefficients used are from Bédard (1994).

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what was found by Muller et al. (1997) by seismic surveying, if one assumes that the seismic Moho represents the crust-mantle boundary.

Inspection of Figure 6-2A shows three main groups of augite-crystals: the JR31-12-28, diabase with very low concentrations; the primitive and intermediate gabbros, which show equilibrium melt REE-patterns nearly parallel to the parental melt; and the very evolved rocks with a more LREE-enriched pattern. Figure 6-3 shows the evolution of the liquid-lines of descent for magmas calculated for 1) fractional crystallization (i.e. Rayleigh fractionation); 2) In-situ crystallization (Langmuir, 1989) where magmas crystallize and create a mush along the walls of the magma-chamber, and varying fractions of the interstitial melt ( $f$ ) returns to the main magma body; and 3) Assimilation-Fractionation-Crystallization (AFC; DePaolo, 1981) by assimilation of a typical Atlantis Bank olivine-gabbro (Table 7-2). The liquid-lines of descent represent 0-90% crystallization relative to the primitive MORB calculated from DMM. The bulk-partition coefficients are calculated assuming fractionation of 35% augite, 15% olivine and 50% plagioclase for all stages of crystallization. Figure 6-3A shows that in-situ crystallization where a small fraction of the interstitial melt is returned to the magma-body, or AFC, where the mass assimilated is similar to the mass crystallized, describe the evolution of the augites better than simple fractional crystallization. Figure 6-3B suggests that none of these models work very well, and that the gabbros have complex, multi-stage evolutions than can be modeled for these processes. The best fitting AFC and In-situ-model are further investigated in Figure 6-4. If AFC or in-situ fractionation controlled the melt evolution, then the range of trace-elements concentrations formed in the gabbros represents >50% - 95% crystallization. This is consistent with the model result developed in Chapter 3, and the melt budget of more than 4.4-km crust produced at Atlantis Bank, in contrast to the conclusions of Robinson et al. (2001).

The very refractory-looking diabase augite is not explained by the any of the fractionation models, even though the augites are fairly evolved with Mg#’s as low as 75 (see Figure 5-11C). In fact, it is almost impossible to produce such a primitive, depleted composition by accumulated melt of a MORB-source mantle. Figure 6-2B shows a melt produced by 30% melting of primitive upper mantle, and although the composition has as low concentration as JR31-12-28, it is nowhere near as depleted in LREE relative to

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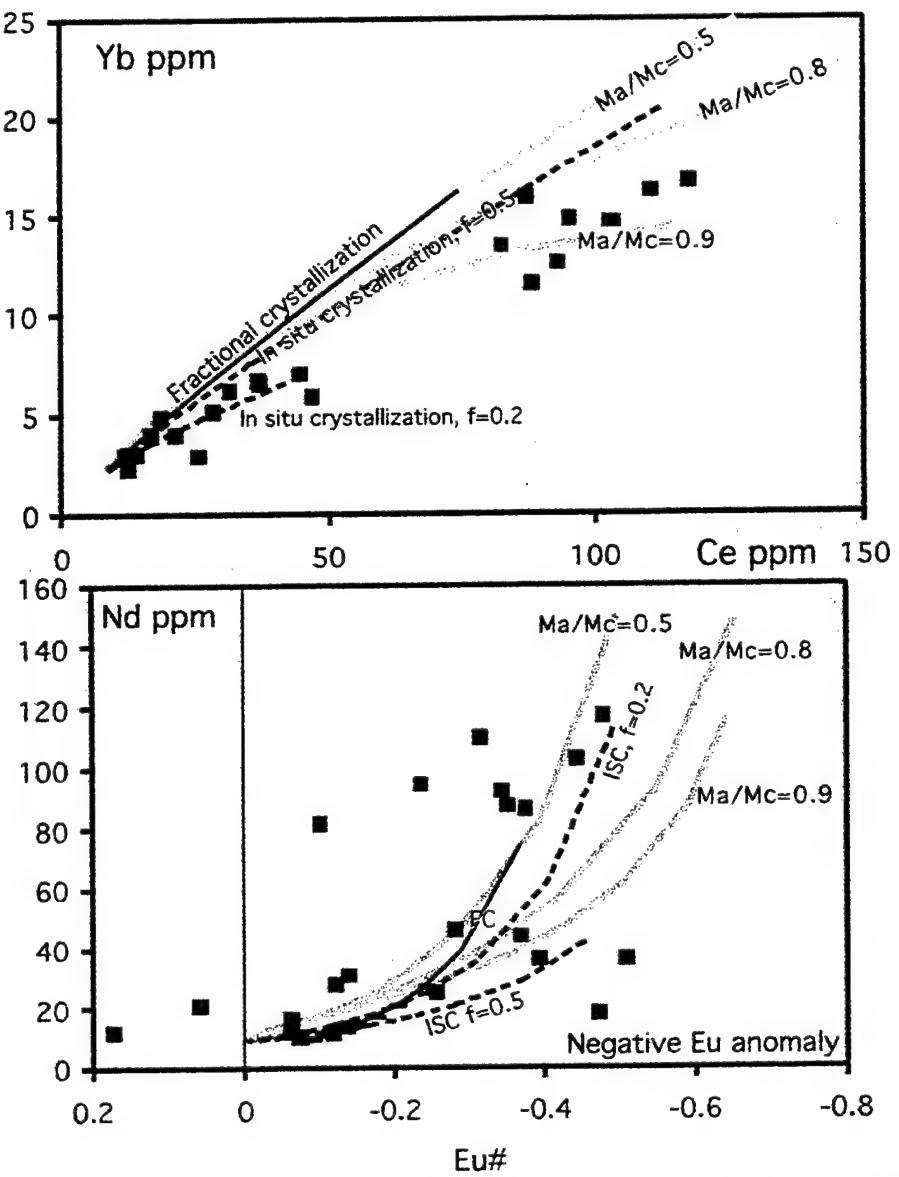


Figure 6-3: Crystallization models for our samples compared to the hypothetical melts in equilibrium with the augites. A: It appears that the AFC-models and in-situ fractionation fit the augite-melts better than a pure fractional crystallization model. "Ma"=Mass assimilated; "Mc"=Mass crystallized; "f" = fraction of melt returned from the mush-zone to the magma chamber;  $\text{Eu}\# = (\text{Eu}^* - (\text{Sm}^* + \text{Gd}^*) \times 0.5) - 1$ , where  $\text{Eu}^*$ ,  $\text{Sm}^*$  and  $\text{Gd}^*$  represent chondrite-normalized values. Negative Eu#'s are negative Eu-anomalies, indicating fractionation of plagioclase from the melt.

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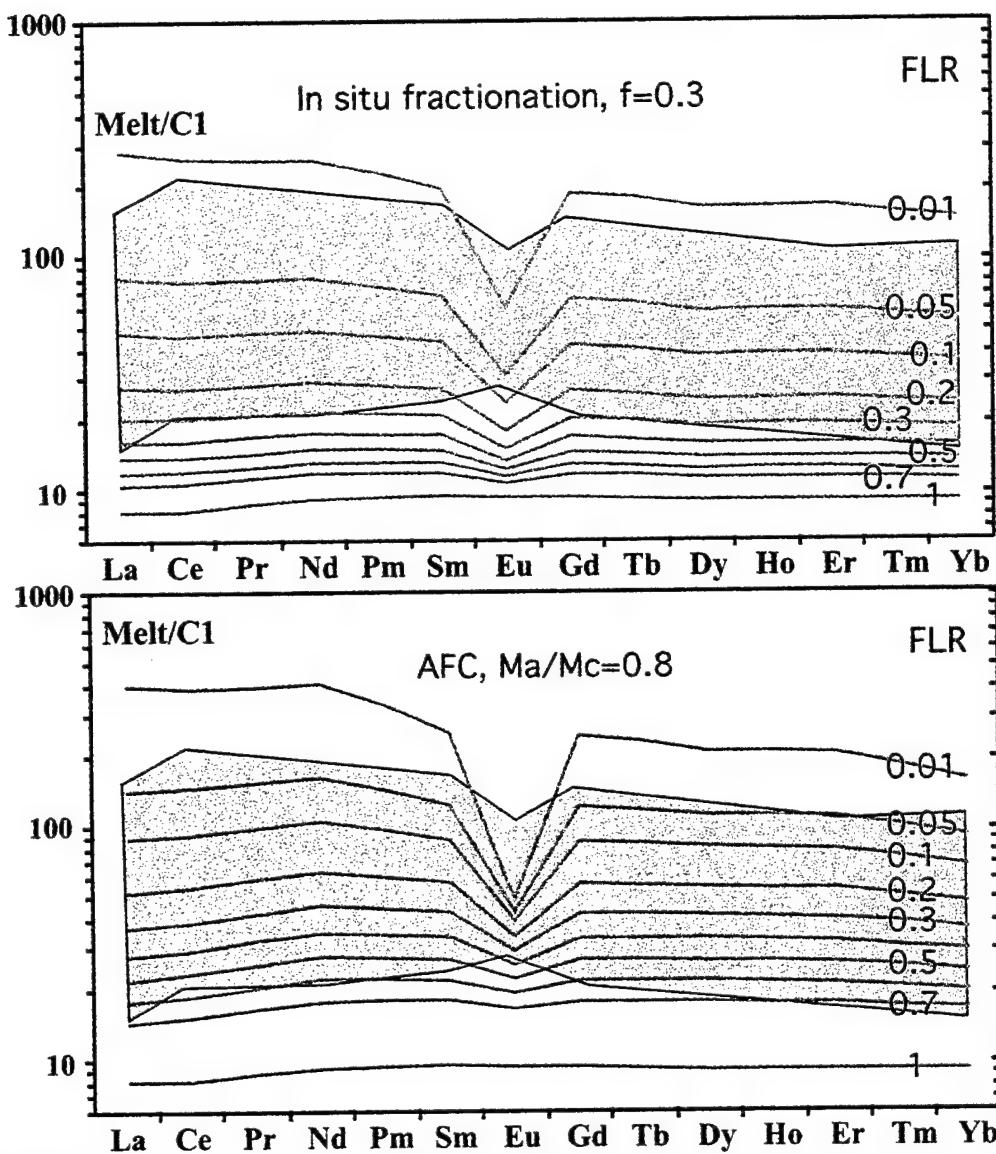


Figure 6-4: REE-element plots for crystallization models (normalized to the C1 chondrite of Taylor and MacLennan (1995)). FLR=Fraction of liquid remaining, the gray lines describe degree of fractionation. FLR(0.05) = 95% crystallized relative to the original melt. The gray field outlines the theoretical melt compositions based on the composition of augite in the gabbros (note that the diabase-data are excluded). A: In-situ fractionation, where only 30% of the interstitial melt is returned from the mush-zone to the central magma. B: The fractionation and crystallization of the DMM-derived melt happens during bulk-assimilation of the average olivine gabbro from Hole 735B (Natland and Dick, 1992). The diagram demonstrates that even if the melt experienced assimilation of pre-existing olivine-gabbros, the gabbros represent melts that first fractionated 50% elsewhere, but then crystallized to near completion.

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HREE. This is not the first diabase augite analyzed with such refractory compositions. Dick and Johnson (1995) found similarly depleted diopside and augite minerals from dikes in Hole 504B near the East Pacific Rise. They suggested that the strongly depleted patterns were due crystal-growth phenomena and heterogeneous nucleation in the early stages of melt crystallization prior to intrusion of the dikes as explored by Shimizu (1983, 1990). In the case examined by Dick and Johnson (1995), however, their augite phenocrysts were strongly sector zoned with respect to trace-elements. This is not the case here. In addition, as stated by Shimizu (1990) the kinetic effects mostly affect the M1 site of clinopyroxene, where the high field-strength elements reside (e.g. Ti, V, Sr, Cr). The M2-site is the host-site for the REE, and is not affected by kinetics of crystallization in the same way. It is likely that the theoretical melt composition derived from the diabase augite-composition are similar to the actual melt composition that produced it, representing a very depleted melt that did not mix with the other mantle-derived melts on ascent. The theoretical melt composition determined from the augite-composition is not as depleted as the “ultra-depleted melt” from Sobolev and Shimizu (1993). However, in this thesis, we have explored the effects of melts interacting with ascending magmas from the viewpoint of the gabbros. We find extreme variations in the magmatic evolution of gabbros, and it is not surprising that dikes, which are melt conduits from the lower to the upper crust, show diverse compositions too.

It can be rightfully argued that Chapter 3 to 5 of this thesis have shown that gabbroic augites are not the ideal candidates for deduction of the evolution of the melts in lower ocean crust environments, as they appear to be out of equilibrium with the rest of the cumulate. This is also evident from Figure 6-3B. The whole-rock trace-element contents of the gabbros represent the cumulate-minerals together with interstitial residual melt-fractions, and may also not directly demonstrate the evolution of the system. Ideally, we would like to support this type of calculation with analyses of all the minerals in each studied sample. Coogan et al. (2000) did show that the melt that produced the plagioclase was different from the one that produced the augite. In addition, the trace-element composition of high-valent ions in primitive olivine (potentially experiencing slower solid-state diffusion than Fe and Mg) can provide information on the evolution of these gabbroic rocks. If detailed geochemical studies are performed in gabbroic samples,

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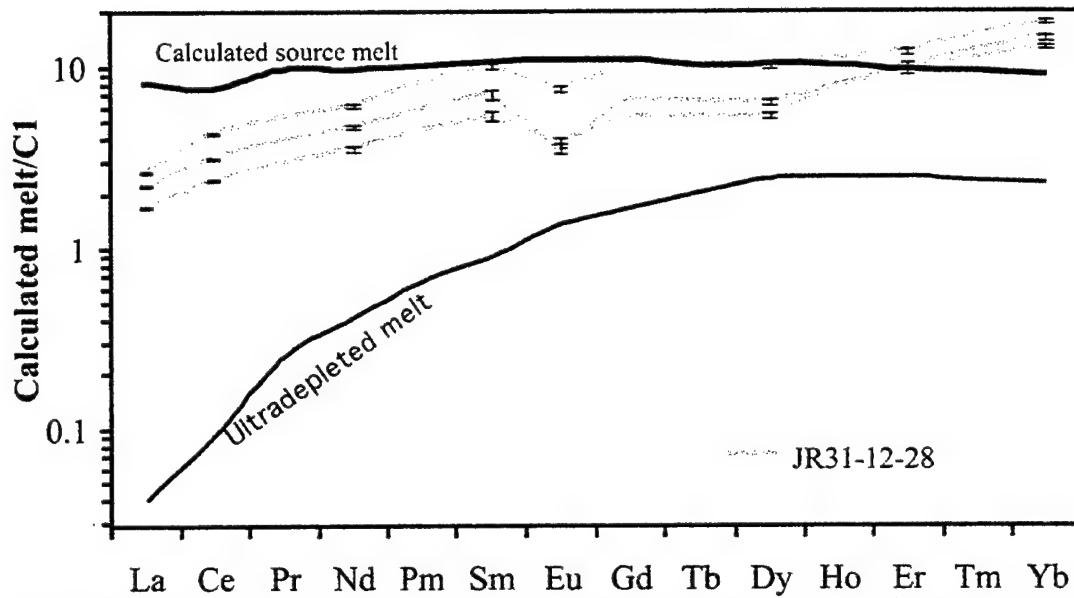


Figure 6-5: The augites-compositions from the diabase-dike represent melt-compositions much more depleted than the gabbroic rocks, and the assumed parental melt. However, the melt-compositions are not as refractive as the Ultra-depleted melt of Sobolev and Shimizu (1993).

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the entire accretion-history, including early precipitation of the initial gabbro, the assimilation of the lower crust into ascending magmas and subsequent crystallization of the hybrid magma might be deconvolved. However, it is clear that the cores and rims of the gabbro augites in this study do show complex evolution that must involve more than a simple fractional crystallization accretion model.

### 5 Conclusion

The trace element compositions of augites in Atlantis Bank gabbros show that they cannot have formed by fractional crystallization of olivine, plagioclase and clinopyroxene alone, as the rocks have increasing LREE/HREE ratios with fractionation. The gabbros may have crystallized by in-situ crystallization, or by significant assimilation of olivine gabbros. Even if the magmas have assimilated host-rocks upon ascent, they still must have crystallized >50% elsewhere before emplacement. The trace-element and REE-compositions show that the gabbros therefore represent melts that had crystallized > 50% -95% relative to a parental magma in equilibrium with the mantle. This agrees well with our previous results using major-elements in Chapter 3, and our estimate for the crustal thickness at Atlantis Bank is still valid.

One diabase-dike has augite-crystals representing a much more depleted melt-composition than the melts that produced gabbros. No known mineral-fractionation can produce the very depleted and refractory REE-pattern seen in this rock and we suggest that this dike represents depleted melts that did not mix with the melts that formed the gabbros sampled on and around Hole 735B.

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	Augite	Plagioclase	Olivine	Parental melt composition	Assimilated olivine gabbro
La	0.0536	0.124	0.003	2.97	1.06
Ce	0.0858	0.117	0.004	7.26	3.21
Sr	0.06	1.8	0.008	97	164
P	0.13	0.08	0.008	742	99
Nd	0.187	0.068	0.01	6.90	3.19
Sm	0.291	0.058	0.02	2.45	1.12
Zr	0.26	0.01	0.02	59.0	24.7
Ti	0.34	0.04	0.02	6523	2420
Eu	0.33	0.37	0.02	0.94	0.59
Gd	0.37	0.04	0.03	3.32	1.66
Tb	0.40	0.03	0.035	0.59	0.30
Dy	0.38	0.03	0.034	3.90	2.02
Y	0.41	0.03	0.033	24.1	11.63
Er	0.39	0.01	0.04	2.37	1.24
Yb	0.43	0.01	0.05	2.23	1.16
Cr	3.8	0.02	1.25	389	224
Ni	2	0.04	10	97	108
V	3.1	0.03	0.03	273	153
Sc	3.9	0	0.16	36.9	36.5

## Evolution of Oceanic Gabbros: In-situ and Ancient Examples

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## Chapter 7

### Residual melt porosities in gabbros from Atlantis Bank

#### Abstract

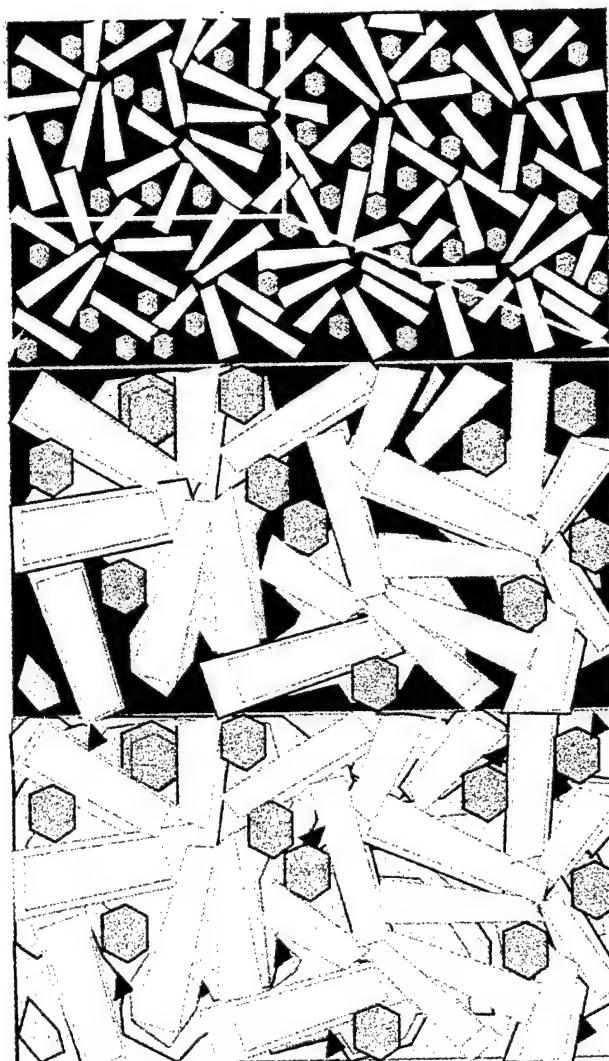
We have developed a method for calculation of trapped melt-fractions in gabbros from the Southwest Indian Ridge from whole-rock analyses. We use published data for basaltic rocks to define the liquid-line of descent for REE, Ti and P, and major-elements. The model is calibrated with the use of in-situ mineral analyses mass-balanced with whole-rock compositions. We find that the trapped melt-fraction of ODP Hole 735B is ~15%. Eu-anomalies are inversely related to the trapped melt-fraction. This suggests that the whole-rock positive Eu-anomalies in Hole 735B, the signature of cumulate plagioclase, are masked by other REE's.

#### 1 Introduction

Lower crustal magmatic rocks are thought to consist of cumulus minerals that precipitated from magma and accumulated along the roof, walls and floors of intrusions. Depending on the closeness of packing of the crystals, melts become trapped in the pore-spaces and may not be able to escape. This "trapped melt"-fraction, or residual melt porosity contributes to the whole-rock composition, possibly modifying pre-existing cumulate minerals, and crystallize interstitially to form the inter-cumulus minerals. Methods to identify and calculate the residual melt-fractions are complicated, poorly constrained, and frequently criticized (Natland et al, 1991; Coogan et al., 2001; Bédard, 1994).

We have directly estimated the amount of trapped melt gabbroic rocks using major-and trace-element mineral and whole-rock compositions of the gabbros and basaltic glasses from the area. We have then used our results to develop a more accurate and internally consistent method for calculating trapped melts at Atlantis Bank simply from whole-rock compositions, using the first method to calibrate it.

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$T_0$ : Formation of plagioclase/olivine network. Melt is in equilibrium with the crystals; melt composition is on the liquid-line of descent of MORB. Large porosity and permeability

$T_1$ : Close-up. This is where melt becomes trapped. Augite is forming between the growing plagioclase-laths. Melt composition is more evolved than typical MORB, but is in equilibrium with the rims of the crystals; There is low permeability, but large porosity.

$T_2$ : Complete solidification. New augite is formed, some by reaction with olivine. The new plagioclase overgrowths are more sodic, accessory-minerals are formed from the small, evolved melt-fraction, and the olivine is re-equilibrated.

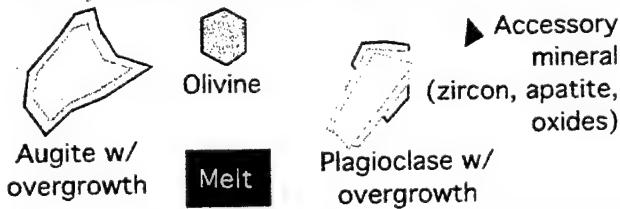


Figure 7-1: The formation of trapped melts. The figure is based on Wager et al. (1960).

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### 2 Results

Our method for estimating the amount of trapped melt in oceanic gabbros is based on the approach by Natland et al. (1991). The theoretical accretion of an olivine-gabbro is outlined in Figure 7-1. At  $T_0$ , plagioclase glomerocrysts form plagioclase-chains in a semi-rigid network. At this stage, the minerals are in equilibrium with the melt, and the network is permeable enough for the melts to flow through it and react with it. At  $T_1$ , permeability is reduced due to overgrowth on the cumulate minerals and increased viscosity of the melt, and the interstitial melt is effectively trapped. The melts in between the minerals at this stage are what is calculated as the “residual melt porosities”. This melt continues to form minerals in-situ, and will re-equilibrate with olivine. As nucleation of mineral-grains requires substantial under-cooling, the melt preferentially crystallizes on pre-existing grains. Therefore, the minerals develop evolved overgrowths, and the melt-pockets become smaller while the melt becomes more enriched in incompatible elements. At  $T_2$ , solidification is complete. The remnants of the small, fractionated melt pockets are observed as the location of the most evolved plagioclase and augite, together with accessory minerals like zircon, magnetite, ilmenite and apatite. These accessory minerals contain most of the magma-budget for elements like Ti, P and REE's. Of the cumulus minerals (olivine, plagioclase and some of the augite), augite has the largest partition-coefficients for these elements. As it is difficult to discern how an augite grew without making very detailed maps of the elemental distribution, we have tried to estimate the composition at the time of melt-entrapment by analyzing both core and rim compositions for augite to get the potential range of the residual melt porosity. Ideally, we would like to have the trace-element compositions of plagioclase also, in order to estimate the internally consistent composition of the melt that fractionated the cumulates.

Natland et al. (1991) used glass-compositions from the Southwest Indian Ridge to estimate liquid-lines of descent for each element ( $\text{TiO}_2$ ,  $\text{P}_2\text{O}_5$  and Zr) with respect to  $\text{MgO}$ . They related the bulk Mg# ( $\text{Mg}/(\text{Mg}+\text{Fe}^{2+})$ ) of a given gabbro to the melt that crystallized it.  $\text{TiO}_2$  in the rock vs. the theoretical amount in the liquid was used to get the fractional dilution of augite by plagioclase and olivine, in order to calculate excess Zr in the rock and thereby the amount of trapped melt in that cumulate.

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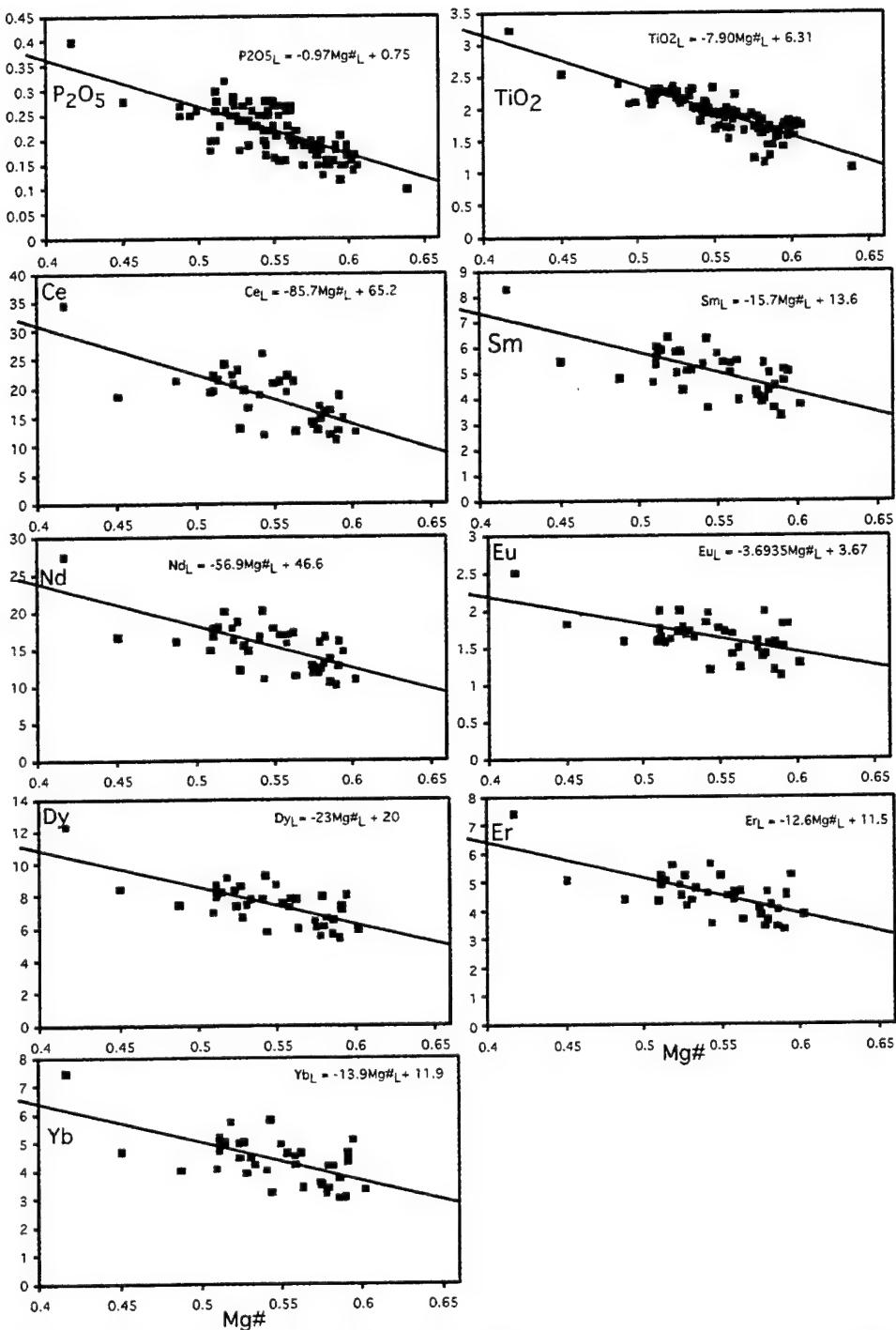


Figure 7-2: Liquid lines of descent for P<sub>2</sub>O<sub>5</sub>, TiO<sub>2</sub> and REE for basalts from Southwest Indian Ridge lavas from 25°-60°E (Mahoney et al., 1992; Le Roex et al., 1989; David et al., 2000). A least-squares fit is made for each element, and the equations are listed on each figure.

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Our model involves four steps:

- 1) It is necessary to find the composition of the interstitial melt at time of entrapment. We determined the liquid line of descent for the regional lavas with respect to  $P_2O_5$ ,  $TiO_2$  and REE (Fig. 7-2; data from PetDB), just like Natland (op. cit) did. Therefore, if we know the Mg# of the melt, we can estimate the incompatible element content of that melt. Note that the trace-element composition at a given Mg# has a large variation and *visa versa*, thus any errors in the Mg#-estimate for the melt may not matter.
- 2) We found the Mg# of the interstitial melt by using the composition of augite in the gabbro, assuming  $Kd_{augite/melt}^{Fe/Mg}$  of 0.23. With respect to the discussions of Chapter 3 and 5, a lower partition-coefficient will give us a higher Mg# of the melt, and will therefore give us slightly larger residual melt fractions. This partition-coefficient therefore provides minimum values. We have performed complete major-, and trace-element mineral composition analyses for all the samples in this study, while Coogan, et al. (2001) provide whole-rock compositions for three of these. We were therefore able to calibrate this well-constrained model to one for the whole-rock compositions. Thus, if the Mg# of the augite is not known, we find that a  $Kd_{rock/melt}^{Fe/Mg}=0.32$  can be used to calculate the Mg# of the melt. This partition-coefficient was found by comparing the Mg#'s of the whole-rocks to those of augite, and fitting a  $Kd_{rock/melt}^{Fe/Mg}$  to what would correspond to  $Kd_{augite/melt}^{Fe/Mg}=0.23$ . The whole-rock calculation will, of course, include the interstitial minerals, but again the range of REE along the liquid-lines of descent is large enough that a Mg# $\pm 5$  is accurate enough to give a satisfying model result.
- 3) We calculated the modes of the rocks where we had complete analyses by mass-balance of the major-element mineral-compositions to the whole-rock composition. We found that the whole-rock norm-algorithm equations from Grove et al. (1992) provide adequate modes for our gabbros.
- 4) The residual melt porosities can therefore be calculated for each element using:

$$\text{Residual porosity} = 100 * \frac{C_{Rock} - C_{Mineral} * Mode_{Mineral}}{C_{Melt}}$$

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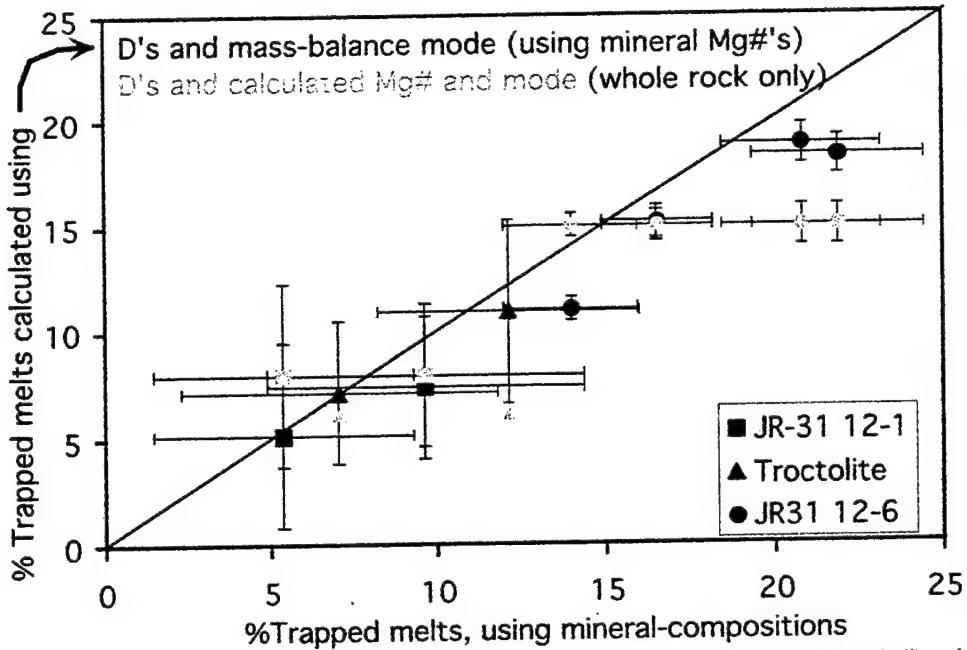


Figure 7-3: A comparison of the calculations for the residual melt porosity ("trapped melts") using the well-constrained method with mineral-compositions and the modes determined by mass-balance, to the whole-rock method (black symbols) and a method using the mass-balance mode and Mg#'s of the minerals and partition-coefficients with the LLD's to get the trace-element mineral-compositions (gray symbols). The mass-balance method gives negative residual-melt porosities for sample JR31 39-8. For each of the other samples (ODP-87-3, JR31 12-1, -6) the porosities are calculated for the cores and rims analyzed in each rock. For ODP-87-3, we used only Ti and P to find the residual melt porosity.

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$Mode_{\text{Mineral}}$  is the calculated mode of augite, and  $C_{\text{rock}}$ ,  $C_{\text{Mineral}}$  and  $C_{\text{Melt}}$  are the concentration in the whole-rock, minerals (olivine, plagioclase and augite), and melt, respectively. For our calculations, it was assumed that olivine and plagioclase contain insignificant amounts of these incompatible trace elements (except for Eu in plagioclase). If the concentration in the augite and plagioclase is not known, it was calculated from the theoretical melts using the partition-coefficients in Table 7-2, Chapter 6 and the Mg# of the whole-rock. The results from the calculations for the rocks from which we had complete mineral-analyses are shown as basalt regressions in Figure 7-2 and Table 7-1 and an example of the calculations is shown in Table 7-2.

### 3 Discussion

We tested our whole-rock procedure for calculating residual melt porosities (using partition-coefficients, calculated modes (Grove et al., 1992) and Mg#'s) by comparing it to the more exact method involving the actual mineral-compositions, and find that the results generally agree for all the trace-elements (Fig. 7-3). The  $1\sigma$  for the average residual melt-porosities using 7 different REE's is generally less than 5%. Unfortunately, our technique is not valid for the evolved oxide gabbros (e.g. JR31 39-8) as the oxides change the Mg# of the whole-rock, and increases the  $\text{TiO}_2$  content dramatically. The technique seems valid for the more primitive olivine-gabbros, however, and the results for the cores and rims using the mineral-composition method bracket the results from the whole-rock method. Our results are in agreement with the technique developed by Bédard (1994).

We have applied our technique to the large whole-rock data-sets from Atlantis Bank. We have, obviously, excluded ilmenite and apatite cumulates, and the felsic veins from these calculations. The results are presented in Figure 7-4. We find that for Hole 735B (excluding the most evolved gabbros) the average trapped melt-fraction is 15%, and the average standard deviation of 7 REE's for each measurement is 3%. We also find an exponential inverse correlation between the amount of trapped melts and Eu-anomalies ( $\text{Eu\#} = (\text{Eu}^* - ([\text{Sm}^* + \text{Gd}^*] \times 0.5) - 1)$ ; Pedersen et al. (1996)), where  $\text{Eu}^*$ ,  $\text{Sm}^*$  and  $\text{Gd}^*$  represent chondrite-normalized values (Taylor and MacLennan, 1985). Positive

**Table 7-1: Residual melt fractions calculated by whole-rock compositions, and by the well-constrained method involving the major and trace-element concentrations of the minerals and whole-rock**

	Well-constrained	JR31 12- method	JR31 12- method	JR31 39- method	ODP-83 ODP-83 7 core 7 rim	JR31 core	JR31 rim 1	JR31 core 1	JR31 12- 6 grain 1	JR31 rim 2	JR31 core 2	JR31 rim	JR31 core
<b>Whole-rock Composition:</b>													
P <sub>2</sub> O <sub>5</sub>	0.02	0.03	0.02	0.02	P <sub>2</sub> O <sub>5</sub>	1590	4282	3207	1876	2593	2424	2212	3056
TiO <sub>2</sub>	0.24	0.54	0.69	0.69	Ti (ppm)	3.36	3.02	3.29	1.65	1.06	1.48	1.25	1.88
Ce	1.74	1.74	1.74	1.74	Ce	1.10	1.10	1.10	1.10	1.10	1.10	1.10	1.09
Nd	1.69	3.26	3.43	3.43	Nd	1.53	4.54	2.86	1.67	2.32	2.21	2.30	1.58
Sm	0.64	1.24	1.44	1.44	Sm	1.09	2.70	1.86	1.26	1.59	1.34	1.53	0.96
Eu	0.43	0.72	0.85	0.85	Eu	0.41	0.55	0.51	0.41	0.61	0.46	0.64	0.45
Dy	1.10	2.24	2.68	2.68	Dy	1.90	4.53	3.41	1.90	3.07	2.33	2.77	1.81
Er	0.71	1.38	1.69	1.69	Er	1.14	2.58	1.99	1.13	1.53	1.35	1.70	1.03
Yb	0.65	1.25	1.56	1.56	Yb	1.24	2.65	1.98	1.21	1.62	1.24	1.62	0.93
Mg#	79.5	75.6	64.2	Mg#, melt	Mg#, melt	89.2	86.9	84.0	86.0	79.0	85.3	82.0	83.0
Mg#, melt	55.4	49.8	36.5	Mg#, melt	Mg#, melt	65.5	60.4	54.7	59.0	46.4	57.2	51.2	52.9
Mode augite	0.16	0.22	0.19	Mode augite	Mode augite	0.11	0.11	0.24	0.24	0.29	0.29	0.29	0.29
<b>Theoretical melt composition:</b>													
P <sub>2</sub> O <sub>5</sub>	0.21	0.40	0.27	P <sub>2</sub> O <sub>5</sub>	P <sub>2</sub> O <sub>5</sub>	0.11	0.16	0.22	0.18	0.30	0.20	0.25	0.24
TiO <sub>2</sub>	1.9	3.4	2.4	TiO <sub>2</sub>	TiO <sub>2</sub>	1.1	1.5	2.0	1.6	2.6	1.8	2.3	2.1
Ce	18	34	23	Ce	Ce	9	13	18	15	25	16	21	20
Nd	4.9	7.9	5.8	Nd	Nd	3.3	4.1	5.0	4.3	6.3	4.6	5.6	5.3
Sm	15	26	18	Sm	Sm	9	12	15	13	20	14	17	17
Eu	1.6	2.3	1.8	Eu	Eu	1.2	1.4	1.6	1.5	2.0	1.6	1.8	1.7
Dy	7.4	11.6	8.6	Dy	Dy	5.6	6.7	8.0	7.0	9.8	7.4	8.7	8.4
Er	4.5	6.8	5.2	Er	Er	3.2	3.9	4.6	4.1	5.7	4.3	5.1	4.8
Yb	4.3	6.8	5.0	Yb	Yb	2.8	3.5	4.3	3.7	5.5	4.0	4.8	4.5
<b>Residual trapped melts, on the basis of:</b>													
P <sub>2</sub> O <sub>5</sub>	8	4	10	P <sub>2</sub> O <sub>5</sub>	P <sub>2</sub> O <sub>5</sub>	9	4	6	10	16	24	19	18
TiO <sub>2</sub>	7	13	15	TiO <sub>2</sub>	TiO <sub>2</sub>	7	10	12	12	18	14	18	14
Ce	5	6	10	Ce	Ce	4	8	12	12	18	15	20	-5
Nd	8	10	14	Nd	Nd	7	10	13	13	19	16	19	-1
Sm	9	13	15	Sm	Sm	7	10	17	17	24	18	24	-2
Eu	3	11	11	Eu	Eu	7	10	17	17	24	18	24	8
Dy	9	15	17	Dy	Dy	4	9	14	14	21	17	23	-4
Er	9	17	17	Er	Er	5	11	17	23	18	25	-1	-1
Yb	9	15	16	Yb	Yb	4	10	14	22	17	24	-3	-5
Average	7	12	14	Average	Average	9	4	5.4	9.7	14.2	21.1	16.8	22.2
<b>10</b>	2	4	3	<b>10</b>	<b>10</b>	1.5	0.9	2.0	2.4	1.7	2.6	6.1	6.9
<b>% Residual trapped melts, on the basis of:</b>													

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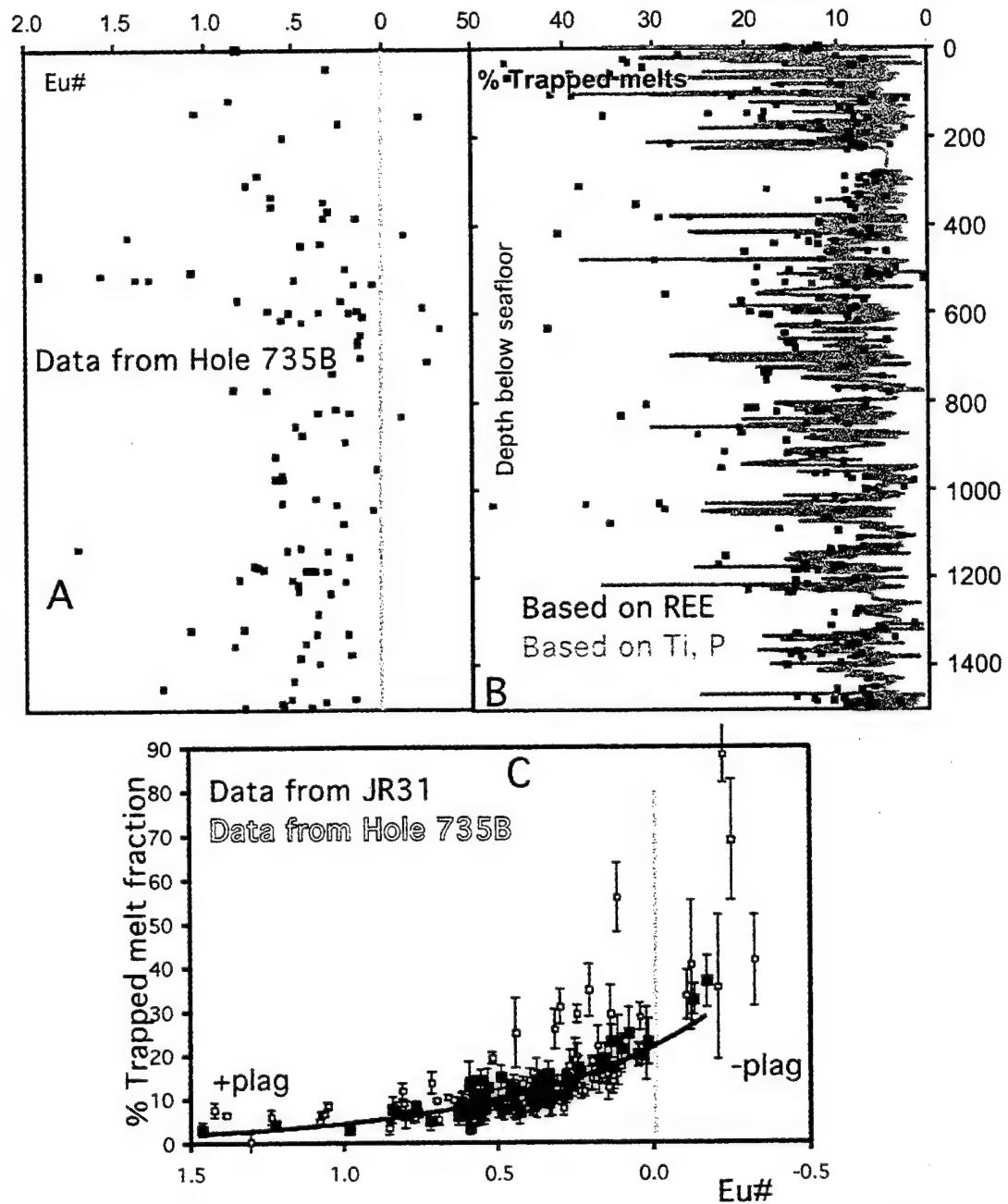


Figure 7-4: Trapped melt fractions from ODP Hole 735B. A: Eu# anomalies for Hole 735B gabbros based on whole rock analyses (Natland and Dick, 2002). B: Residual-melt porosities calculated from whole-rock composition for Hole 735B gabbros based on the average results for REE, and for Ti and P. C: The correlation between Eu# and the residual melt porosity in Hole 735B and in the gabbros around Atlantis Bank.

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and negative Eu-anomalies are thought to indicate accumulated or removed plagioclase in a rock, respectively. In Hole 735B, however, it seems that the Eu#'s are more a function of the extent of trapped melt than they are of plagioclase accumulation, and can therefore act as a proxy for residual melt porosity.

### 4 Conclusion

We have developed a technique for calculating residual melt porosity for oceanic gabbros at the Southwest Indian Ridge from whole-rock analyses of major-elements with rare-earth elements, Ti and P. The model is calibrated with the use of in-situ mineral analyses. We find that the trapped melt-fraction of Hole 735B is ~15%. The Eu-anomalies are inversely related to the trapped melt-fraction, possibly indicating that the Eu# in Hole 735B is masked by the abundance of other REE's in the whole-rock compositions, as is often seen in diabase-dikes.

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Table 7-2: An example of the residual melt fraction calculation for JR31-12-1 calculated for Dy.

<b>Method of calculation when the mineral-compositions are available</b>	
<b>Calculation</b>	<b>Result</b>
Determine the Mg# of the melt on the basis of augite compositions: $Mg\#_{melt} = \frac{Mg\#_{augite} * 0.23}{1 - Mg\#_{augite} + Mg\#_{augite} * 0.23}$	For the augite: Mg# core 0.86 => Mg# melt = 0.59 Mg# rim 0.84 => Mg# melt = 0.55
Determine the Dy-content for the melt on the basis of the Mg# $Dy_{melt} = -23 * Mg\#_{melt} + 20$	Dy in the melt (core) = 7 ppm Dy in the melt (rim) = 8 ppm
Need to know: Augite norm based on mass-balance = 0.24 Measured Dy in core 1.9 ppm Measured Dy in rim 3.4 ppm Whole-rock content of Dy = 1.1 ppm	
Residual porosity = $100 * \frac{C_{Rock} - C_{Mineral} * Mode_{Mineral}}{C_{Melt}}$	Residual porosity for the Core = $100 * (1.1 - (1.9 * 0.24)) / 7 = 9.2\%$ Rim = $100 * (1.1 - (3.4 * 0.24)) / 8 = 4.1\%$
<b>Method of calculation when only the whole-rock compositions are available</b>	
$Mg\#_{melt} = \frac{Mg\#_{rock} * 0.32}{1 - Mg\#_{rock} + Mg\#_{rock} * 0.32}$	Mg# (rock) = 0.795, => Mg# melt = 0.55
From Grove et al. (1992):  Recalculate whole - rock composition from weight % oxides to mol %, and Al, Fe <sup>3+</sup> , K, Na and P as : $AlO_{1.5}, K_{0.5}O, Na_{0.5}O$ and $PO_{2.5}$ . Then combine molar MgO and FeO to FMO and $K_{0.5}O, Na_{0.5}O$ to Alk. $Sum = SiO_2 - CaO - 2 * Alk + Cr_2O_3 + TiO_2 + 2 * PO_{2.5} + FeO_{1.5}$ Quartz = $[SiO_2 - 0.5 * FMO - 1.5 * CaO - 0.25 * AlO_{1.5} - 2.75 * Alk + 0.5 * Cr_2O_3 + 0.5 * TiO_2 + 2.5 * PO_{2.5}] / Sum$ Plagioclase = $4 * [AlO_{1.5} + NaO_{0.5} - KO_{0.5}] / [2 * Sum]$ Olivine = $2 * [FMO + 0.5 * (AlO_{1.5} - Alk) - CaO - TiO_2 - Cr_2O_3 + 1.667 * PO_{2.5}] / [2 * Sum]$ Augite = $3 * [CaO - 0.5 * AlO_{1.5} + 0.5 * Alk - 1.667 * PO_{2.5}] / Sum$ IlmeniteHematiteChromite = $1.5 * [Cr_2O_3 + TiO_2 + FeO_{1.5}] / Sum$ Orthoclase = $4 * [KO_{0.5}] / Sum$ Apatite = $6 * PO_{2.5} / Sum$ Normalize all the minerals to 1 to get the mode of the rock.  Calculated mode of augite based on whole-rock = 0.16	
Theoretical Dy of melt (as calculated above): 7.36	Mineral Dy composition using partition-coefficient from Table 6-2: $Dy_{Augite} = 0.4 * Dy_{melt} = 3$
Residual trapped melt (as calculated above): $100 * (1.1 - (3 * 0.16)) / 7.36 = 8.2\%$	

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6. Abstract (Limit: 200 words)  This study is a geochemical investigation into the accretion of lower oceanic crust and processes of shallow melt-rock reaction at mid-ocean ridges. Geochemical signatures from the 480-My Lyngen Gabbro indicate the complex formed from hydrous supra-subduction zone magmas.  Minerals in gabbros from Atlantis Bank, Southwest Indian Ridge, an <i>in-situ</i> example of lower ocean-crust, formed by near-fractional crystallization at mid-crustal pressures, modeled by the MELTS algorithm. The lower crust is more evolved than the lavas, representing mantle-derived melts fractionated 50-95%. >770-m of additional primitive cumulates remains below the 1500-m deep Hole 735B or within the underlying mantle.  The Atlantis Bank gabbros minerals have internal disequilibrium. Melt-rock interaction may have caused this. Dissolution-experiments for plagioclase-olivine and plagioclase-augite pairs were performed at 1180°-1330°C and 20-min-24hrs. Dissolution occurs rapidly and out of equilibrium, the dissolution rates dependent on ΔT above the solidus. Coarse-grained rocks will heat internally when enclosed in hot magma, melt and disaggregate. The hybrid magma crystallizes more refractory-looking minerals than the melts that precipitated the original gabbroic rocks. Assimilation increases the Na content and decreases the Fe content of the melt that digests gabbro, thus hybrid basaltic glasses will falsely reflect a lower degree and pressure of mantle melting.				
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